

# From Climate to Kolmogorov – Simulations Spanning Upper Ocean Scales

Baylor Fox-Kemper

with

Sean Haney (ATOC), Adrean Webb (APPM), Scott Bachman (ATOC), Katie McCaffrey (ATOC)  
Keith Julien (CU-Boulder), Peter Hamlington (CU-Boulder), Luke Van Roekel (Northland  
College), Peter Sullivan (NCAR), Ramsey Harcourt (UW), Eric D'Asaro (UW), Jim McWilliams  
(UCLA), Mark Hemer (CSIRO)

URI GSO Physical Oceanography Seminar

Sponsors: NSF 1245944, 0934737, 0825614, NASA NNX09AF38G



Now with Rhode Island  
Ingredients!

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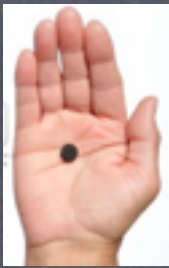
Baylor Fox-Kemper  
Brown University

with

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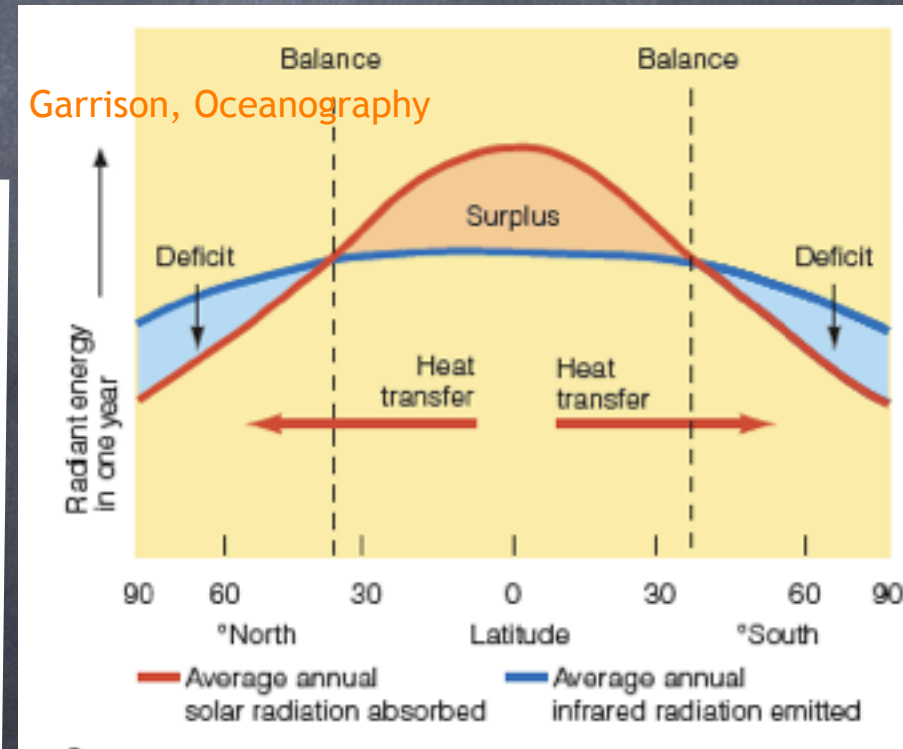
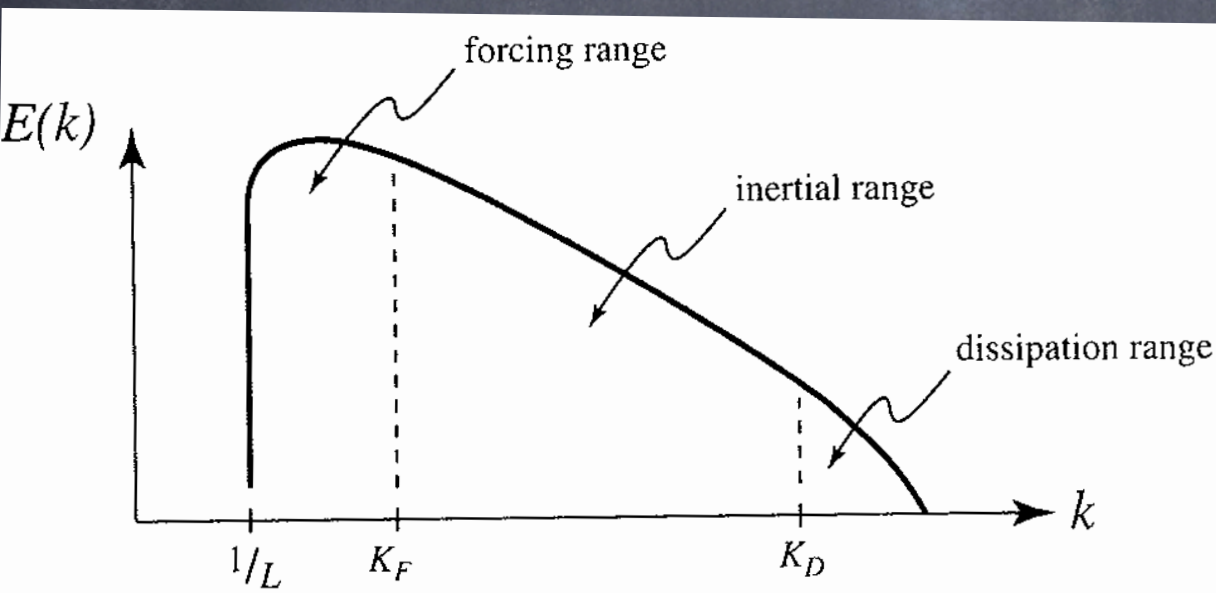
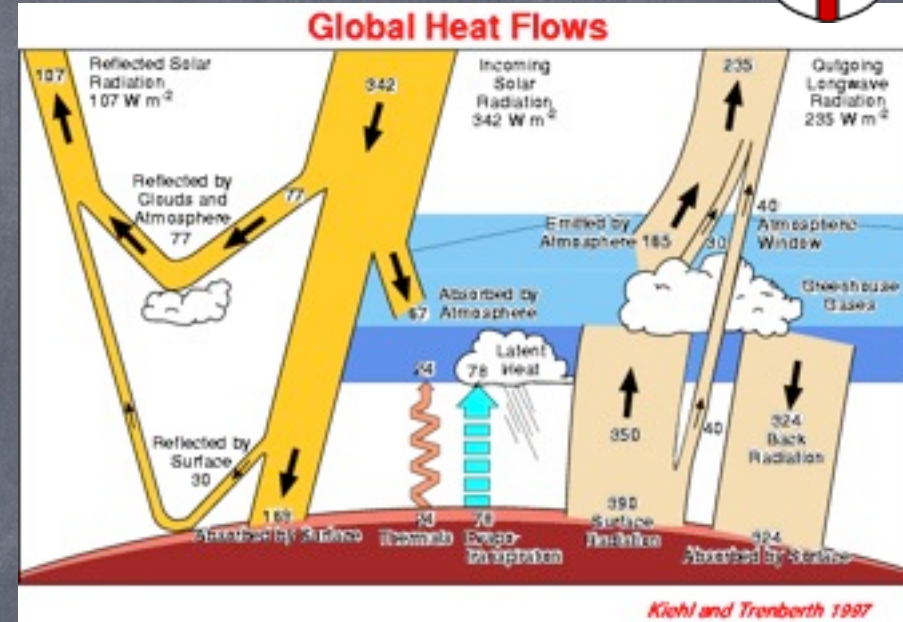
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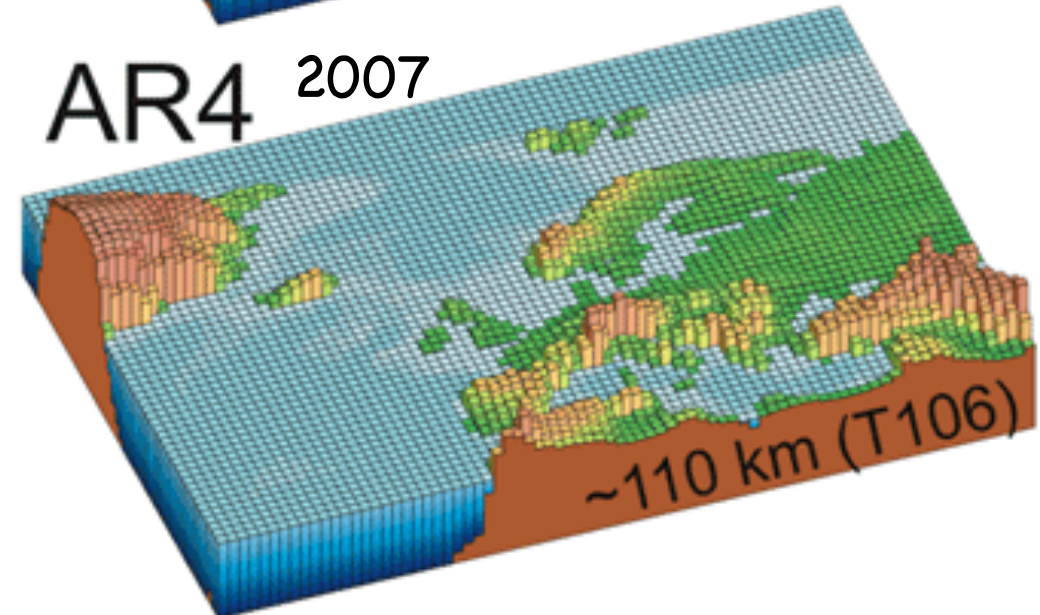
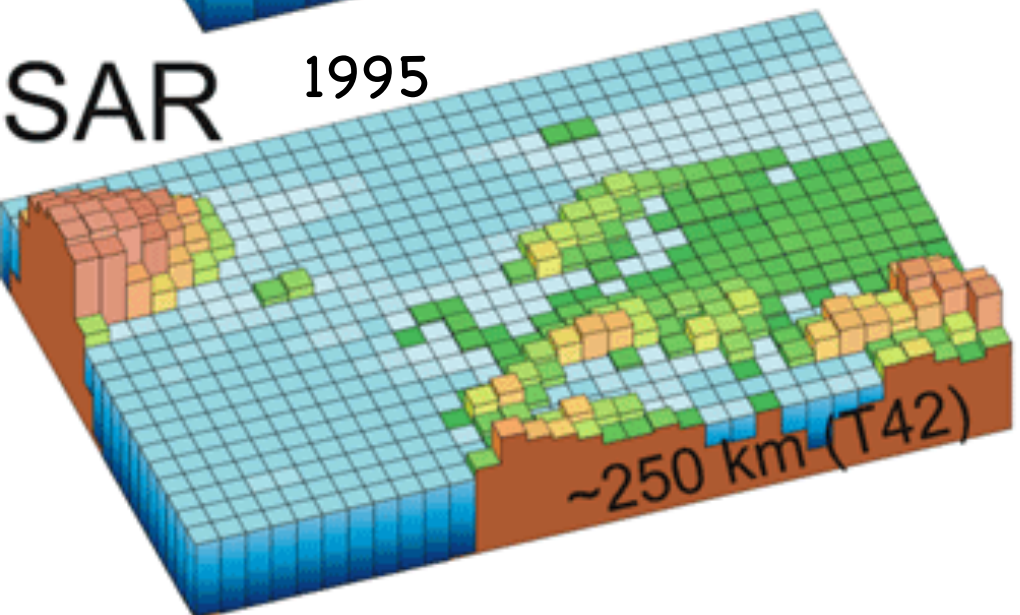
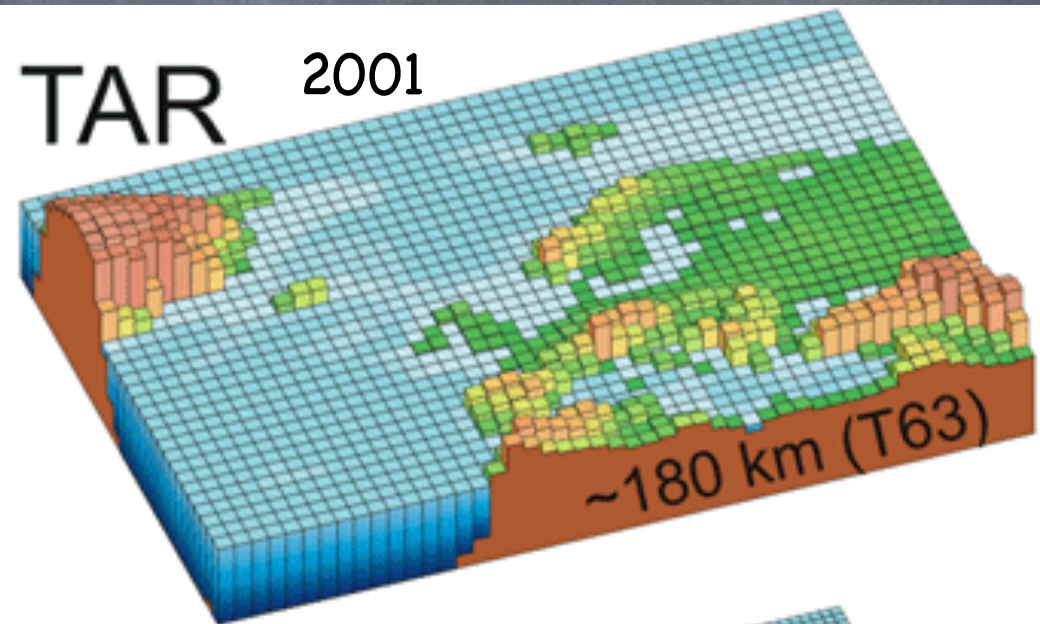
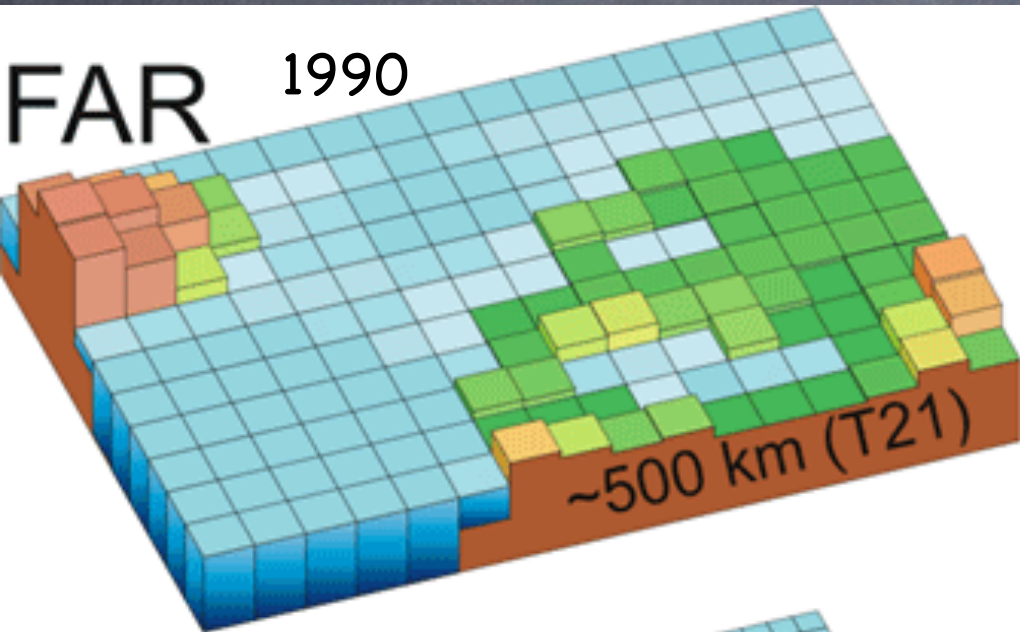
The Earth's Climate System is driven by the Sun's light (minus outgoing infrared) on a global scale



Dissipation concludes turbulence cascades to scales about a billion times smaller



