

An aerial photograph of a large body of water, possibly a lake or a wide river. The water is a deep blue color with some white foam or ice visible. In the center of the image, there is a small, bright white object that appears to be a small boat or a piece of ice. The background is a light blue sky with some faint clouds.

Baylor Fox-Kemper (Brown University)

with Scott Bachman (DAMTP)

ECCO Workshop

Jan 22-24, 2014

Sponsors: NSF 1258907, 1245944, 0934737, 0825614, NASA NNX09AF38G

Scale-aware subgrid closures for models that partly resolve the mesoscale and submesoscale

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with Scott Bachman (DAMTP)

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A Global Parameterization of Mixed Layer Eddy Flow & Scale Aware Restratification validated against simulations

B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg, M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels. Parameterization of mixed layer eddies. III: Implementation and impact in global ocean climate simulations. *Ocean Modelling*, 39:61-78, 2011.

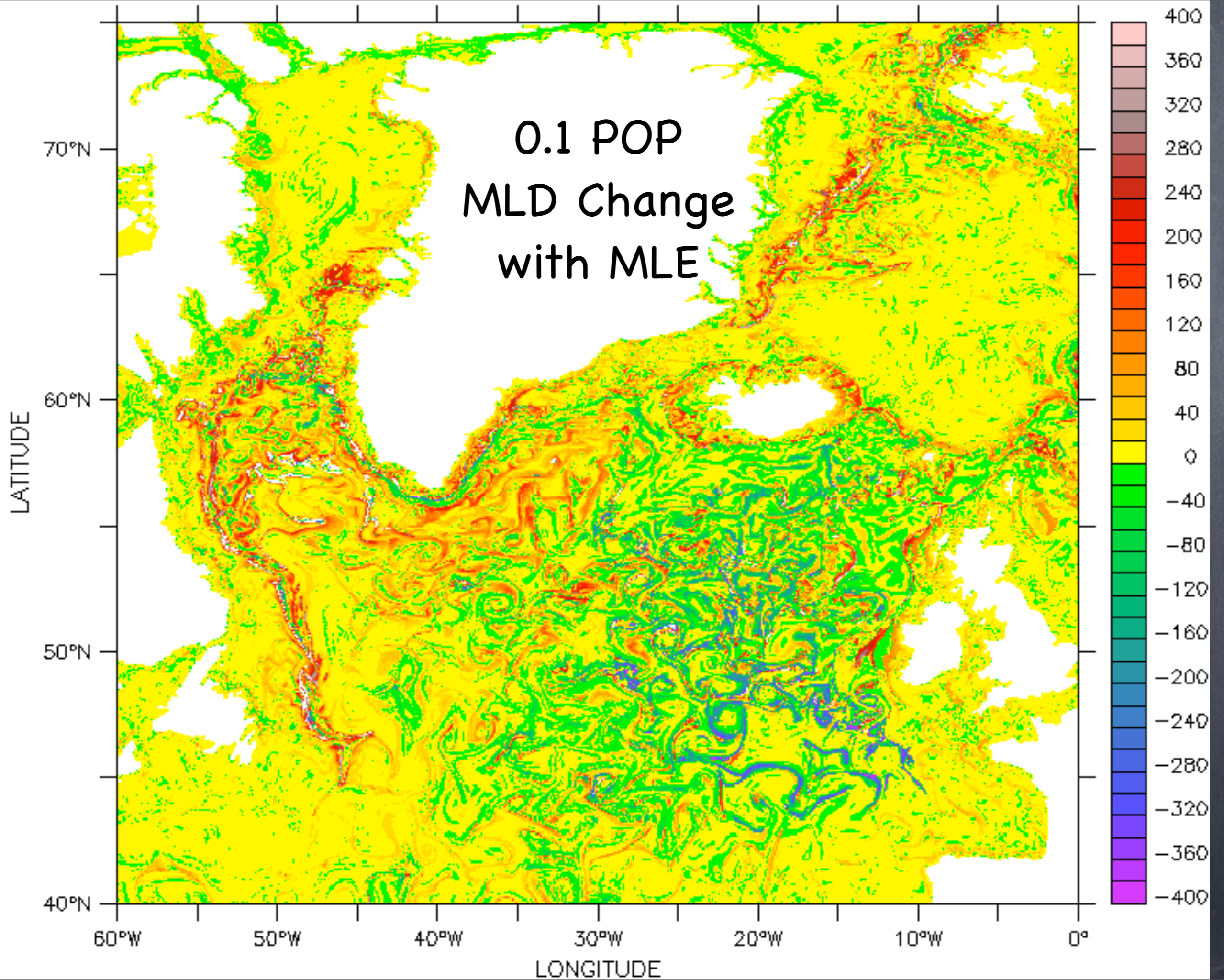
$$\overline{\mathbf{u}'b'} \equiv \Psi \times \nabla \bar{b}$$

$$\Psi = \left[\begin{array}{c} \Delta x \\ L_f \end{array} \right] \frac{C_e H^2 \mu(z)}{\sqrt{f^2 + \tau^{-2}}} \nabla \bar{b} \times \hat{\mathbf{z}}$$

Compare to the original **singular, unrescaled** version

$$\Psi = \left| \frac{C_e H^2 \mu(z)}{|f|} \nabla \bar{b} \times \hat{\mathbf{z}} \right.$$

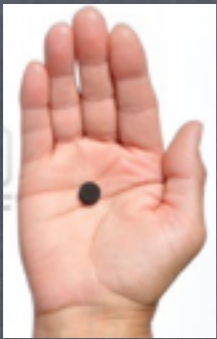
New version **handles the equator**, and **averages over many fronts**



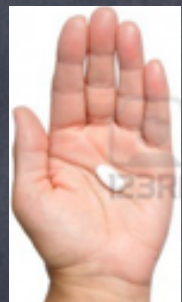
Submeso closure is already scale-aware!

- You can use the global/gcm form at any resolution.
- It turns itself off as the ML def. radius becomes resolved.

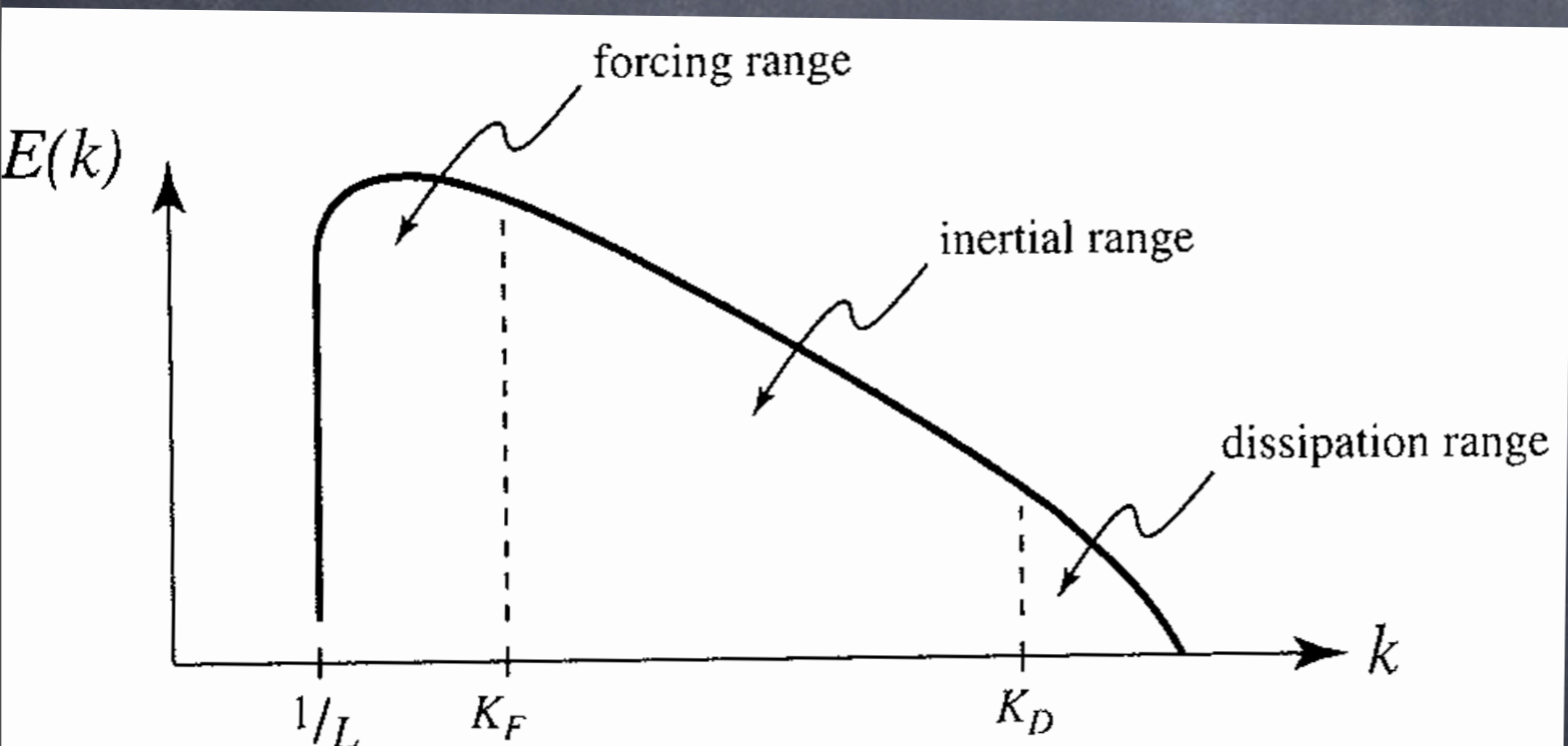
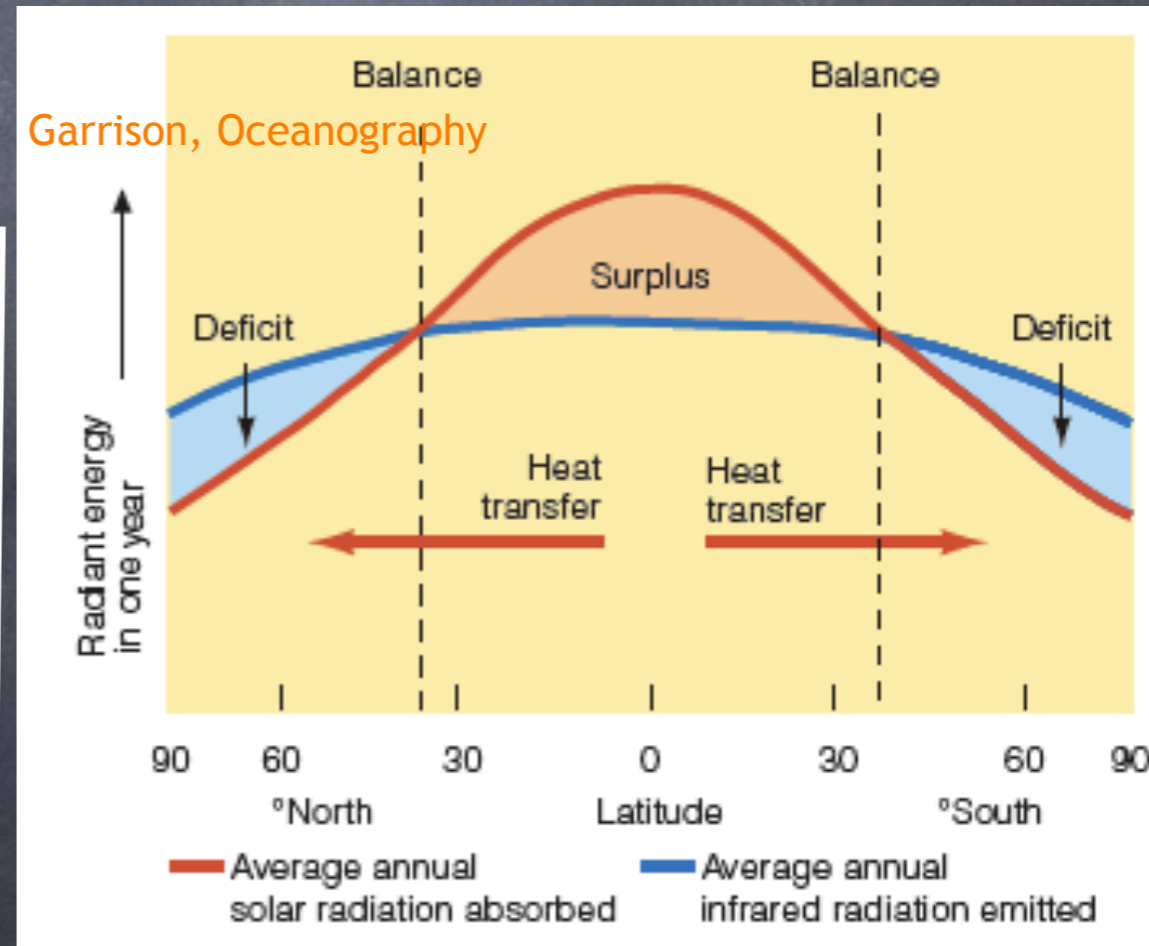
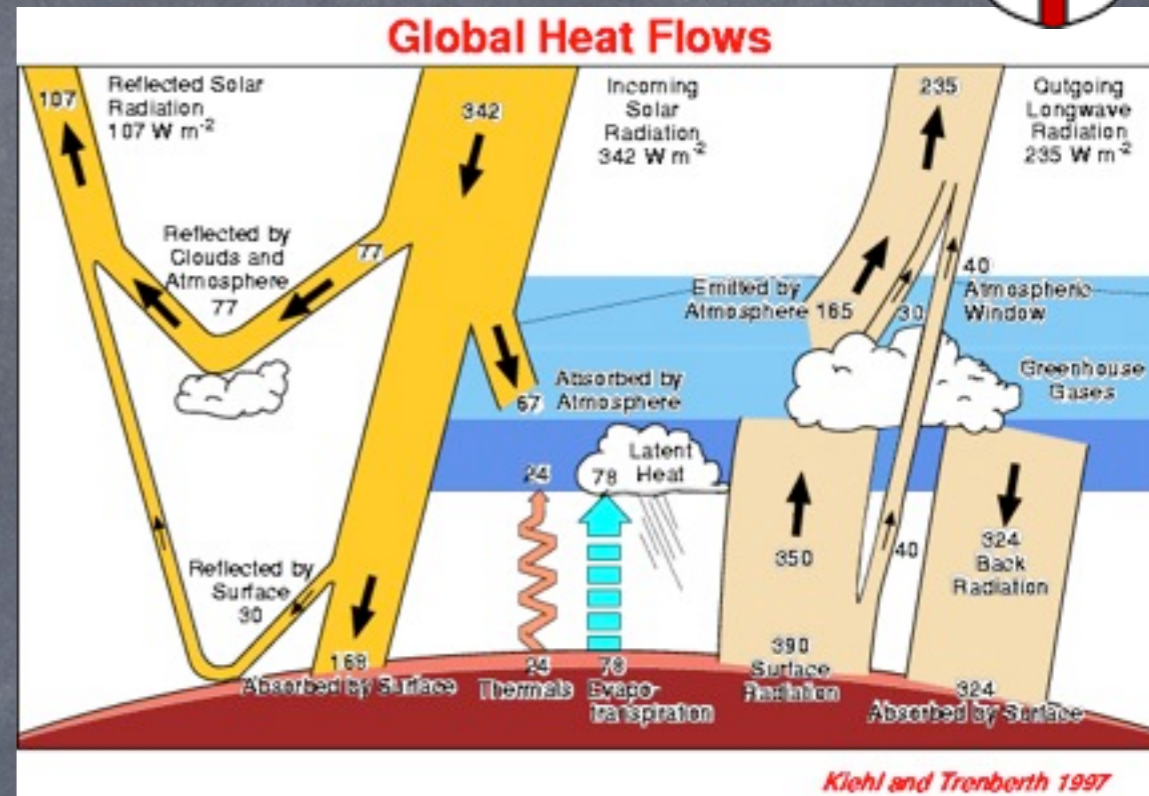
$$\Psi = \left[\frac{\Delta x}{L_f} \right] \frac{C_e H^2 \mu(z)}{\sqrt{f^2 + \tau^{-2}}} \nabla \bar{b} \times \hat{\mathbf{z}}$$



