

Effects of Ocean Surface Waves: on Turbulence, Climate, and Frontogenesis



Expanding on past work with:

Jim McWilliams (UCLA), Peter Hamlington (CU-Boulder), Eric D'Asaro & Ramsey Harcourt (UW), Luke Van Roekel (LANL), Qing Li (Brown), Sean Haney (CU), Adrean Webb (TUMST), Keith Julien (CU-APPM), Greg Chini (UNH), Peter Sullivan (NCAR), Mark Hemer (CSIRO)

Baylor Fox-Kemper
with **Nobuhiro Suzuki**
(Brown University)

Wed, 20/04/16
13:45-14:00

EGU, 20/4/16

Sponsors: NSF 1258907,

Gulf of Mexico Research Initiative

We Will Examine the Effects of Surface Waves on:

- Boundary Layer Turbulence
(wave-driven or Langmuir Turbulence)
- Climate through Langmuir Turbulence
(via MLD changes)
- Submesoscale Fronts & Instabilities
within the Mixed Layer
(Stokes forces and Langmuir coupling)

Wave-Averaged Equations

following Holm (96), Lane et al. (07), McWilliams & F-K (13), and Suzuki & F-K (16)

$$\underbrace{v_j^L}_{\text{Lagrangian}} = \underbrace{v_j}_{\text{Eulerian}} + \underbrace{v_j^S}_{\text{Stokes}}$$

Coupling Depends on Stokes drift-WAVE effects in YELLOW

Boundary conditions, plus:

$$Ro [v_{i,t} + v_j^L v_{i,j}] + \frac{M_{Ro}}{Ri} w v_{i,z} + \boxed{\epsilon_{izj} v_j^L} = \text{(Lagrangian) geostrophic} \quad -M_{Ro} \pi_{,i} + \frac{Ro}{Re} v_{i,jj}$$

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b} - \epsilon v_j^L v_{j,z}^S + \frac{\alpha^2}{Re Ri} w_{,jj}$$

hydrostatic

$$b_t + v_j^L b_{,j} + \frac{M_{Ro}}{Ro Ri} w b_z = \frac{1}{Pe} b_{,jj}$$

$$v_{j,j} + \frac{M_{Ro}}{Ro Ri} w_z = 0$$

$$\boxed{\epsilon = \frac{V^S H}{f L H_s}}$$

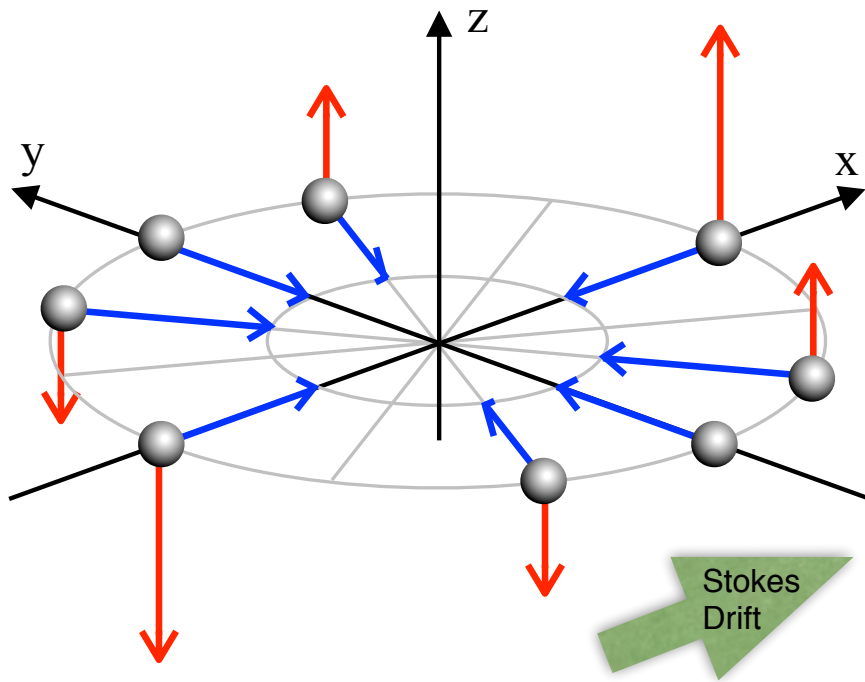
$$Re = \frac{UL}{\nu} \quad Ro = \frac{U}{fL} \quad Ri = \frac{N^2}{(U_{,z})^2} \quad \alpha = H/L \quad M_{Ro} \equiv \max(1, Ro)$$

J. C. McWilliams and BFK. Oceanic wave-balanced surface fronts and filaments. Journal of Fluid Mechanics, 730:464-490, 2013.

N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. JGR-Oceans, December 2015. Submitted.

3 Wave Effects, 3: Stokes Shear Force and the CL2 mechanism for Langmuir circulations

Flow directed along Stokes shear=downward force



← : Stokes-shear force ● : water parcel
← : turbulent velocity

$$\epsilon = \frac{V^s H}{f L H_s}$$

"wavy hydrostatic" if

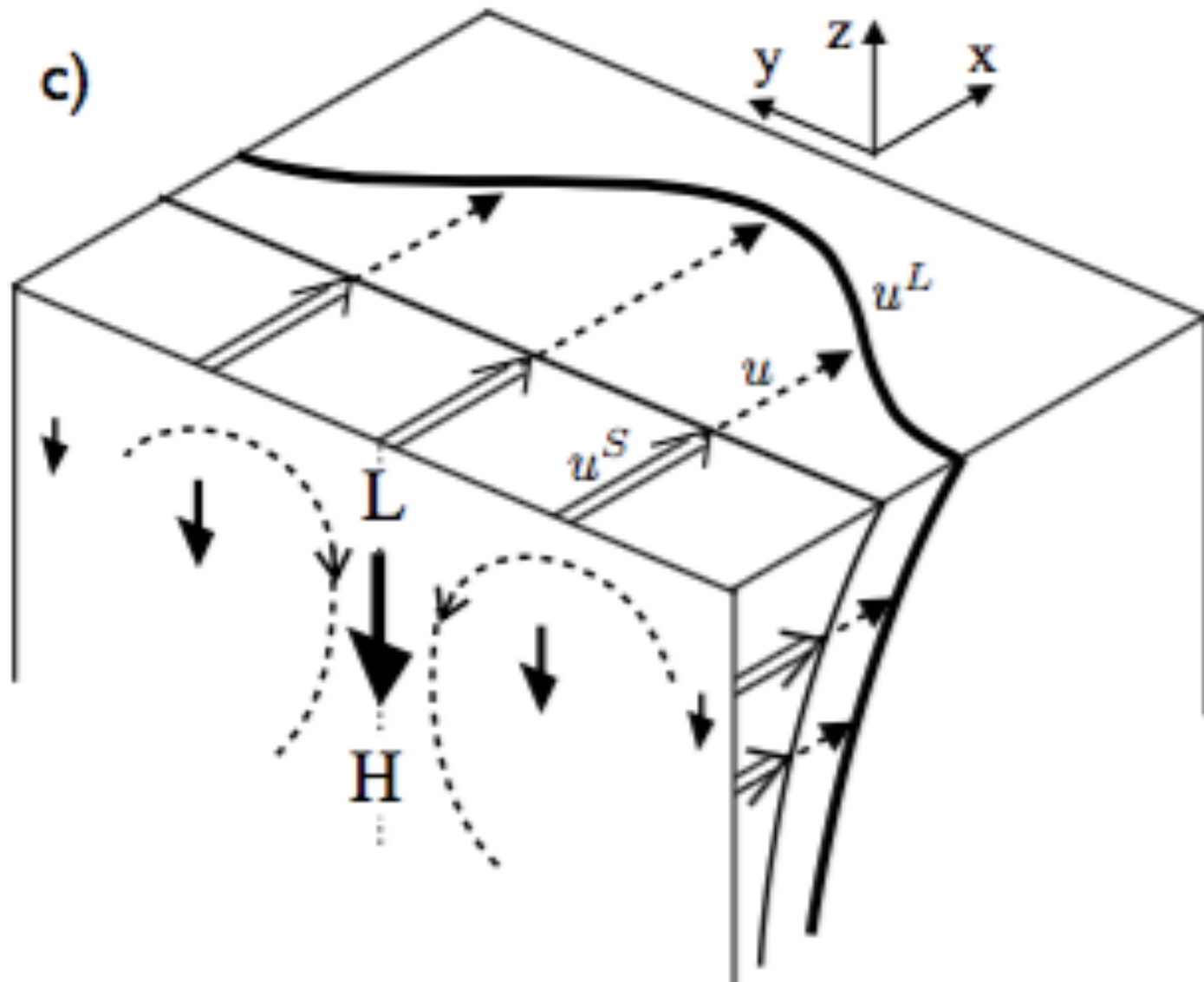
hydrostatic $\epsilon \gg 1$

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b} \boxed{-\epsilon v_j^L v_{j,z}^s} + \frac{\alpha^2}{Re Ri} w_{,jj}$$

image:
Thorpe, 0



Figure 1
windrows
practice th
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within the



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Typical effect: Downward Force for down-Flow Stokes Drift

$$\frac{\alpha^2}{Re} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{RoRe} w w_{,z} \right] = -\pi_{,z} + b - \epsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{ReRe} w_{,jj}$$

To quantify Langmuir Turb. effects on climate: 3 WAYS

- 1) From OBSERVATIONS, estimate wave effects on key parameters ($\langle w^2 \rangle$, sources of energy) using scalings from Large Eddy Simulations. MODEL INDEPENDENT
- 2) OFFLINE 1d mixing with waves parameterized, mixing into observed Argo profiles, reanalysis winds, waves, cooling. ROBUST TO MODEL ERRORS
- 3) In a climate model, *add in a wave forecast model as a new component in addition to atmosphere, ocean, ice, etc.*, use this to drive parameterizations of wave mixing in ocean component. FEEDBACKS PRESENT

No Retuning! ALL coefficients from LES

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E. A. D'Asaro, J. Thomson, A. Y. Shcherbina, R. R. Harcourt, M. F. Cronin, M. A. Hemer, and BFK. Quantifying upper ocean turbulence driven by surface waves. Geophysical Research Letters, 41(1):102-107, January 2014.

Langmuir important



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S. E. Belcher, A. A. L. M. Grant, K. E. Hanley, BFK, L. Van Roekel, P. P. Sullivan, W. G. Large, A. Brown, A. Hines, D. Calvert, A. Rutgersson, H. Petterson, J. Bidlot, P. A. E. M. Janssen, and J. A. Polton. A global perspective on Langmuir turbulence in the ocean surface boundary layer. Geophysical Research Letters, 39(18):L18605, 9pp, September 2012.