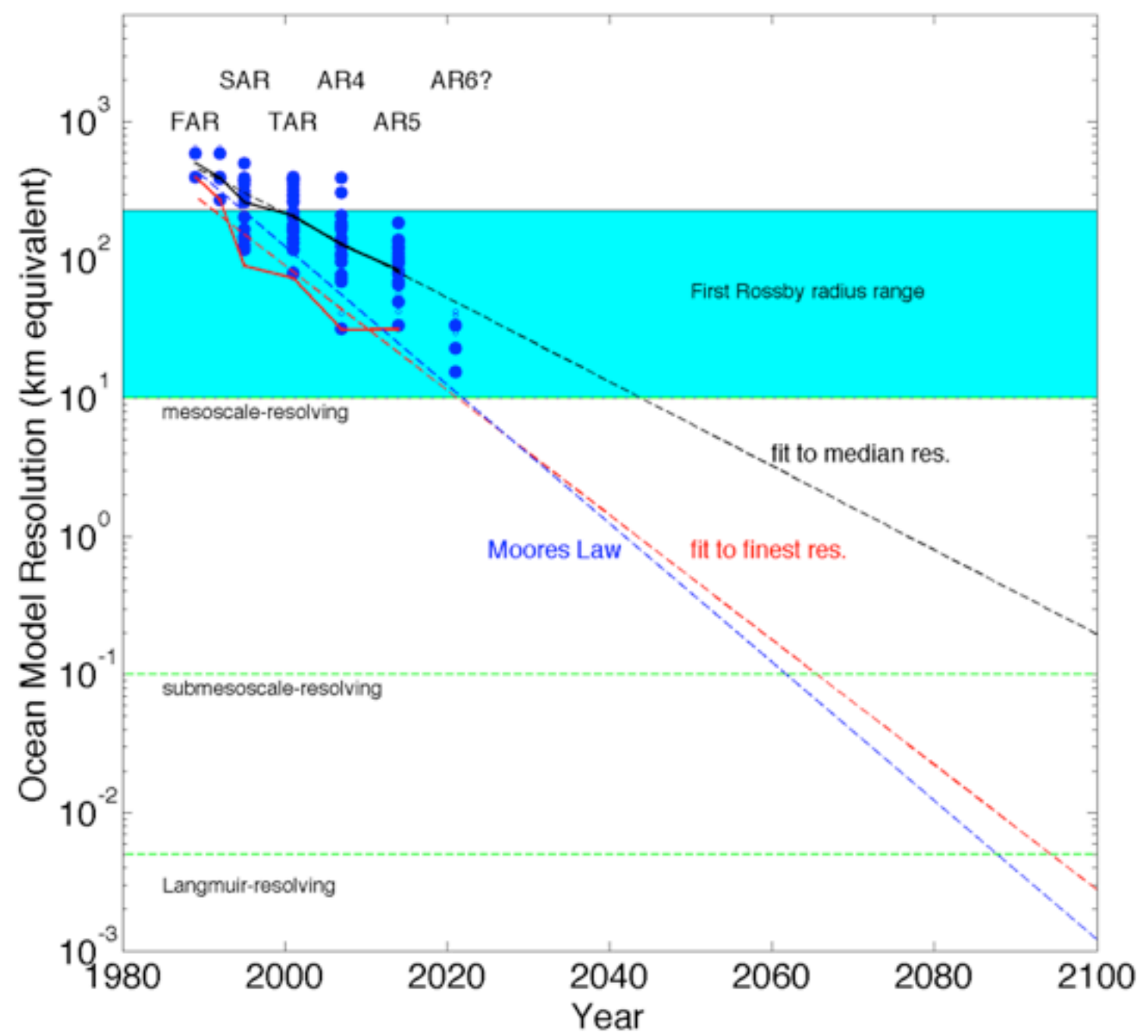
# Climate model resolution has been an issue...



And will be an issue for centuries to come!



Resolution of Ocean Component of Coupled IPCC models



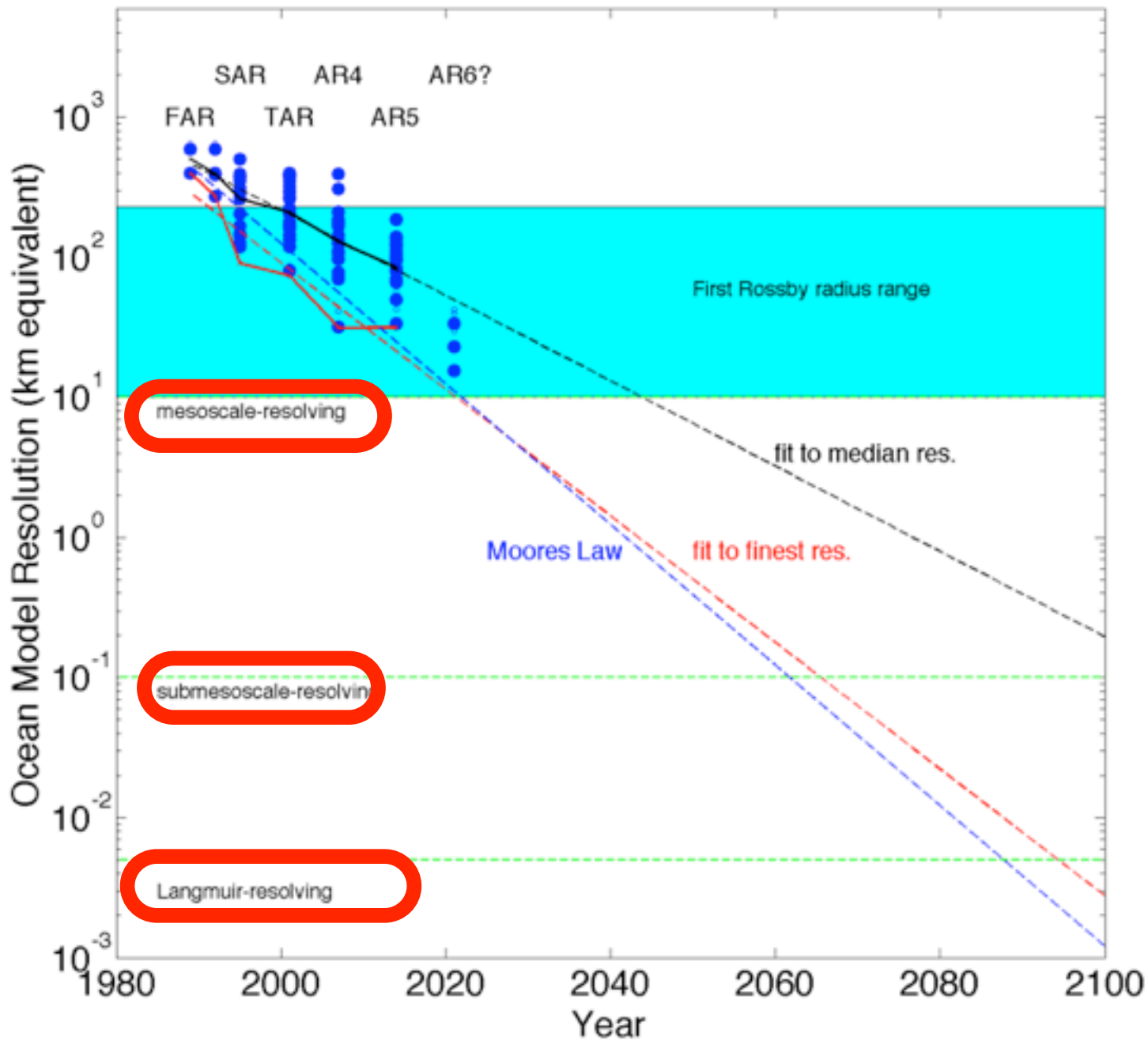
What are the processes capable of spanning scales?

What unresolved physics should we worry about?

And will be an issue for centuries to come!



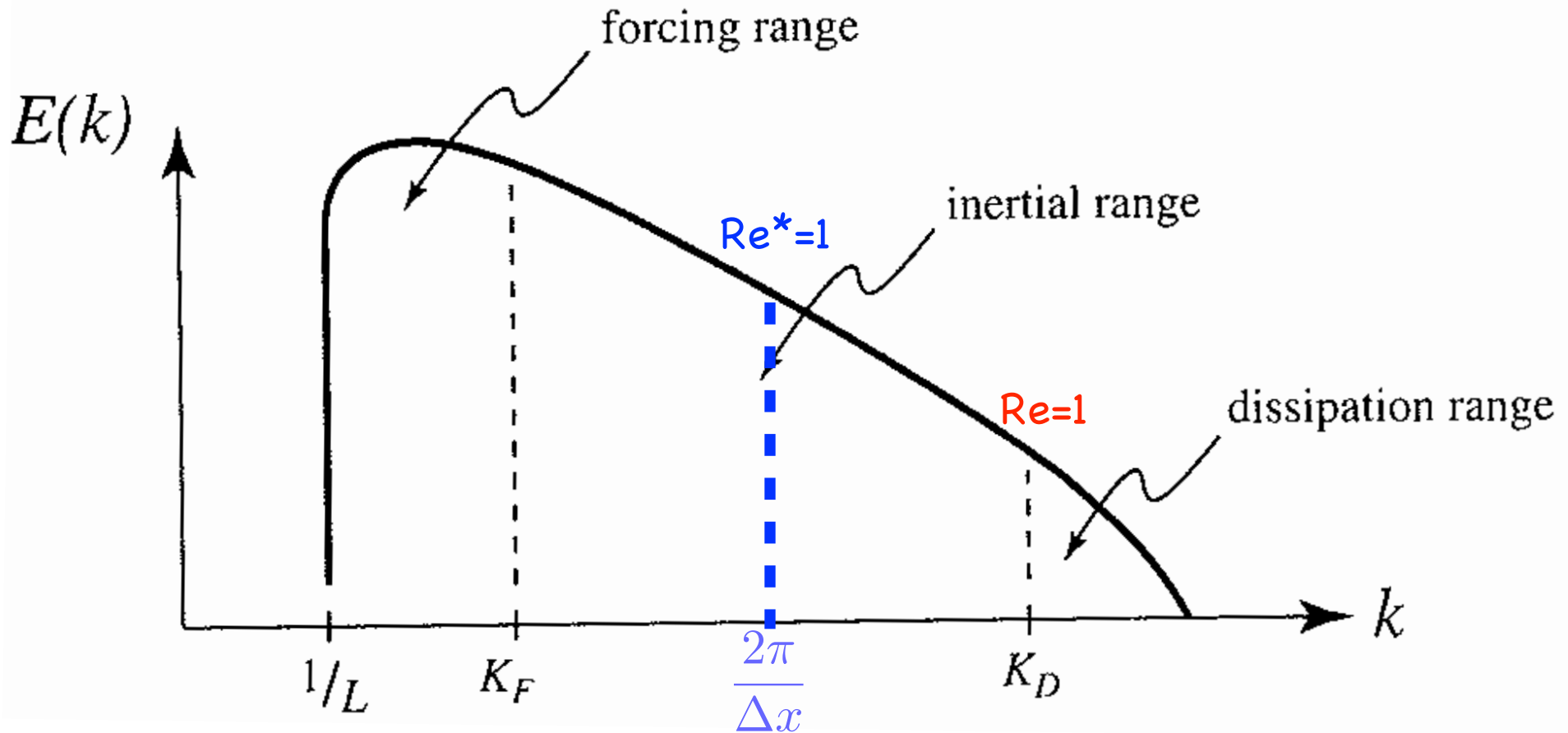
Resolution of Ocean Component of Coupled IPCC models



What are the processes capable of spanning scales?

What unresolved physics should we worry about?

# Truncation of Cascades in models



1963: Smagorinsky Devises Viscosity Scaling,  
Energy Flow is Preserved,

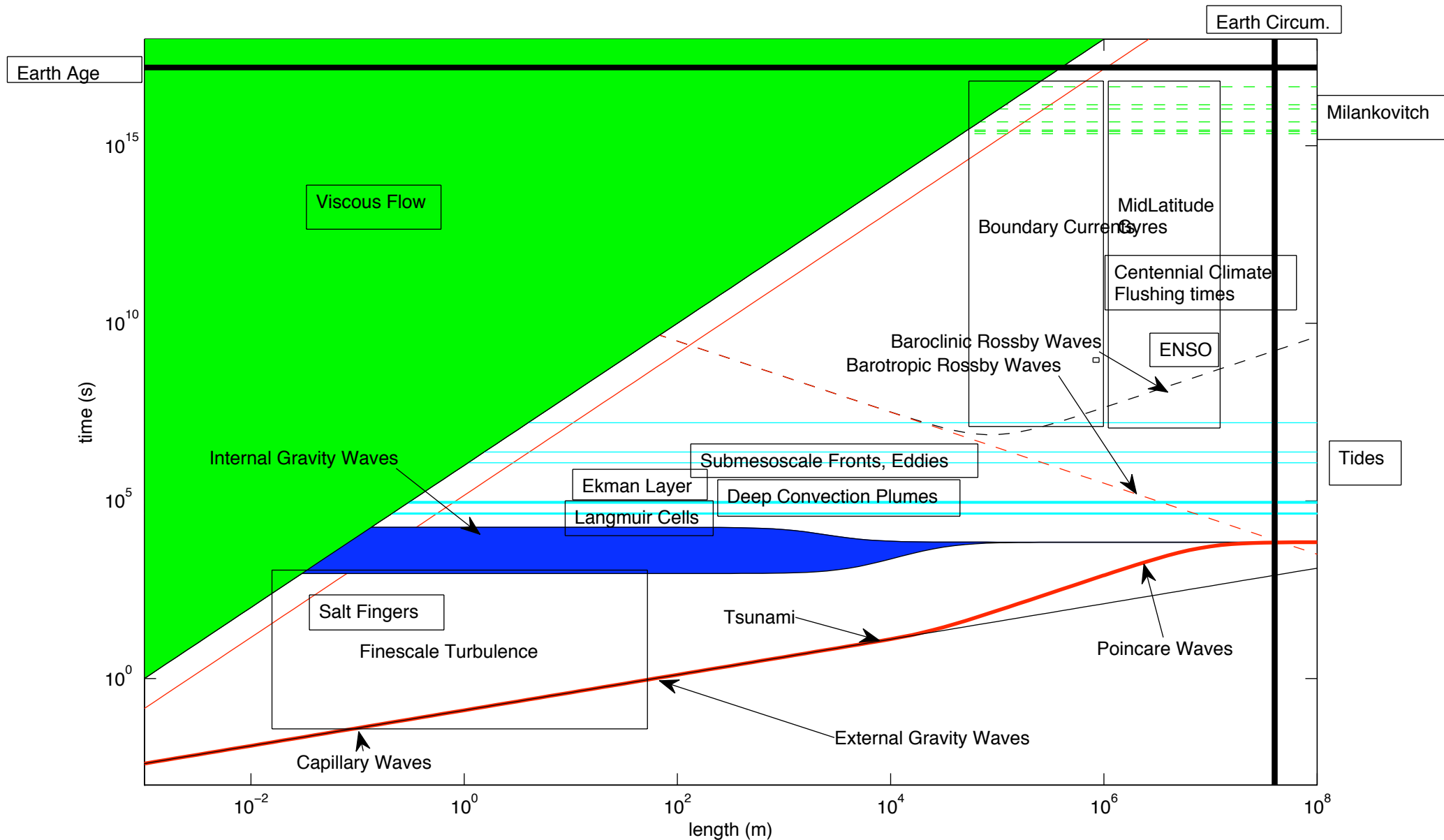
but order-1 gridscale Reynolds #:  $Re^* = UL/\nu_*$

$$\nu_{*h} = \left( \frac{\Upsilon_h \Delta x}{\pi} \right)^2 \sqrt{\left( \frac{\partial u_*}{\partial x} - \frac{\partial v_*}{\partial y} \right)^2 + \left( \frac{\partial u_*}{\partial y} + \frac{\partial v_*}{\partial x} \right)^2}$$

Gridscale-  
dependent



# The Ocean is Vast & Diverse: just one spectral cascade?





# So, what to do?

- Climate modelling requires that we truncate the model grid at coarse resolution (albeit improving slowly)
- Whatever resolution we can afford will leave some physics unresolved or partially-resolved, so we need subgrid closures!
- The vast & diverse scales of motion in the ocean suggest that we cannot use a one-size-fits-all approach, e.g., a turbulent cascade of 3d turbulence
- So, we have to invent new subgrid closures repeatedly, parameterizing processes important at each gridscale

# What is a subgrid model?



- Express the **coarse-grain averages** of quantities (including the subgrid effects), e.g.:

$$\overline{\frac{\partial \tau}{\partial t}} \quad \overline{\frac{\partial u}{\partial x}} \quad \overline{\frac{\partial u \tau}{\partial x}}$$

- As a function of the **resolved coarse-grain fields**

$$\overline{\frac{\partial \tau}{\partial t}} = \frac{\partial \bar{\tau}}{\partial t} \quad \overline{\frac{\partial u}{\partial x}} = \frac{\partial \bar{u}}{\partial x} \quad \overline{\frac{\partial u \tau}{\partial x}} = \frac{\partial \bar{u} \bar{\tau}}{\partial x} + \overline{\frac{\partial u' \tau'}{\partial x}}$$

- Note that **nonlinear** terms require **special treatment**
- These couple different scales, small talks to large

# Climate: What is important?

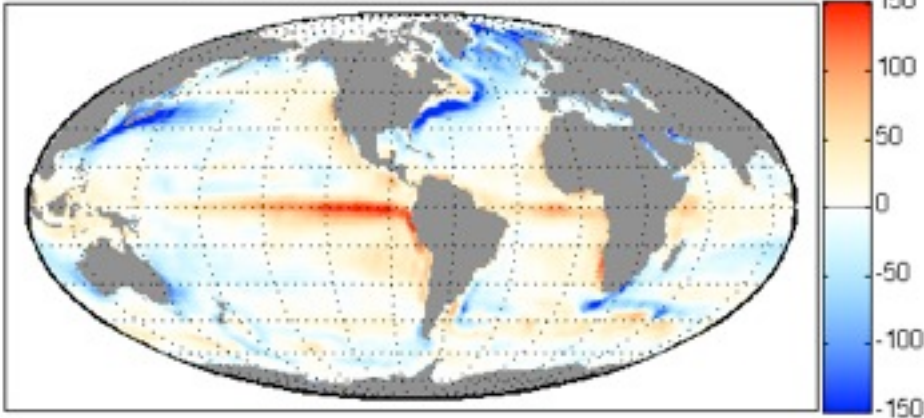


- To approximate absorption, reemission, and redistribution of the Sun's energy across the globe
- Need atmospheric chemistry (greenhouse gasses) & clouds for absorption & reemission
- Need ocean (surface) as it exchanges
  - sensible heat
  - latent heat (evaporation, freezing, precipitation)
  - gasses
  - momentum
- Plus, ocean transports heat itself!

