

From Climate to Kolmogorov - Simulations Spanning Upper Ocean Scales

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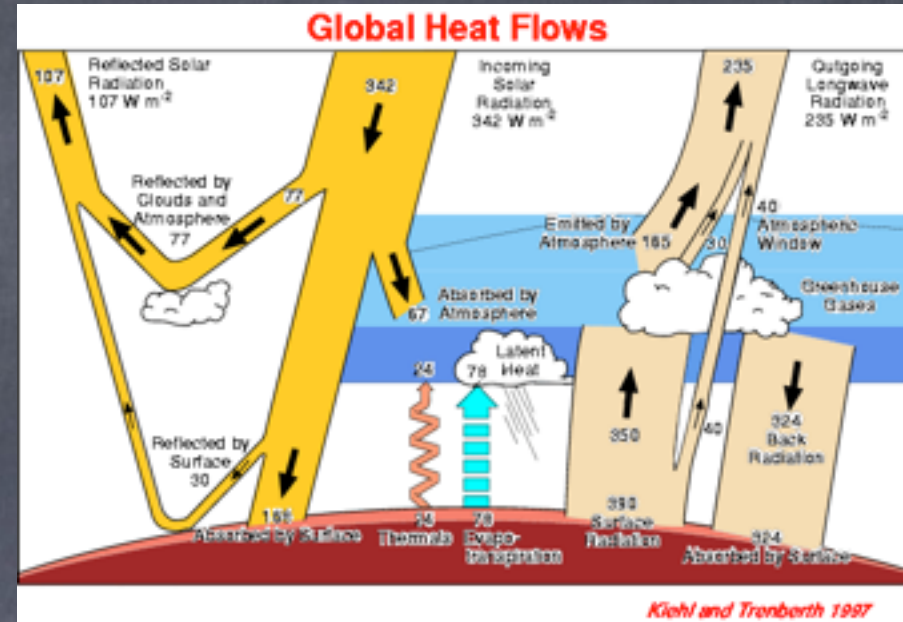
Keith Julien (CU-APPM),

Luke Van Roekel (Northland), Peter Sullivan (NCAR), Jim McWilliams (UCLA), Mark Hemer (CSIRO)

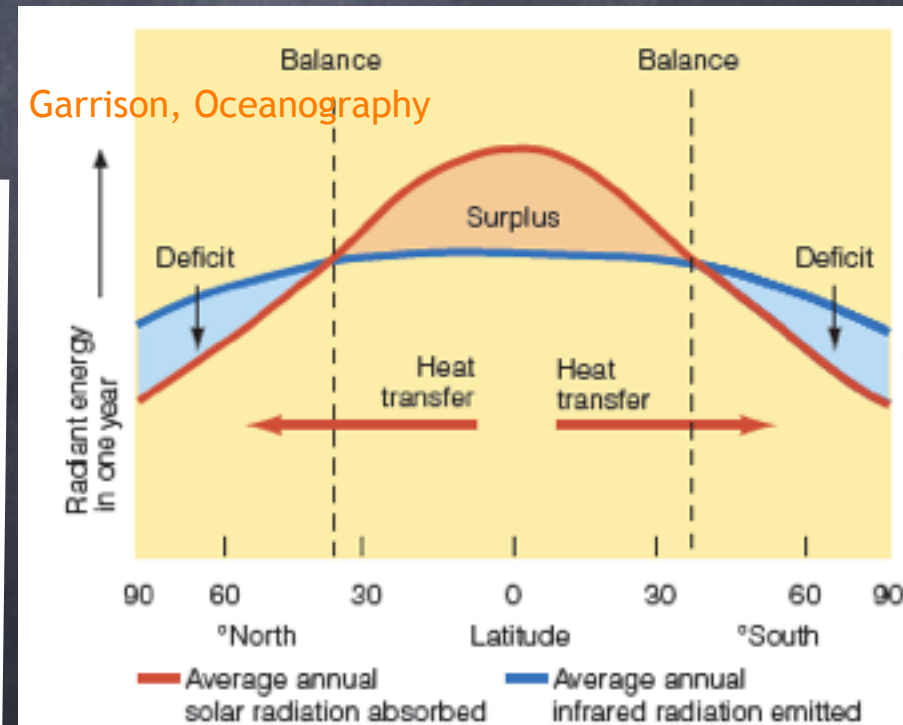
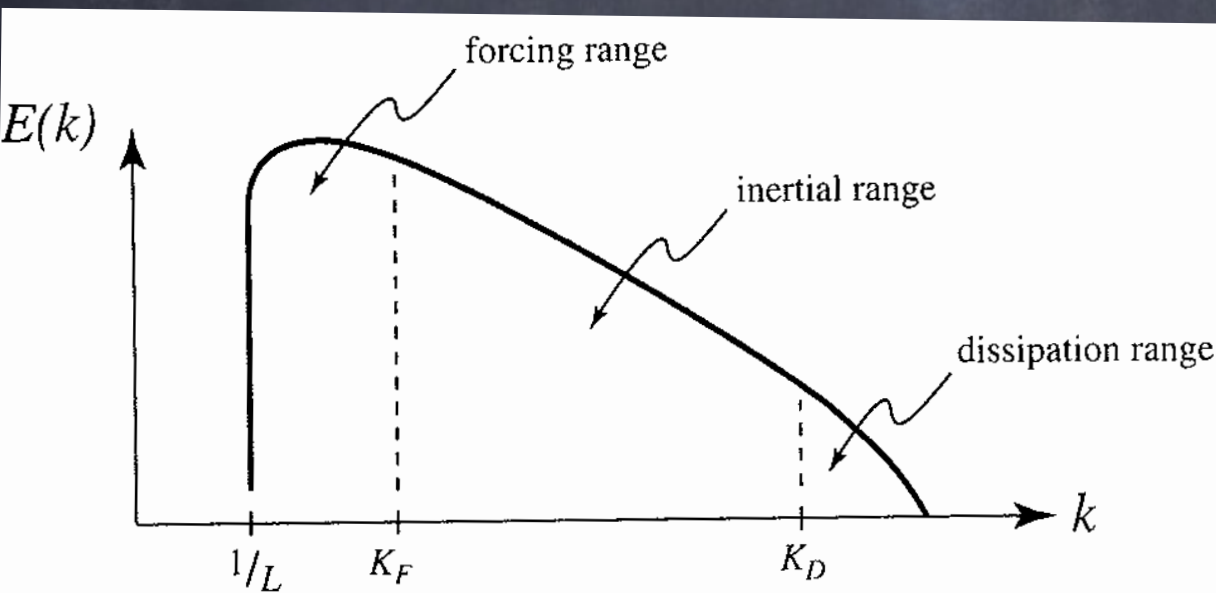
Frontiers in Computation

Sponsors: NSF 1245944, 0934737, 0825614, NASA NNX09AF38G

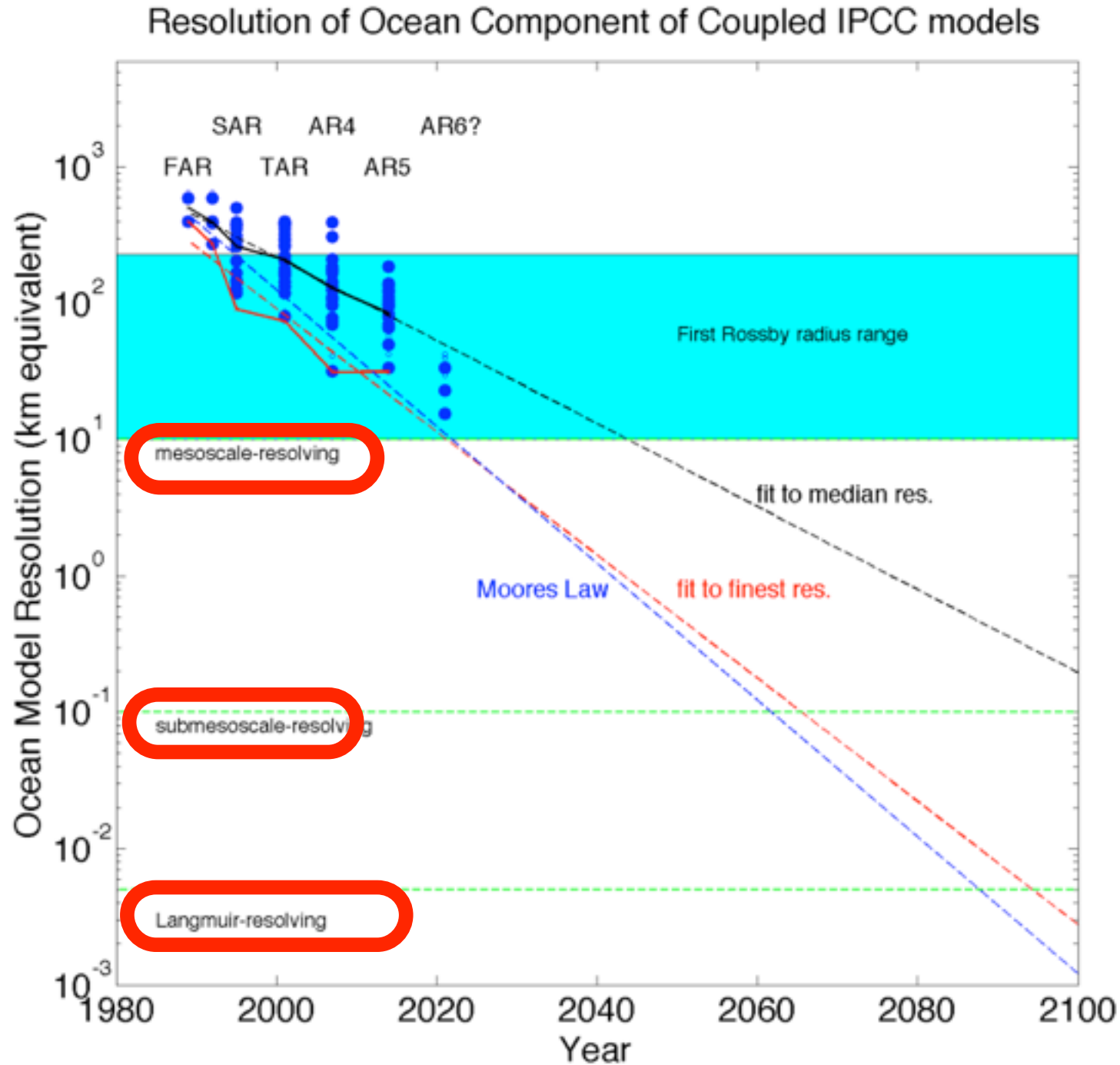
The Earth's Climate System is driven by the Sun's light (minus outgoing infrared) on a global scale