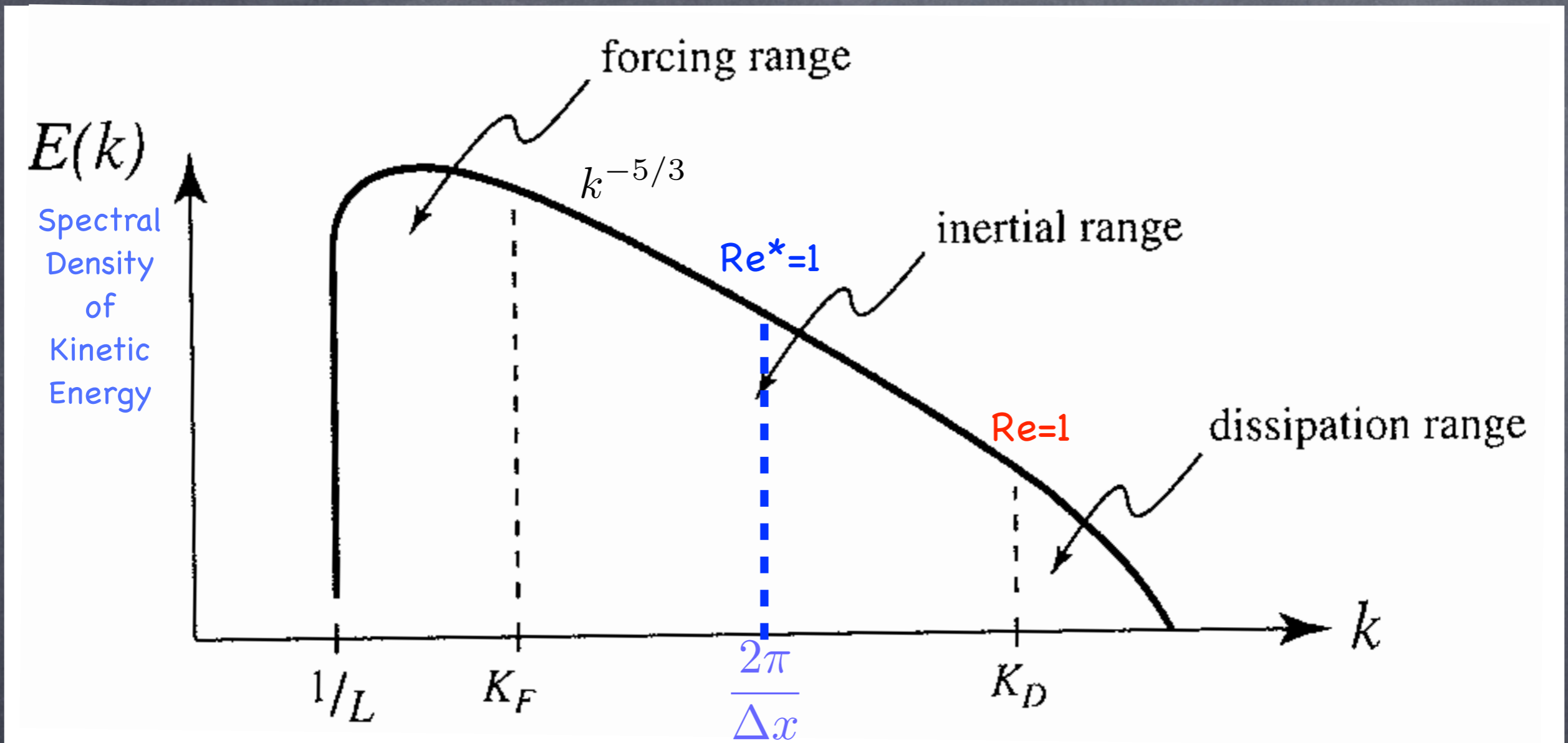
The Earth's Climate System is driven by the Sun's light (minus outgoing infrared) on a global scale



Dissipation concludes turbulence cascades to scales about a billion times smaller



3D Turbulence Cascade

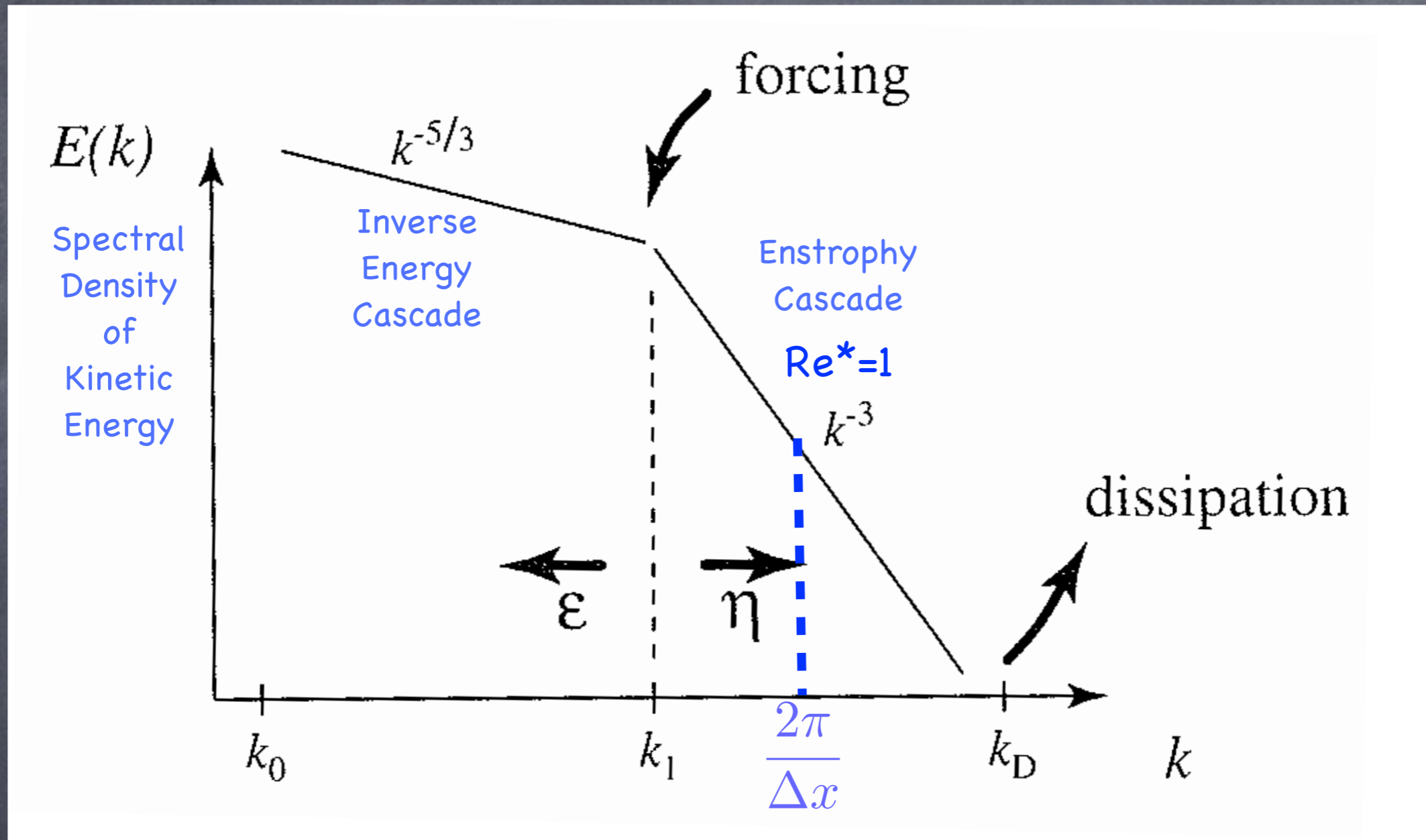


1963: Smagorinsky Scale & Flow Aware Viscosity Scaling,
 So the Energy Cascade is Preserved,
 but order-1 gridscale Reynolds #: $Re^* = UL/\nu_*$

$$\mathbf{v}_{*h} = \left(\frac{\Upsilon_h \Delta x}{\pi} \right)^2 \sqrt{\left(\frac{\partial u_*}{\partial x} - \frac{\partial v_*}{\partial y} \right)^2 + \left(\frac{\partial u_*}{\partial y} + \frac{\partial v_*}{\partial x} \right)^2}$$