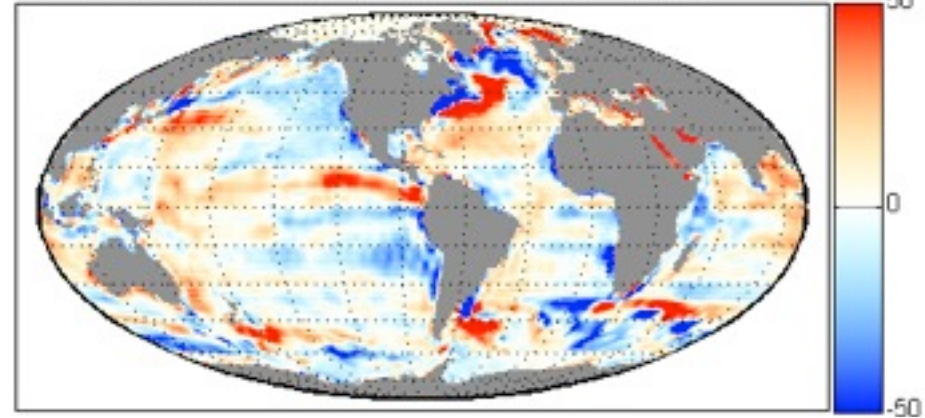


# CCSM4 Air-Sea Flux Errors vs. Data (Large & Yeager 09)

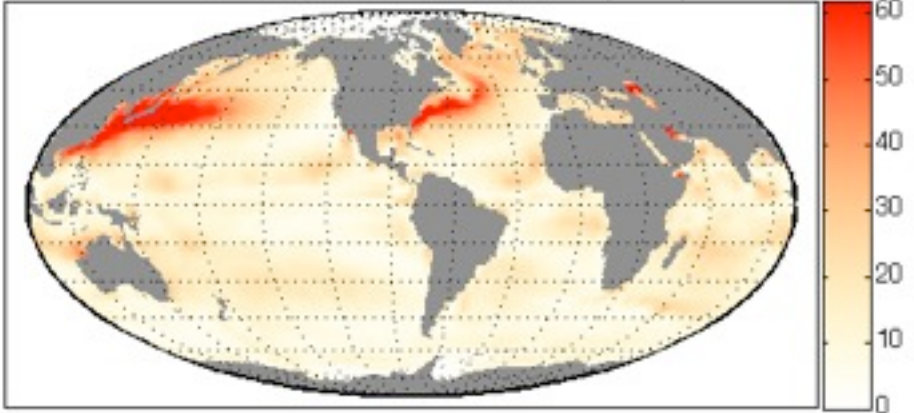
Mean of 1986-2005 CORE  $Q_{es}$  ( $W/m^2$ )



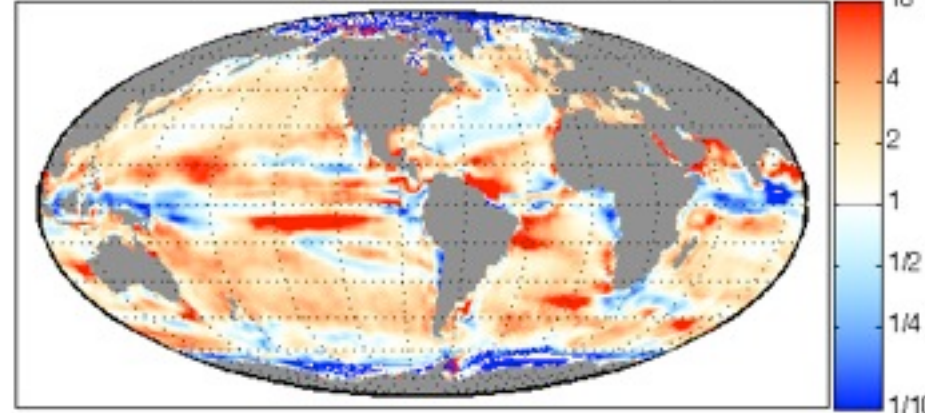
1986-2005 CCSM4-CORE  $Q_{es}$  bias, mean:1.5, rms:23 ( $W/m^2$ )



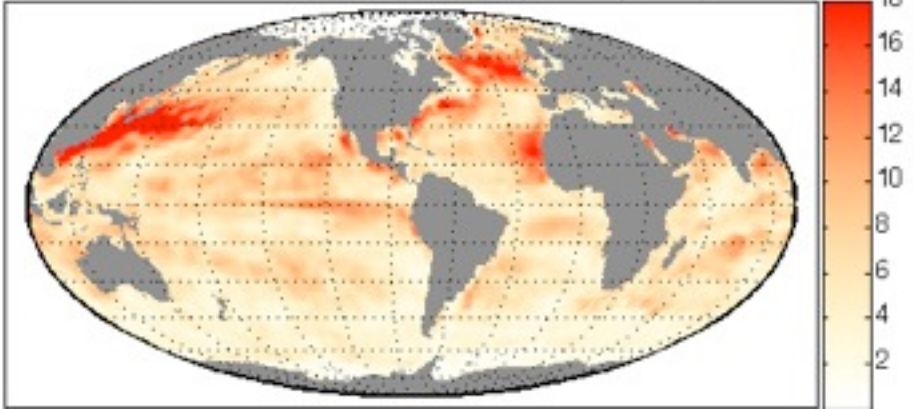
St. Dev. of CORE annual evaporation ( $W/m^2$ )



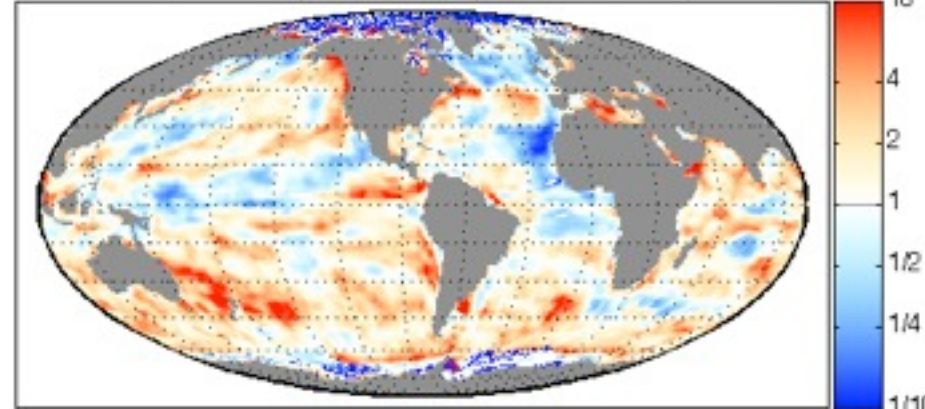
Variance ratio (CCSM4/CORE) of annual evaporation



St. Dev. of CORE interannual evaporation ( $W/m^2$ )



Variance ratio (CCSM4/CORE) of interannual evaporation



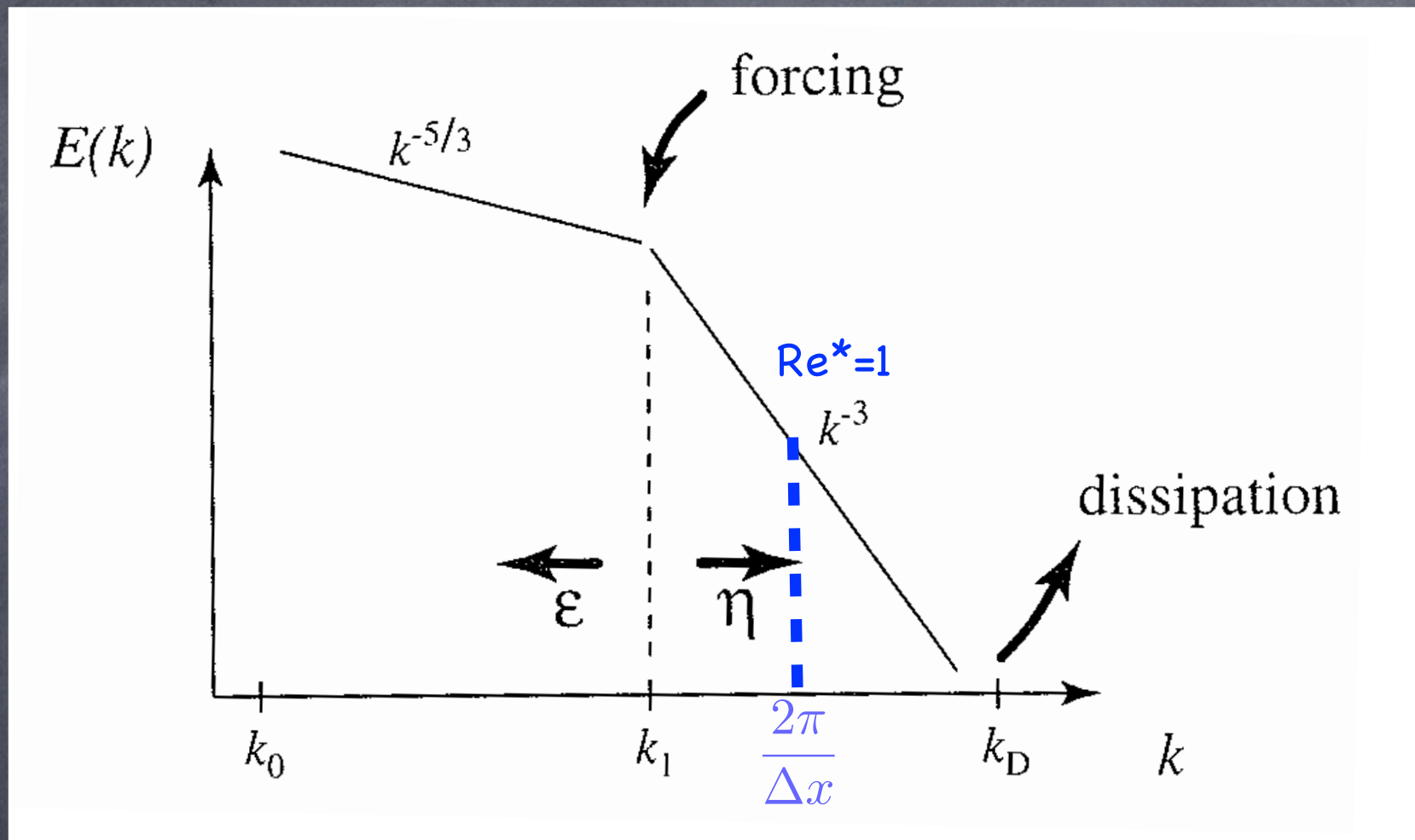
Mean  
Annual 9-15mo  
Interannual 2-7yr

# Except... Ocean Turbulence isn't 3d Turbulence at the mesoscale

- The ocean is wide (10,000km)
- But not deep (4km)
  - Motions in upper 1km
- Motions are largely 2d
- The layer of blue paint on a globe has roughly the right aspect ratio!



# 2d Turbulence Differs

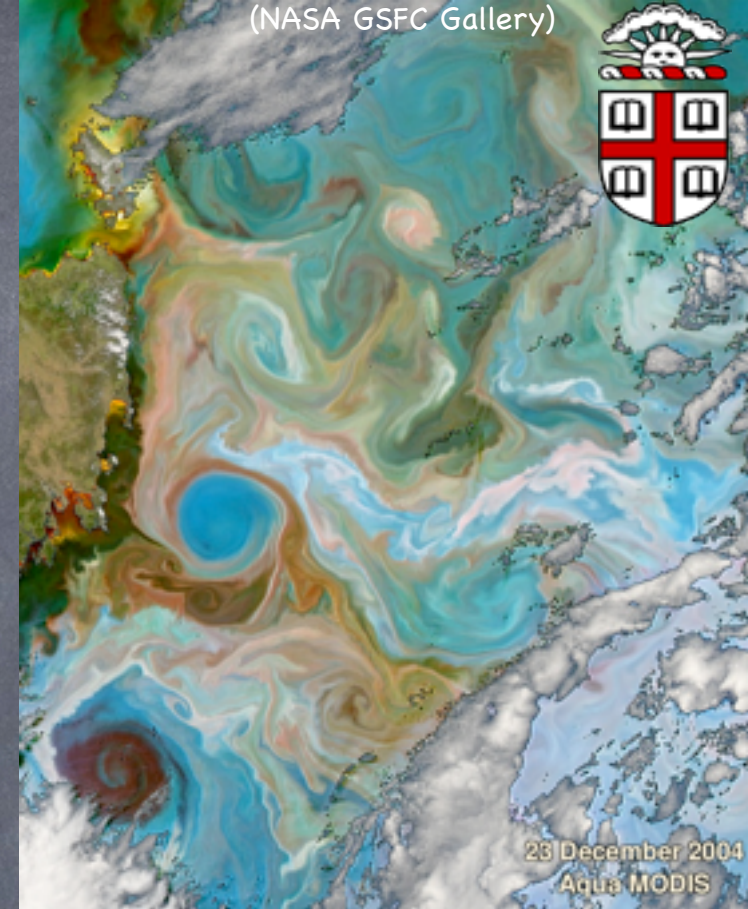


1996: Leith Devises Viscosity Scaling,  
So that the Enstrophy Flow is Preserved

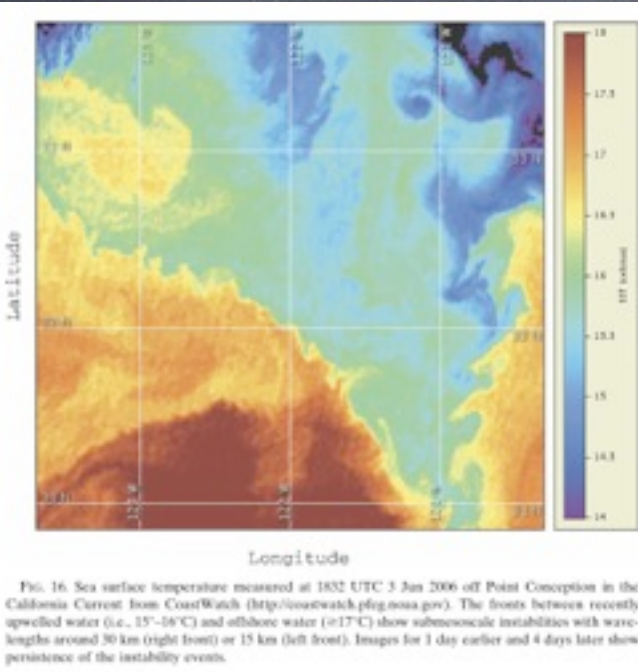
$$\nu_* = \left( \frac{\Lambda \Delta x}{\pi} \right)^3 \left| \nabla_h \left( \frac{\partial u_*}{\partial y} - \frac{\partial v_*}{\partial x} \right) \right|.$$

# The Character of the Mesoscale

← 100 km



(Capet et al., 2008)