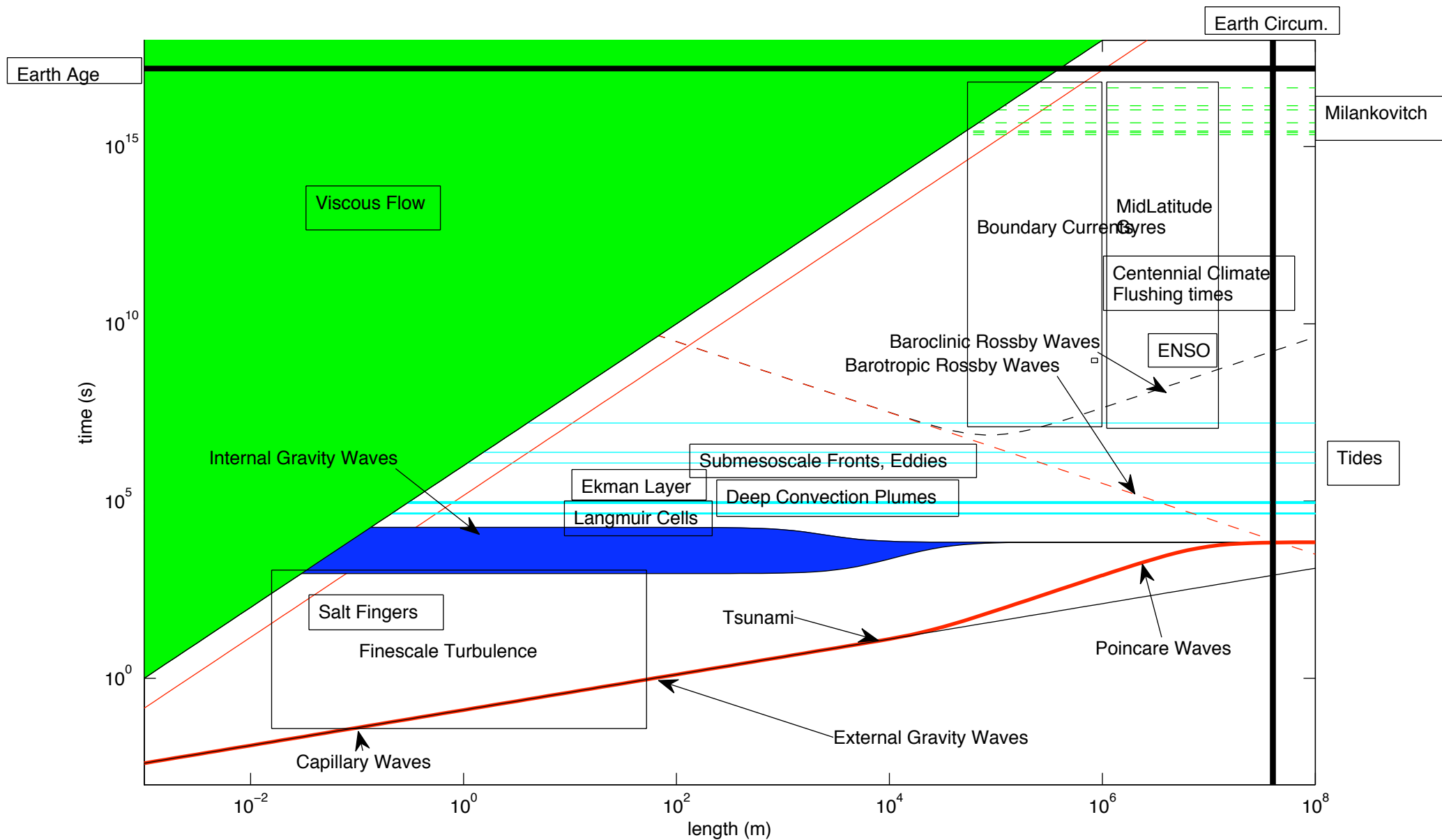
Dissipation concludes turbulence cascades on scales about a billion times smaller



Resolution will be an issue for centuries to come!

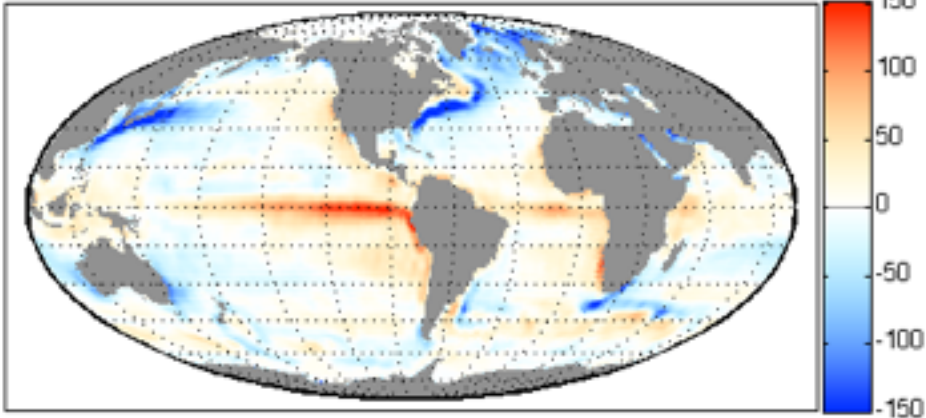


The Ocean is **Vast & Diverse**: just one spectral cascade?

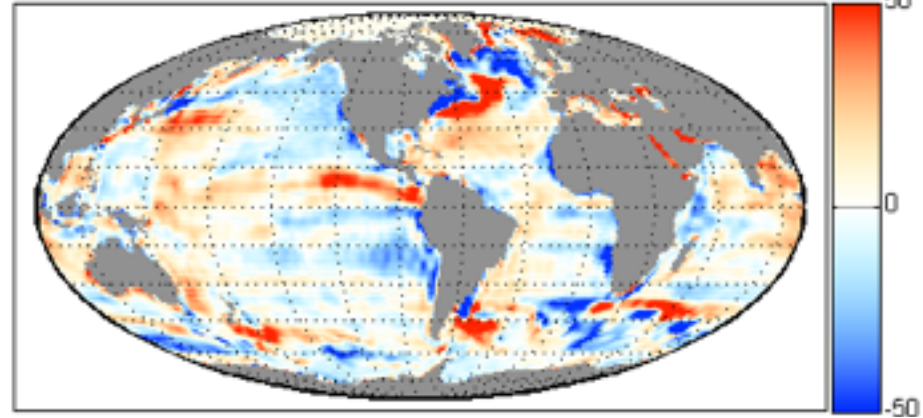


Air-Sea Flux Errors vs. Data (Large & Yeager 09)

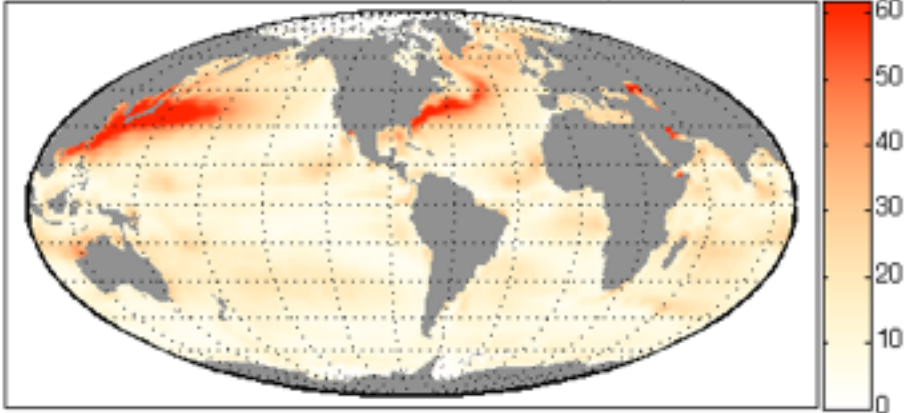
Mean of 1986-2005 CORE Q_{as} (W/m^2)



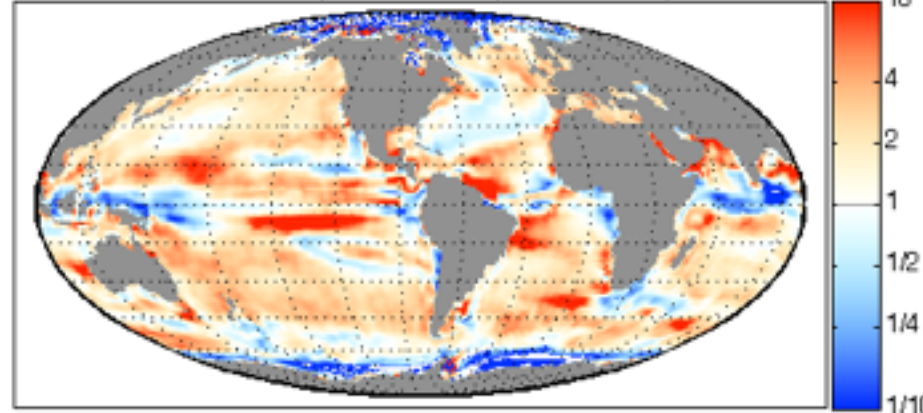
1986-2005 CCSM4-CORE Q_{as} bias, mean:1.5, rms:23 (W/m^2)



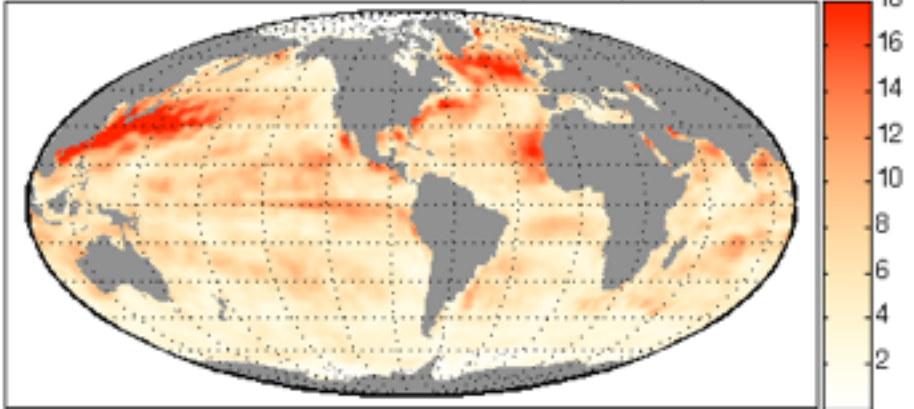
St. Dev. of CORE annual evaporation (W/m^2)



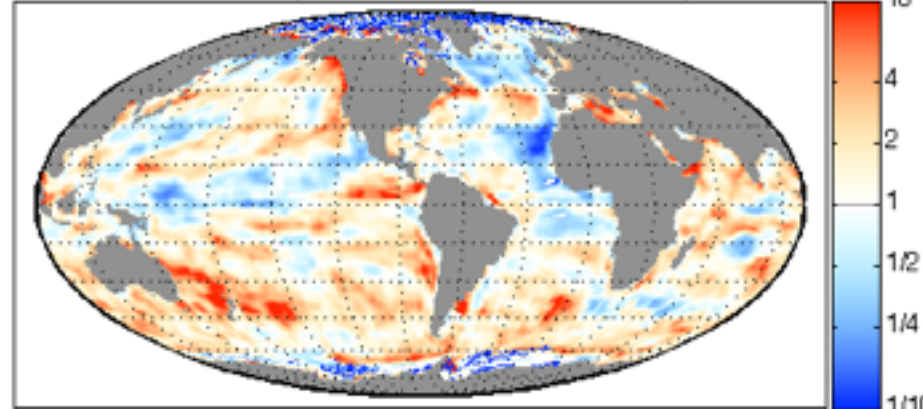
Variance ratio (CCSM4/CORE) of annual evaporation



St. Dev. of CORE interannual evaporation (W/m^2)