2D Turbulence Differs

R. Kraichnan, 1967 JFM



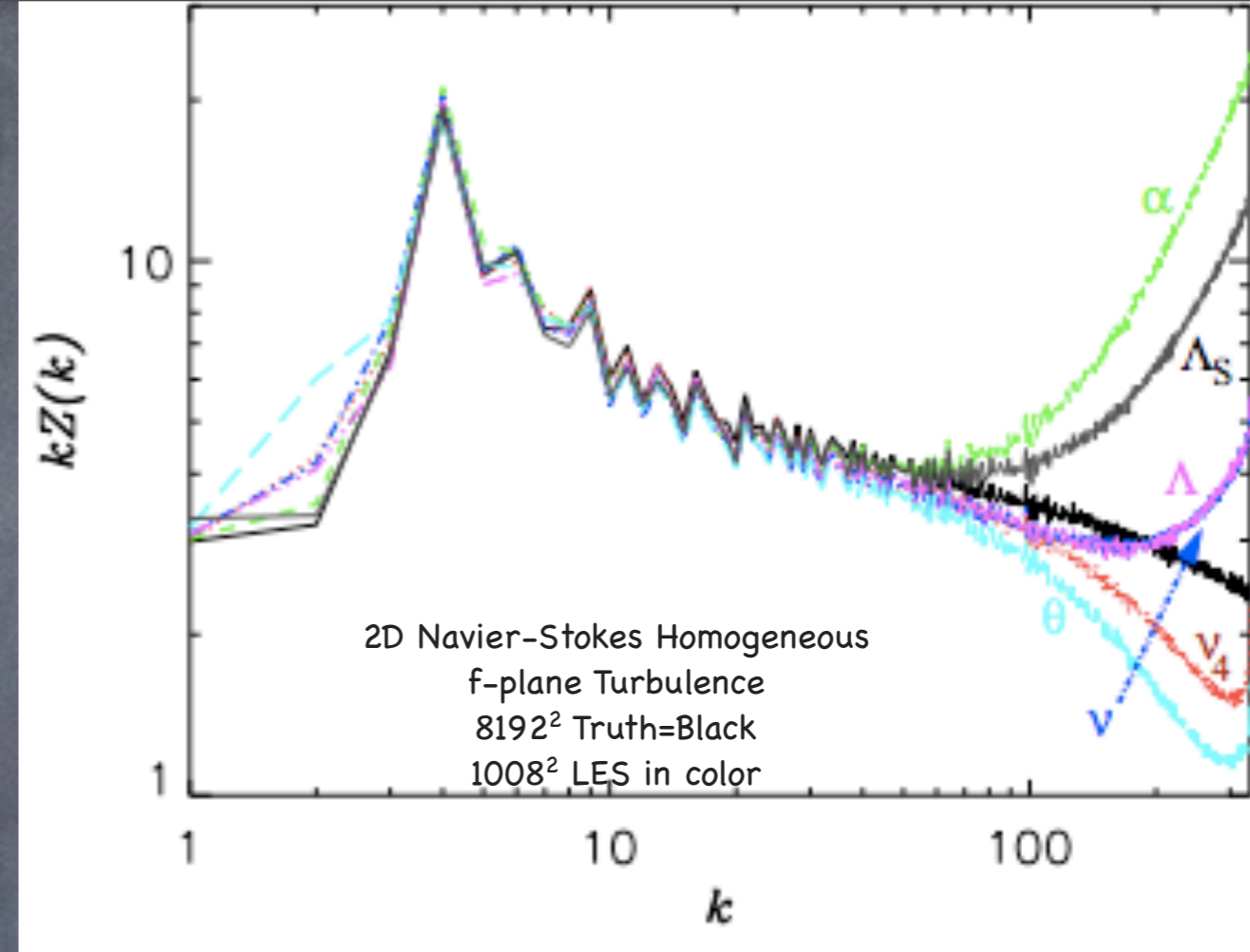
1996: Leith Devises Viscosity Scaling,
So that the Enstrophy (vorticity²) Cascade is Preserved

$$\mathbf{v}_* = \left(\frac{\Lambda \Delta x}{\pi} \right)^3 \left| \nabla_h \left(\frac{\partial u_*}{\partial y} - \frac{\partial v_*}{\partial x} \right) \right|$$

Barotropic or
stacked layers

Some MOLES Truncation Methods In Use 2d (SWE) test

- Harmonic/Biharmonic/Numerical
 - Many. Often not scale- or flow-aware
 - Griffies & Hallberg, 2000, is one aware example
- Fox-Kemper & Menemenlis, 2008. ECCO2.
 - Leith Viscosity (2d Enstrophy Scaling)
- Chen, Q., Gunzburger, M., Ringler, T., 2011
 - Anticipated Potential Vorticity of Sadourny
- San, Staples, Iliescu (2011, 2013)
 - Approximate Deconvolution Method
- Stochastic & Statistical Parameterizations
 - Other session going on now in Y10



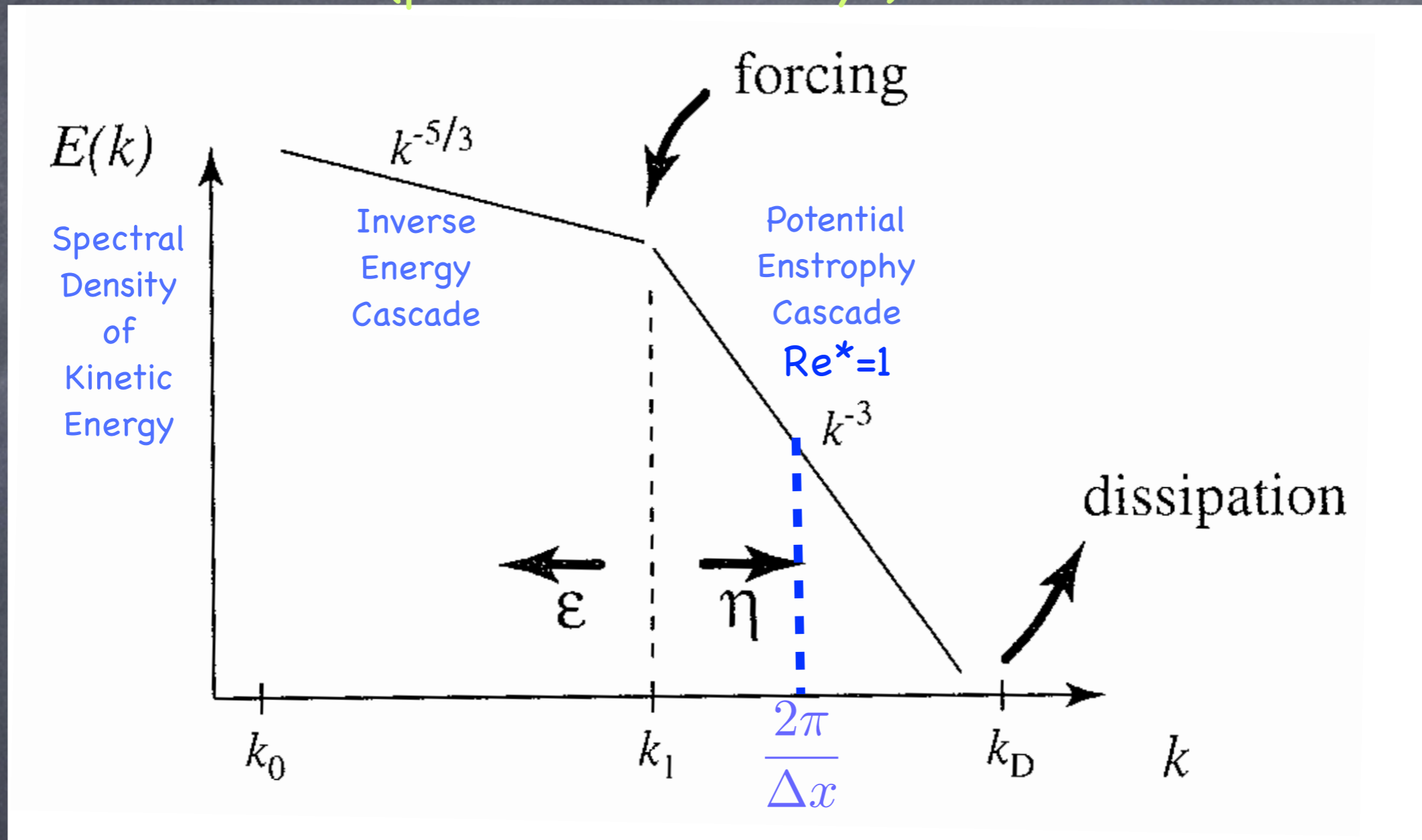
Graham & Ringler, 2013 Ocean Modelling

In this comparison,
untuned Leith beats:
tuned harmonic,
tuned biharmonic,
Smagorinsky,
LANS-alpha, &
Anticipated PV

See also Ramachandran et al, 2013
Ocean Modelling for SMOLES

QG Turbulence: Pot'l Enstrophy cascade (potential vorticity²)

J. Charney, 1971 JAS



**F-K & Menemenlis '08: Revise Leith Viscosity Scaling,
So that diverging, vorticity-free, modes are also damped**

$$\mathbf{v}_* = \left(\frac{\Delta x}{\pi}\right)^3 \sqrt{\Lambda^6 |\nabla_h q_{2d}|^2 + \Lambda_d^6 |\nabla_h (\nabla_h \cdot \mathbf{u}_*)|^2}$$

B. Fox-Kemper and D. Menemenlis. Can large eddy simulation techniques improve mesoscale-rich ocean models? In M. Hecht and H. Hasumi, editors, Ocean Modeling in an Eddying Regime, volume 177, pages 319-338. AGU Geophysical Monograph Series, 2008.