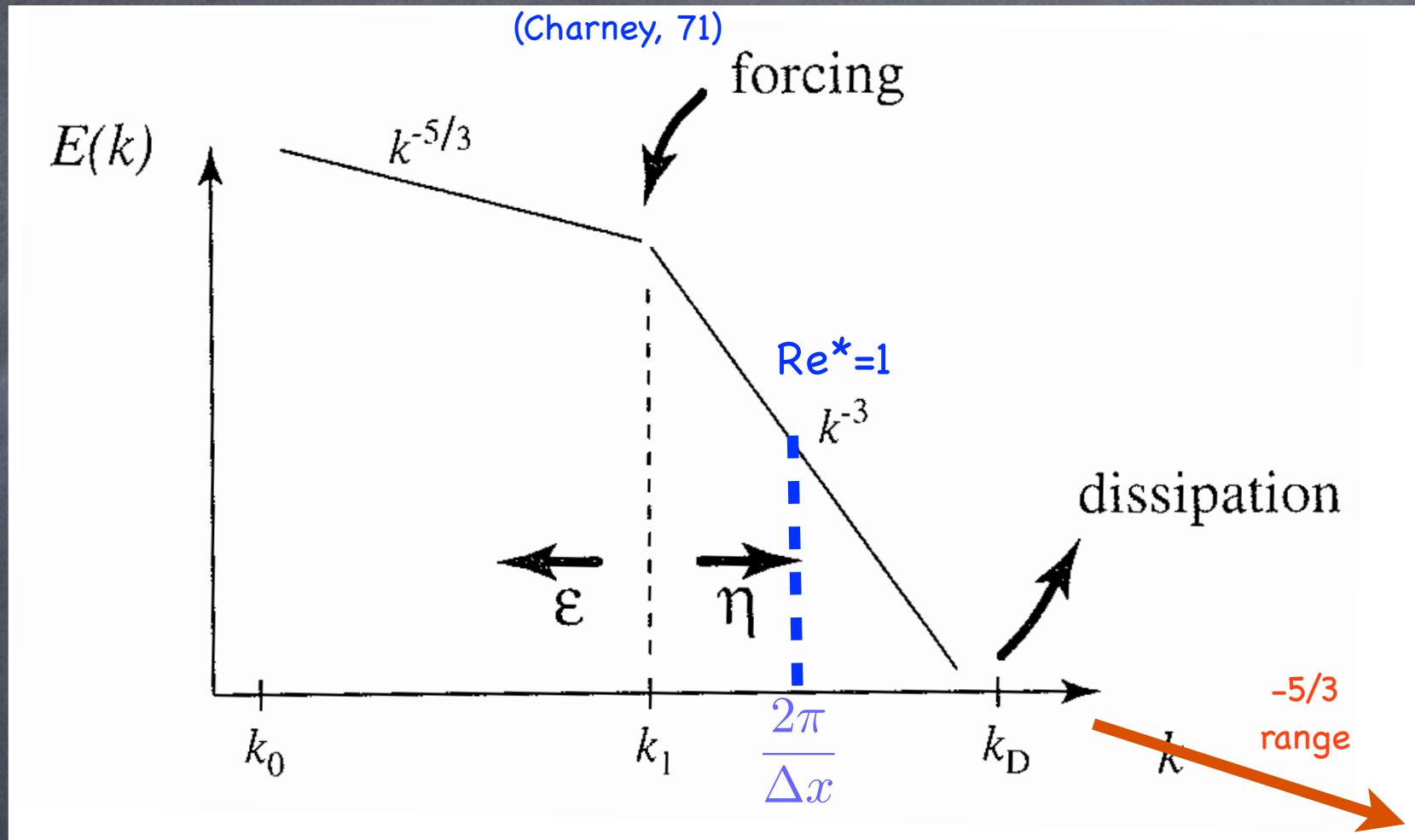
- Boundary Currents
- Eddies
- $Ro=O(0.1)$
- $Ri=O(1000)$
- Full Depth
- Eddies strain to produce Fronts
- 100km, months

Eddy processes mainly **baroclinic & barotropic instability**. Parameterizations of baroclinic instability (GM, Visbeck...).





Mesoscale Turb. has Pot'l Enstrophy cascade, but divergent



F-K & Menemenlis Revise Leith Viscosity Scaling,  
So that diverging, vorticity-free, modes are also damped

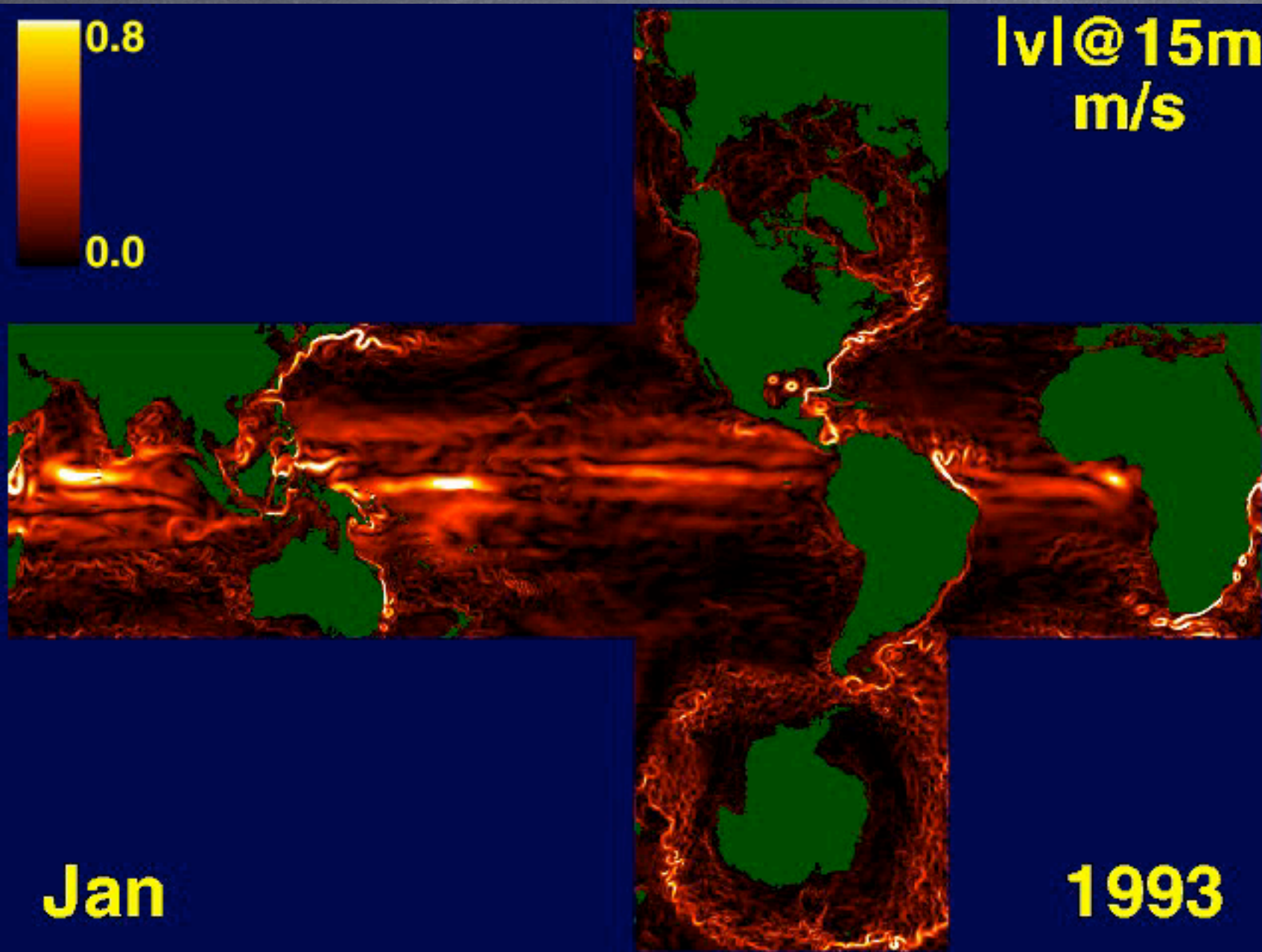
$$\nu_* = \left(\frac{\Delta x}{\pi}\right)^3 \sqrt{\Lambda^6 |\nabla_h q_{2d}|^2 + \Lambda_d^6 |\nabla_h (\nabla_h \cdot \mathbf{u}_*)|^2}$$

B. Fox-Kemper and D. Menemenlis. Can large eddy simulation techniques improve mesoscale-rich ocean models? In M. Hecht and H. Hasumi, editors, Ocean Modeling in an Eddy Regime, volume 177, pages 319-338. AGU Geophysical Monograph Series, 2008.

# ECCO2: Estimating the Current Climate of the Ocean



Phase II uses this viscosity scaling.



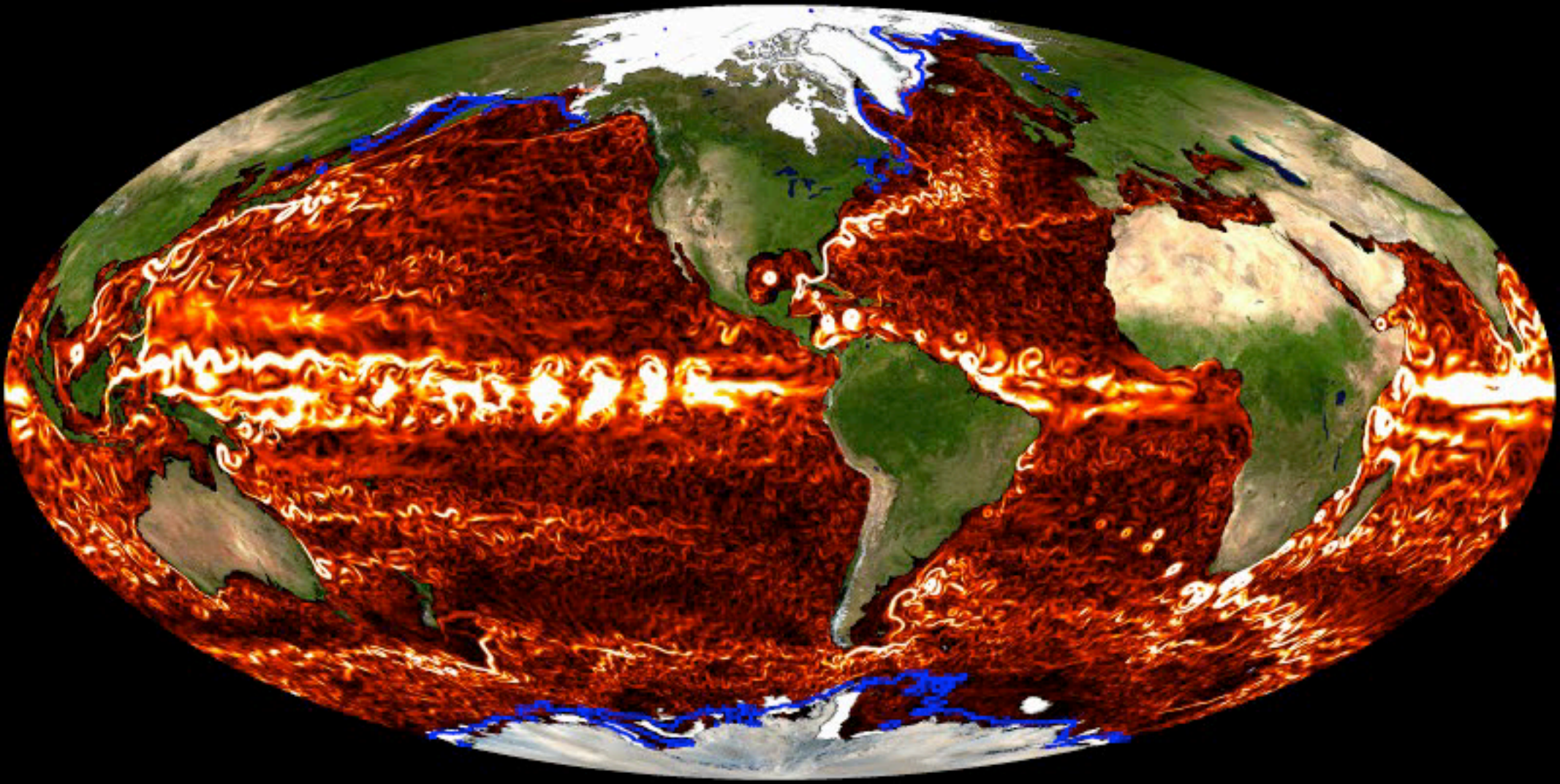
Note how no obvious artifacts occur at the grid joints, despite variations in grid scale, etc.

0.0

0.5

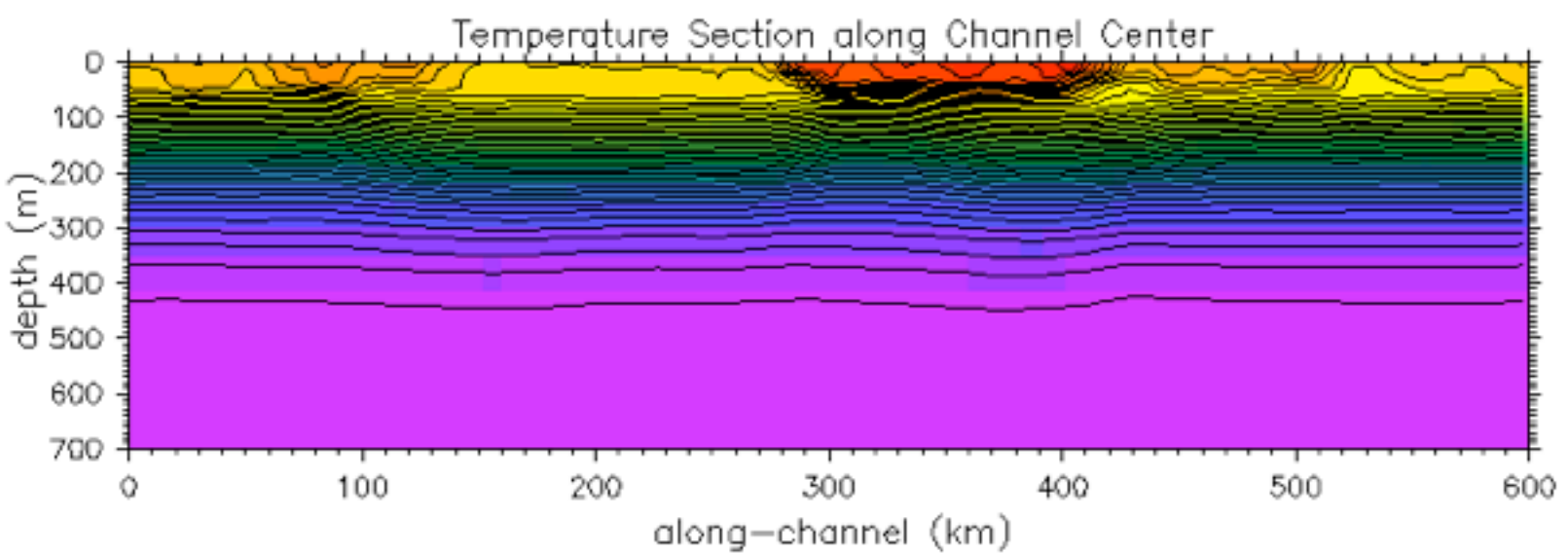
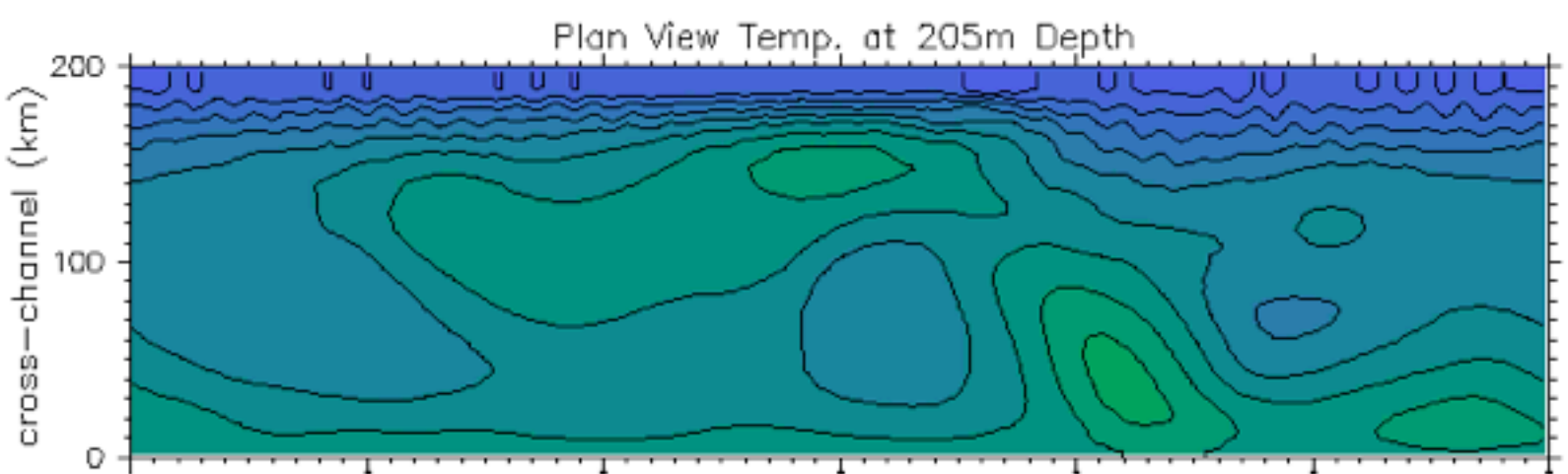
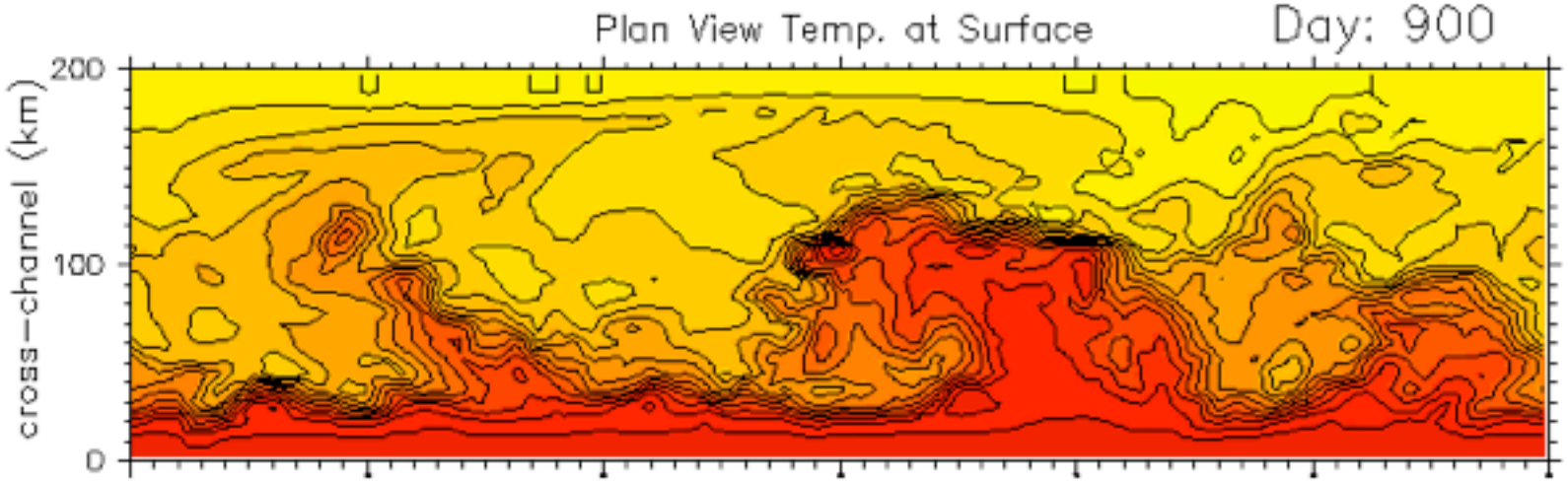