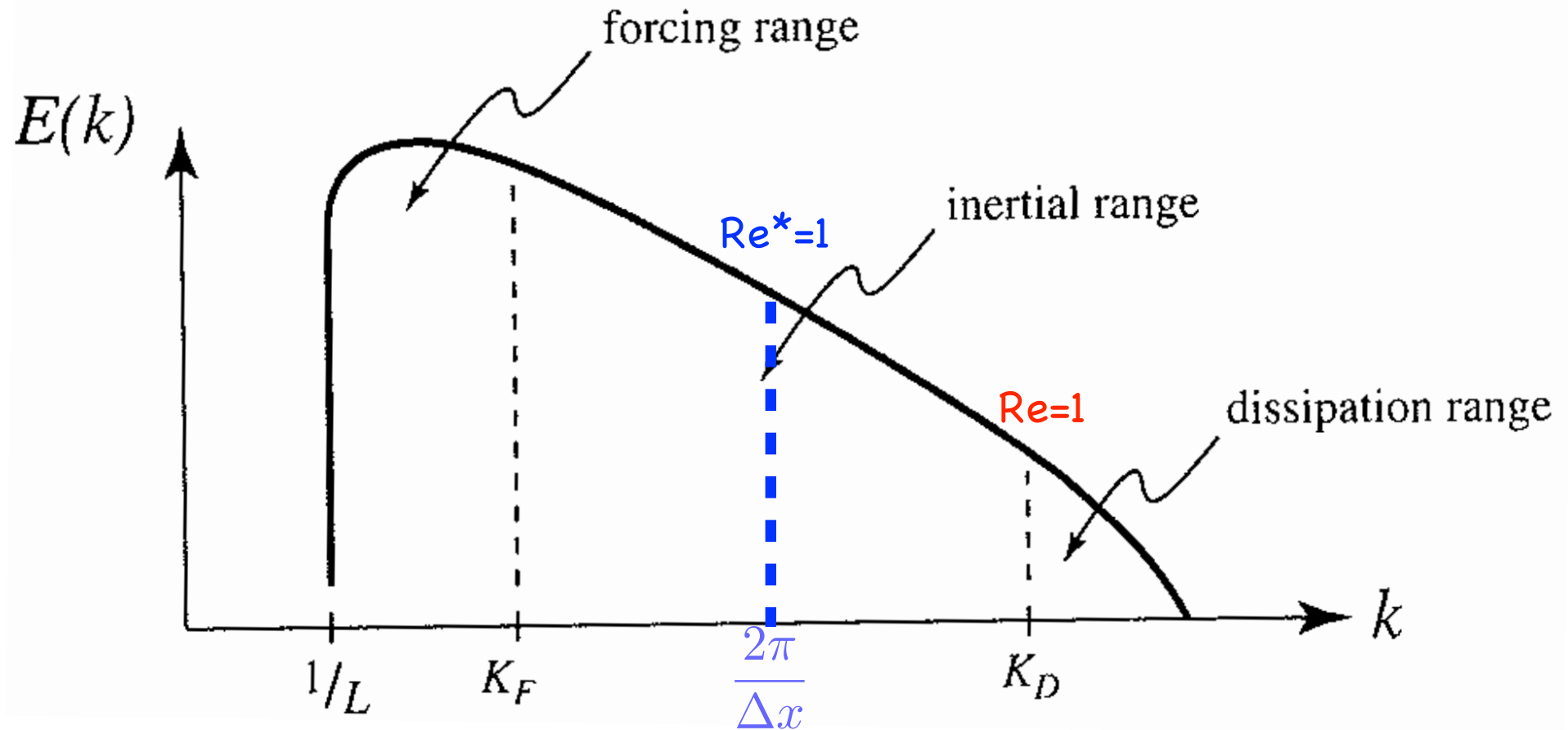


Variance ratio (CCSM4/CORE) of interannual evaporation



Mean
Annual
Interannual
9-15mo
2-7yr

Truncation of Cascades



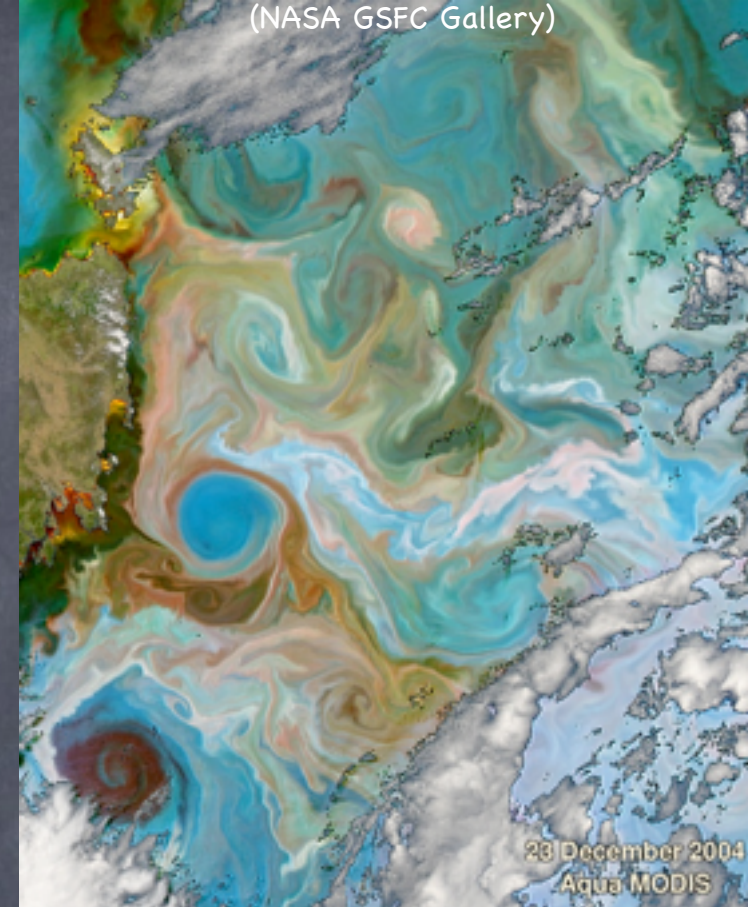
1963: Smagorinsky Devises Viscosity Scaling,
So that the Energy Flow is Preserved,

but order-1 gridscale Reynolds #: $Re^* = UL/\nu_*$

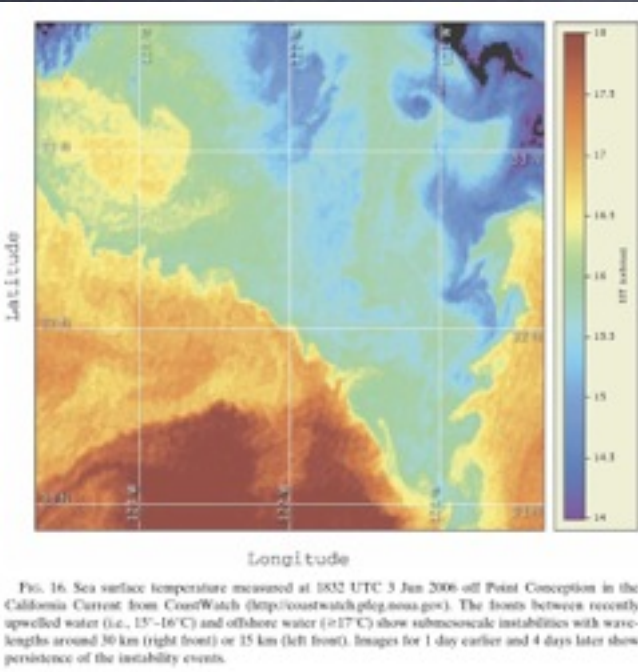
$$\nu_{*h} = \left(\frac{\Upsilon_h \Delta x}{\pi} \right)^2 \sqrt{\left(\frac{\partial u_*}{\partial x} - \frac{\partial v_*}{\partial y} \right)^2 + \left(\frac{\partial u_*}{\partial y} + \frac{\partial v_*}{\partial x} \right)^2}$$

The Character of the Mesoscale

←
100
km



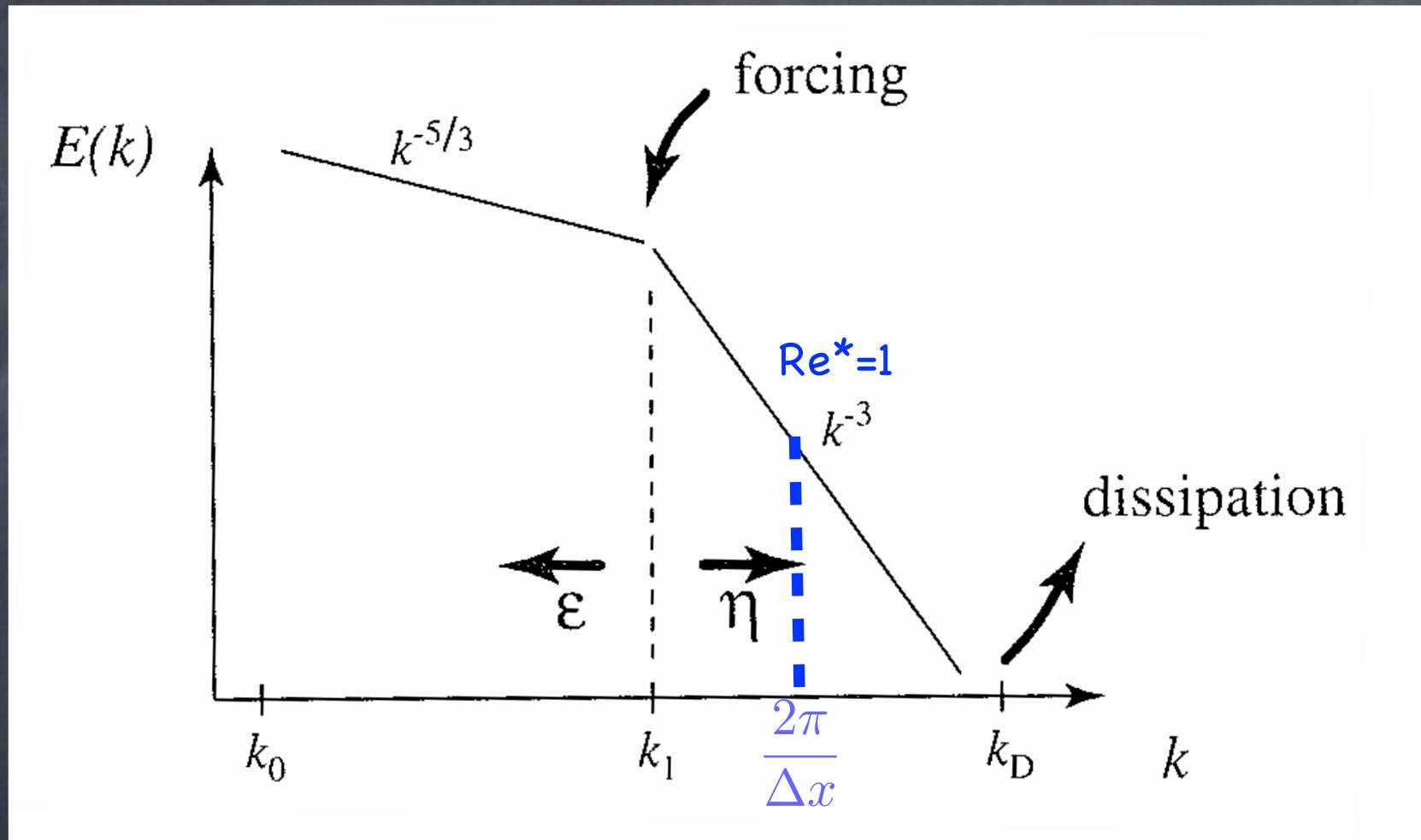
(Capet et al., 2008)



- Boundary Currents
- Eddies
- $Ro=O(0.1)$
- $Ri=O(1000)$
- Full Depth
- Eddies strain to produce Fronts
- 100km, months

Eddy processes mainly **baroclinic & barotropic instability**. Parameterizations of baroclinic instability (GM, Visbeck...).

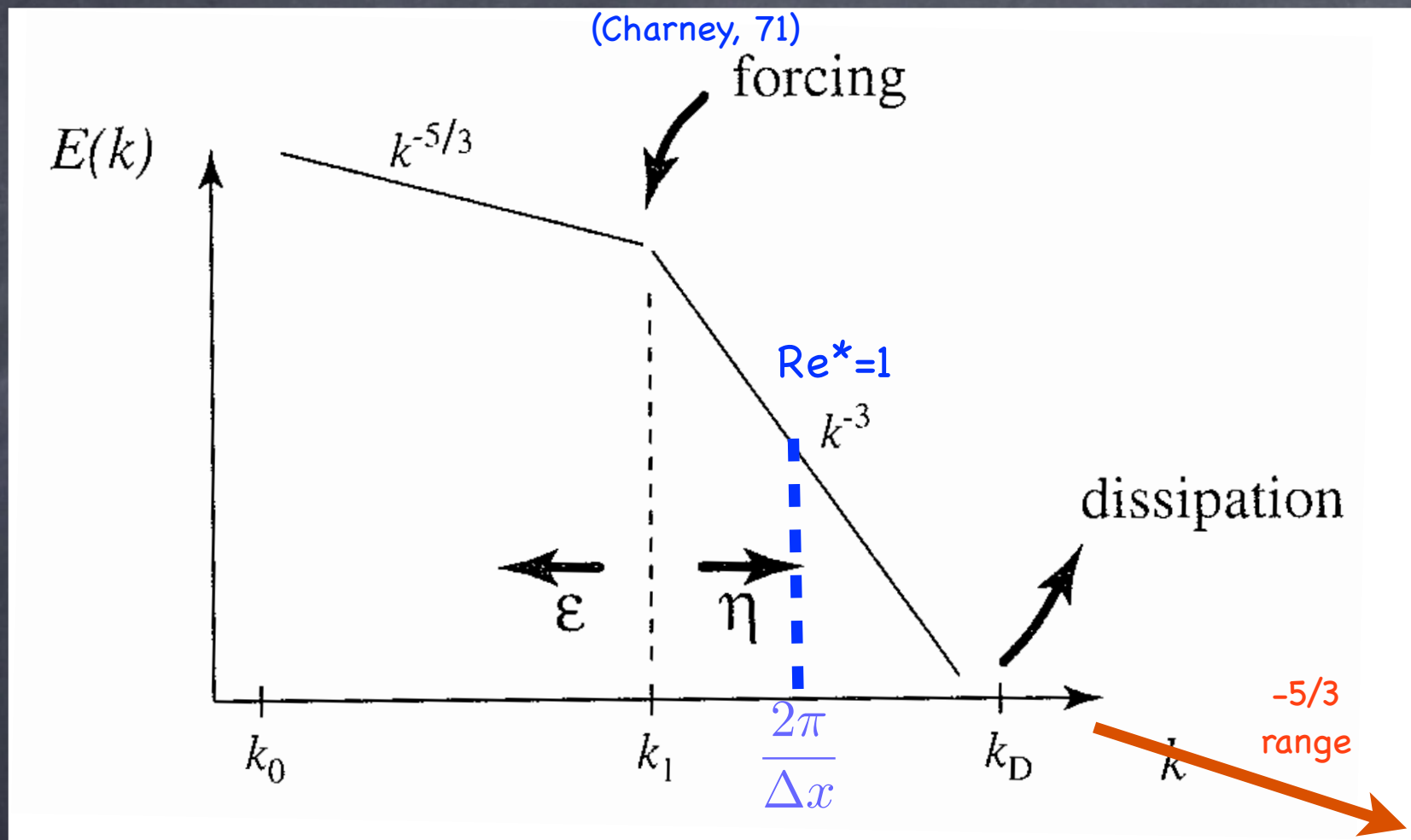
2d Turbulence Differs



1996: Leith Devises Viscosity Scaling,
So that the Enstrophy Flow is Preserved

$$\nu_* = \left(\frac{\Lambda \Delta x}{\pi} \right)^3 \left| \nabla_h \left(\frac{\partial u_*}{\partial y} - \frac{\partial v_*}{\partial x} \right) \right|.$$