Langmuir important



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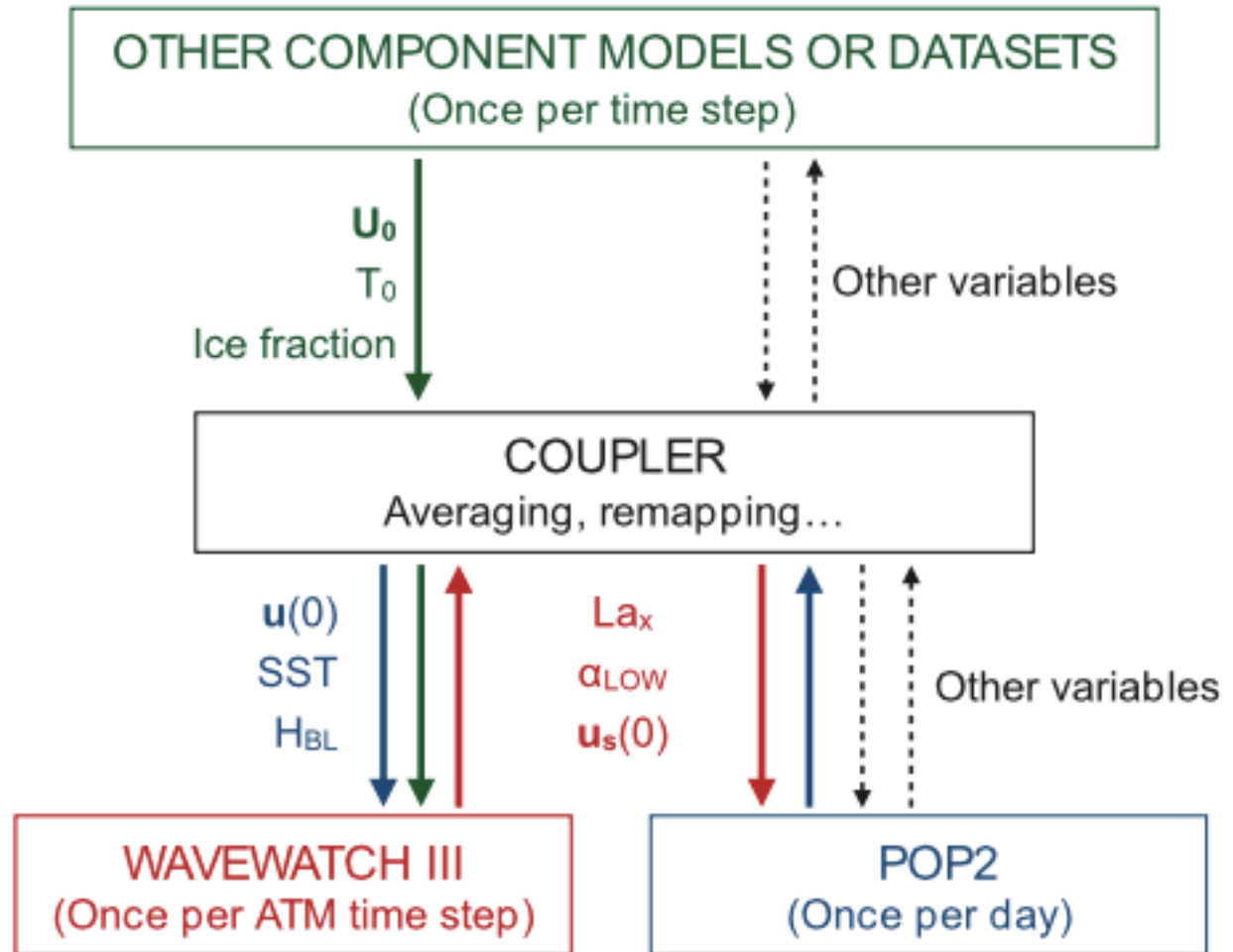
U. S. Department of Commerce
National Oceanic and Atmospheric Administration
National Weather Service
National Centers for Environmental Prediction
5200 Auth Road
Camp Springs, MD 20746

Technical Note

User manual and system documentation of
WAVEWATCH III™ version 3.14 †

Hendrik L. Tolman †

Environmental Modeling Center
Marine Modeling and Analysis Branch



Langmuir Mixing in Climate: Boundary Layer Depth Improved

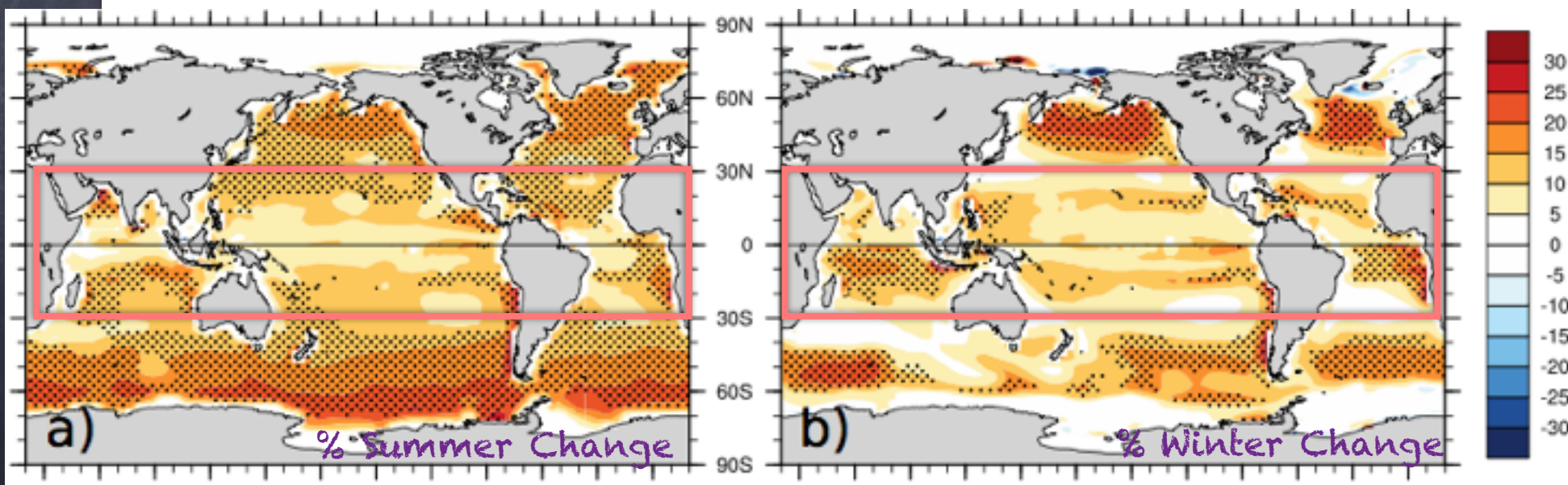
Case	Summer			Winter		
	Global	South of 30°S	30°S-30°N	Global	South of 30°S	30°S-30°N
CTRL	10.62±0.27 ^a (13.40±0.19) ^b	17.24±0.48 (21.73±0.32)	5.38±0.14 (6.71±0.09)	43.85±0.38 (45.50±0.40)	57.19±0.76 (56.53±0.59)	12.57±0.28 (16.16±0.29)
MS2K	15.37	15.47	17.03	119.91	171.92	40.31
SS02	36.79	63.83	7.54	99.32	164.34	17.39
VR12-AL	9.06	13.47	6.49	40.45	50.33	14.52
VR12-MA	8.73±0.30 (11.83±0.29)	12.65±0.47 (18.13±0.62)	6.61±0.22 (7.52±0.16)	40.99±0.37 (42.02±0.39)	51.78±0.65 (50.78±0.67)	14.23±0.30 (15.67±0.35)
VR12-EN	8.95	10.52	8.91	41.94	52.98	19.58

Control

Competition

3 versions of Van Roekel et al

dotted when statistically significant



L. P. Van Roekel, BFK, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. *Journal of Geophysical Research-Oceans*, 117:C05001, 22pp, May 2012.

Q. Li, A. Webb, BFK, A. Craig, G. Danabasoglu, W. G. Large, and M. Vertenstein. Langmuir mixing effects on global climate: WAVEWATCH III in CESM. *Ocean Modelling*, August 2015. in press.

So, we'll quantify Langmuir effects on climate

- 1) From OBSERVATIONS, estimate wave effects on key parameters ($\langle w^2 \rangle$, sources of energy) using scalings from Large Eddy Simulations. MODEL INDEPENDENT Langmuir important ✓
- 2) OFFLINE 1d mixing with waves parameterized, mixing into observed Argo profiles, reanalysis winds, waves, cooling. ROBUST TO MODEL ERRORS Langmuir important ✓
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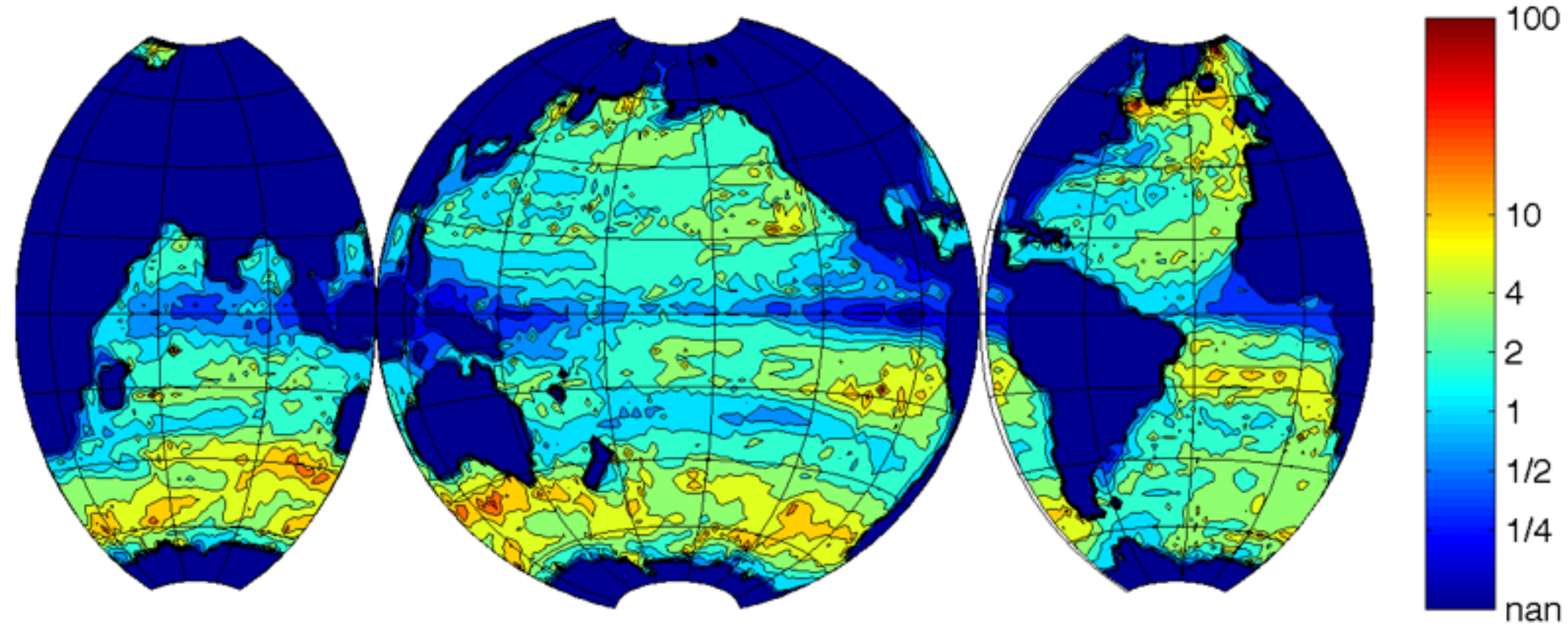
No Retuning! ALL coefficients from LES

Mid-way Conclusions

- Stokes forces may accelerate upper ocean mixing, leading to a wind-wave-convective turbulence driven partly by Stokes forces: Langmuir turbulence
- Three effects of Stokes drift are important: Stokes Advection, Stokes Coriolis, and Stokes Shear Force
- The Stokes Shear Force enhances downward and upward velocities in boundary layer turbulence.
- Including Langmuir mixing in climate models improves the climate model MLD, T, and uptake of CFCs.
- All papers at: fox-kemper.com/pubs

Stokes force directly affects larger scales?