$|v|@15m$   
m/s



ECCO : 1992 - 2002  
c6ncp10

Jan 1992



Big, Deep  
(meso)

interact  
with

Little,  
Shallow  
(submeso)

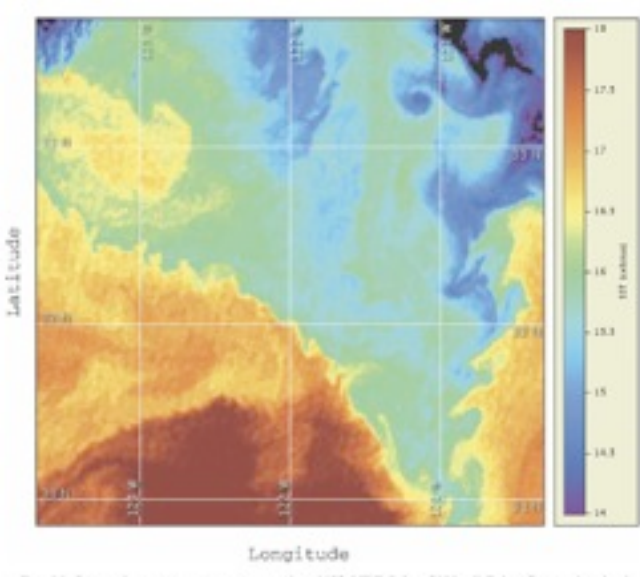
B. Fox-Kemper, R. Ferrari,  
and R. W. Hallberg.  
Parameterization of mixed  
layer eddies. Part I: Theory  
and diagnosis. *Journal of  
Physical Oceanography*,  
38(6):1145-1165, 2008.



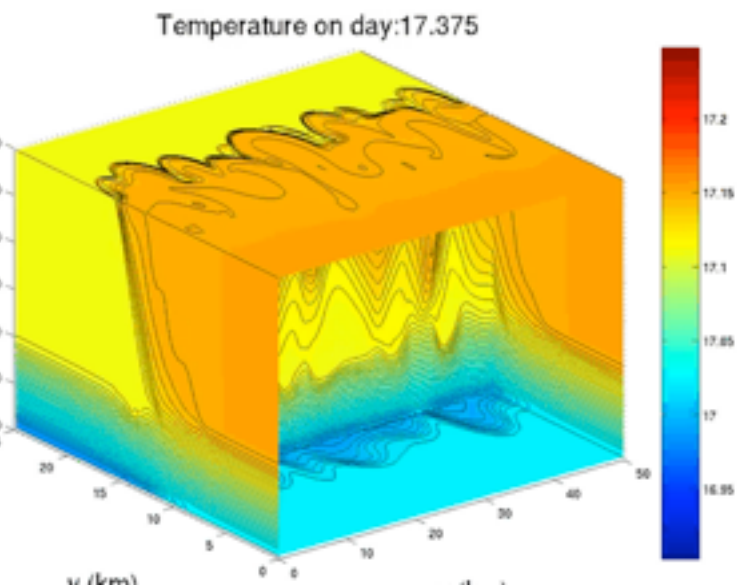
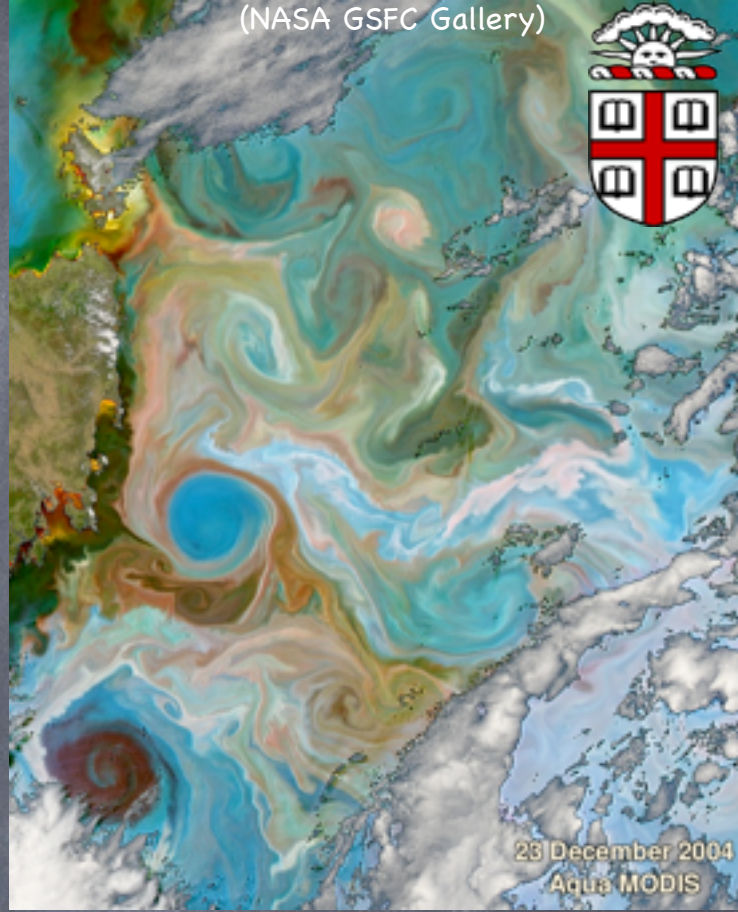
10 km

# The Character of the Submesoscale

(Capet et al., 2008)



- Fronts
- Eddies
- $Ro=O(1)$
- $Ri=O(1)$
- near-surface
- 1-10km, days



Eddy processes often **baroclinic instability**

Parameterizations of submesoscale baroclinic instability?

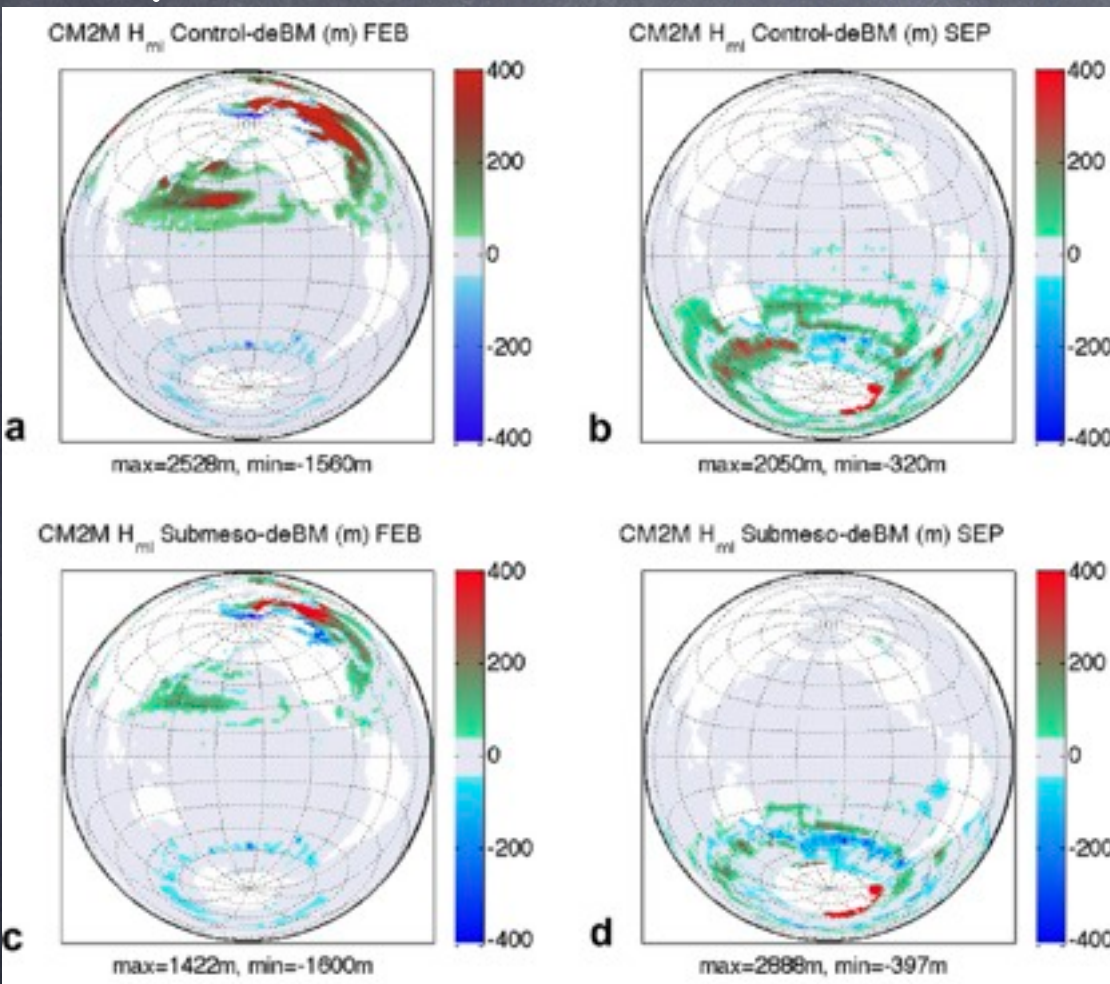
B. Fox-Kemper, R. Ferrari, and R. W. Hallberg. Parameterization of mixed layer eddies. Part I: Theory and diagnosis. *Journal of Physical Oceanography*, 38(6):1145-1165, 2008.

S. Bachman and B. Fox-Kemper. Eddy parameterization challenge suite. I: Eady spindown. *Ocean Modelling*, 2013. In press.

# Physical Sensitivity of Ocean Climate to MLE: **Ocean Climate** to MLE:

## Mixed Layer Eddy Restratification

### Improves Mixed Layers

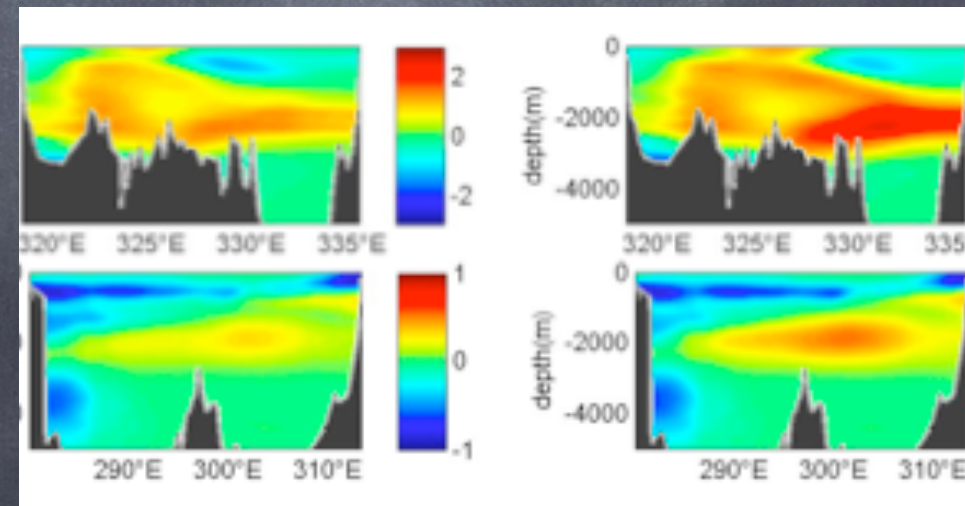


Bias  
w/o  
MLE

Improves CFCs  
(water masses)

Bias with MLE

Bias w/o MLE

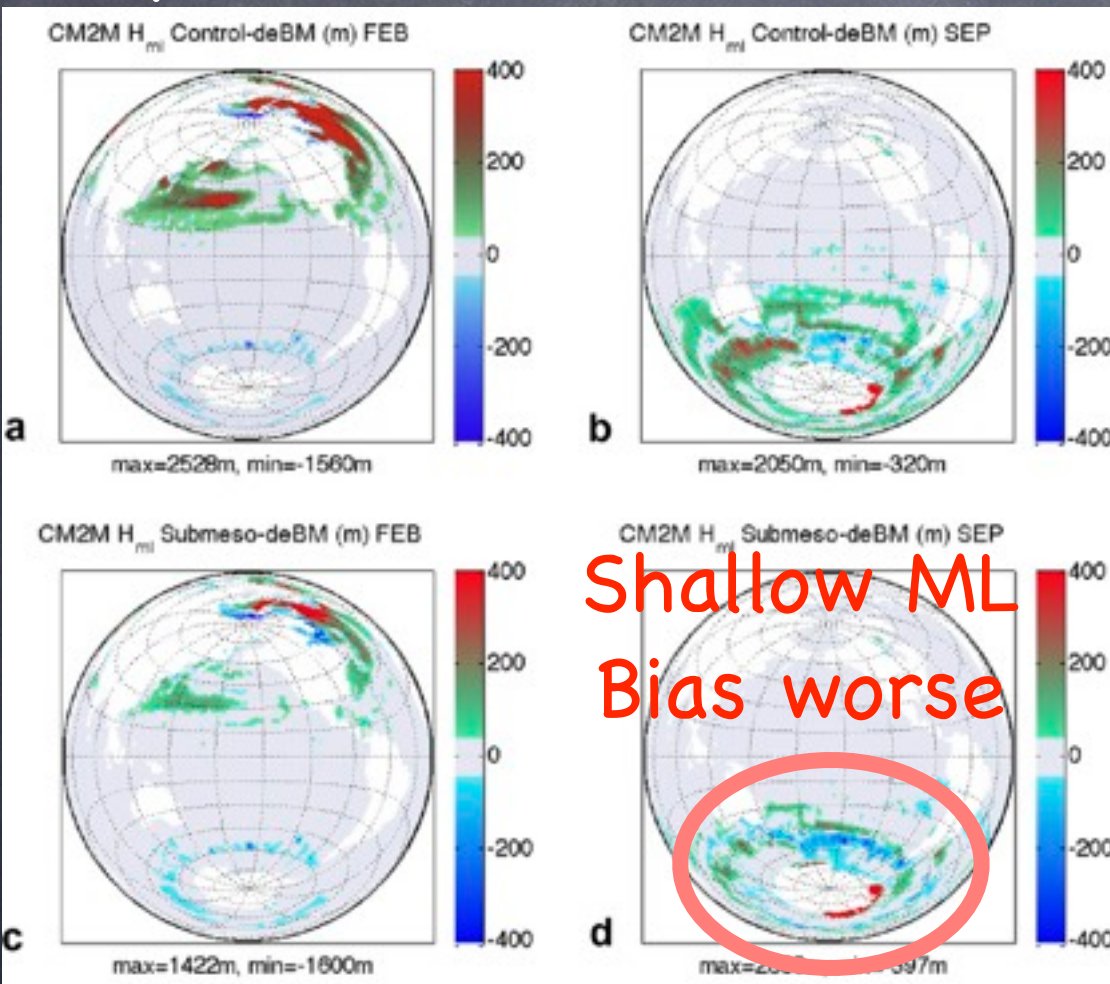


B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg, M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels.  
Parameterization of mixed layer eddies. III: Implementation and impact in global ocean climate simulations. *Ocean Modelling*, 39:61-78, 2011.