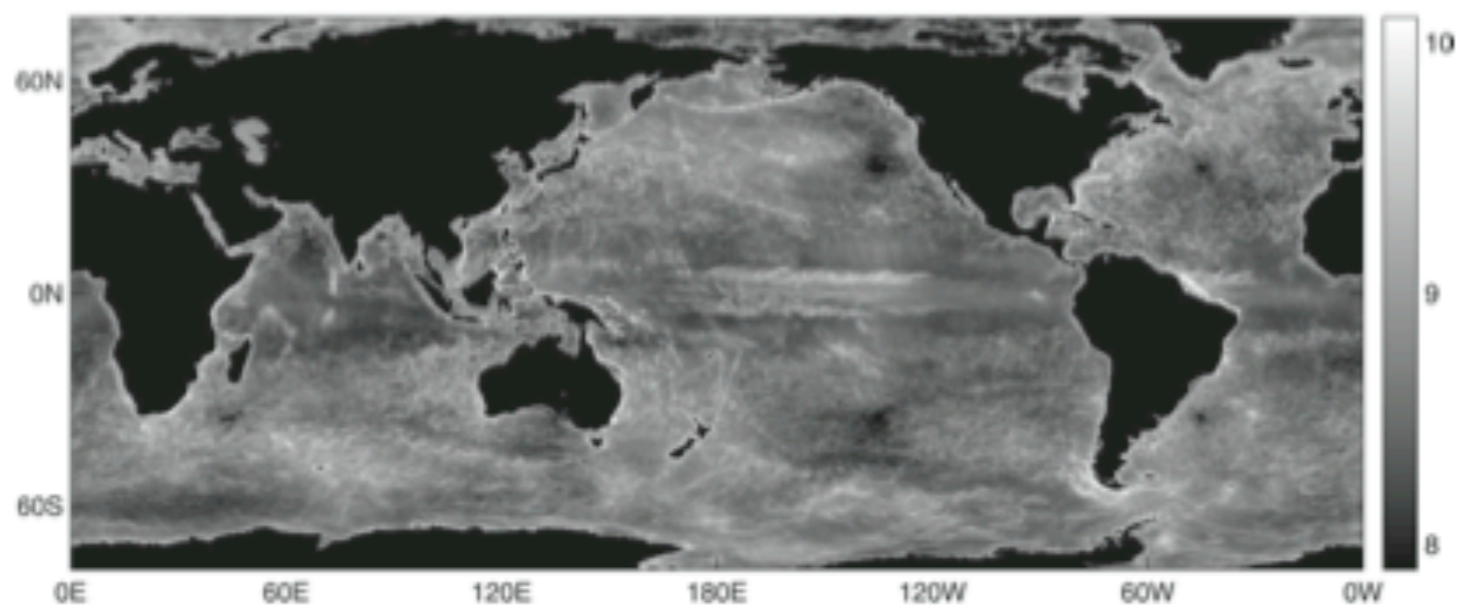
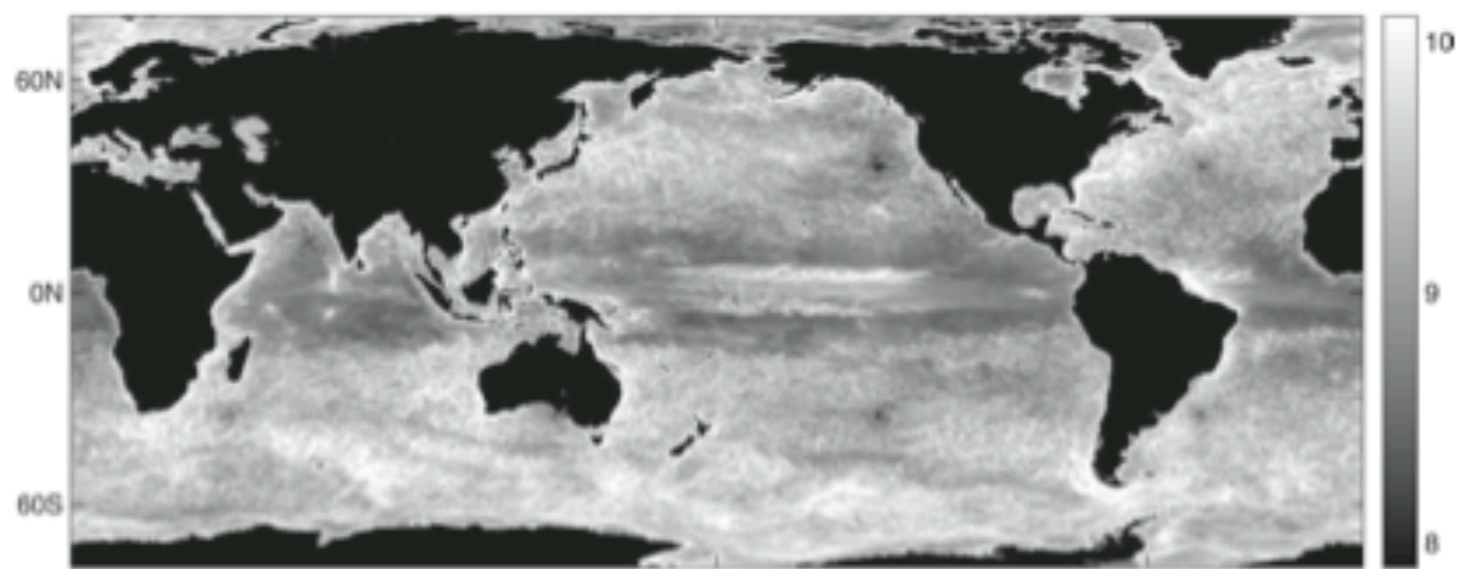
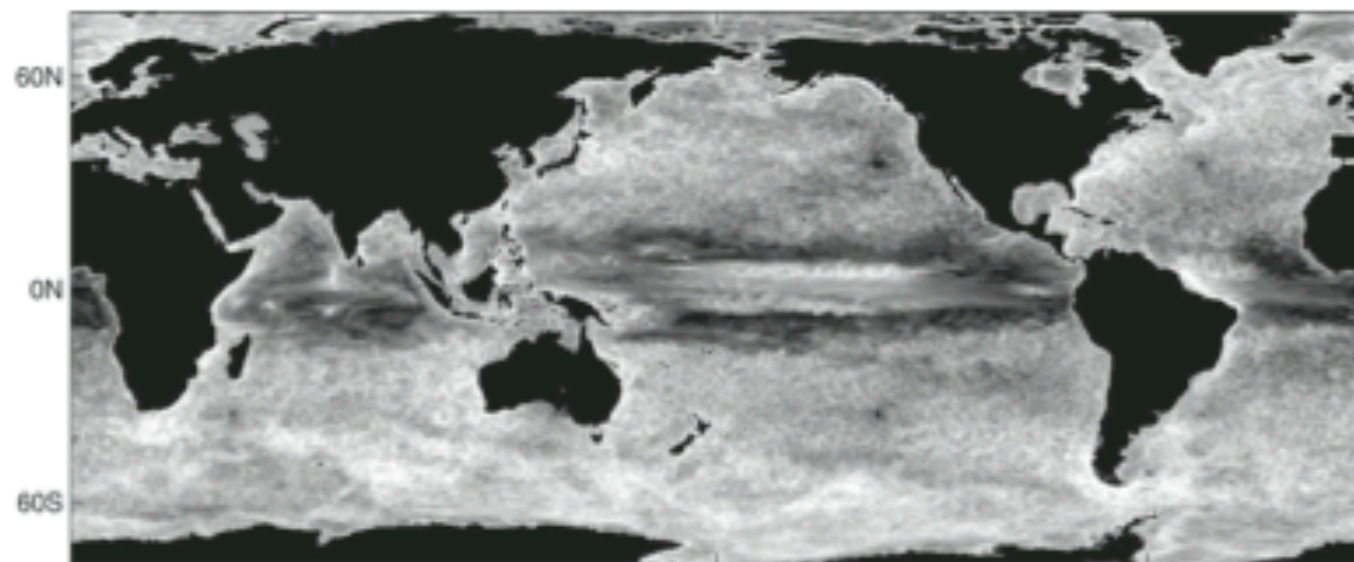


Figure 3. Monthly mean biharmonic viscosity, ν_b , in the model's surface level for December 2001. Units are $m^4 s^{-1}$ and color scale displays $\log_{10}(\nu_b)$. Top panel is from the *LeithOnly* integration. Middle panel is from the *LeithPlus* integration. Bottom panel shows the divergent modification of the *LeithPlus* integration.

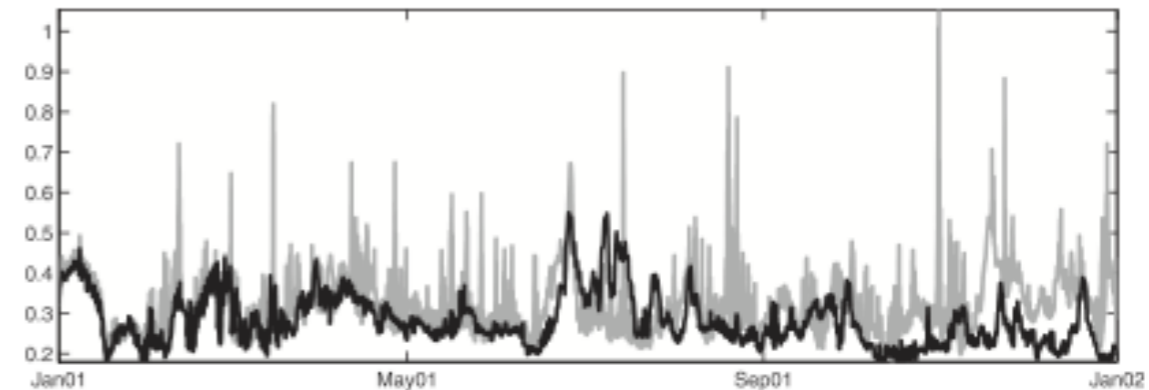
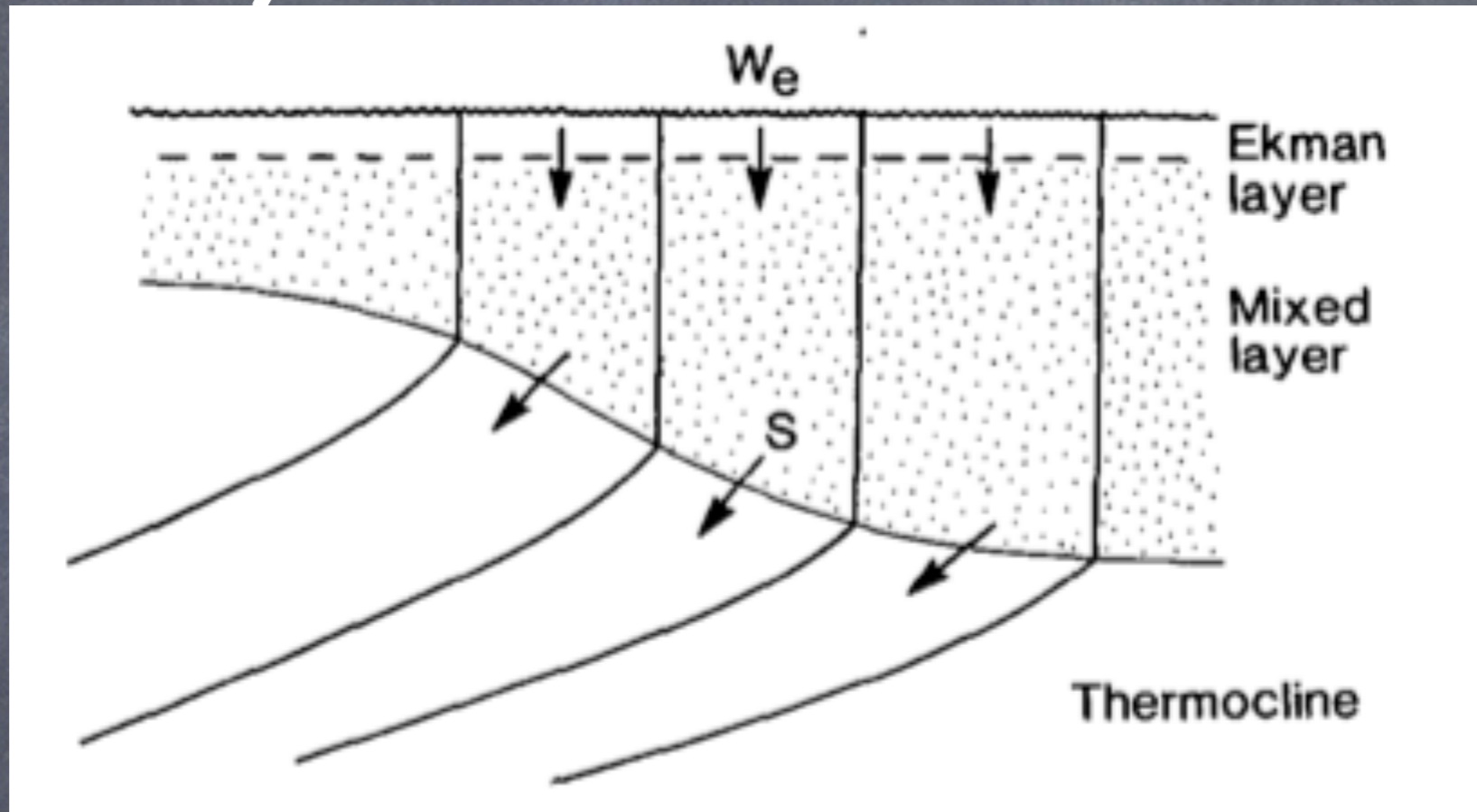


Figure 4. Maximum Courant number, $w\Delta t/\Delta z$, for vertical advection. Gray line is from the *LeithOnly* integration, and black line is from the *LeithPlus* integration.

B. Fox-Kemper and D. Menemenlis. Can large eddy simulation techniques improve mesoscale-rich ocean models? In M. Hecht and H. Hasumi, editors, *Ocean Modeling in an Eddying Regime*, volume 177, pages 319-338. AGU Geophysical Monograph Series, 2008.

Is 2D Turbulence a good proxy for neutral flow?



Yes:

No:

Nurser & Marshall, 1991 JPO

- For a few eddy time-scales QG & 2D AGREE (Bracco et al. '04)
- Barotropic Flow--Obvious 2d analogue

- Bolus Fluxes--Divergent 2d flow
- Sloped, not horiz.
- Surface Effects?