ε/Ro



$$\frac{\varepsilon}{Ro} = \frac{V_s H f L}{f L H_s V} = \frac{V_s H}{V H_s}$$

$$\varepsilon = \frac{V^s H}{f L H_s} \quad Ro = \frac{U}{f L}$$

Langmuir-Submesoscale Front Interactions?

LES of Langmuir turbulence with a
submesoscale temperature front

Use NCAR LES model to solve Wave-
Averaged Eqtns.

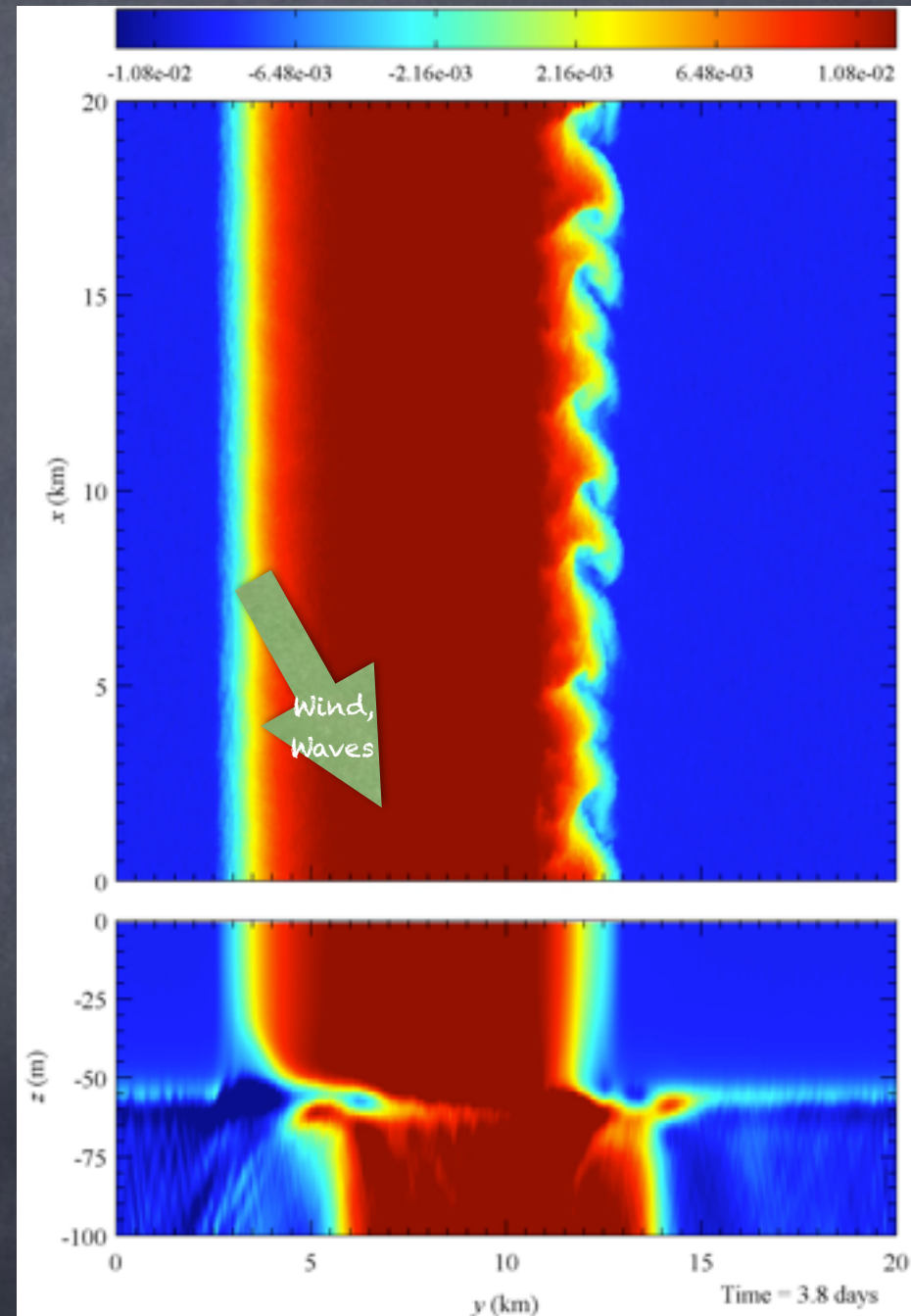
2 Versions: 1 With Waves & Winds
1 With only Winds

Computational parameters:

Domain size: 20km x 20km x -160m

Grid points: 4096 x 4096 x 128

Resolution: 5m x 5m x -1.25m



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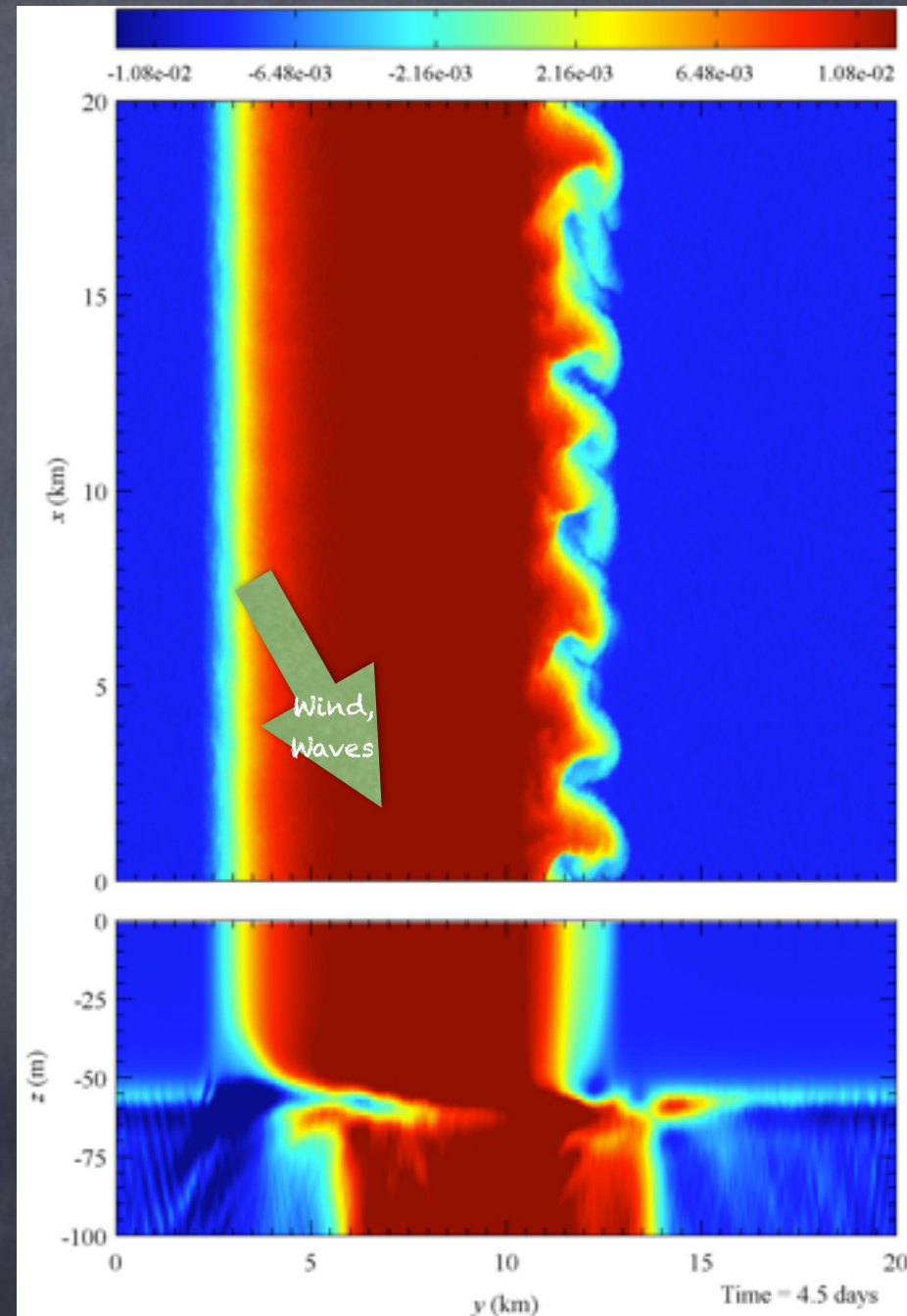
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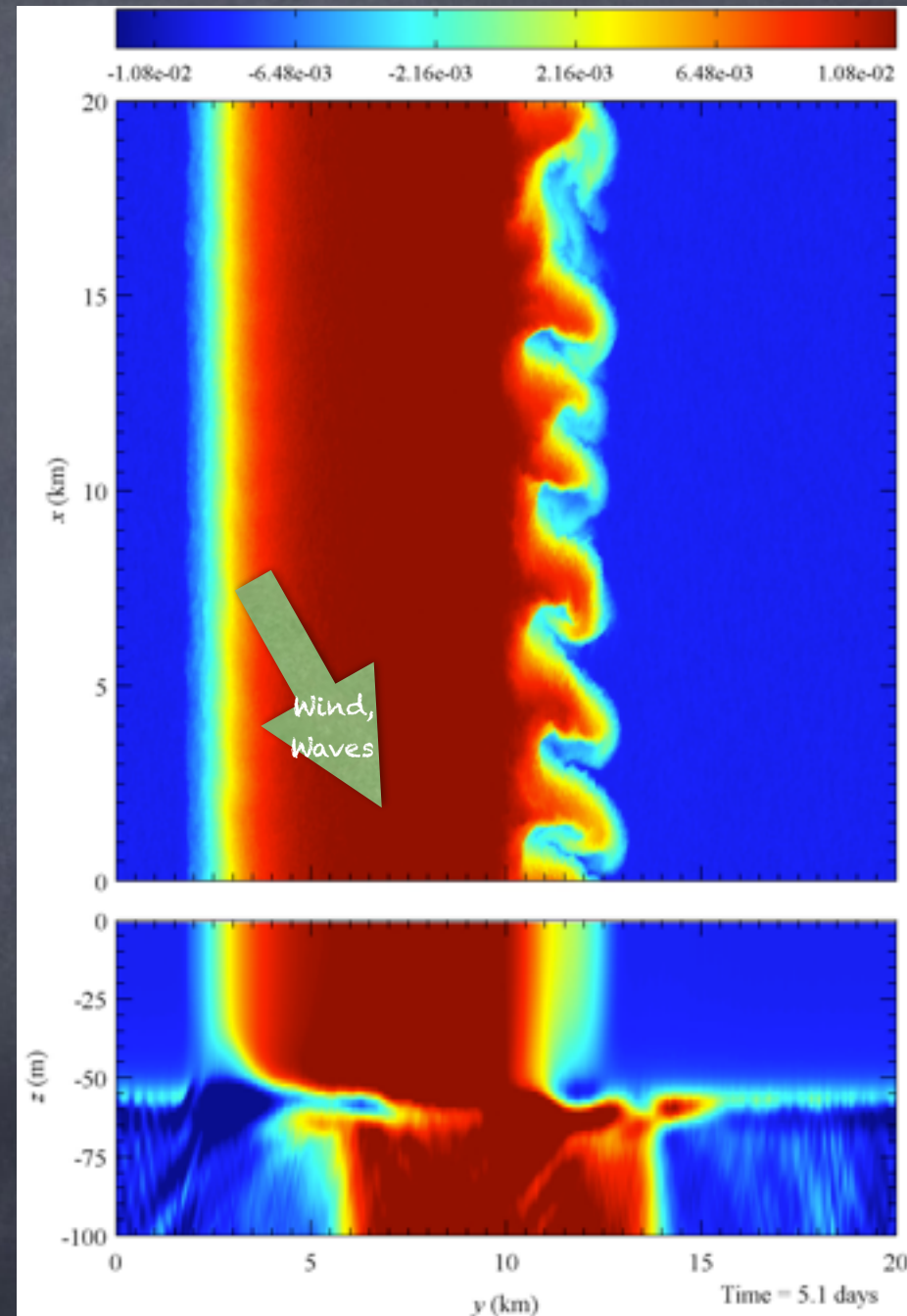
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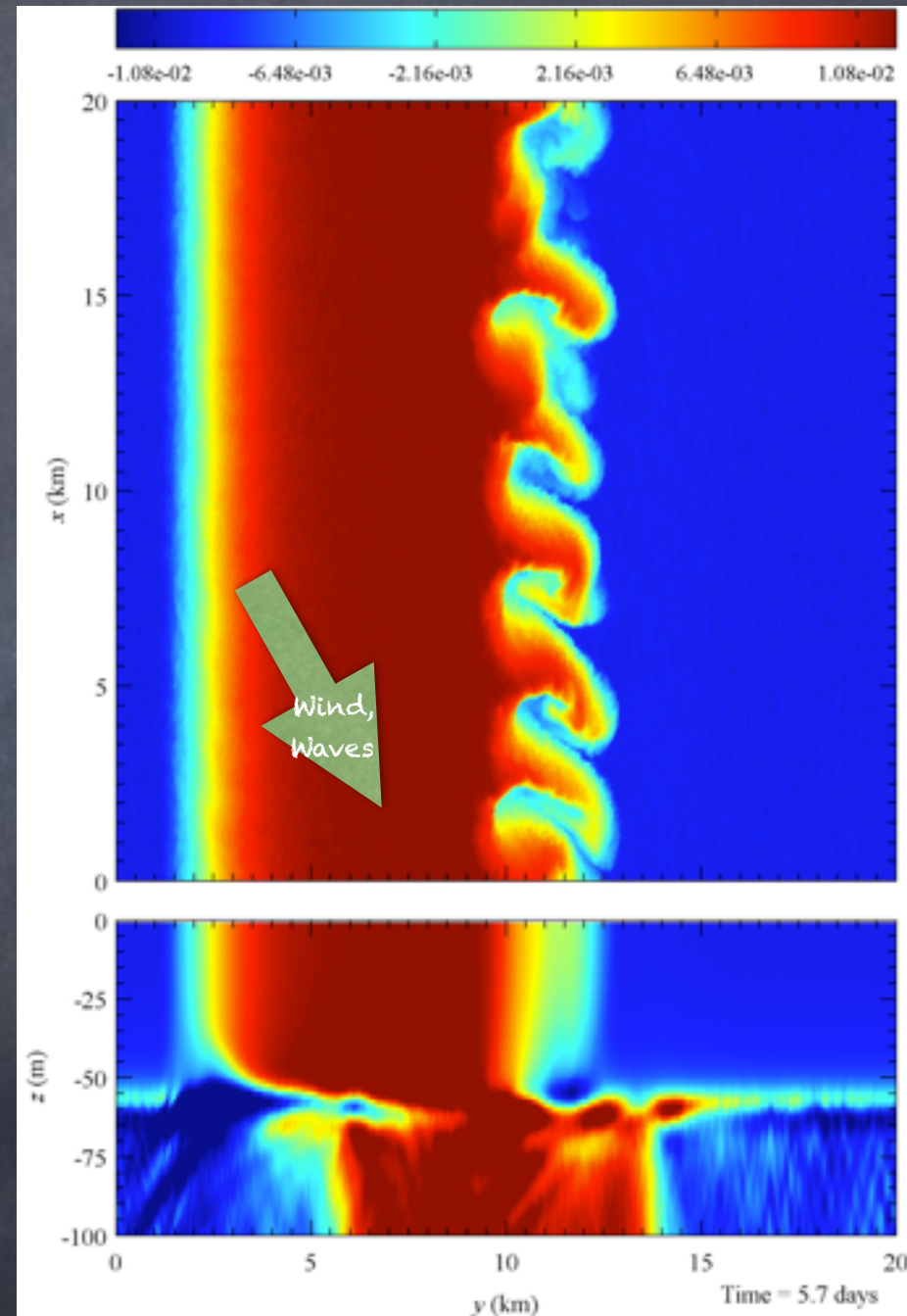