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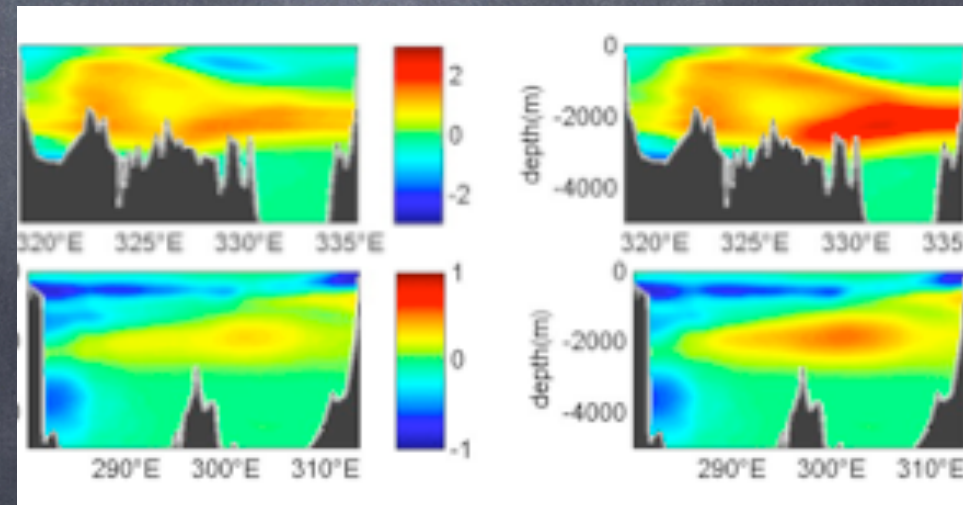


Bias  
w/o  
MLE

Improves CFCs  
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Bias with MLE

Bias w/o MLE



B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg, M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels.  
Parameterization of mixed layer eddies. III: Implementation and impact in global ocean climate simulations. *Ocean Modelling*, 39:61-78, 2011.

# The Character of the Langmuir Scale

- Near-surface
- Langmuir Cells & Langmuir Turb.
- $Ro \gg 1$
- $Ri < 1$ : Nonhydro
- 1-10m
- 10s to mins
- $w, u = O(10\text{cm/s})$
- Stokes drift
- Eqtns: Craik-Leibovich
- Params: McWilliams & Sullivan, 2000, etc.

Image: NPR.org,  
Deep Water  
Horizon Spill



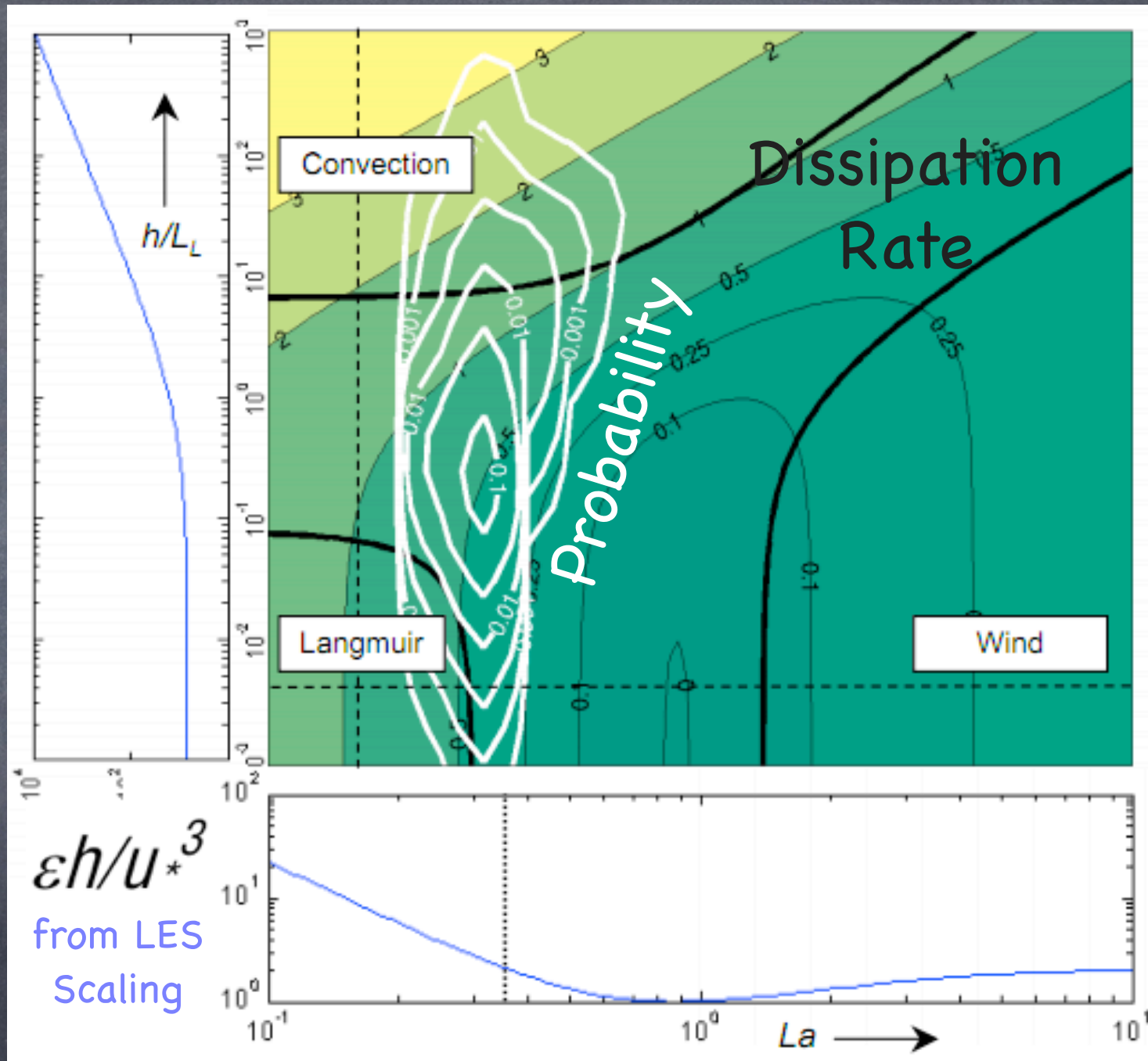


Combining Data w/  
LES scalings:

Where does energy  
for Southern Ocean  
mixing come from?

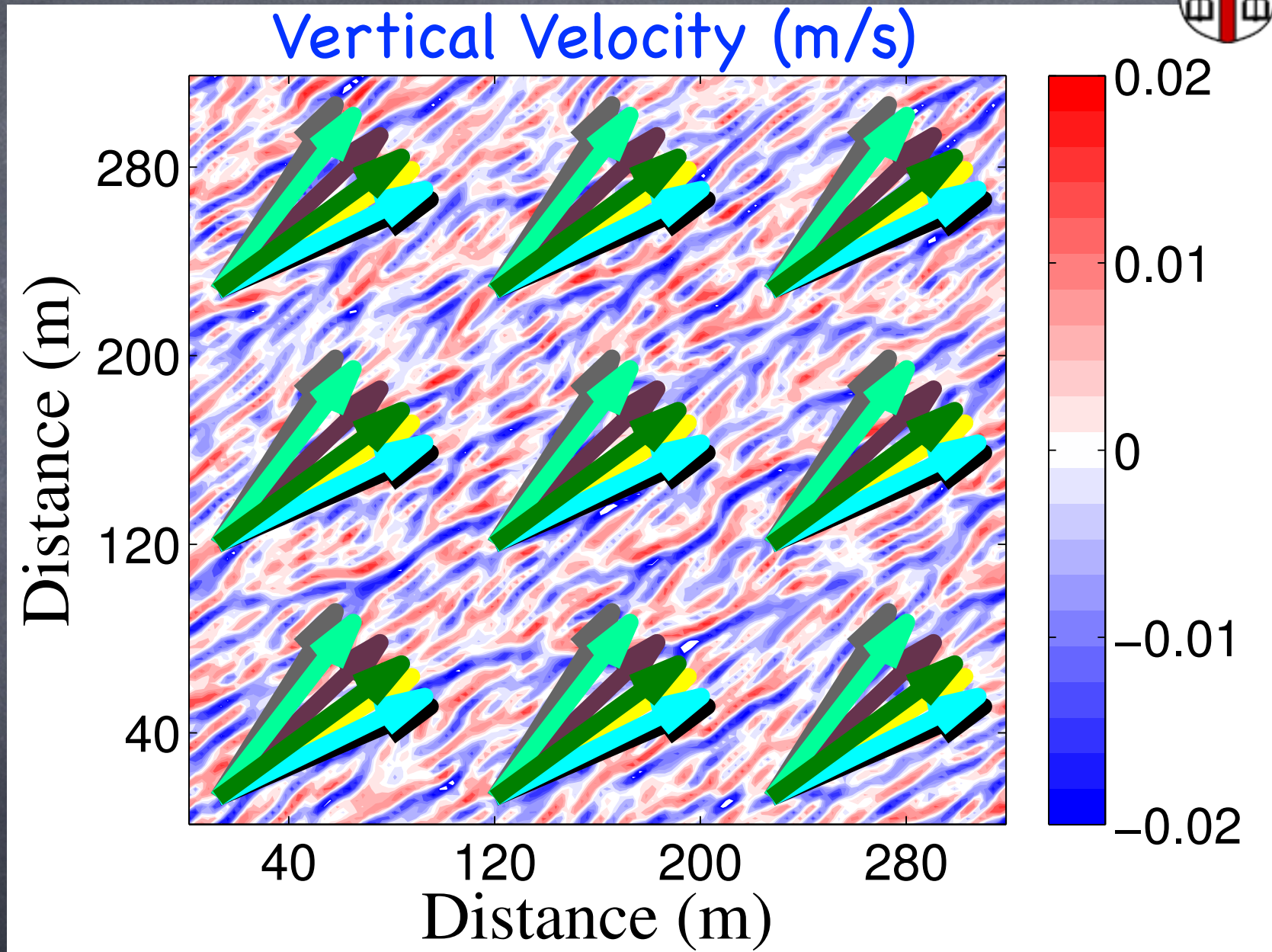
Langmuir  
(wave-driven) and  
Convective

But, we don't have  
wave-driven mixing  
in climate models!!!



S. E. Belcher, A. A. L. M. Grant, K. E. Hanley, B. Fox-Kemper, L. Van Roekel, P. P. Sullivan, W. G. Large, A. Brown, A. Hines, D. Calvert, A. Rutgersson, H. Petterson, J. Bidlot, P. A. E. M. Janssen, and J. A. Polton. A global perspective on Langmuir turbulence in the ocean surface boundary layer. *Geophysical Research Letters*, 39(18):L18605, 9pp, 2012.

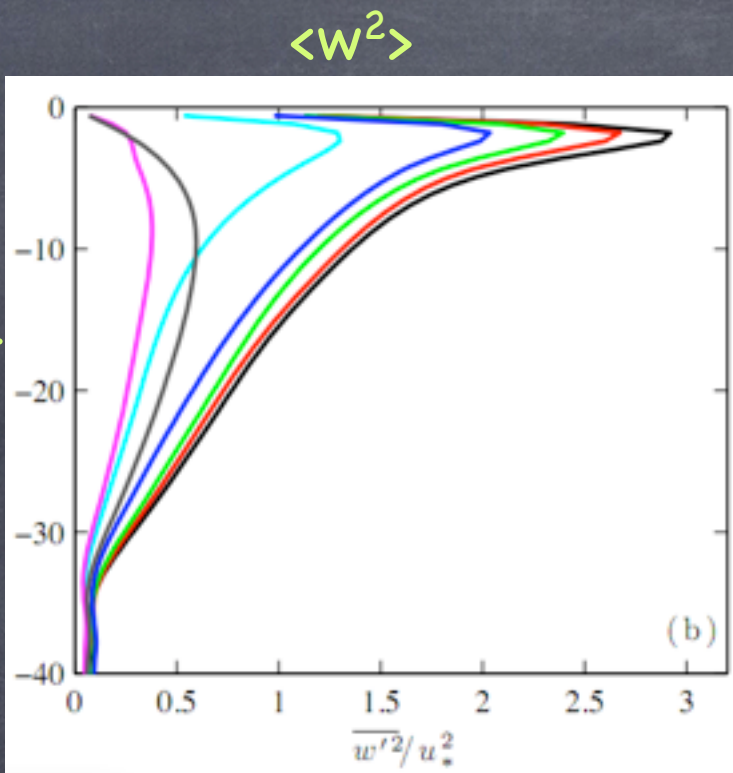
# Real World Forcing: Misaligned Wind & Waves



L. Van Roekel, B. Fox-Kemper, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. *Journal of Geophysical Research-Oceans*, 2012. In press.



depth



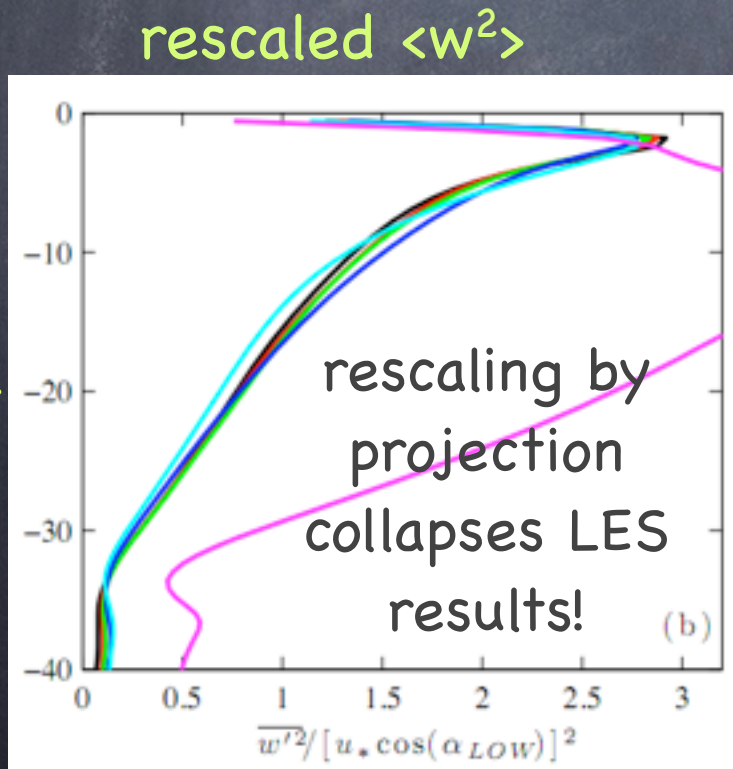
Generalized Turbulent Langmuir No.,  
Projection of  $u^*$ ,  $u_s$  into Langmuir Direction

$$\frac{\langle \overline{w'^2} \rangle_{ML}}{u_*^2} = 0.6 \cos^2(\alpha_{LOW}) [1.0 + (3.1 La_{proj})^{-2} + (5.4 La_{proj})^{-4}],$$

$$La_{proj}^2 = \frac{|u_*| \cos(\alpha_{LOW})}{|u_s| \cos(\theta_{ww} - \alpha_{LOW})},$$

$$\alpha_{LOW} \approx \tan^{-1} \left( \frac{\sin(\theta_{ww})}{\frac{u_*}{u_s(0)\kappa} \ln \left( \left| \frac{H_{ML}}{z_1} \right| \right) + \cos(\theta_{ww})} \right)$$

depth



A scaling for LC strength & direction!

L. P. Van Roekel, B. Fox-Kemper, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. Journal of Geophysical Research-Oceans, 117:C05001, 22pp, 2012.