Some Asymptotic Limits, Following McWilliams '85

- Very Small Fr,
but large Ro:

- stacked 2d
layers

- Very Small H/L,
small b:

- Barotropic 2d
Turbulence

- Ro, Fr, beta $\ll 1$,

- & Ro/Fr=Ld/L=1,

- QG, SQG

The following equations result

$$Ro \left[\frac{\partial \mathbf{V}_h}{\partial t} + \mathbf{V}_H \cdot \nabla \mathbf{V}_H + Fr^2 \max(Ro^{-1}, 1) w \frac{\partial}{\partial z} \mathbf{V}_H \right] + \max(1, Ro) \nabla_H \pi + \left(1 + \frac{\beta y}{f_0} \right) \mathbf{k} \times \mathbf{V}_H = 0, \quad (2.44)$$

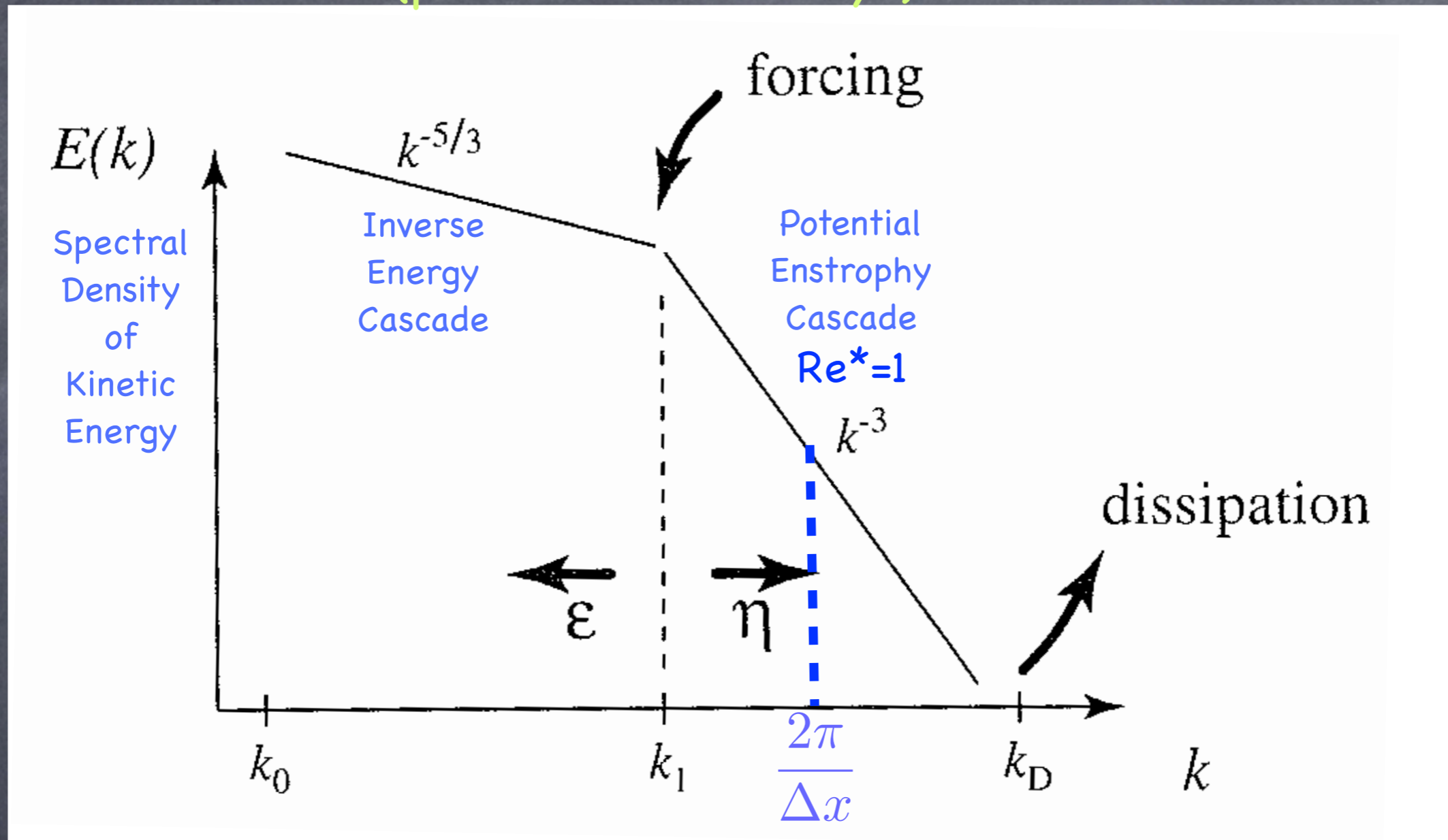
$$Fr^2 \frac{H^2}{L^2} \left[\frac{\partial w}{\partial t} + \mathbf{V}_H \cdot w + Fr^2 \max(Ro^{-1}, 1) w \frac{\partial}{\partial z} w \right] + \frac{\partial}{\partial z} \pi - b = 0, \quad (2.45)$$

$$\frac{\partial b}{\partial t} + \mathbf{V}_H \cdot \nabla b + Fr^2 \max(Ro^{-1}, 1) w \frac{\partial}{\partial z} b + w \frac{N^2}{N_0^2} = 0, \quad (2.46)$$

$$\nabla \cdot \mathbf{V}_H + Fr^2 \max(Ro^{-1}, 1) \frac{\partial}{\partial z} w = 0. \quad (2.47)$$

QG Turbulence: Pot'l Enstrophy cascade (potential vorticity²)

J. Charney, 1971 JAS

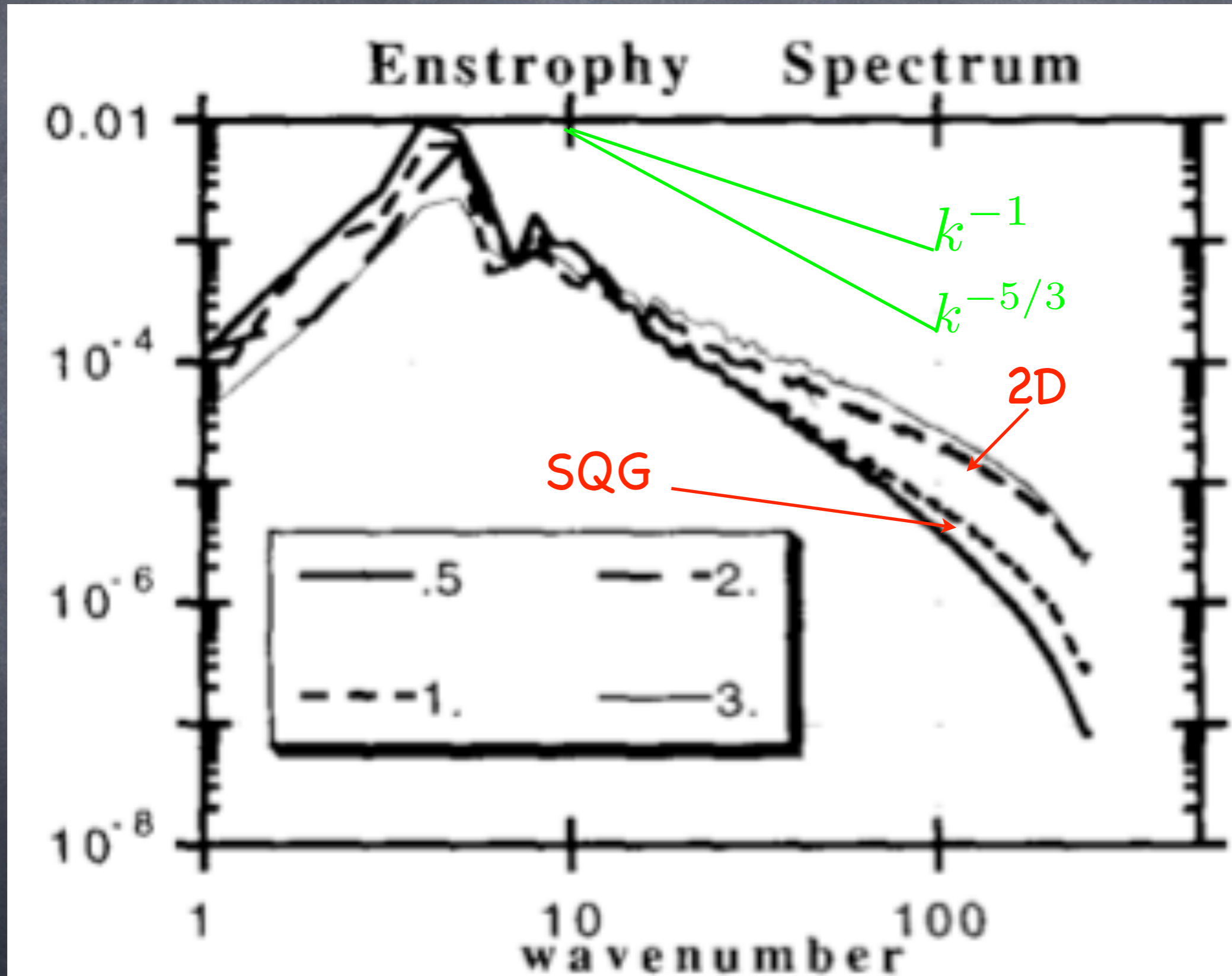


$$\nu_{qg} = \left(\frac{\Lambda_{qg} \Delta x}{\pi} \right)^3 |\nabla q_{qg}| = \left(\frac{\Lambda_{qg} \Delta x}{\pi} \right)^3 \left| \nabla_h \left[\beta y + \nabla_h^2 \psi + \frac{\partial}{\partial z} \left(\frac{f_0^2}{N^2} \frac{\partial \psi}{\partial z} \right) \right] \right|.$$

$$\nabla_h \nabla_h^2 \psi = \nabla_h \left[\frac{\partial u_*}{\partial y} - \frac{\partial v_*}{\partial x} \right]$$

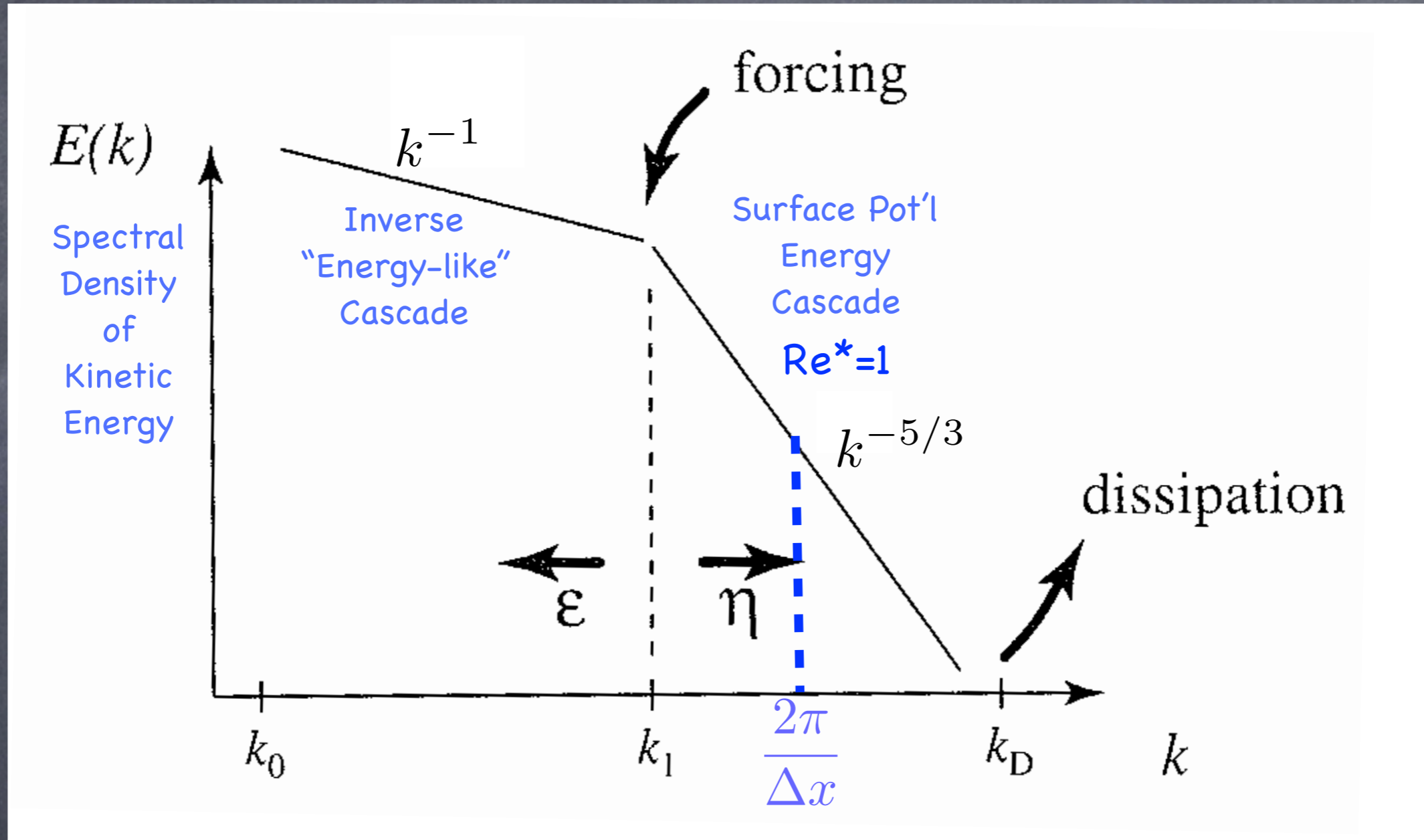
$$\nu_{qg} = \kappa_{Redi} = \kappa_{GM} = \left(\frac{\Lambda_{qg} \Delta x}{\pi} \right)^3 |\nabla q_{qg}|.$$