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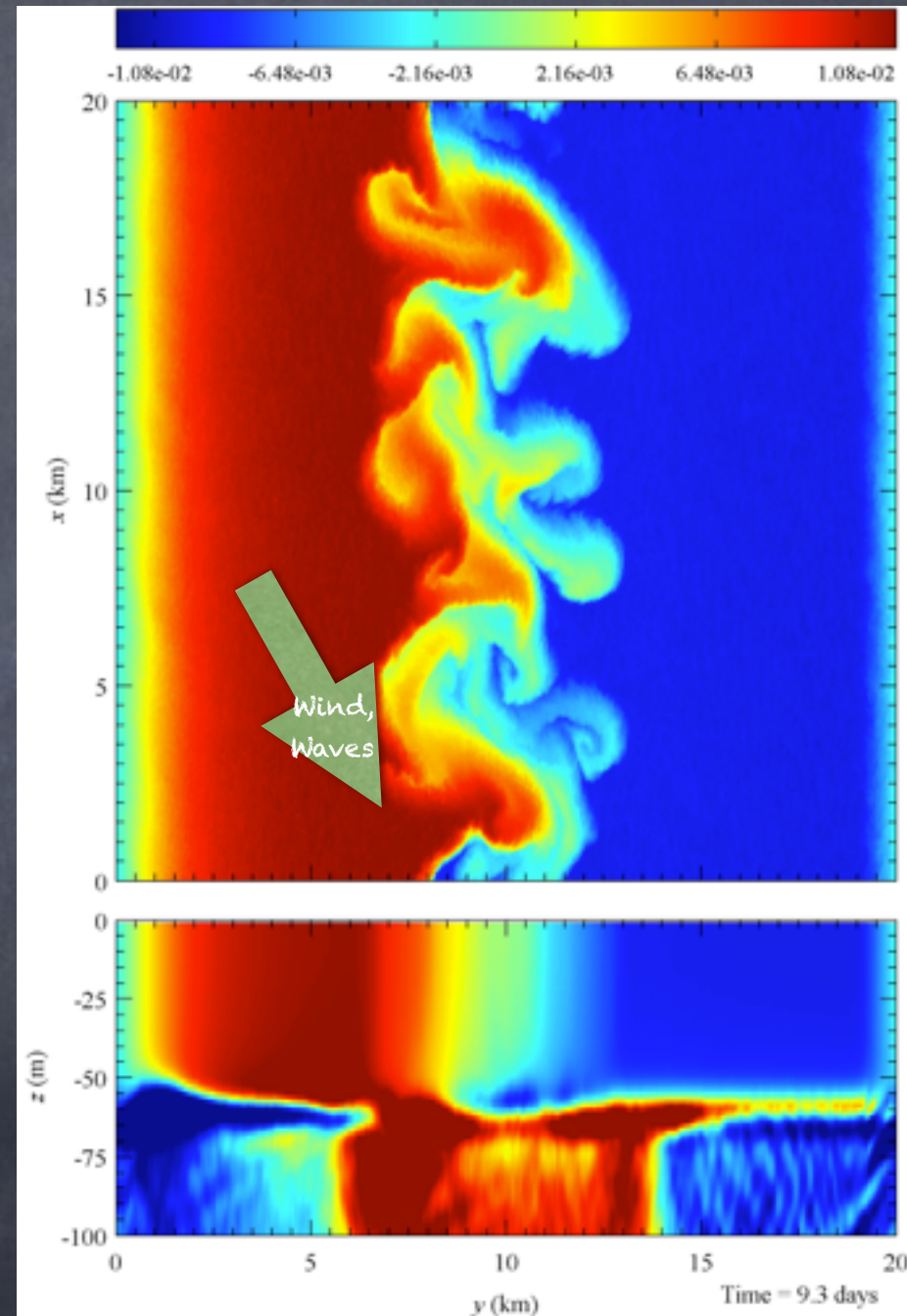
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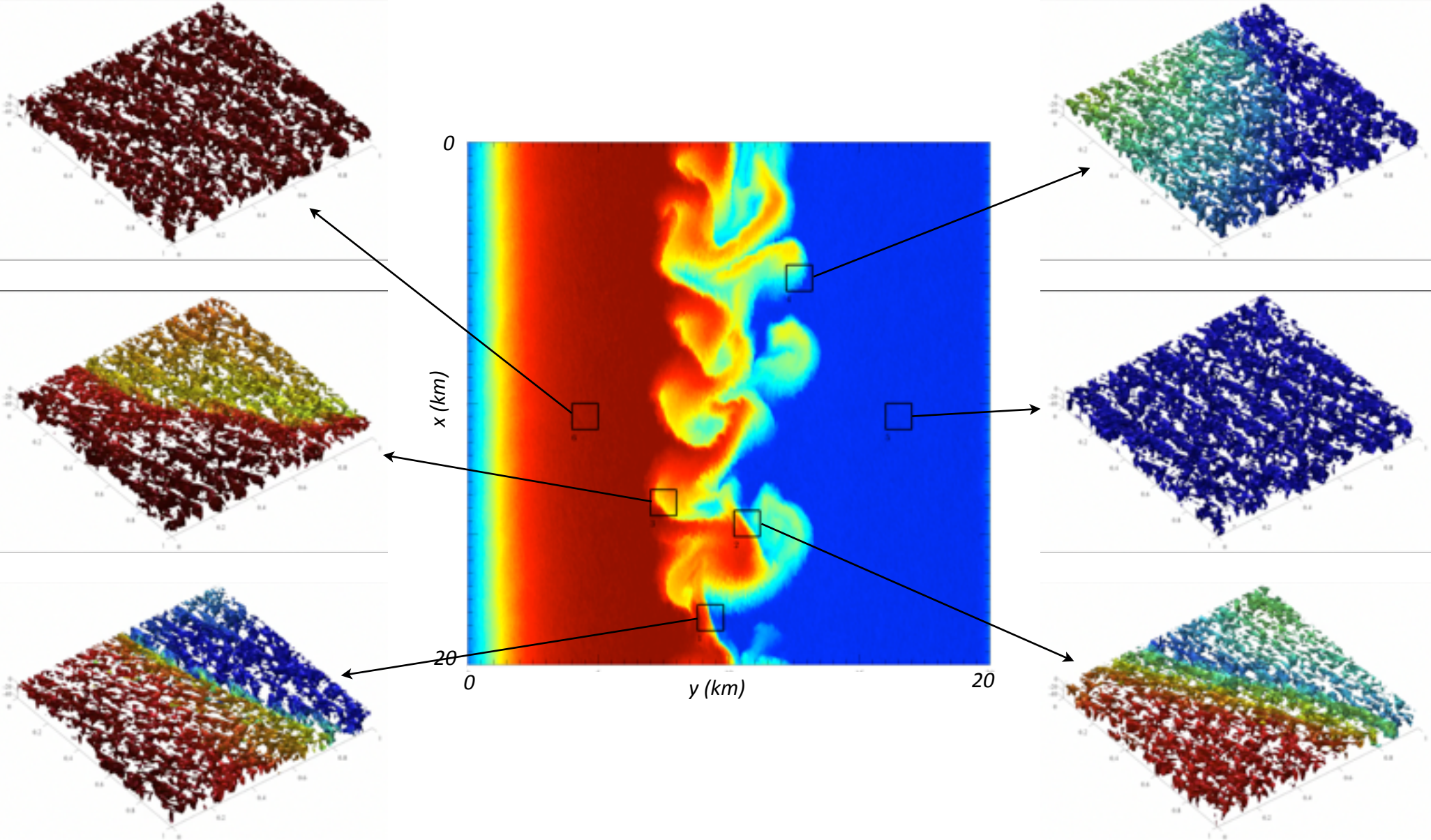
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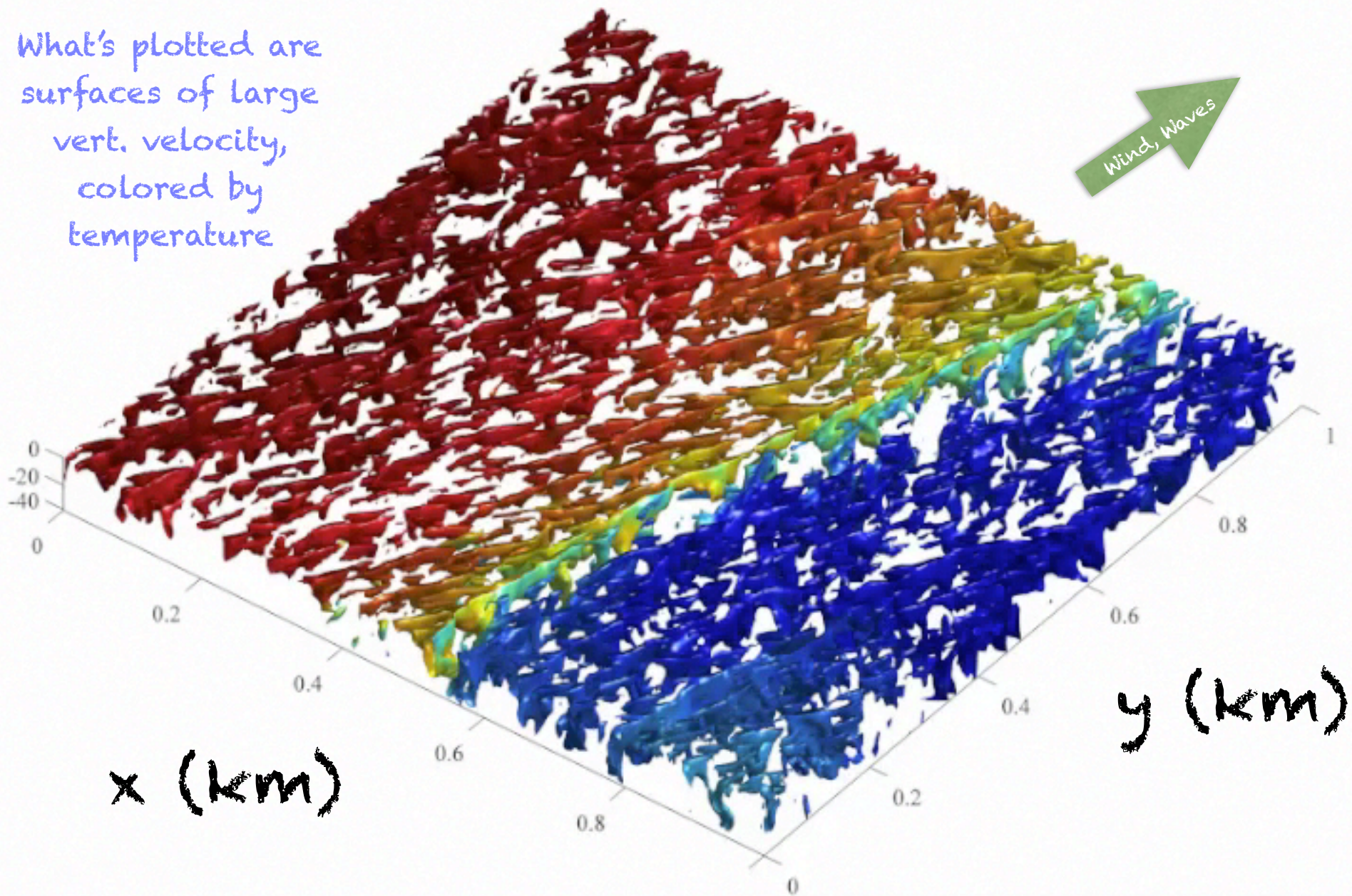
Resolution: 5m x 5m x -1.25m