# Estimated Mixing with Harcourt Second Moment Closure Model:

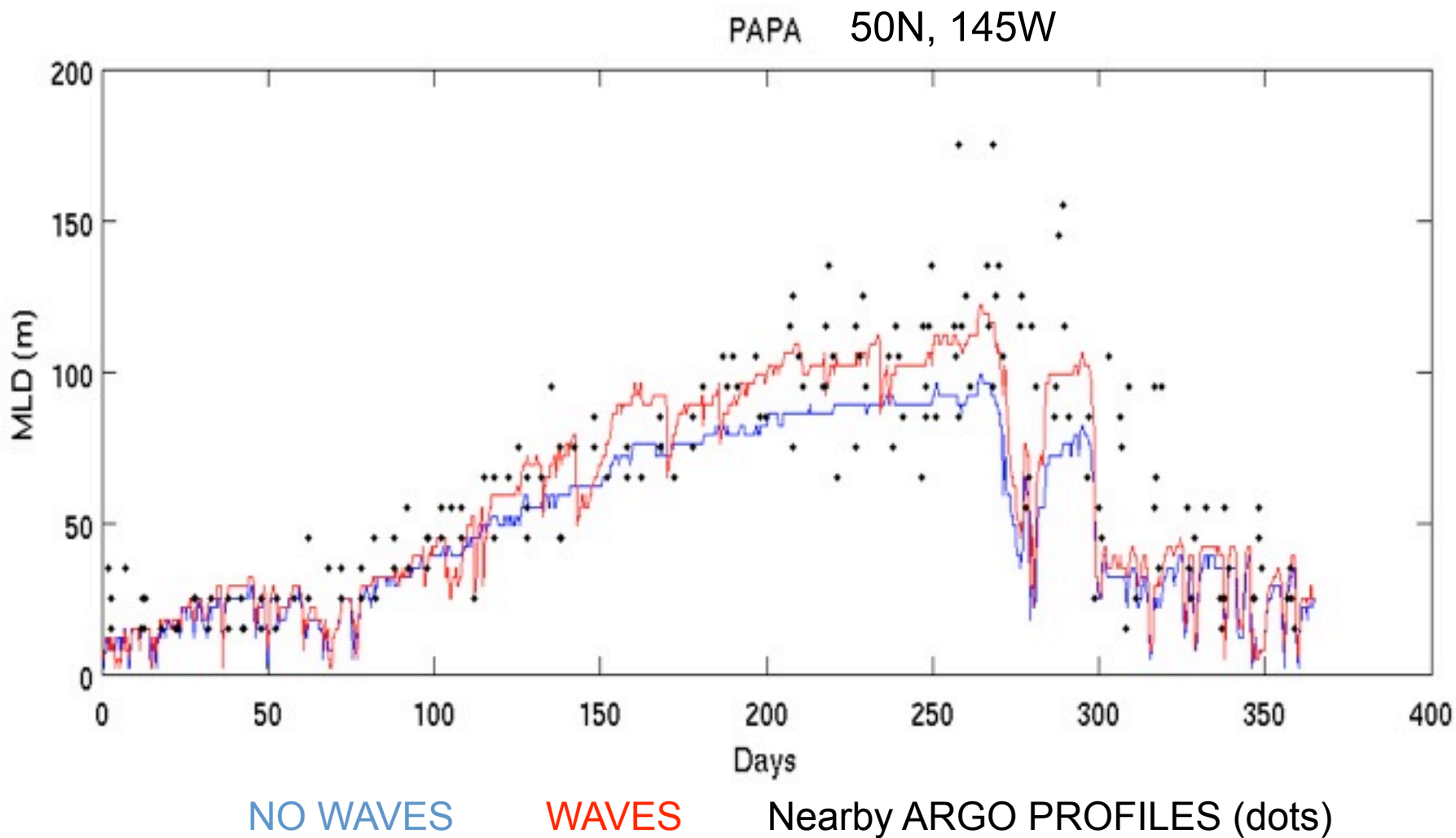
- Expands on implementation of Langmuir mixing of Kantha & Clayson (and Mellor–Yamada, etc).
- Allows waves to affect diffusivity and vertical momentum flux as well as energy sources...
- Consistent treatment of wave–wind misalignment
- Was run for 1-year, using Large & Yeager (2004) “Normal Year” and matching WaveWatch–III–simulated waves.
- ARGO profiles near summer solstice initial condition.
- Pure 1D mixing vs. solar & fluxes only -- No ocean circulation or eddy restratification, etc.

R. R. Harcourt. A second moment closure model of Langmuir turbulence. *Journal of Physical Oceanography*, 2013, in press.

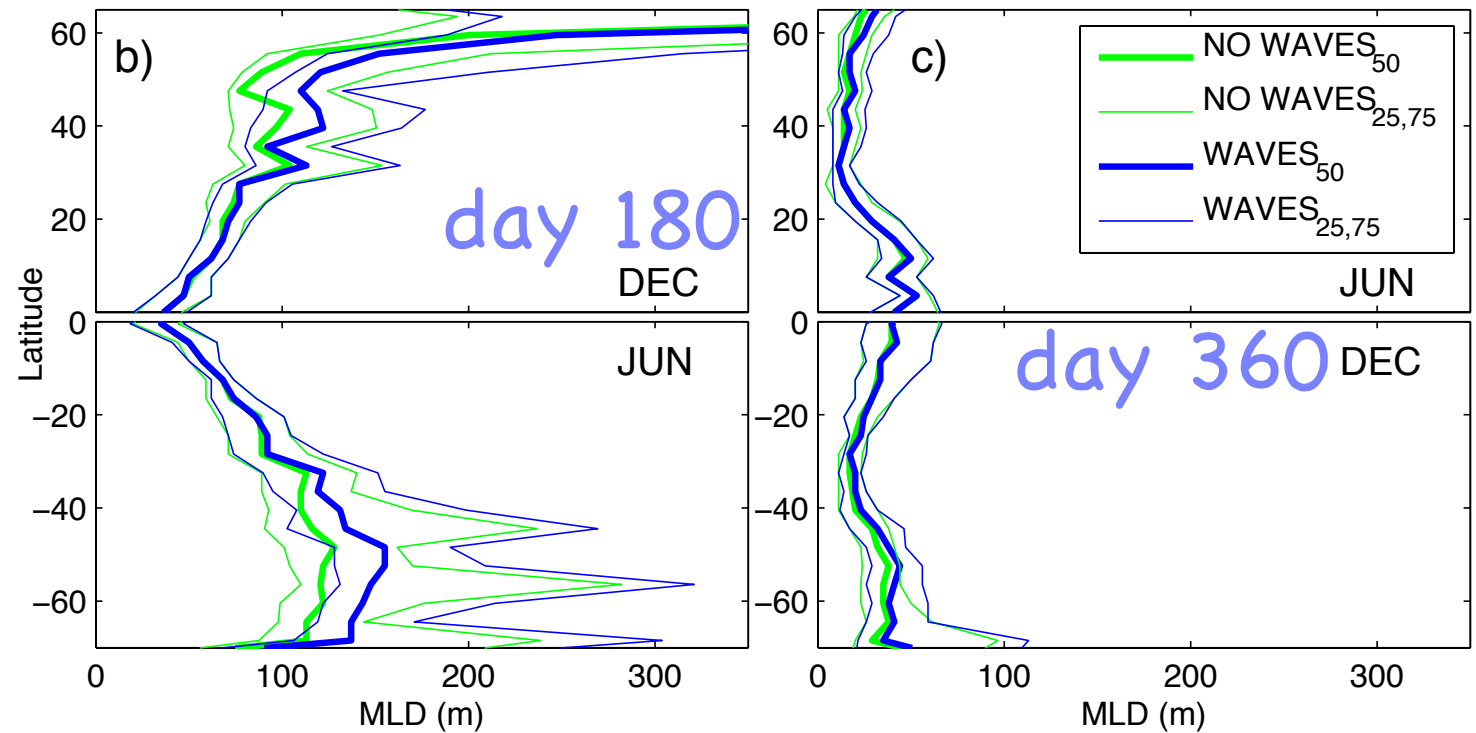
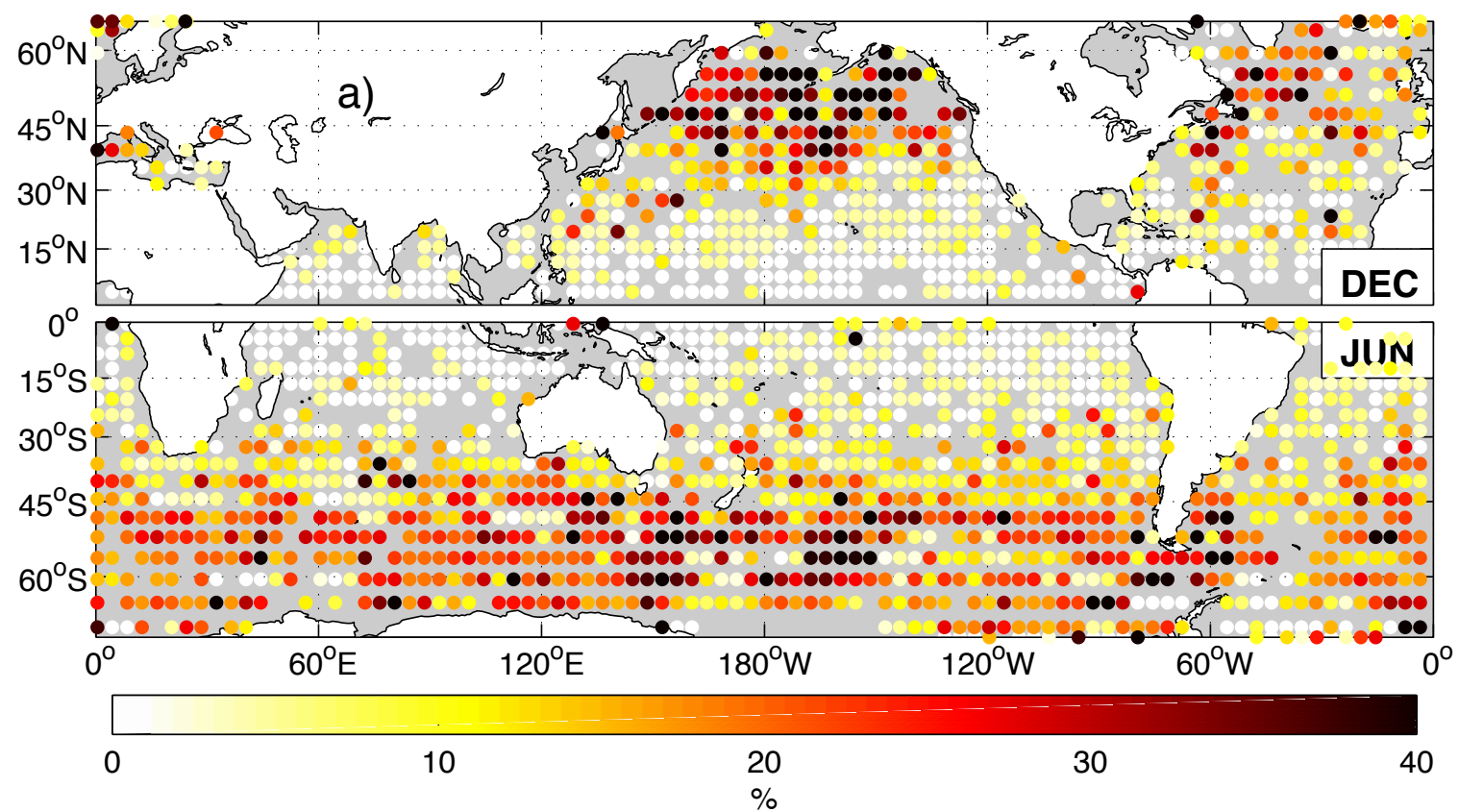
# Harcourt--Evolving Mixed Layer Depth Results (Simulated by M. Hemer, CSIRO)



Potential for improvement! (not all look so good...)



Including  
Wave-driven  
Mixing  
Deepens the  
Wintertime  
and S. Ocean  
Mixed Layer!



M. A. Hemer, B. Fox-Kemper,  
& R. R. Harcourt. Quantifying  
the effects of wind waves on  
the coupled climate system, in  
prep. 2012.

Well that was fun...



Lots of small turbulence  
affects the upper ocean on  
seasonal timescales!

But, why stop now?

How do these small processes  
affect each other?

# What about Langmuir-Submeso Interactions?

Movie: P.  
Hamlington



Perform large eddy simulations (LES) of  
Langmuir turbulence with a  
submesoscale temperature front

Use NCAR LES model to solve Craik-  
Leibovich equations (Moeng, 1984,  
McWilliams et al, 1997)

$$\frac{\partial \rho}{\partial t} + \mathbf{u}_L \cdot \nabla \rho = \text{SGS}$$

$$\nabla \cdot \mathbf{u} = 0$$

$$\frac{\partial \mathbf{u}}{\partial t} + (\boldsymbol{\omega} + f\hat{\mathbf{z}}) \times \mathbf{u}_L = -\nabla \pi - \frac{g\rho\hat{\mathbf{z}}}{\rho_0} + \text{SGS}$$

Computational parameters:

Domain size: 20km x 20km x -160m

Grid points: 4096 x 4096 x 128

Resolution: 5m x 5m x -1.25m



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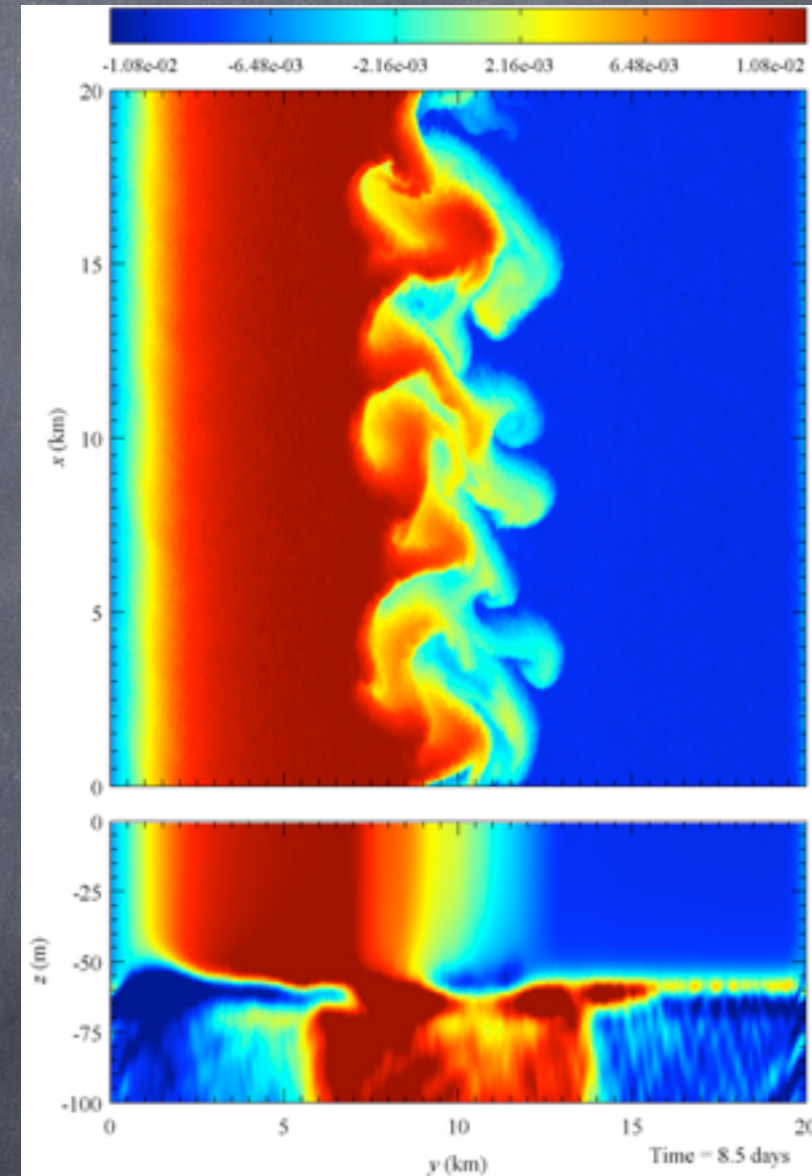
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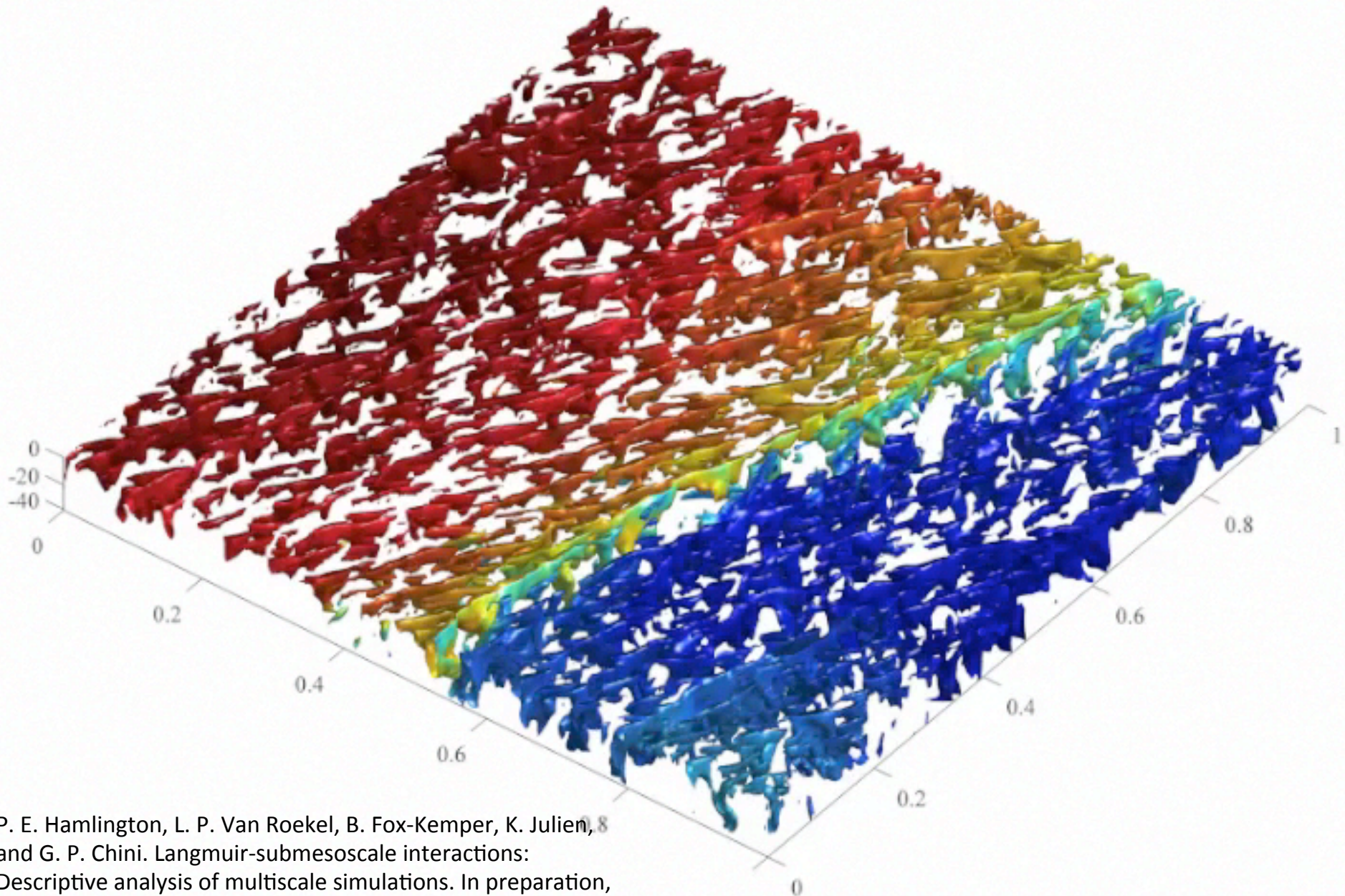
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Movie: P.  
Hamlington