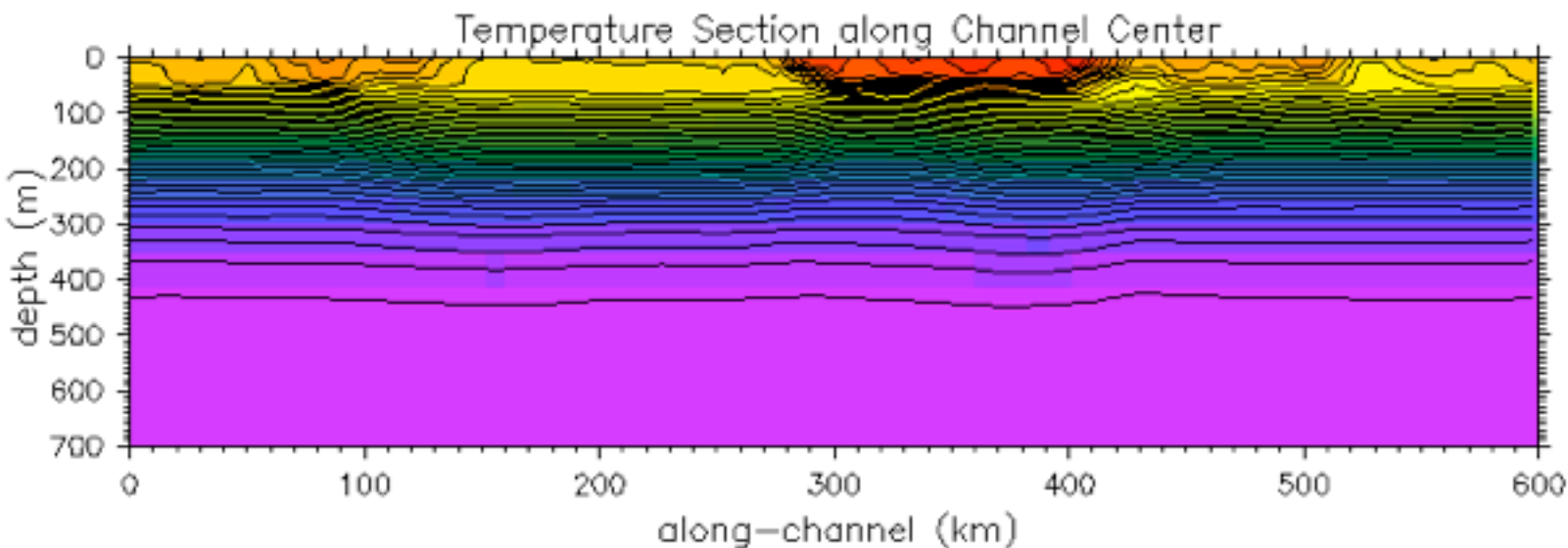
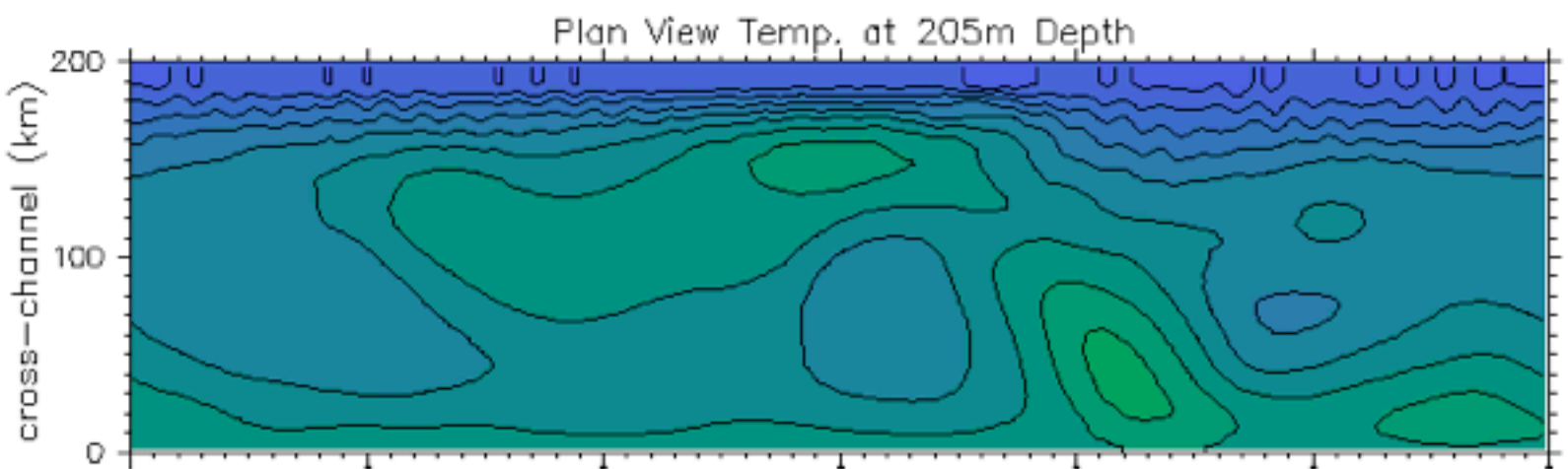
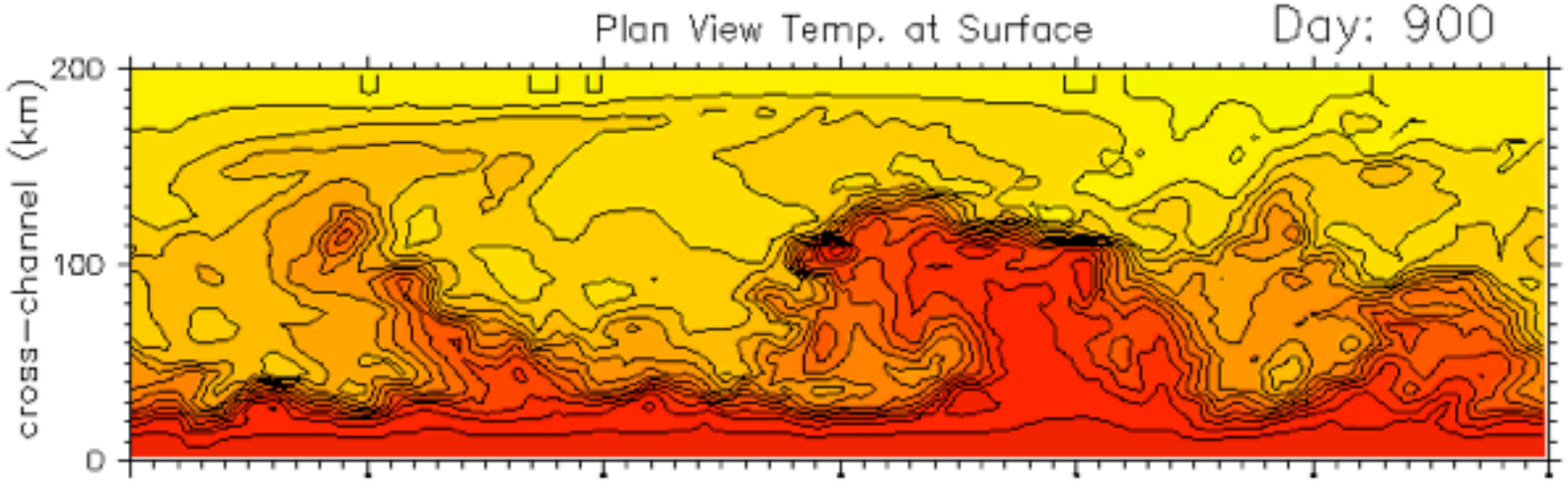
MOLES Turbulence Like Pot'l Enstrophy cascade, but divergent



F-K & Menemenlis Revise Leith Viscosity Scaling,
So that diverging, vorticity-free, modes are also damped

$$\nu_* = \left(\frac{\Delta x}{\pi}\right)^3 \sqrt{\Lambda^6 |\nabla_h q_{2d}|^2 + \Lambda_d^6 |\nabla_h (\nabla_h \cdot \mathbf{u}_*)|^2}$$

B. Fox-Kemper and D. Menemenlis. Can large eddy simulation techniques improve mesoscale-rich ocean models? In M. Hecht and H. Hasumi, editors, Ocean Modeling in an Eddy Regime, volume 177, pages 319-338. AGU Geophysical Monograph Series, 2008.



Big, Deep
(meso)

interact
with

Little,
Shallow
(submeso)

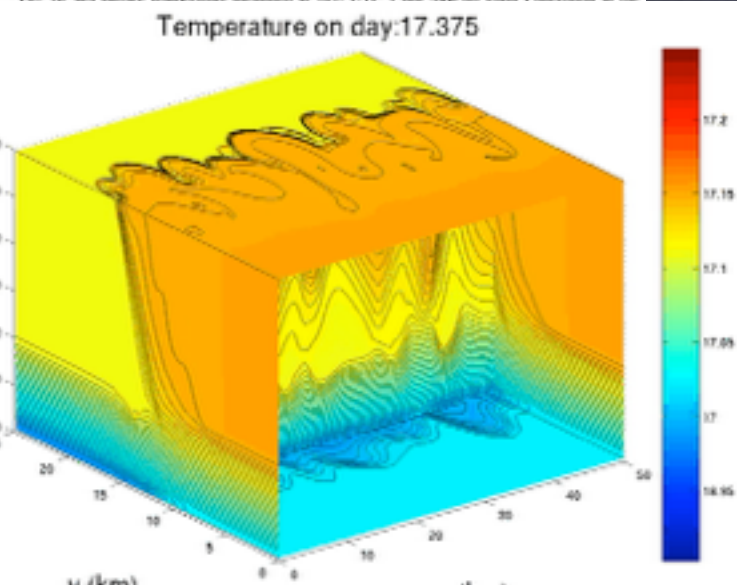
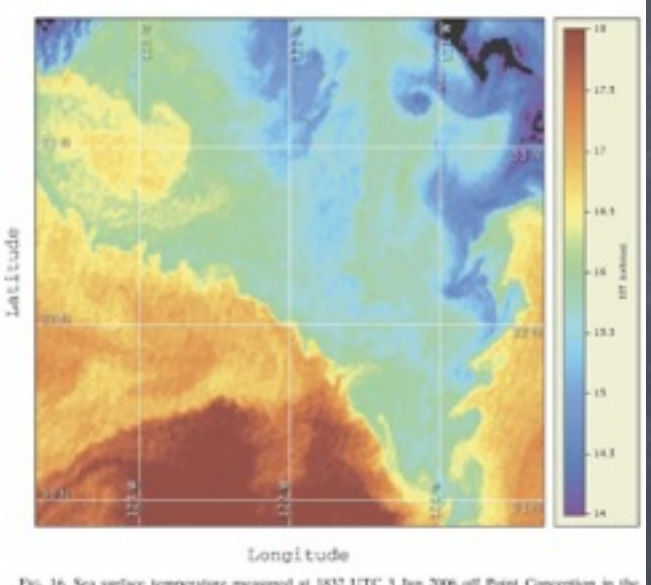
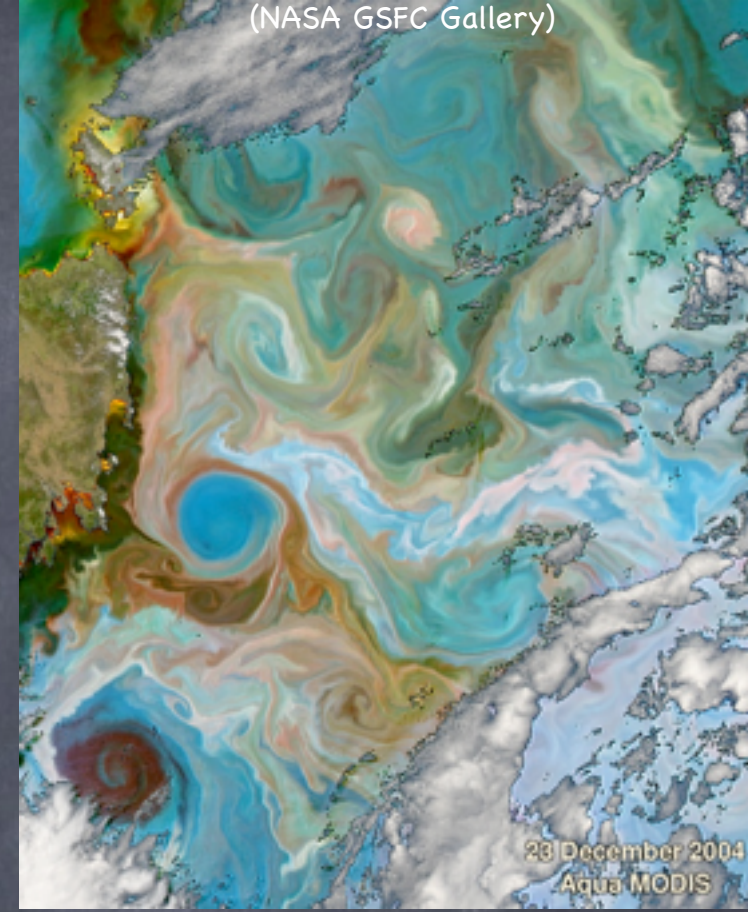
B. Fox-Kemper, R. Ferrari,
and R. W. Hallberg.
Parameterization of mixed
layer eddies. Part I: Theory
and diagnosis. *Journal of
Physical Oceanography*,
38(6):1145-1165, 2008.