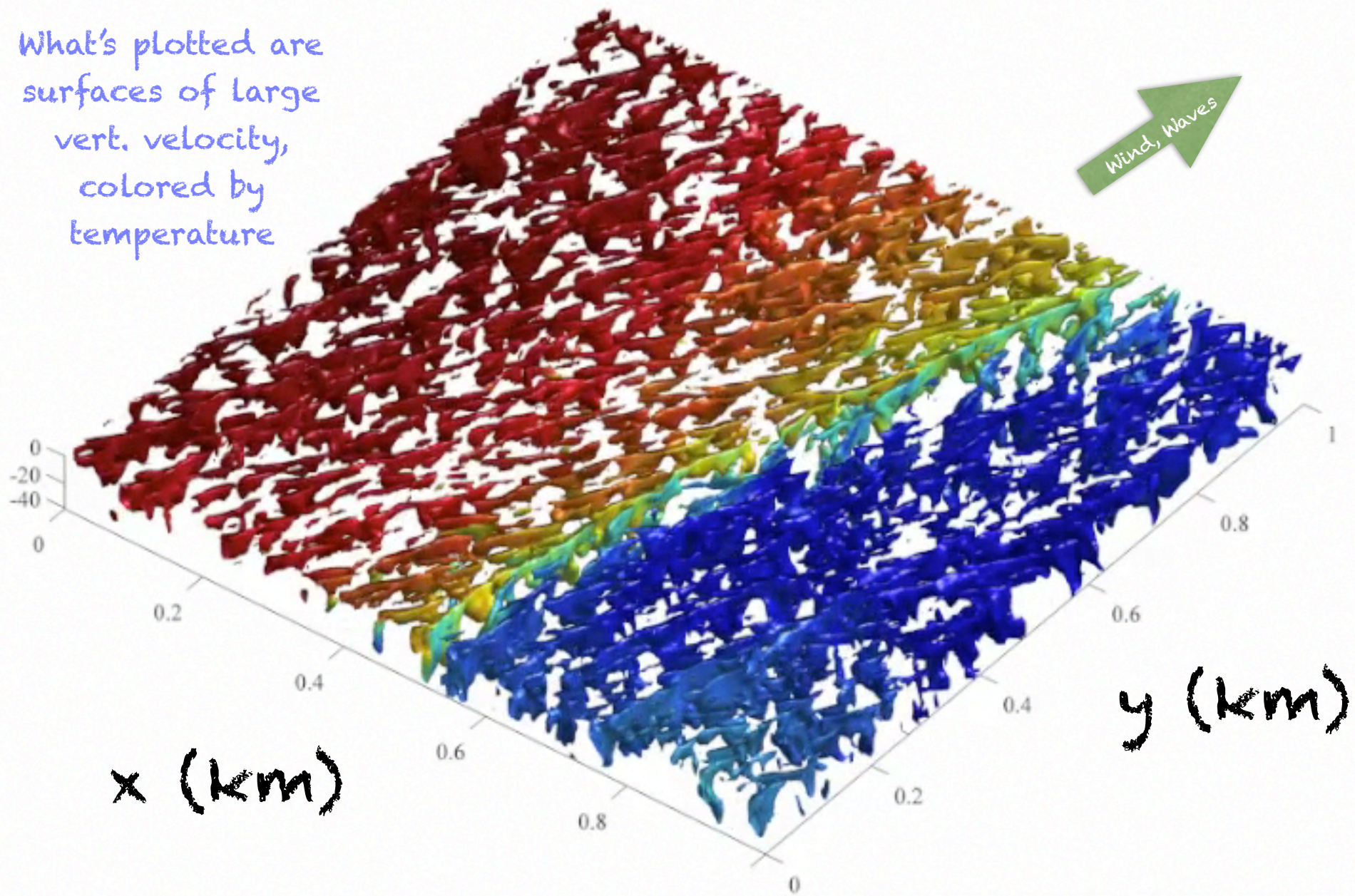
Diverse types of interaction



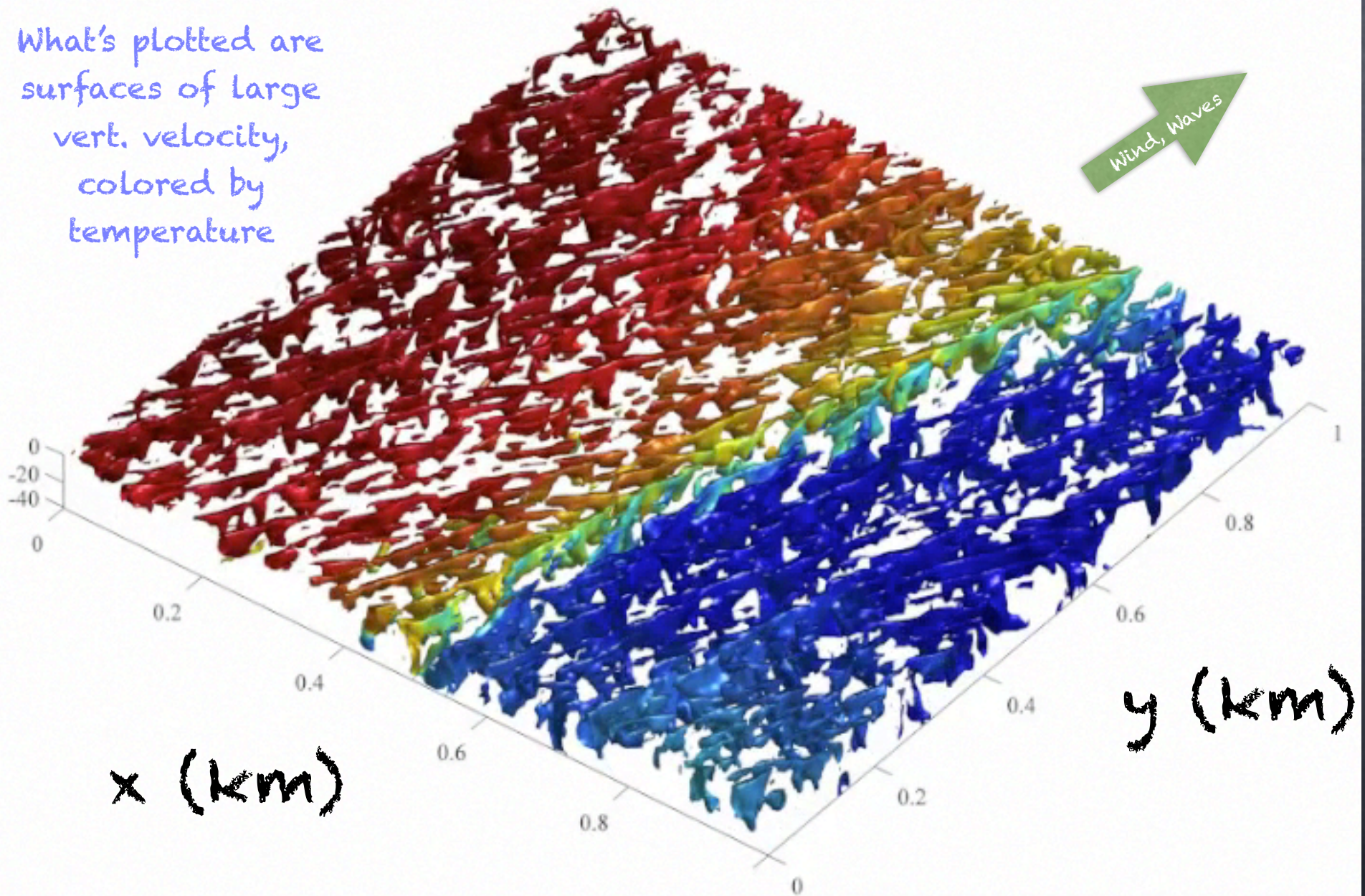
What's plotted are
surfaces of large
vert. velocity,
colored by
temperature



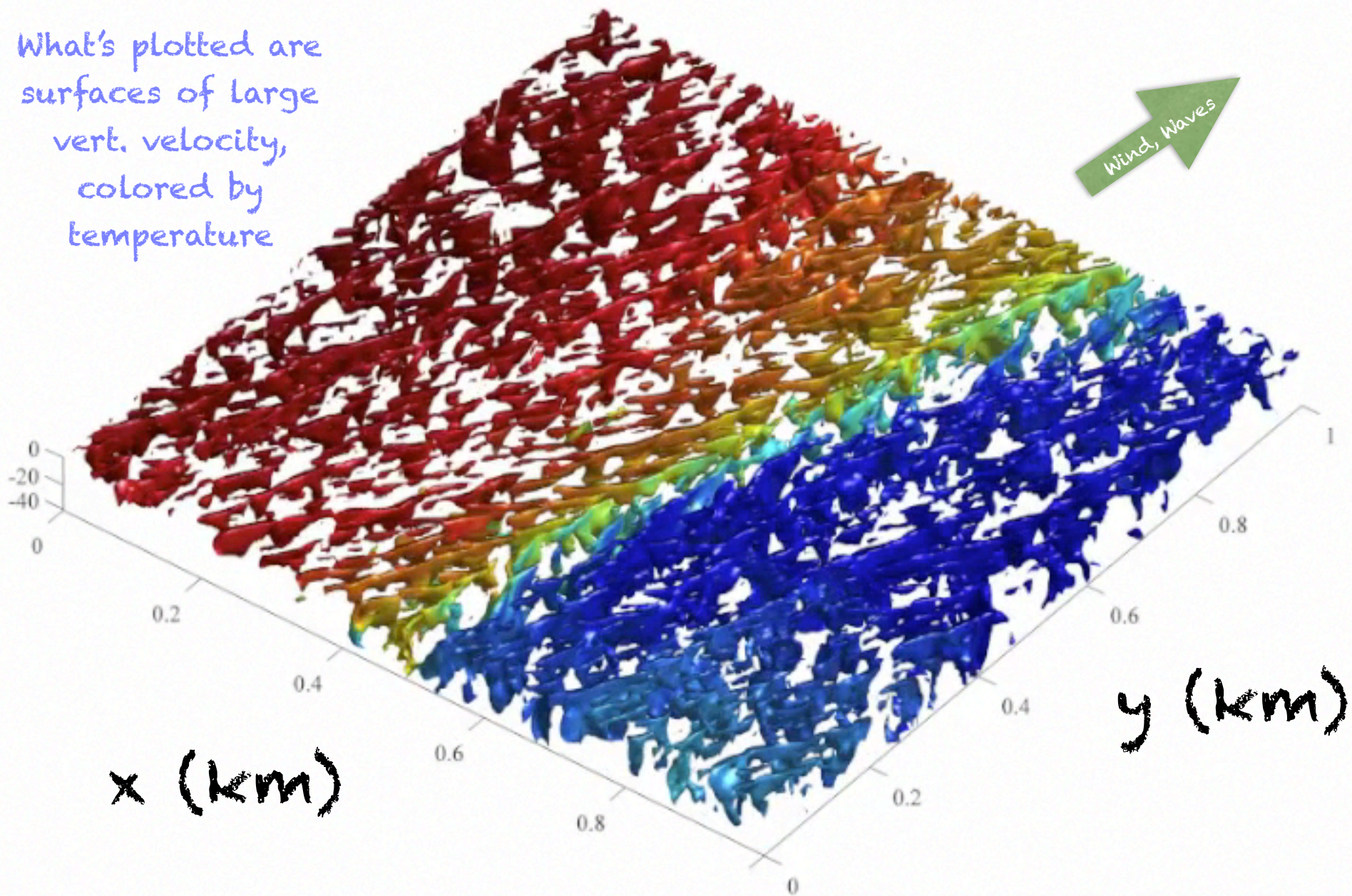
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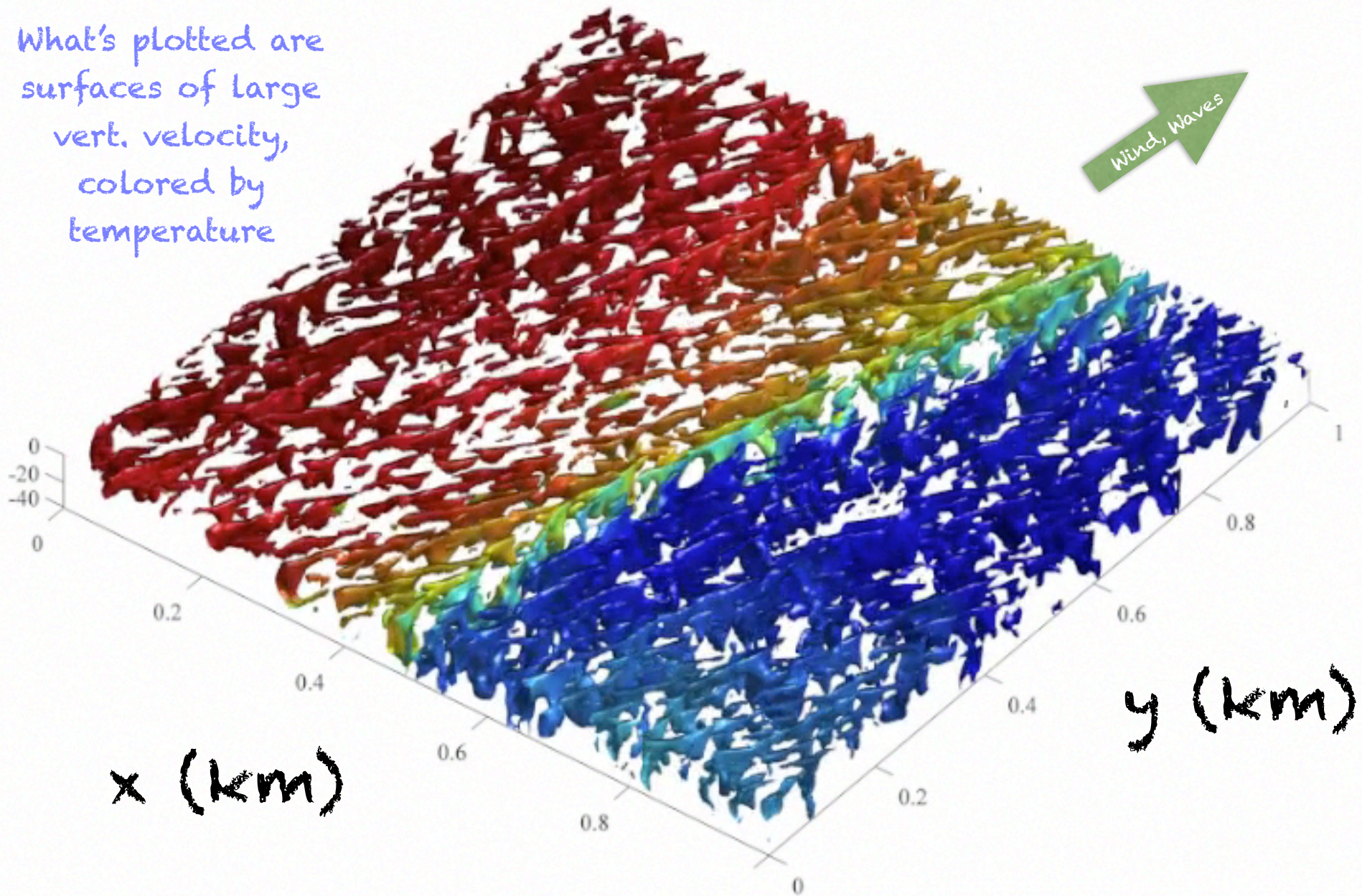
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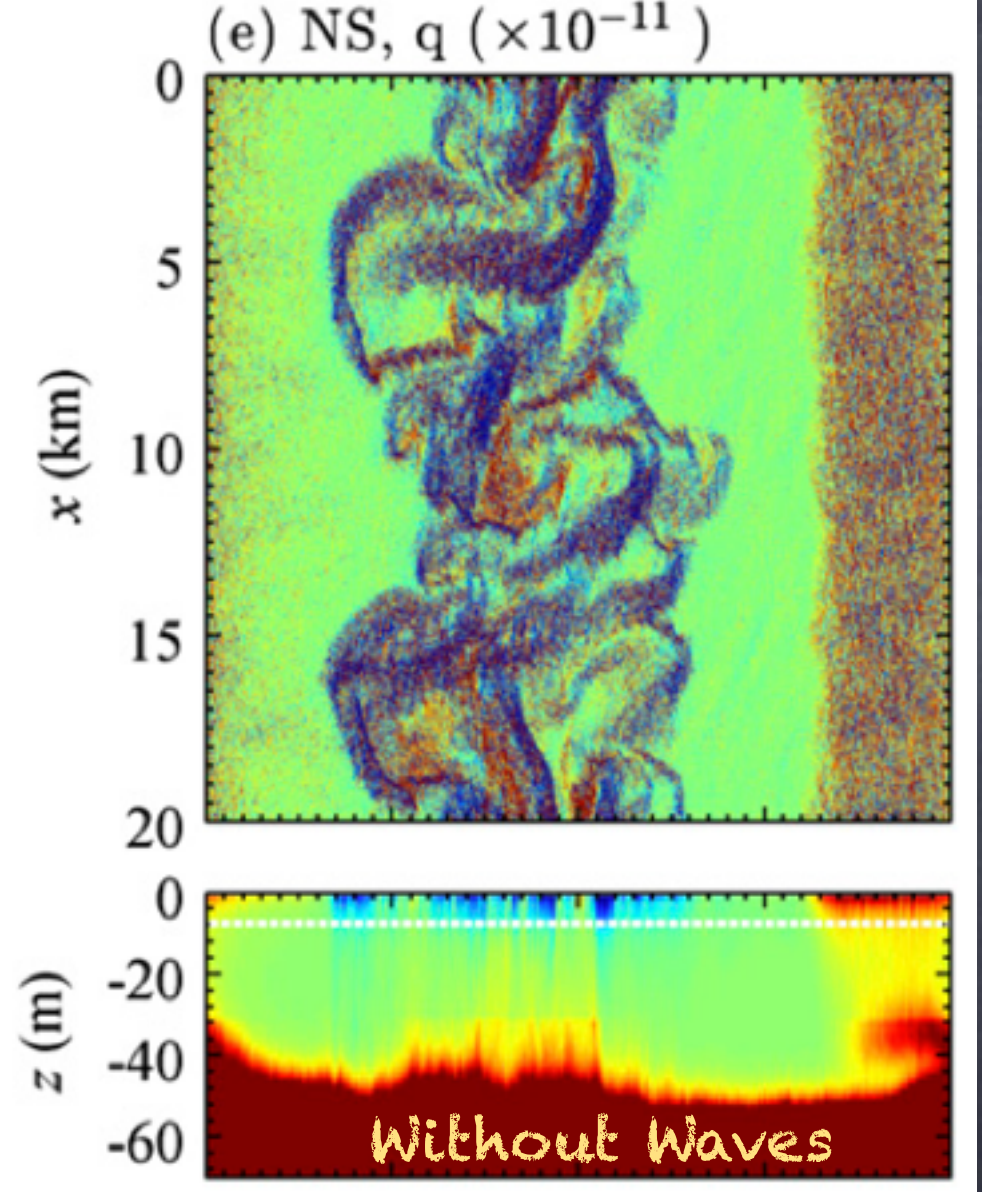
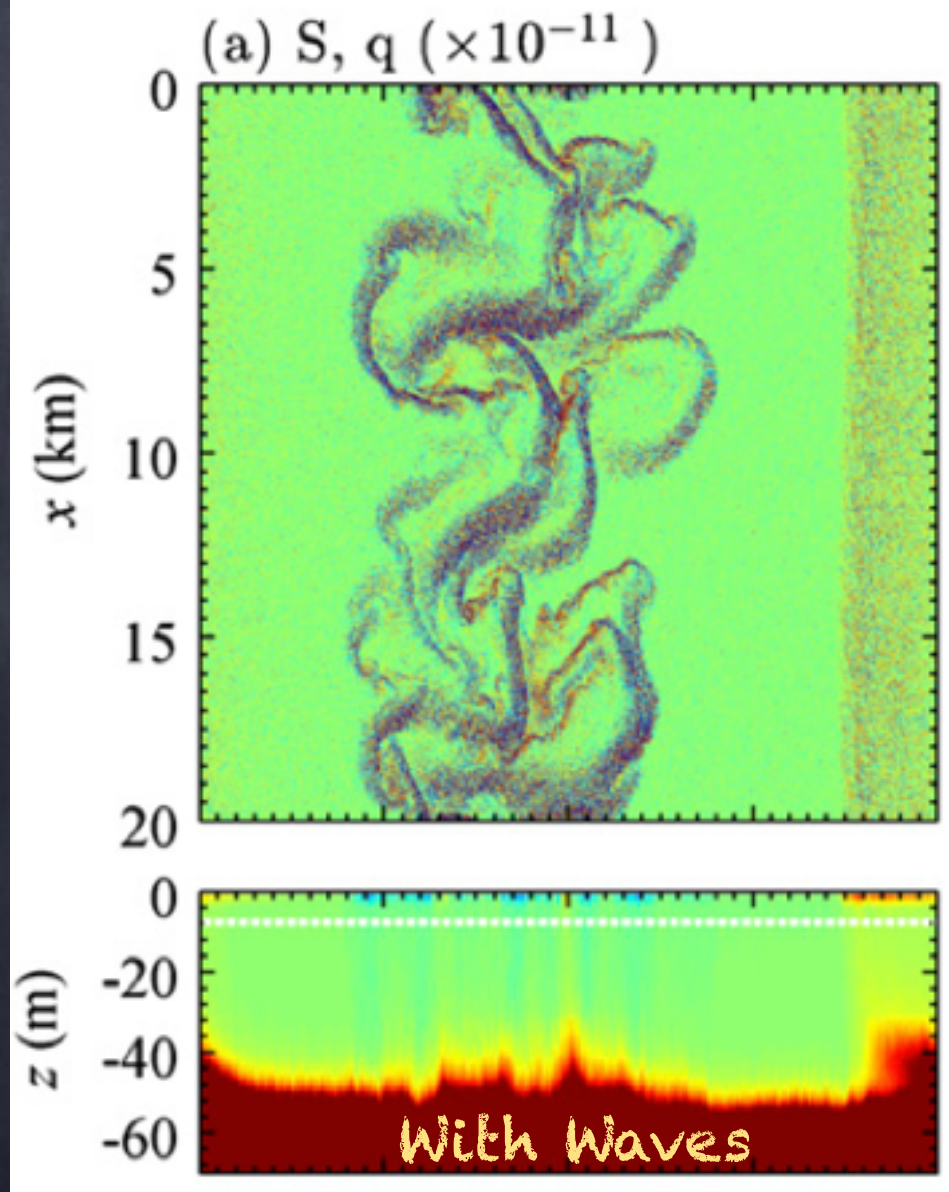