In real stratified flows, things are a bit more complex than in 2d. Even more than QG... Surface Effects may dominate



SQG Turbulence: Surface Buoyancy & Velocity cascade--scales surface horiz. diffusivity only

W. Blumen, 1978 JAS
 Held et al 1995, JFM.
 Smith et al. 2002, JFM



Smag-Like
 (Inverse):
 Leith-Like
 (Direct):

$$\kappa_* = \left(\frac{\Upsilon \Delta x}{\pi} \right)^{4/3} \left| \frac{1}{f} \nabla_h b \right|^{2/3}$$

$$\kappa_* = \left(\frac{\Lambda \Delta x}{2\pi} \right)^{3/2} \left[-\frac{\partial}{\partial z} |\nabla_h \psi|^2 \right]^{1/2}$$

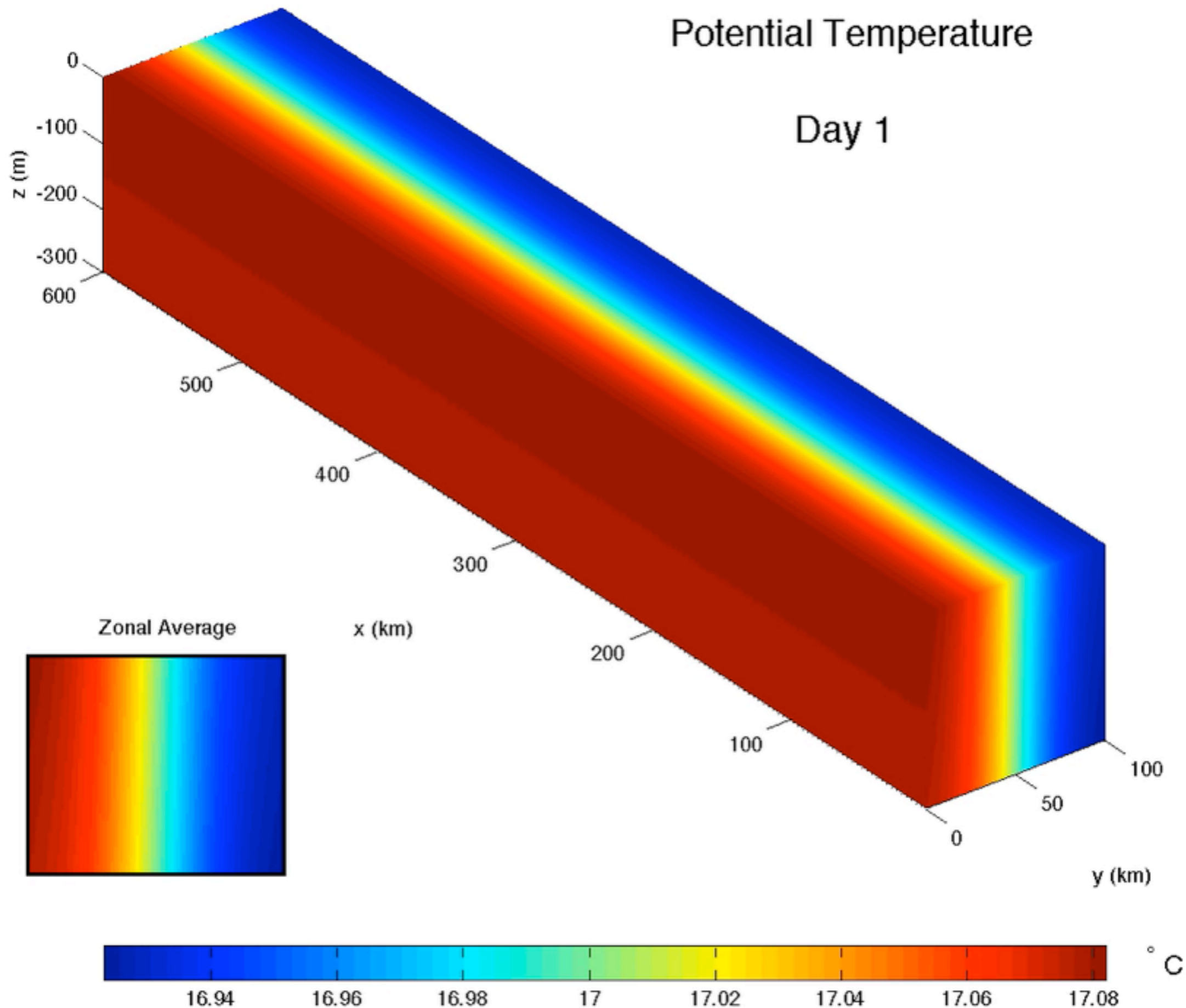
And QG pot'l enstrophy Leith is ... working in MITgcm

- Scott Bachman (DAMTP) has implemented this QG Leith closure in the MITgcm
 - Both Germano Dynamic and Fixed Coefficient
- Sets viscosity=diffusivity=GM coefficient
- Both are stable and robust, very similar (is dynamical needed?)
- Both work better than Smagorinsky, smoother spectrum to grid scale (to be shown next).
- But, we don't yet understand the spectral behavior of all test cases. 2d barotropic, QG, & SQG, equatorial are coexistent..

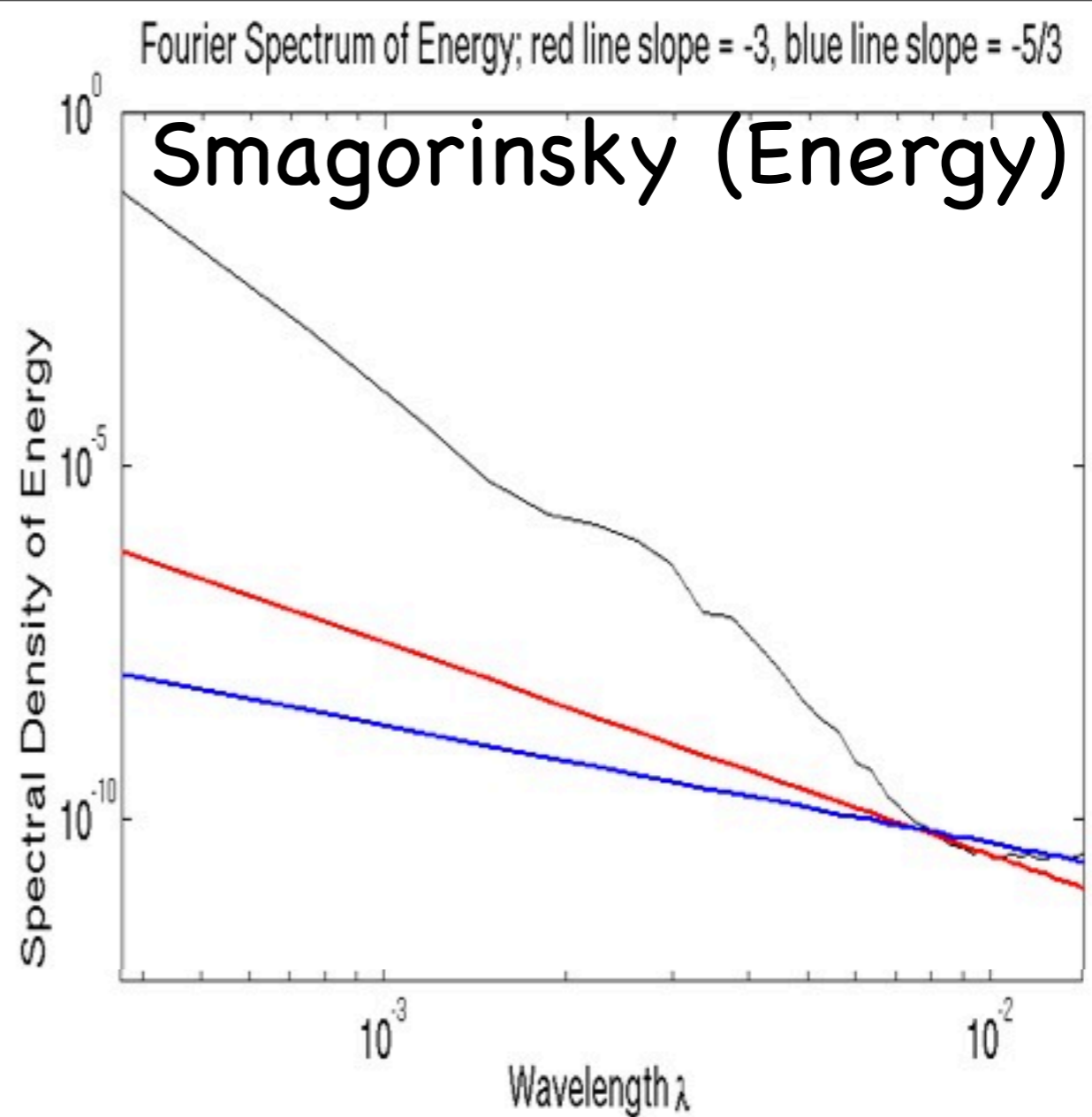
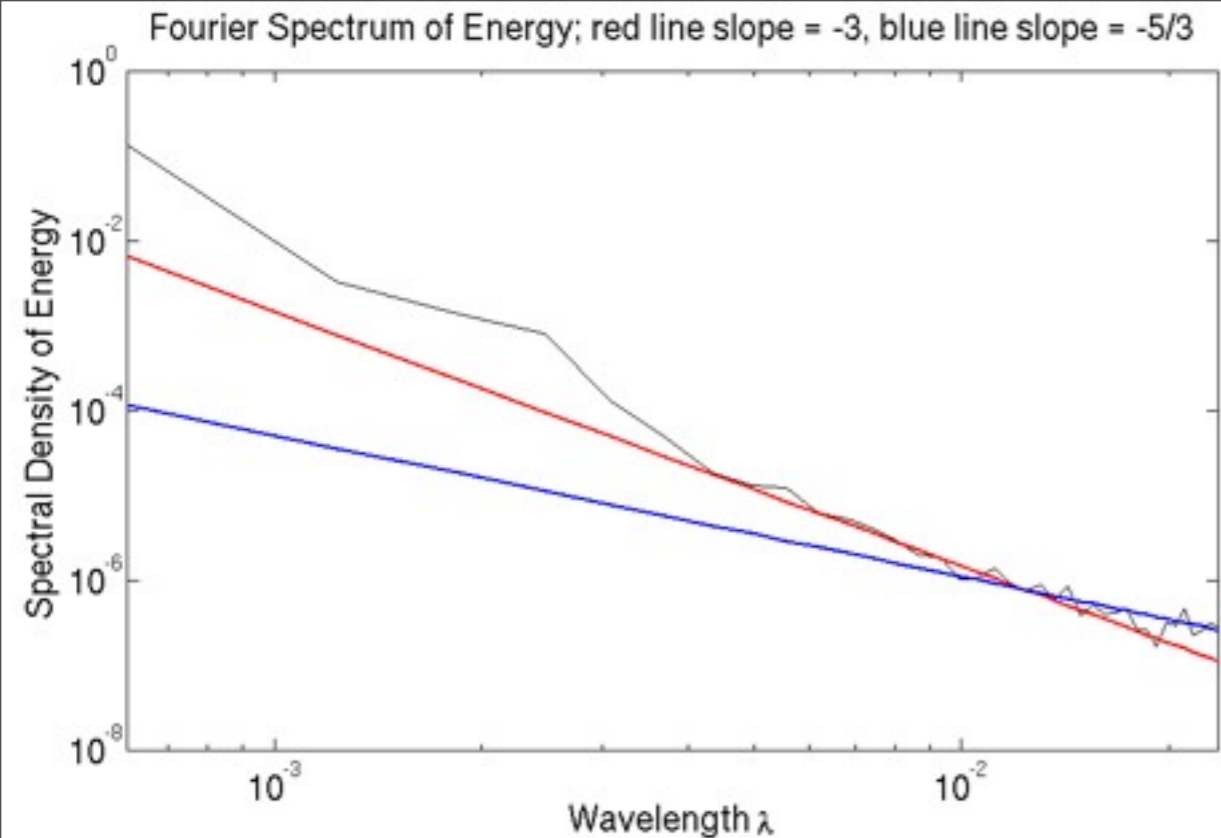
Movie: S. Bachman

Potential Temperature

Day 1



S. Bachman and B. Fox-Kemper. Eddy parameterization challenge suite. I: Eady spindown. *Ocean Modelling*, 64:12-28, 2013.