10 km

The Character of the Submesoscale

(Capet et al., 2008)

- Fronts
- Eddies
- $Ro=O(1)$
- $Ri=O(1)$
- near-surface
- 1-10km, days



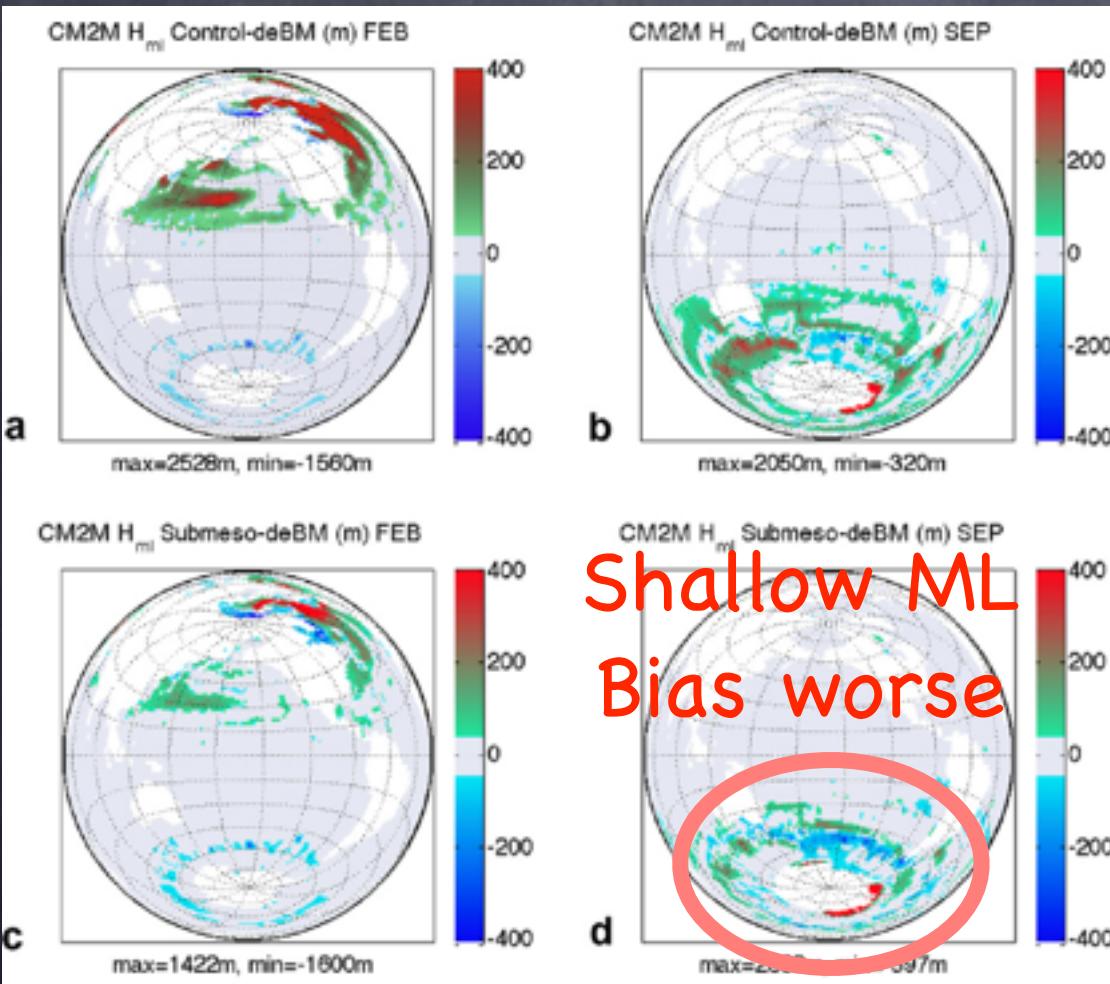
Eddy processes often **baroclinic instability**

Parameterizations of submesoscale baroclinic instability?

B. Fox-Kemper, R. Ferrari, and R. W. Hallberg. Parameterization of mixed layer eddies. Part I: Theory and diagnosis. Journal of Physical Oceanography,

S. Bachman and B. Fox-Kemper. Eddy parameterization challenge suite. I: Eady spindown. Ocean Modelling, 2013. In press.

Physical Sensitivity of Ocean Climate to MLE: Mixed Layer Eddy Restratification

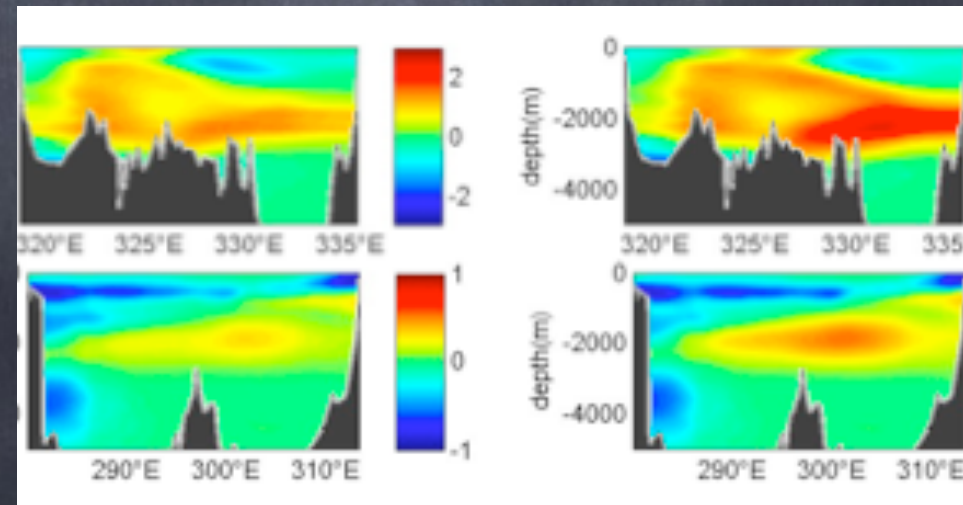


Bias
w/o
MLE

Improves CFCs
(water masses)

Bias with MLE

Bias w/o MLE



B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg, M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels.
Parameterization of mixed layer eddies. III: Implementation and impact in global ocean climate simulations. *Ocean Modelling*, 39:61-78, 2011.

The Character of

the Langmuir Scale

- Near-surface
- Langmuir Cells & Langmuir Turb.
- $Ro \gg 1$
- $Ri < 1$: Nonhydro
- 1-10m
- 10s to mins
- $w, u = O(10\text{cm/s})$
- Stokes drift
- Eqtns: Craik-Leibovich
- Params: McWilliams & Sullivan, 2000, etc.

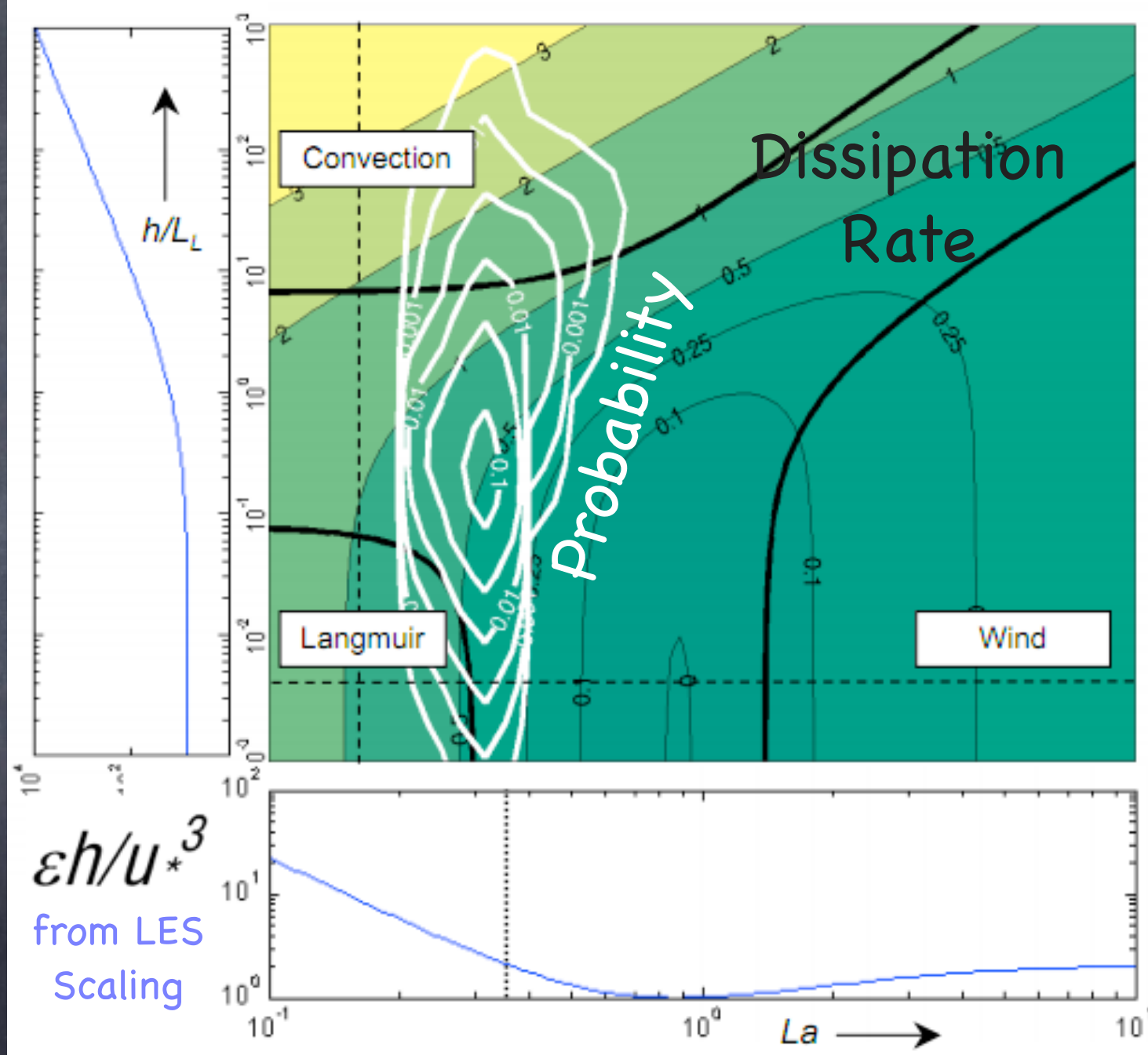
Image: NPR.org,
Deep Water
Horizon Spill

image:
Leibovich, 83



Figure 1a. Illustration of Langmuir circulations showing notation used in this review and surface and subsurface motions.