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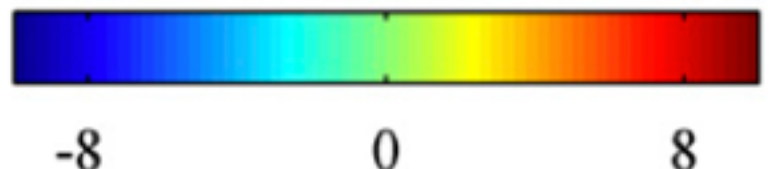




So, if $fQ < 0$ indicates likely regions of symmetric instability—Surface Waves STRONGLY affect SI!

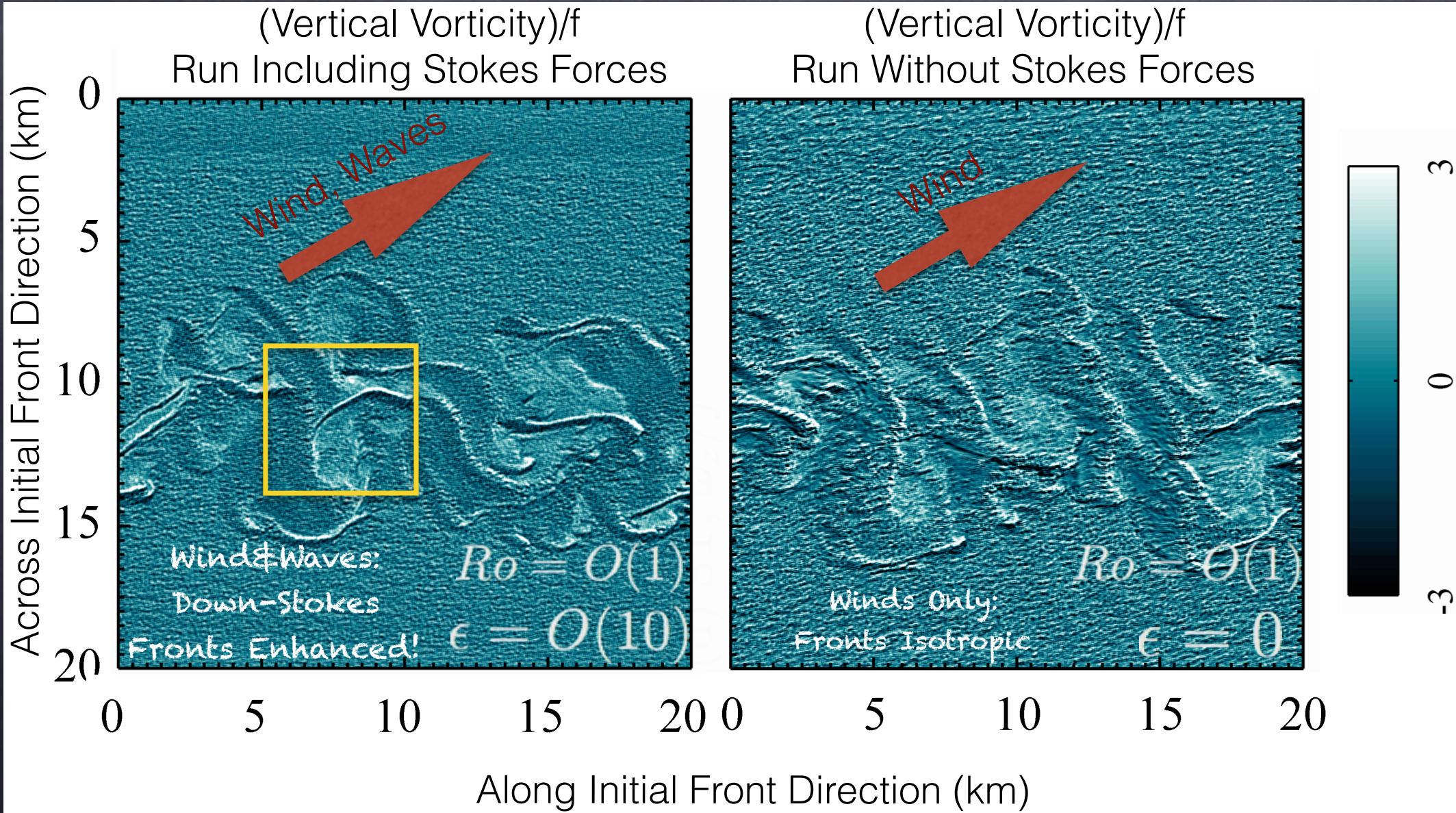
P. E. Hamlington, L. P. Van Roekel, BFK, K. Julien, and G. P. Chini. Langmuir-submesoscale interactions: Descriptive analysis of multiscale frontal spin-down simulations. *Journal of Physical Oceanography*, 44(9):2249-2272, September 2014

S. Haney, BFK, K. Julien, and A. Webb. Symmetric and geostrophic instabilities in the wave-forced ocean mixed layer. *JPO* 45:3033-3056, 2015.