P. E. Hamlington, L. P. Van Roekel, B. Fox-Kemper, K. Julien, and G. P. Chini. Langmuir-submesoscale interactions: Descriptive analysis of multiscale simulations. In preparation, 2013.

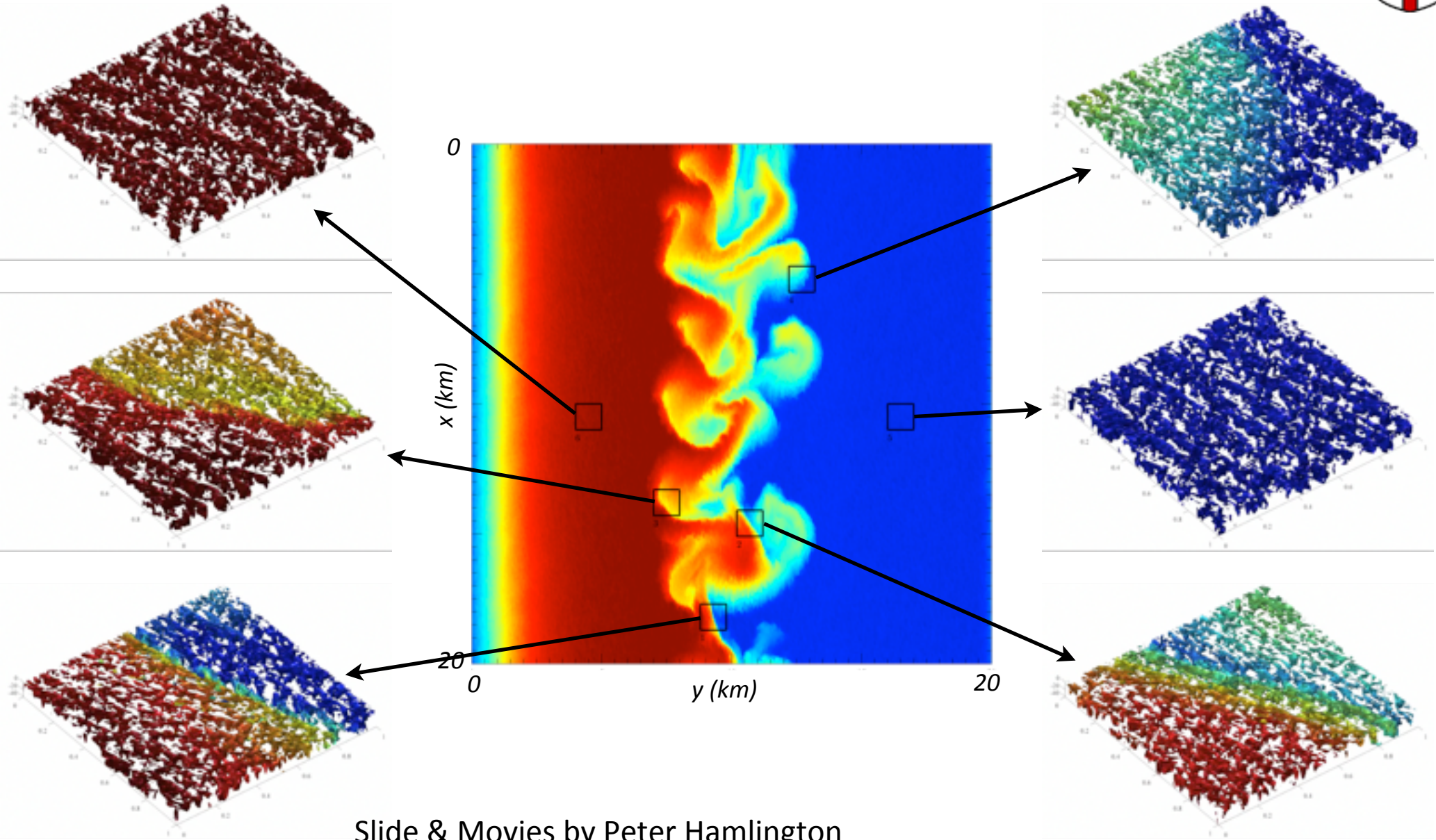
# Zoom: Submeso-Langmuir Interaction!



P. E. Hamlington, L. P. Van Roekel, B. Fox-Kemper, K. Julien,  
and G. P. Chini. Langmuir-submesoscale interactions:  
Descriptive analysis of multiscale simulations. In preparation,



# Diverse types of interaction



Slide & Movies by Peter Hamlington



But, why stop now?

How do these small processes  
affect each other?

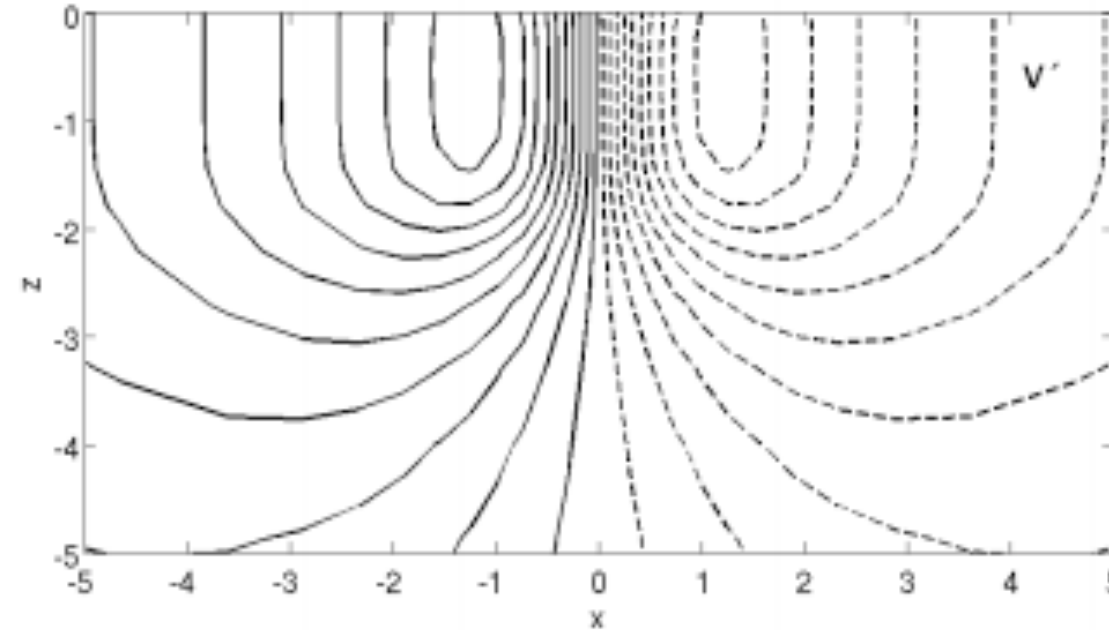
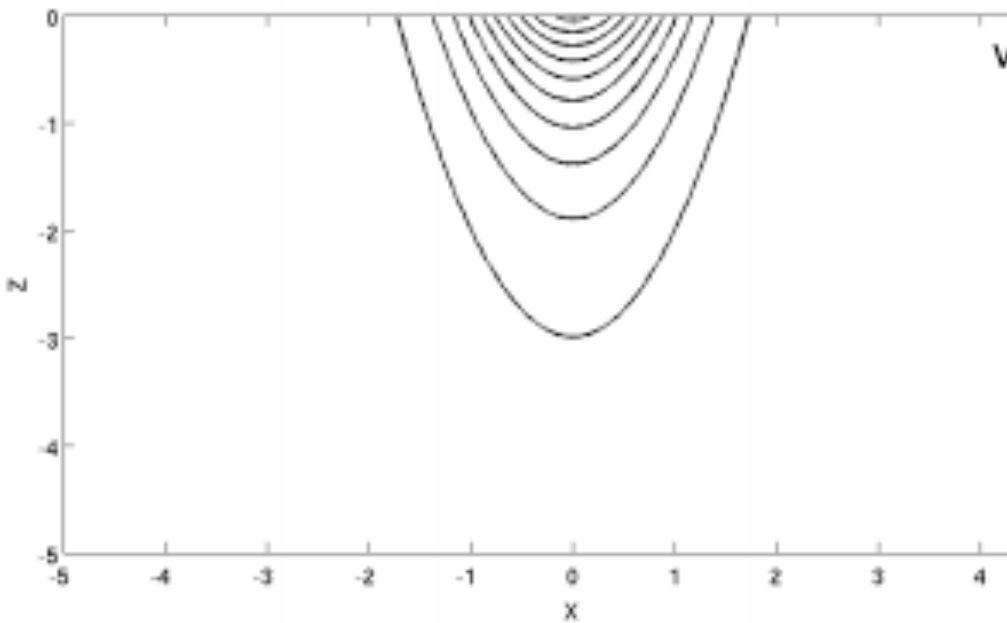
Weakly, except near strong  
fronts... stay tuned...

OK, Langmuir Turbulence  
can interact with larger  
scales...

What about the surface  
waves themselves, do  
they interact?



# Waves (Stokes Drift Vortex Force) → Submeso, Meso **Balanced Flows Change!**



Initial Submeso Front

Contours: 0.1

Perturbation on that scale  
due to waves (Stokes Force)

Contours: 0.014



Many more wave-climate effects to come... stay tuned!

FREE ATMOSPHERE

ATMOSPHERIC BOUNDARY LAYER

OCEANIC BOUNDARY LAYER

Processes: 2012

Solar & Longwave Radiation



Sensible and Latent Heat

Marine Aerosol Production

Sea-state dependent drag



Kinetic Energy Fluxes



Heat Fluxes

Mass Fluxes



Whitecapping

Bubble injection by breaking waves

Albedo



MIZ Extent

Stokes drift

Wave driven and Langmuir turbulence

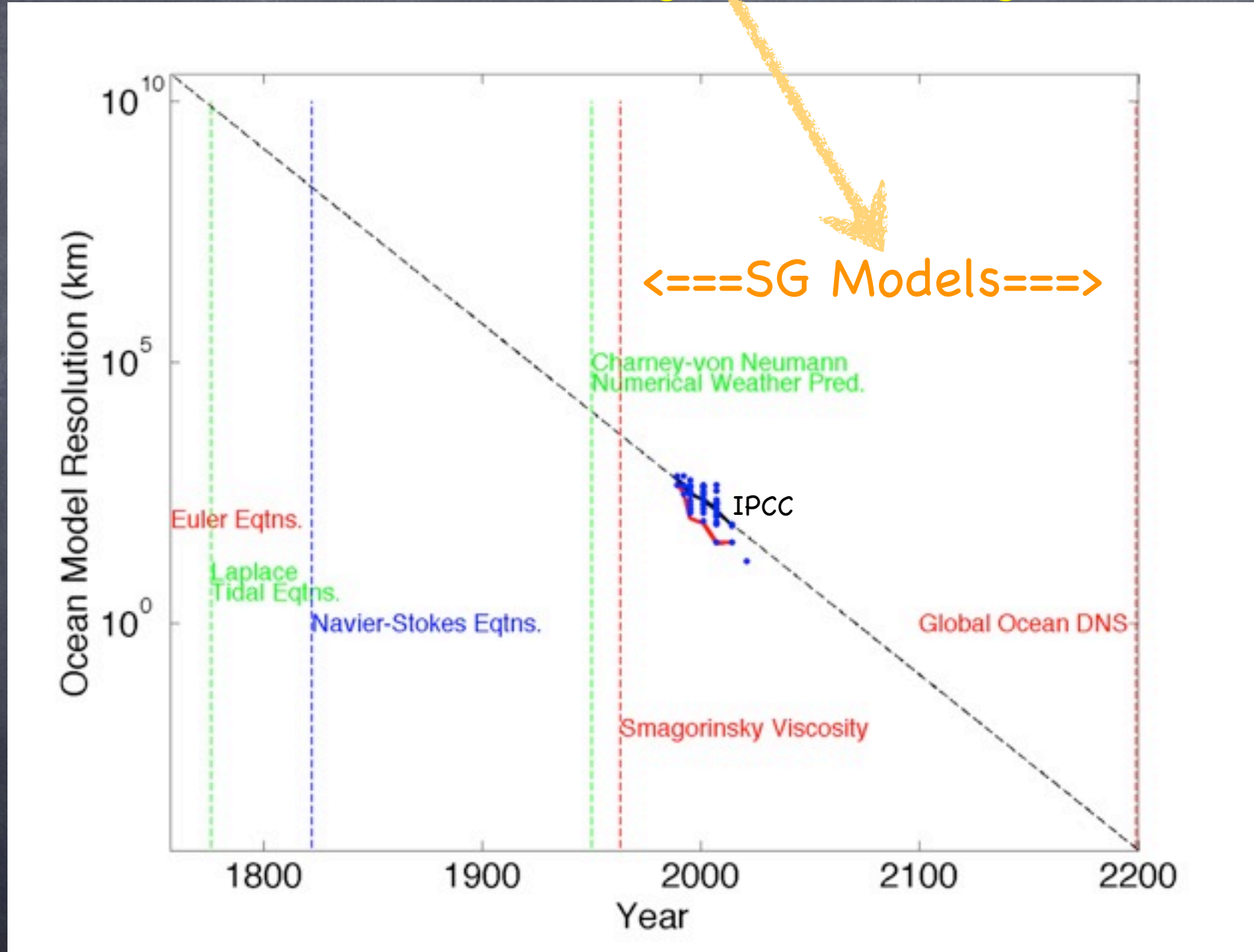
Wind-wave dependent processes in the coupled climate system  
Towards coupled wind-wave-AOGCM models

# Conclusions

- Climate modeling is challenging partly due to the vast and diverse scales of fluid motions
- In the upper ocean, horizontal scales as big as basins, and as small as meters contribute non-negligibly to the air-sea exchange
- Process models, especially those spanning a whole or multiple scales, are a powerful tool in studying these connections and improving subgrid models.
- Based on present rates of increase of computing power, we will need these subgrid models for at least another century!



# Extrapolate for historical perspective: The Golden Era of Subgrid Modeling is Now!



All papers at: [fox-kemper.com/research](http://fox-kemper.com/research)