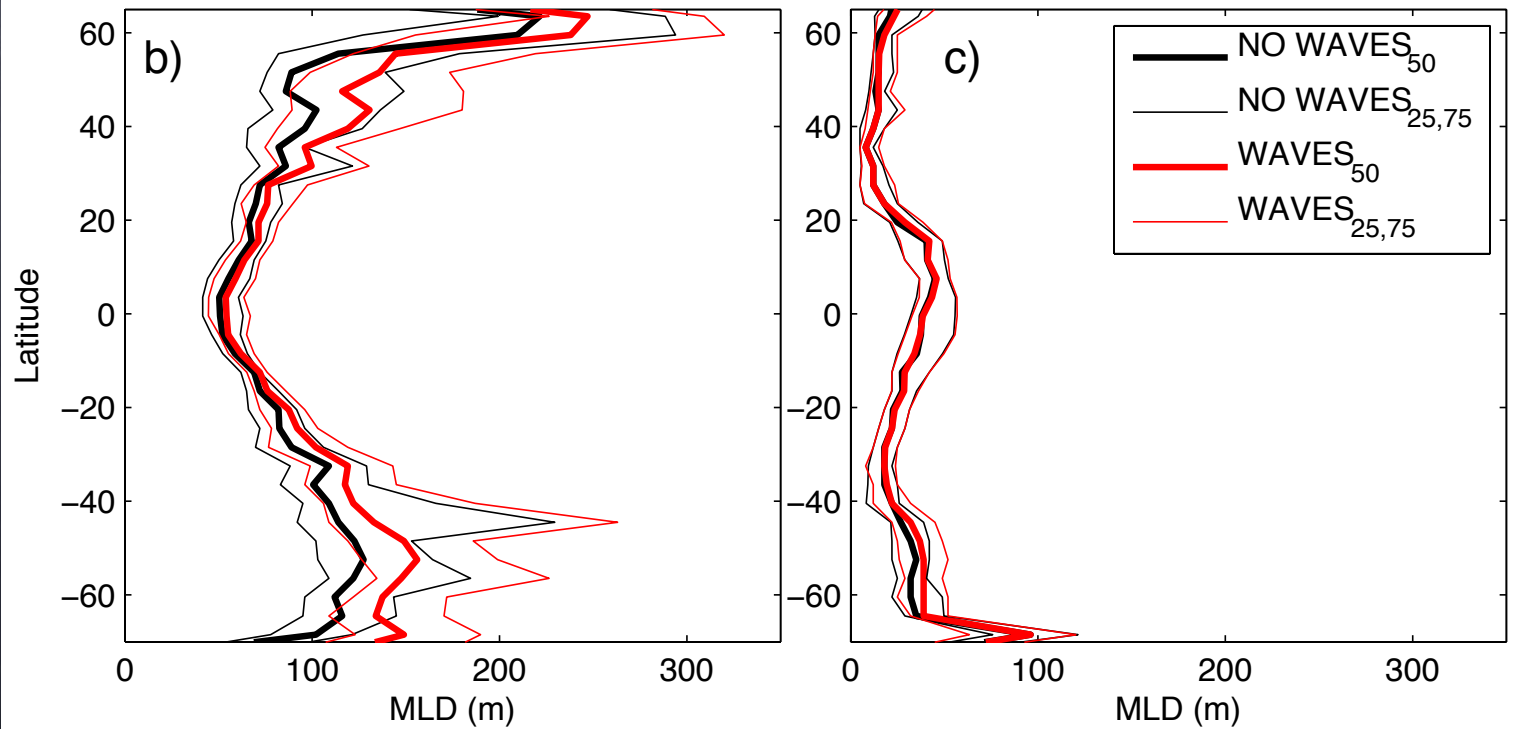
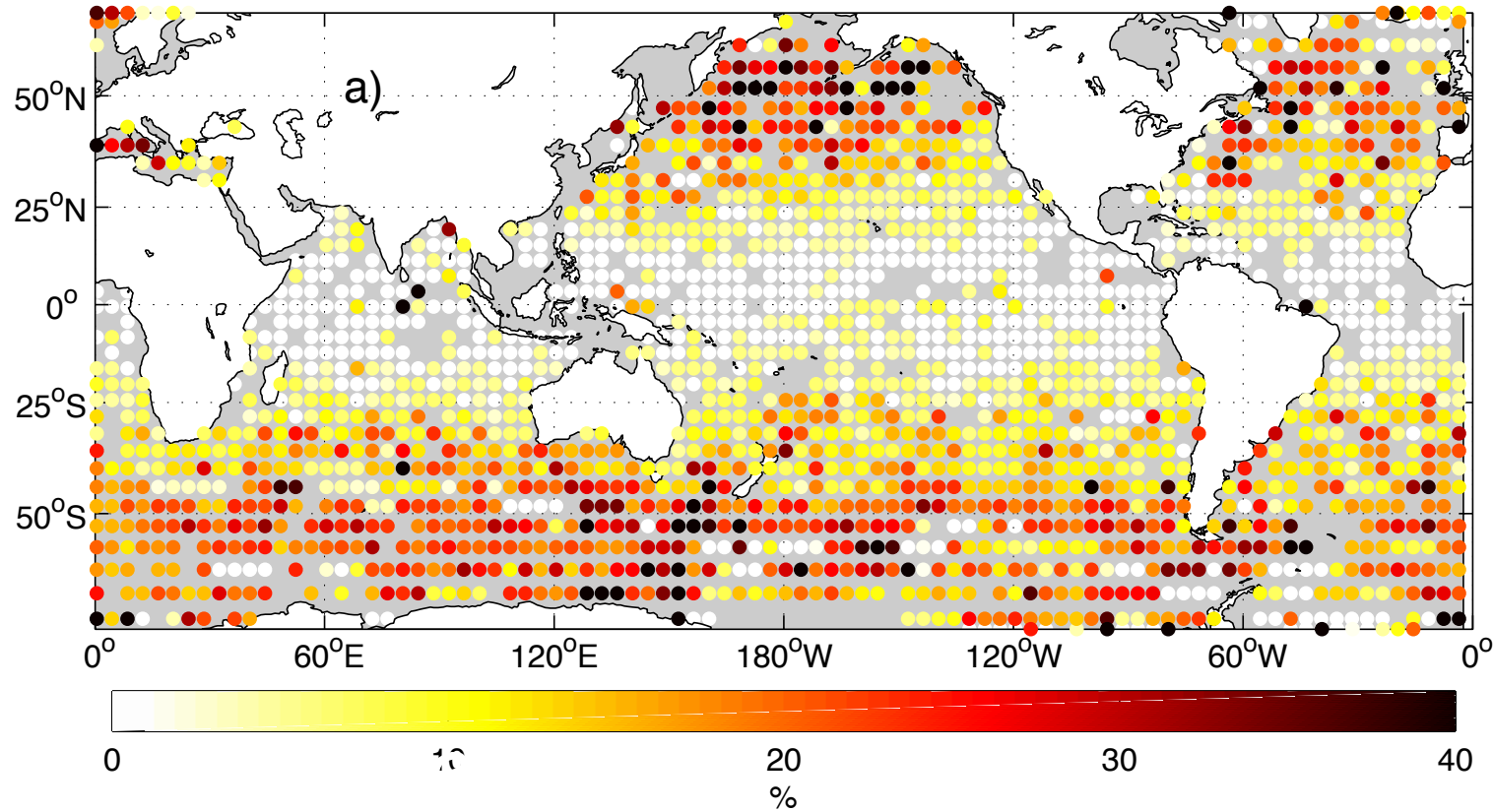
Data + LES,
 Southern Ocean
 mixing energy:
 Langmuir (Stokes-
 drift-driven) and
 Convective



S. E. Belcher, A. A. L. M. Grant, K. E. Hanley, B. Fox-Kemper, L. Van Roekel, P. P. Sullivan, W. G. Large, A. Brown, A. Hines, D. Calvert, A. Rutgersson, H. Petterson, J. Bidlot, P. A. E. M. Janssen, and J. A. Polton. A global perspective on Langmuir turbulence in the ocean surface boundary layer. *Geophysical Research Letters*, 39(18):L18605, 9pp, 2012.

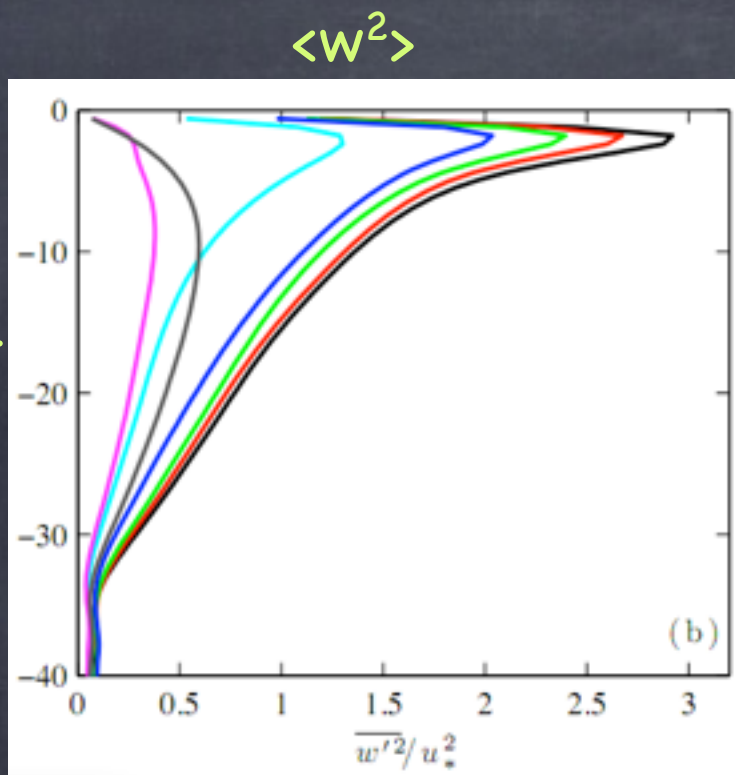
Including Wave-driven Mixing Deepens the Mixed Layer!

Fig: M. Hemer

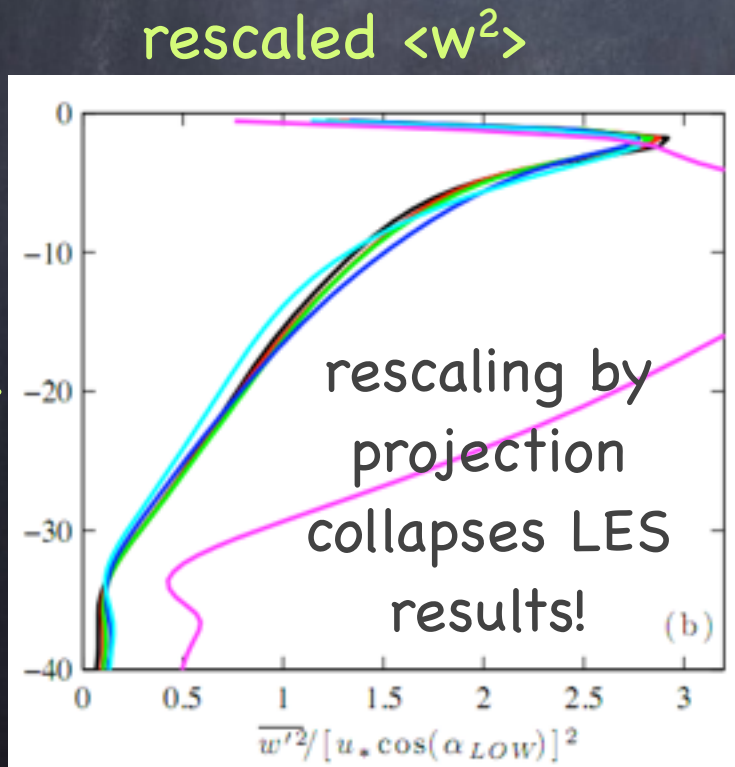


M. A. Hemer, B. Fox-Kemper,
& R. R. Harcourt. Quantifying
the effects of wind waves on
the coupled climate system, in
prep. 2012.

depth



depth



Generalized Turbulent Langmuir No.,
Projection of u^* , u_s into Langmuir Direction

$$\frac{\langle \overline{w'^2} \rangle_{ML}}{u_*^2} = 0.6 \cos^2(\alpha_{LOW}) [1.0 + (3.1 La_{proj})^{-2} + (5.4 La_{proj})^{-4}],$$

$$La_{proj}^2 = \frac{|u_*| \cos(\alpha_{LOW})}{|u_s| \cos(\theta_{ww} - \alpha_{LOW})},$$

$$\alpha_{LOW} \approx \tan^{-1} \left(\frac{\sin(\theta_{ww})}{\frac{u_*}{u_s(0)\kappa} \ln \left(\left| \frac{H_{ML}}{z_1} \right| \right) + \cos(\theta_{ww})} \right)$$

A scaling for LC
strength & direction!