QG Leith (Pot'l Enstrophy)

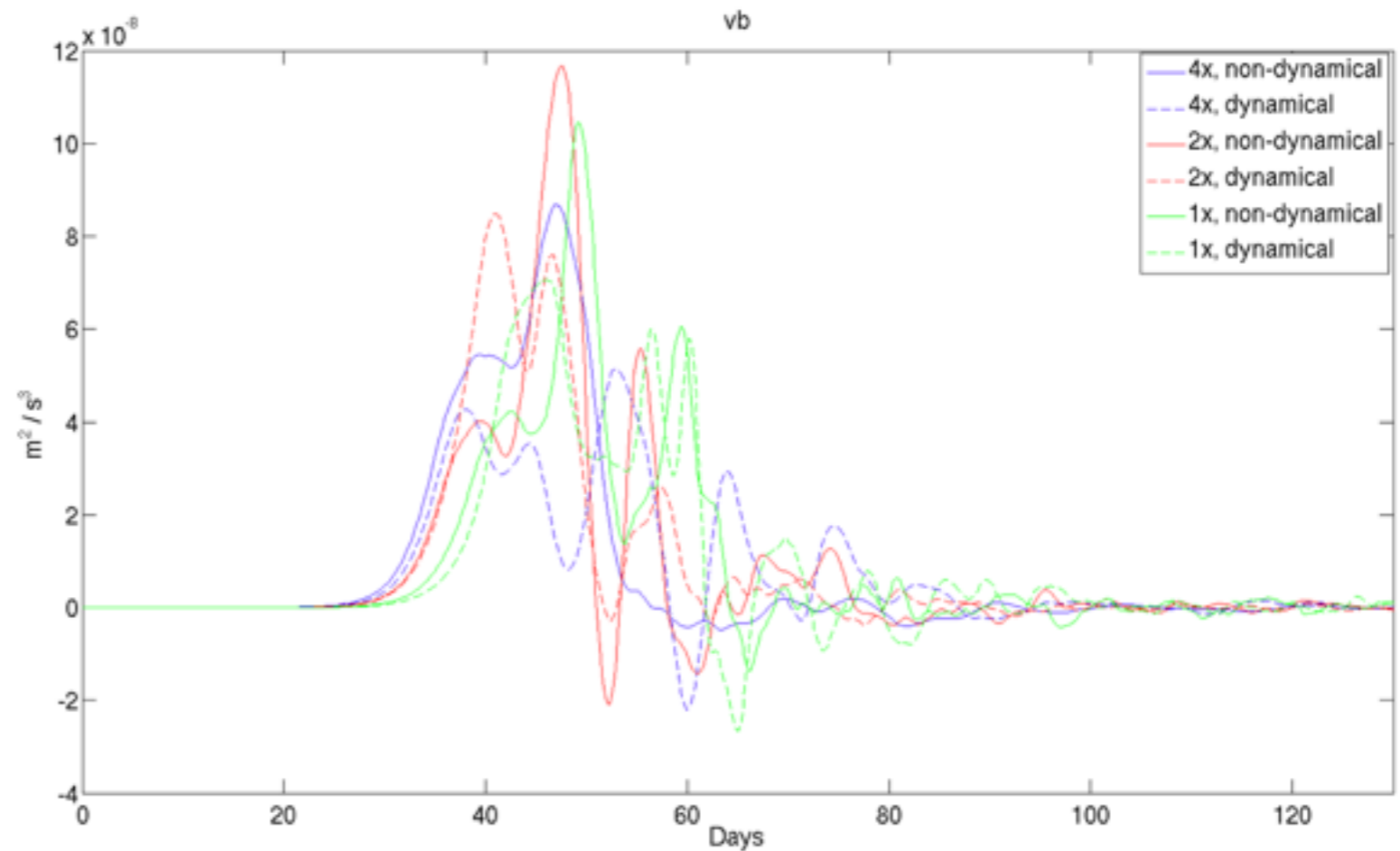
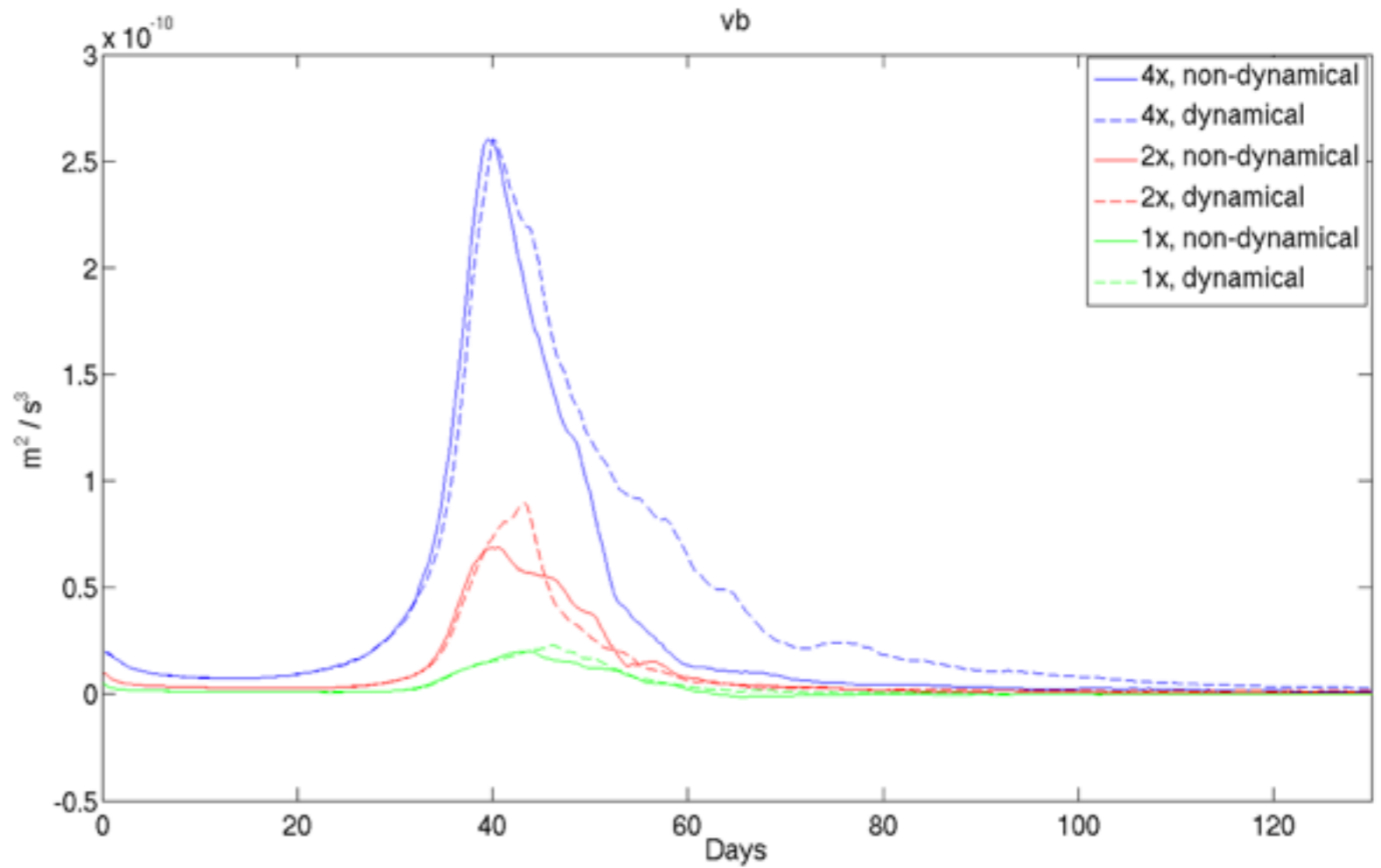
Comparing the spectrum in QG Leith against another (inappropriate) LES closure, we see:

- 1) Better adherence to expected spectrum
- 2) Less "ski jump" near gridscale
- 3) Effects of choice *not limited* to small scales, slope in Smag. is too steep across whole range!

Fluxes:
Horizontal Buoyancy
 $\langle vb \rangle$

Parameterized:

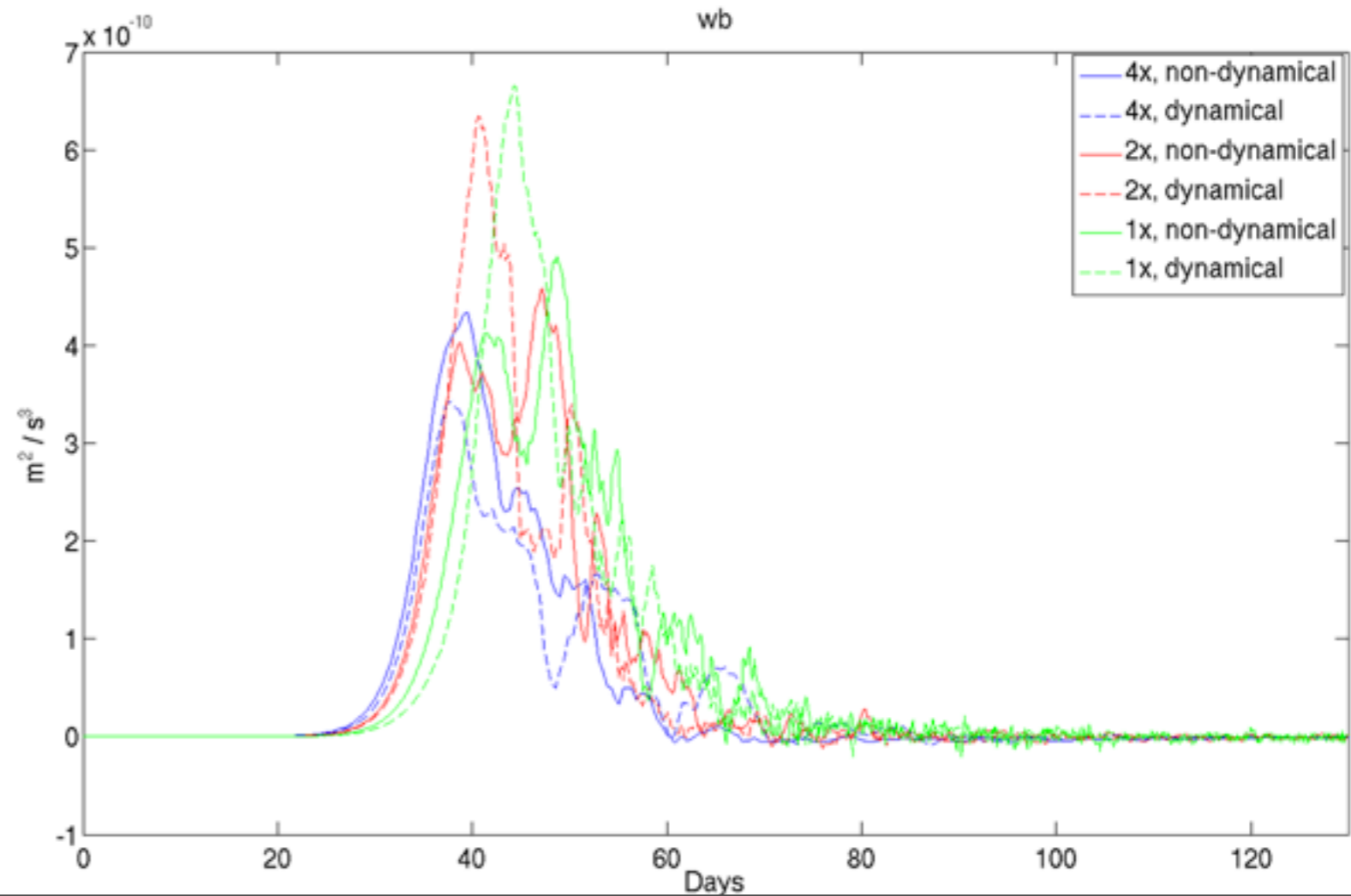
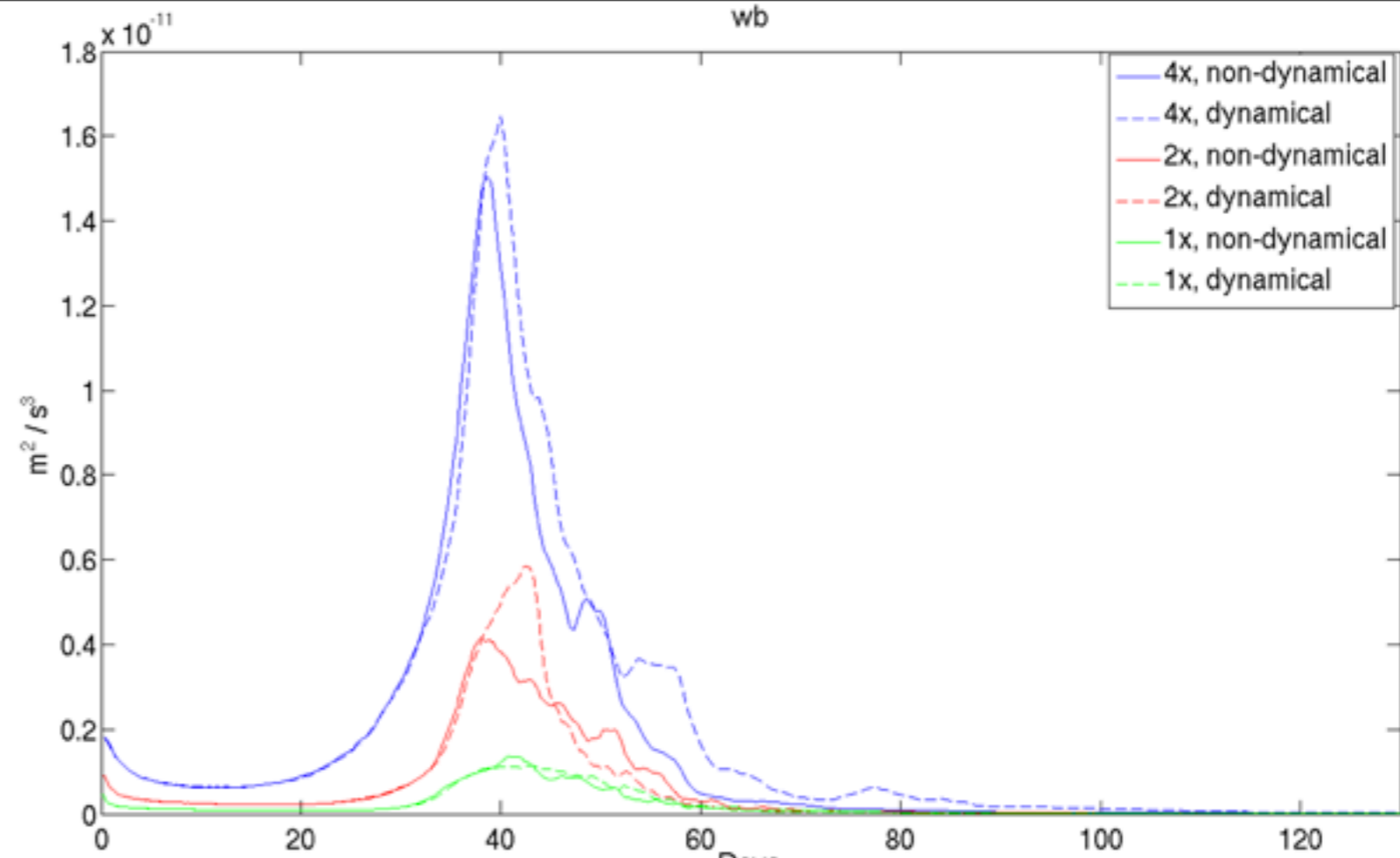
Total:



Fluxes:
Vertical Buoyancy
 $\langle wb \rangle$

Parameterized:

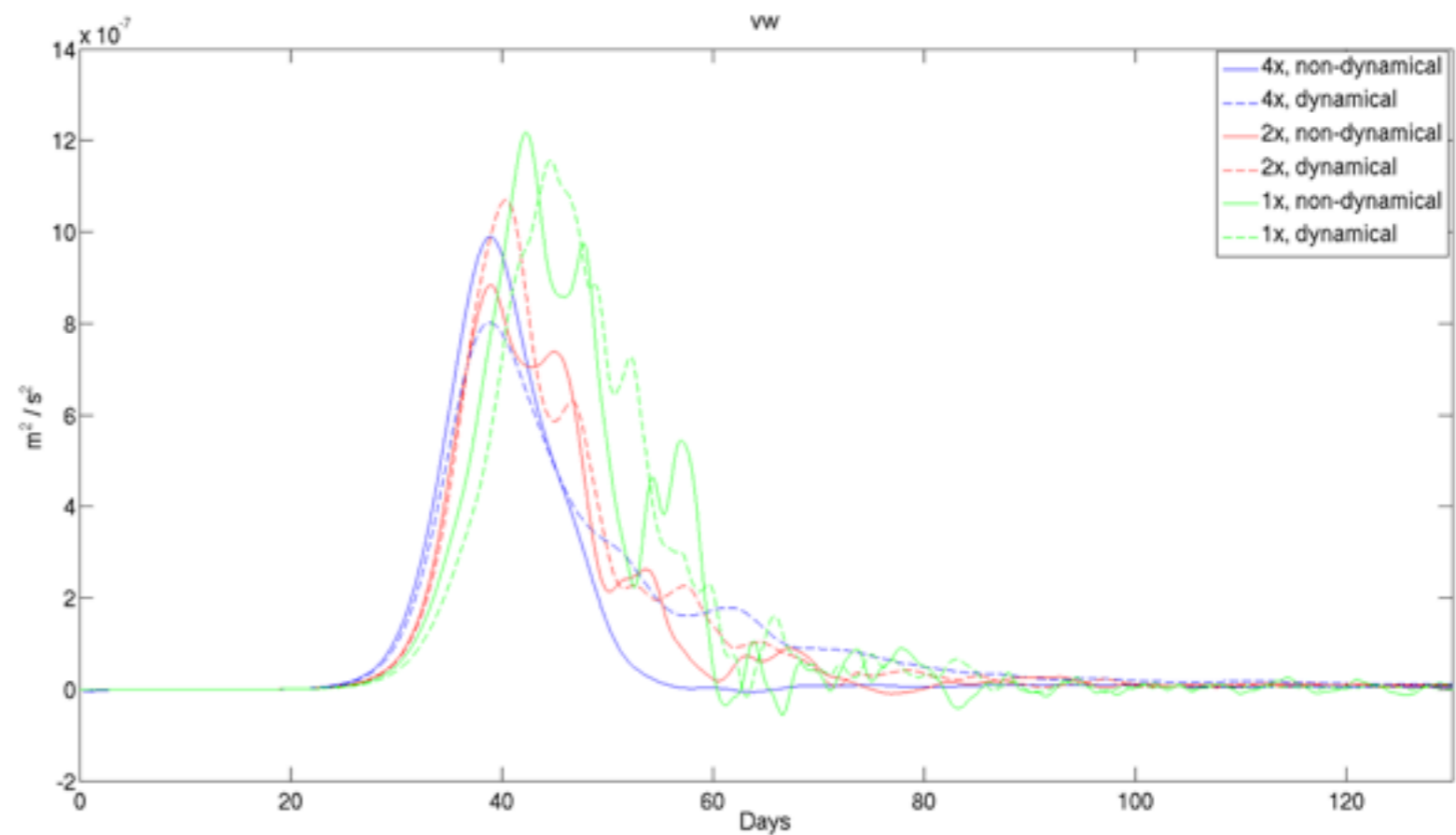
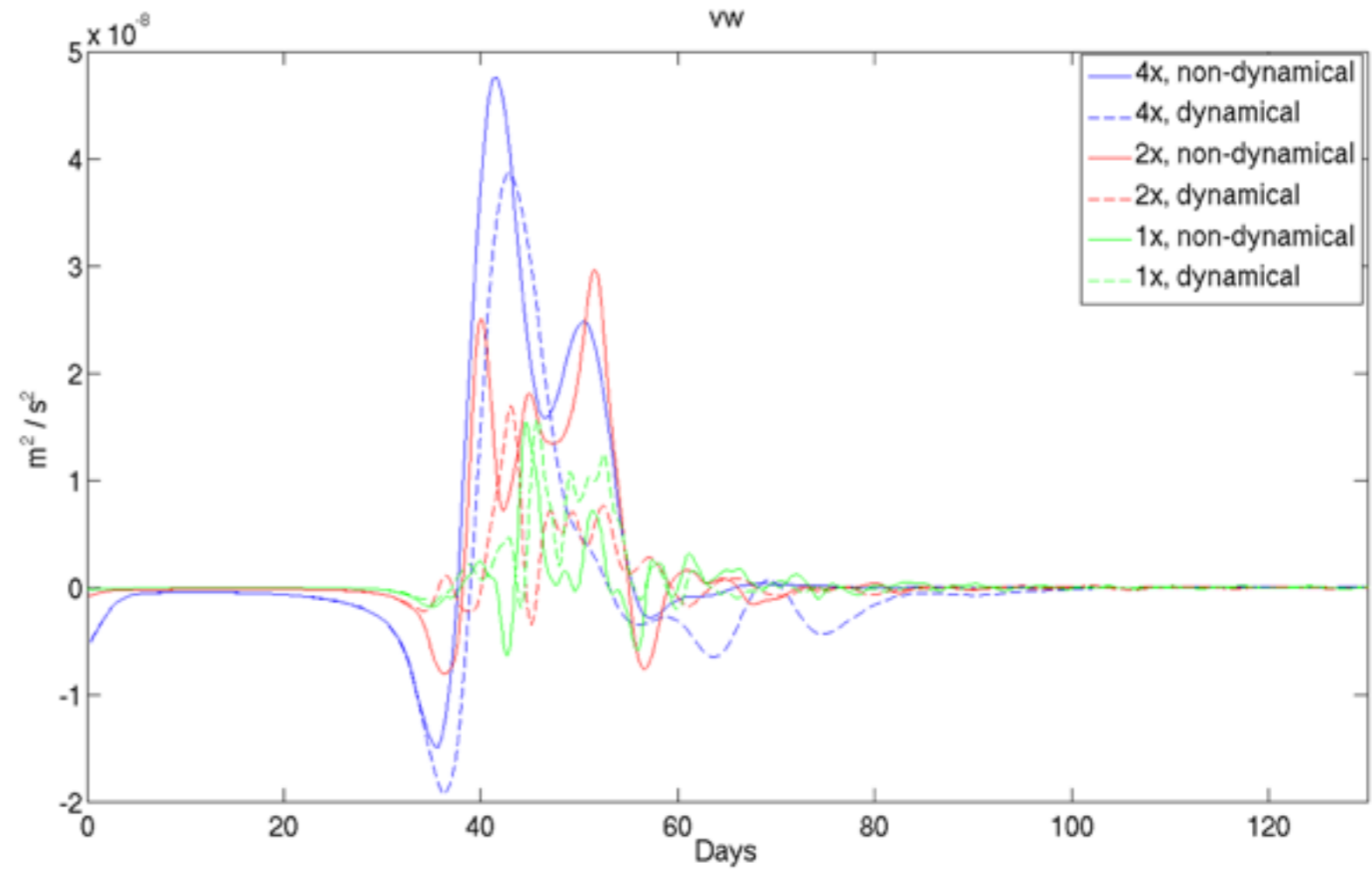
Total:



Fluxes:
Momentum
<vw>

Parameterized:

Total:



A Prescription for Parameterization...

Needs to be checked in various regimes

- QG Leith & Potential Vorticity to generate #1 viscosity
- 2D Leith & Barotropic Vorticity to generate #2 viscosity
- SQG Leith & Surf. Buoyancy to generate #3 diffusivity
- Take $\max(\#1, \#2, \#3)$ as viscosity, Redi diffusivity, *and* as GM transfer coeff.
Nearly suggested by Roberts & Marshall, 98, JPO
- Note: Unlike Eddy-Free closures, e.g., Visbeck et al (97), Eddy-Rich closures take advantage of resolved eddies & instabilities, only need a boost from eddy-permitting to eddy-resolving (and for numerical stability)