L. P. Van Roekel, B. Fox-Kemper, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. *Journal of Geophysical Research-Oceans*, 117:C05001, 22pp, 2012.

What about Langmuir-Submeso Interactions?

Movie: P. Hamlington

Perform large eddy simulations (LES) of Langmuir turbulence with a submesoscale temperature front

Use NCAR LES model to solve Craik-Leibovich equations (Moeng, 1984, McWilliams et al, 1997)

$$\frac{\partial \rho}{\partial t} + \mathbf{u}_L \cdot \nabla \rho = \text{SGS}$$

$$\nabla \cdot \mathbf{u} = 0$$

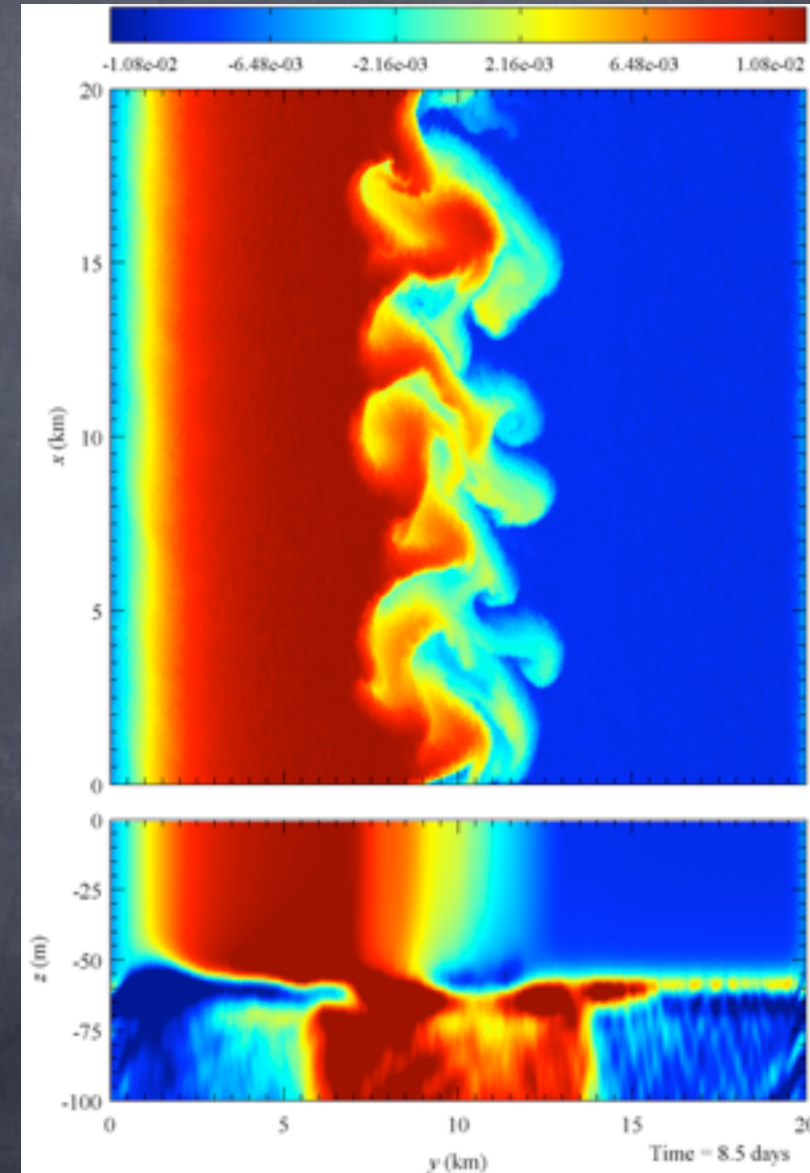
$$\frac{\partial \mathbf{u}}{\partial t} + (\boldsymbol{\omega} + f\hat{\mathbf{z}}) \times \mathbf{u}_L = -\nabla \pi - \frac{g\rho\hat{\mathbf{z}}}{\rho_0} + \text{SGS}$$

Computational parameters:

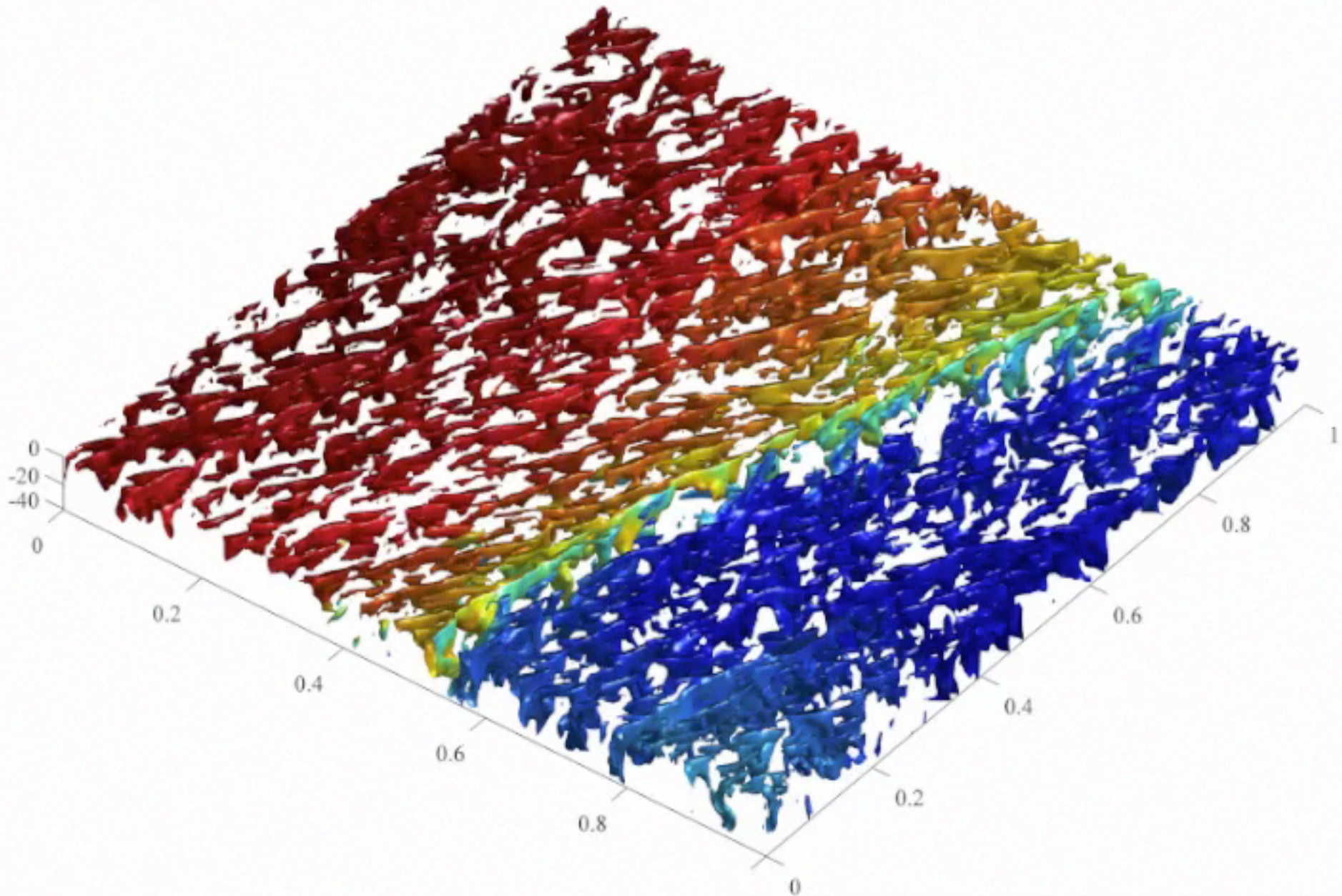
Domain size: 20km x 20km x -160m

Grid points: 4096 x 4096 x 128

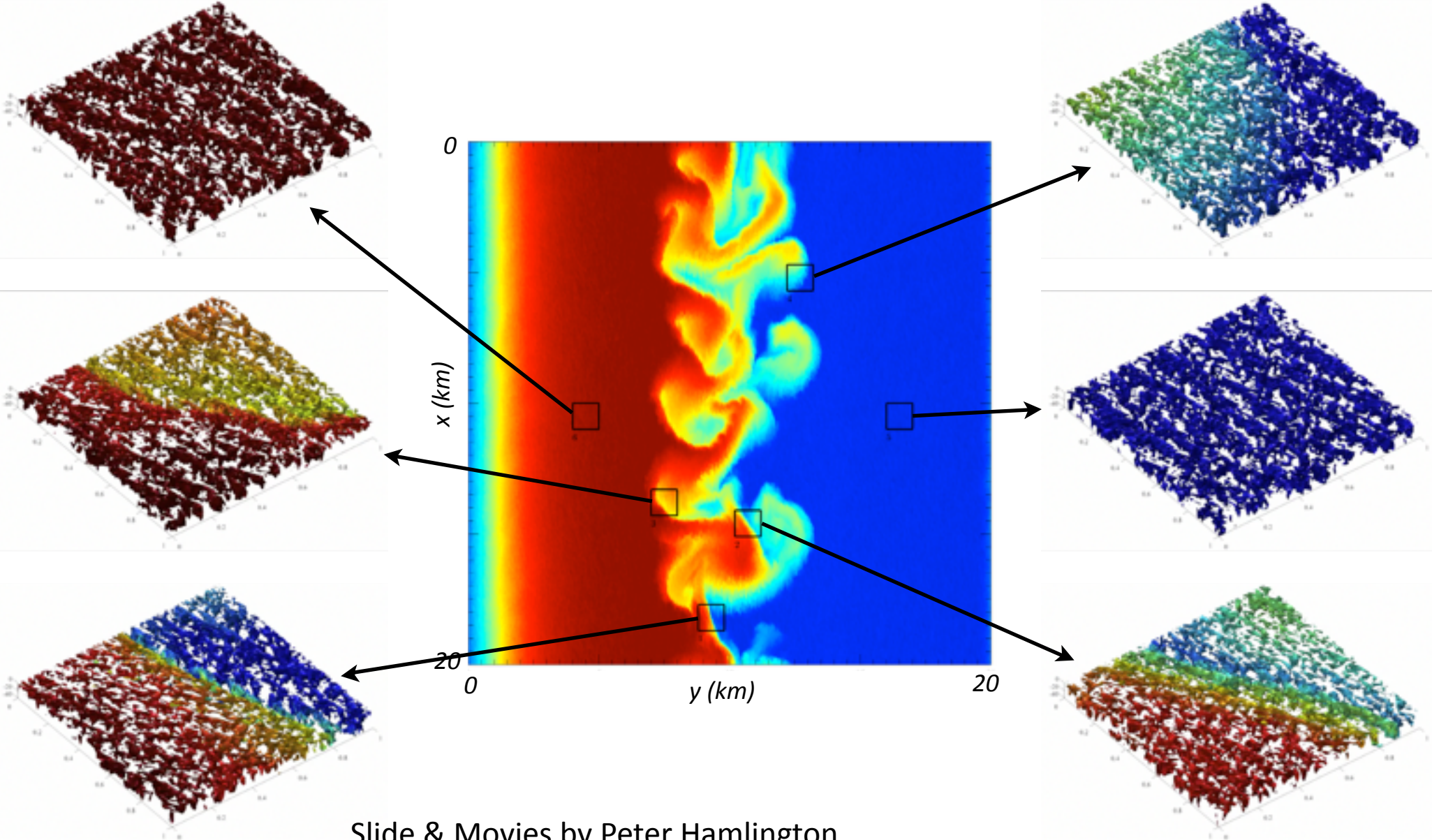
Resolution: 5m x 5m x -1.25m



Zoom: Submeso-Langmuir Interaction!

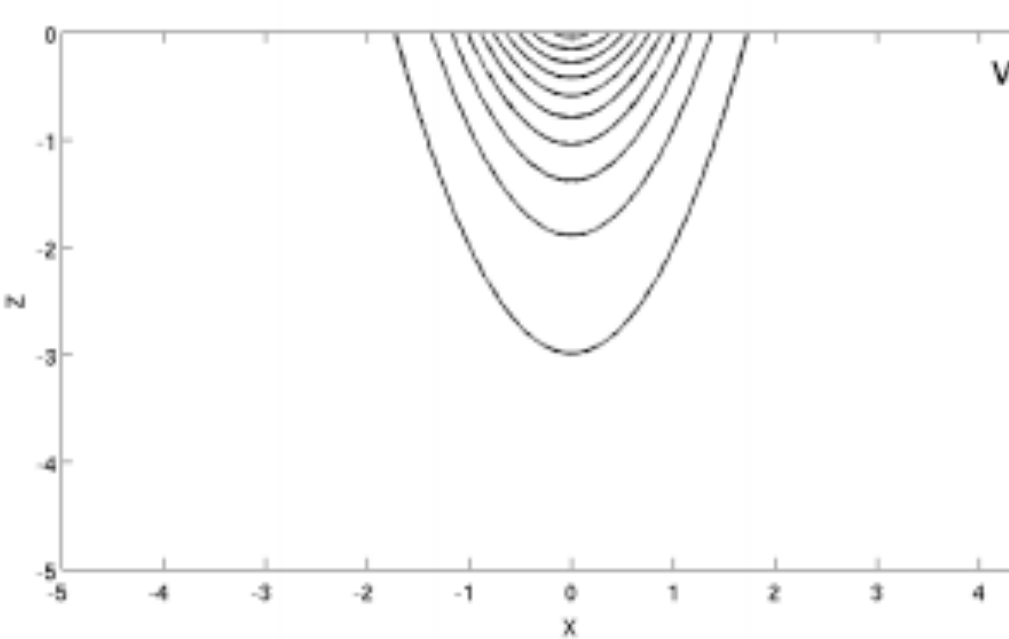


Diverse types of interaction



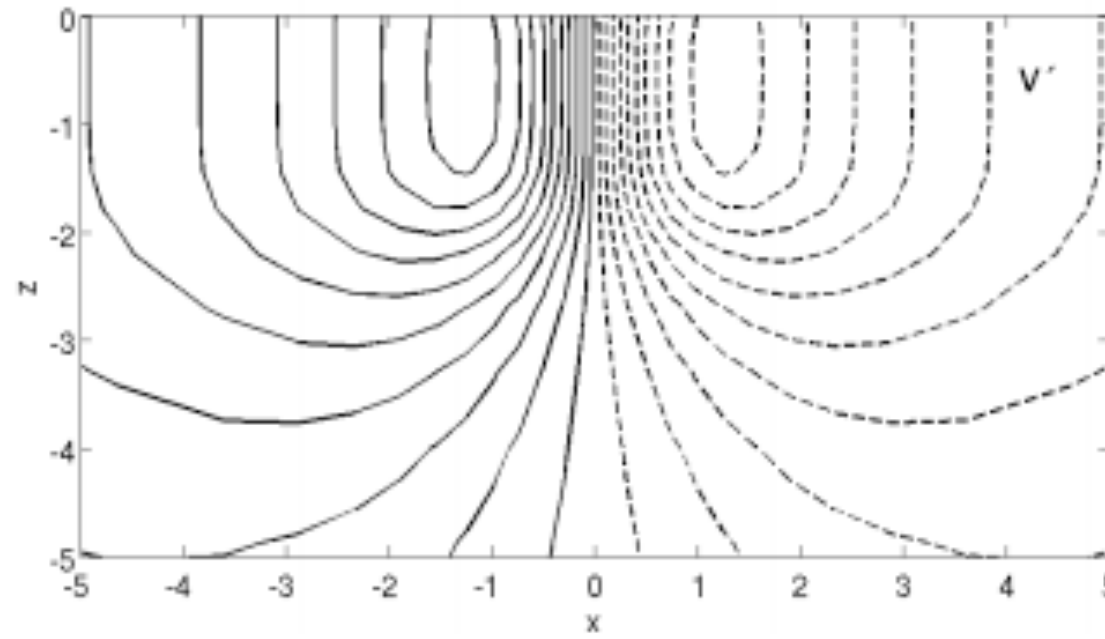
Slide & Movies by Peter Hamlington

Waves (Stokes Drift Vortex Force) \rightarrow Submeso, Meso



Initial Submeso flow

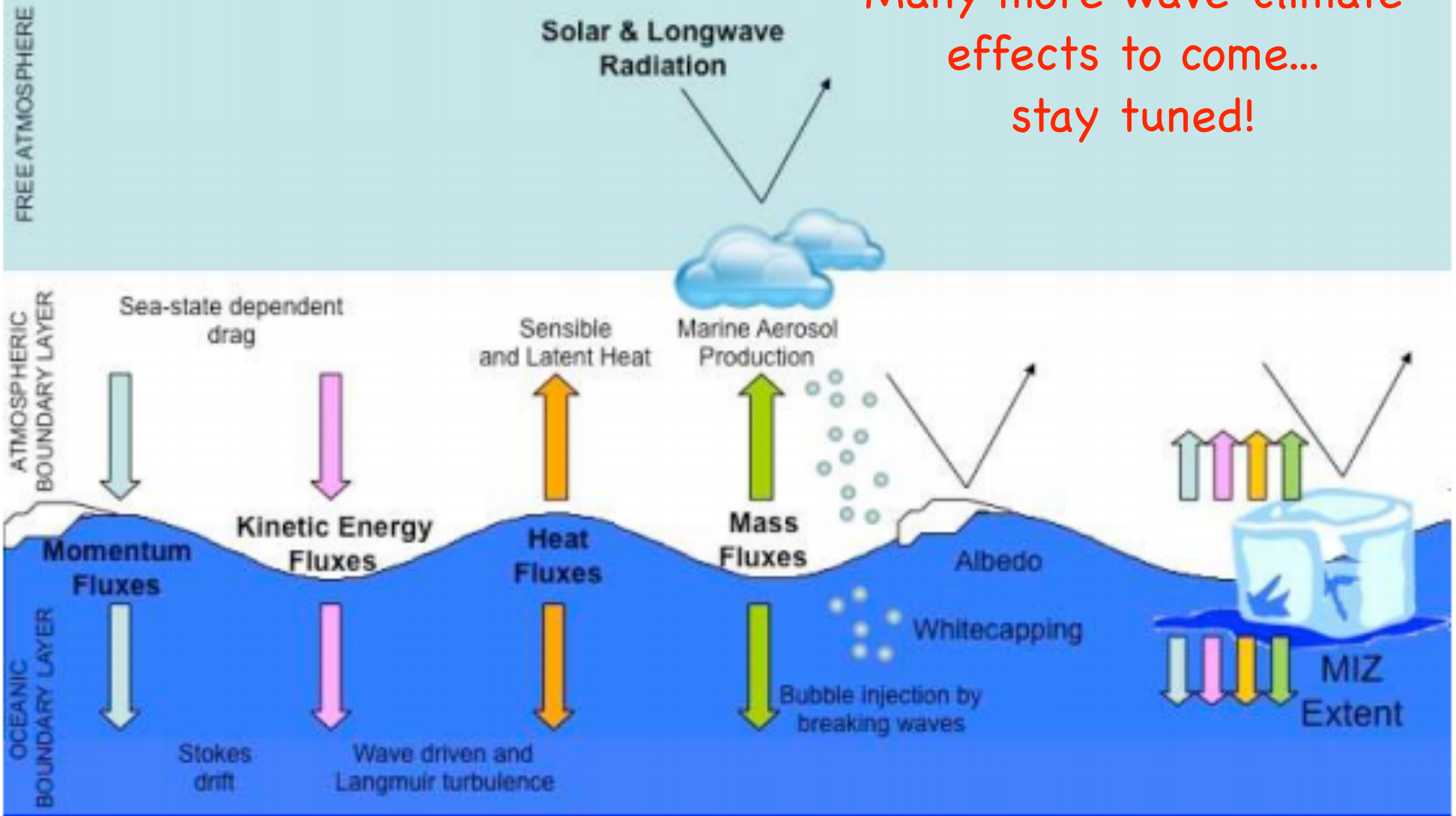
Contours: 0.1



Perturbation on that scale
due to waves

Contours: 0.014

Many more wave-climate effects to come... stay tuned!

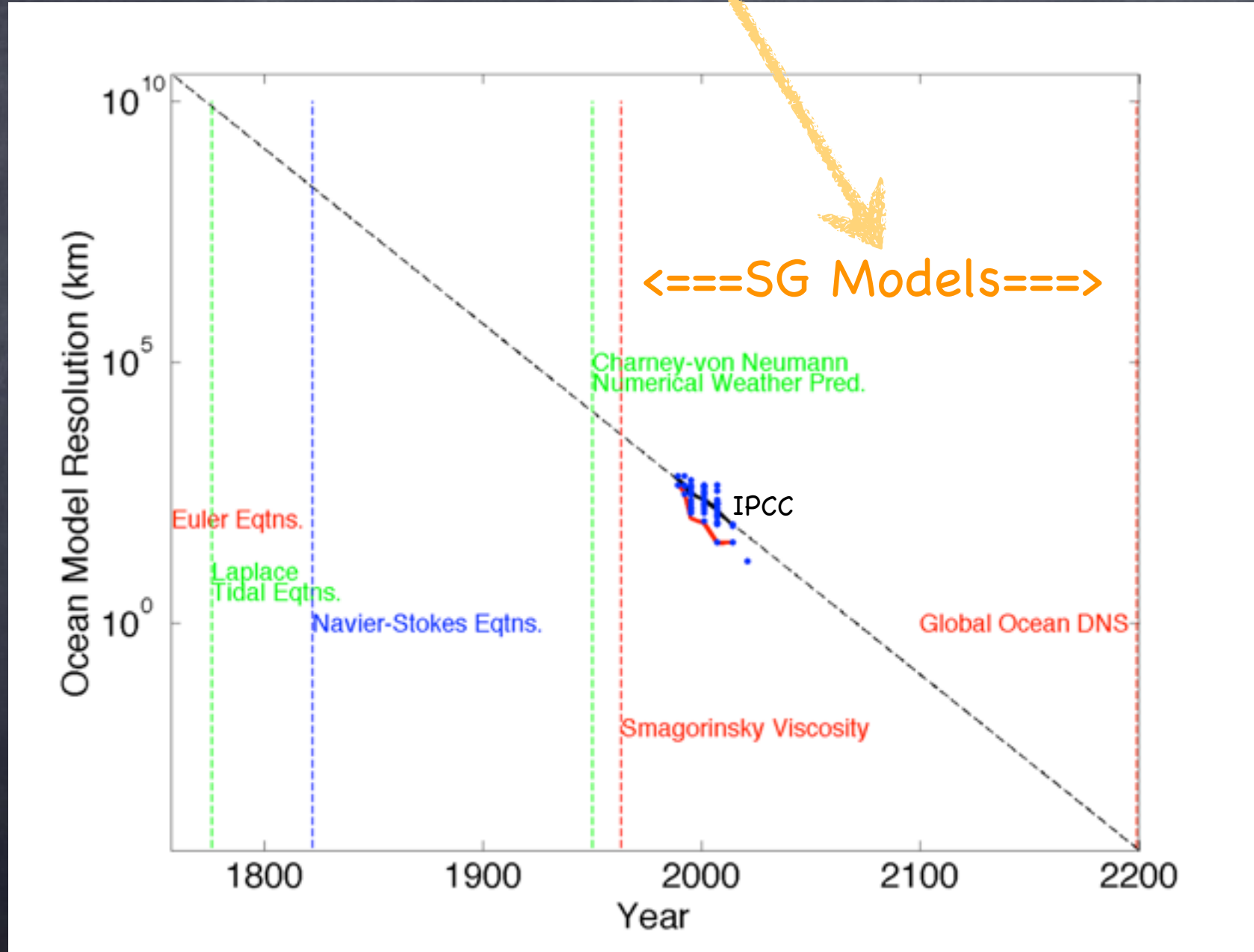


Wind-wave dependent processes in the coupled climate system
Towards coupled wind-wave-AOGCM models

Conclusions

- Climate modeling is challenging partly due to the vast and diverse scales of fluid motions
- In the upper ocean, horizontal scales as big as basins, and as small as meters contribute non-negligibly to the air-sea exchange
- Process models, especially those spanning a whole or multiple scales, are a powerful tool in studying these connections and improving subgrid models.

Extrapolate for historical perspective: The Golden Era of Subgrid Modeling is Now!



All papers at: fox-kemper.com/research