Are Fronts and Filaments different with Stokes shear force?

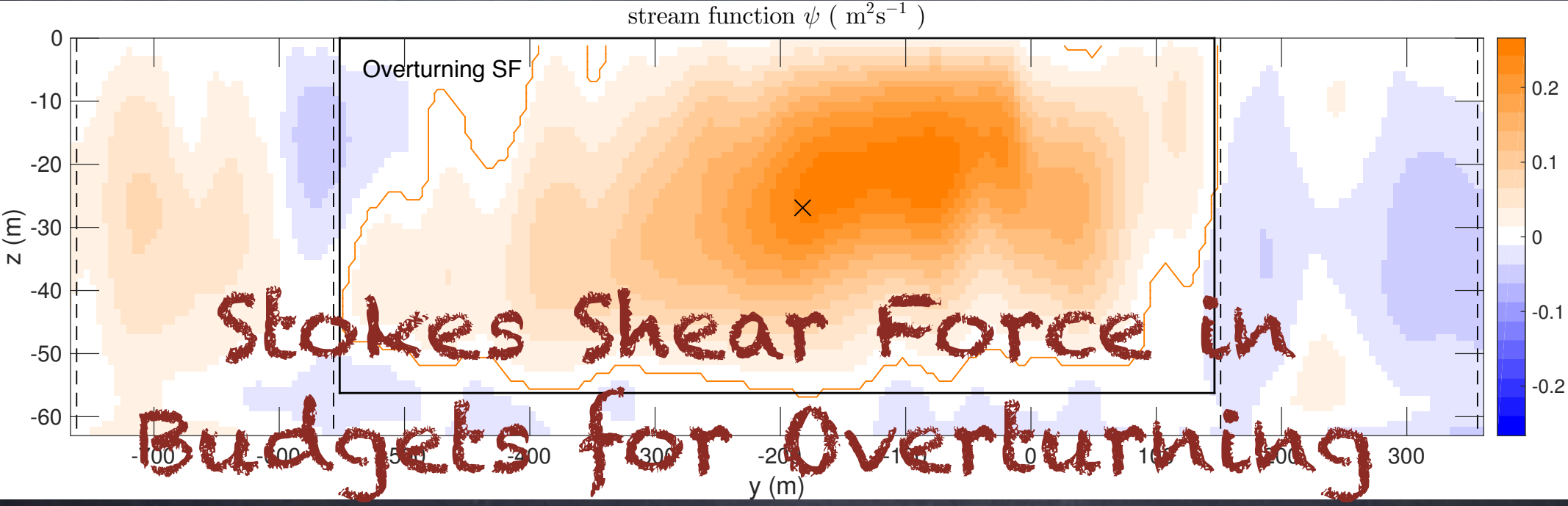
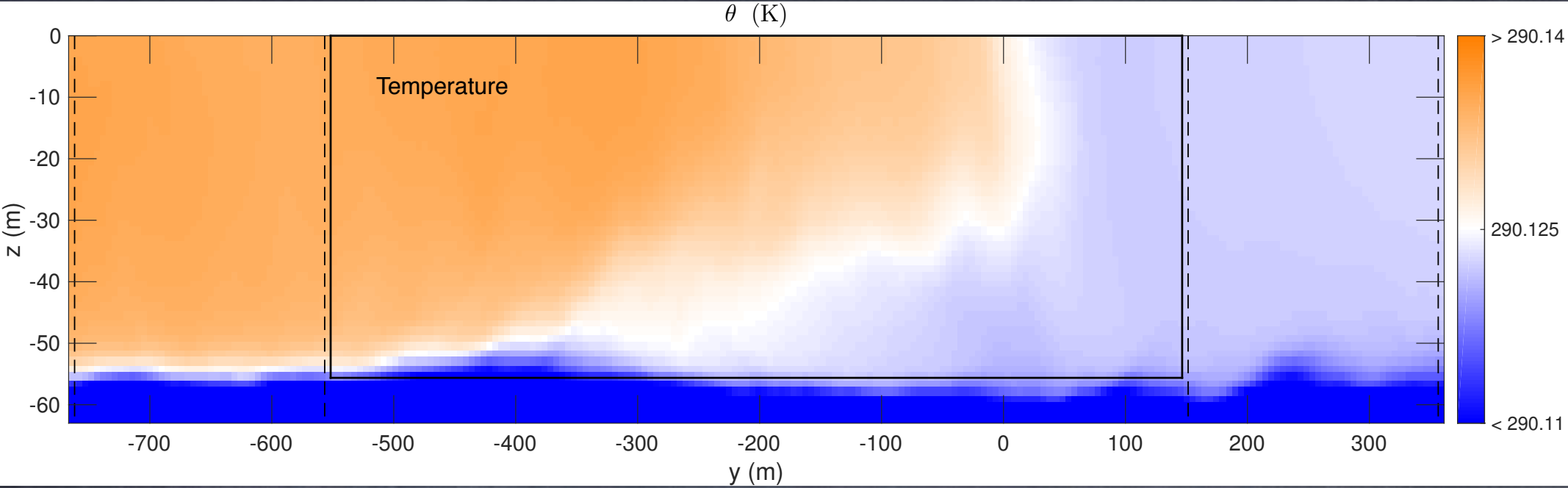
$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = -\pi_{,z} + b - \varepsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{Re Ri} w_{,jj}$$



N. Suzuki, BFK, P. E. Hamlington, and L. P. Van Roekel. Surface waves affect frontogenesis. JGR-Oceans, December 2015. Submitted.

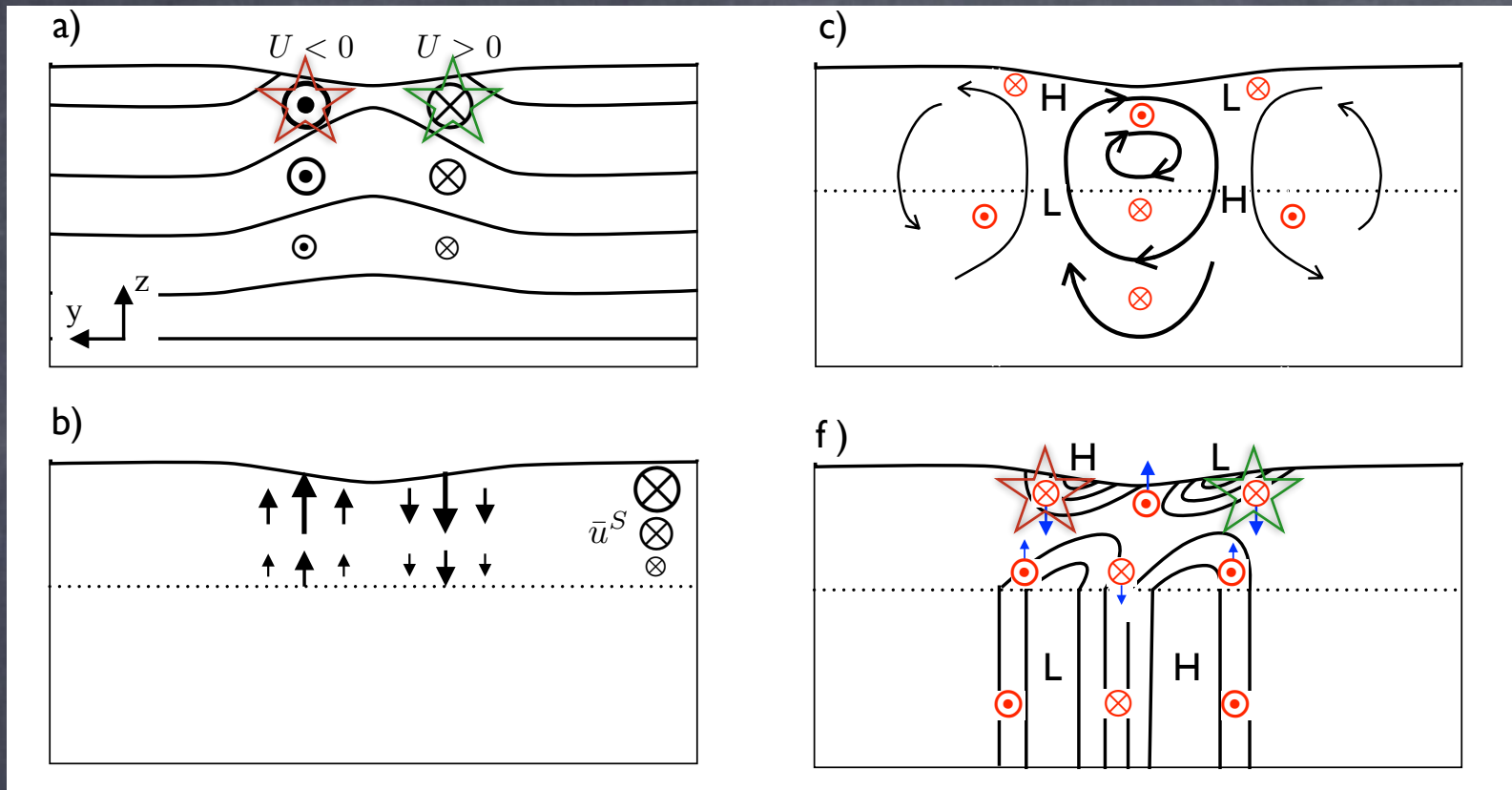
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2nd Largest Source for Overturning Kinetic Energy (4.4% of budget)

Stokes Shear Force Affects Fronts and Filaments



Enhances Fronts for Down-Front Stokes
 Opposes Fronts for Up-Front Stokes

$$\frac{\alpha^2}{Ri} \left[w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = -\pi_{,z} + b - \epsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{Re Ri} w_{,jj}$$

Conclusions

- In the upper ocean, horizontal scales as big as basins, and as small as meters contribute non-negligibly to the air-sea exchange and climate
- Interesting transition occurs on the submeso to Langmuir scale boundary, as nonhydro. & ageostrophic effects begin to dominate
- Langmuir mixing scalings consistent with LES & Obs., reduce climate model biases in MLD, T, CFCs vs. observations by 5-25%.
- The 25-45% forcing effects of the Stokes Shear force on submesoscale fronts & dynamics are under-appreciated.
- All papers at: fox-kemper.com/pubs

Oxford University Press is pleased
to announce the launch of their
new open access publication,
*Dynamics and Statistics of the
Climate System.*

EDITOR-IN-CHIEF:

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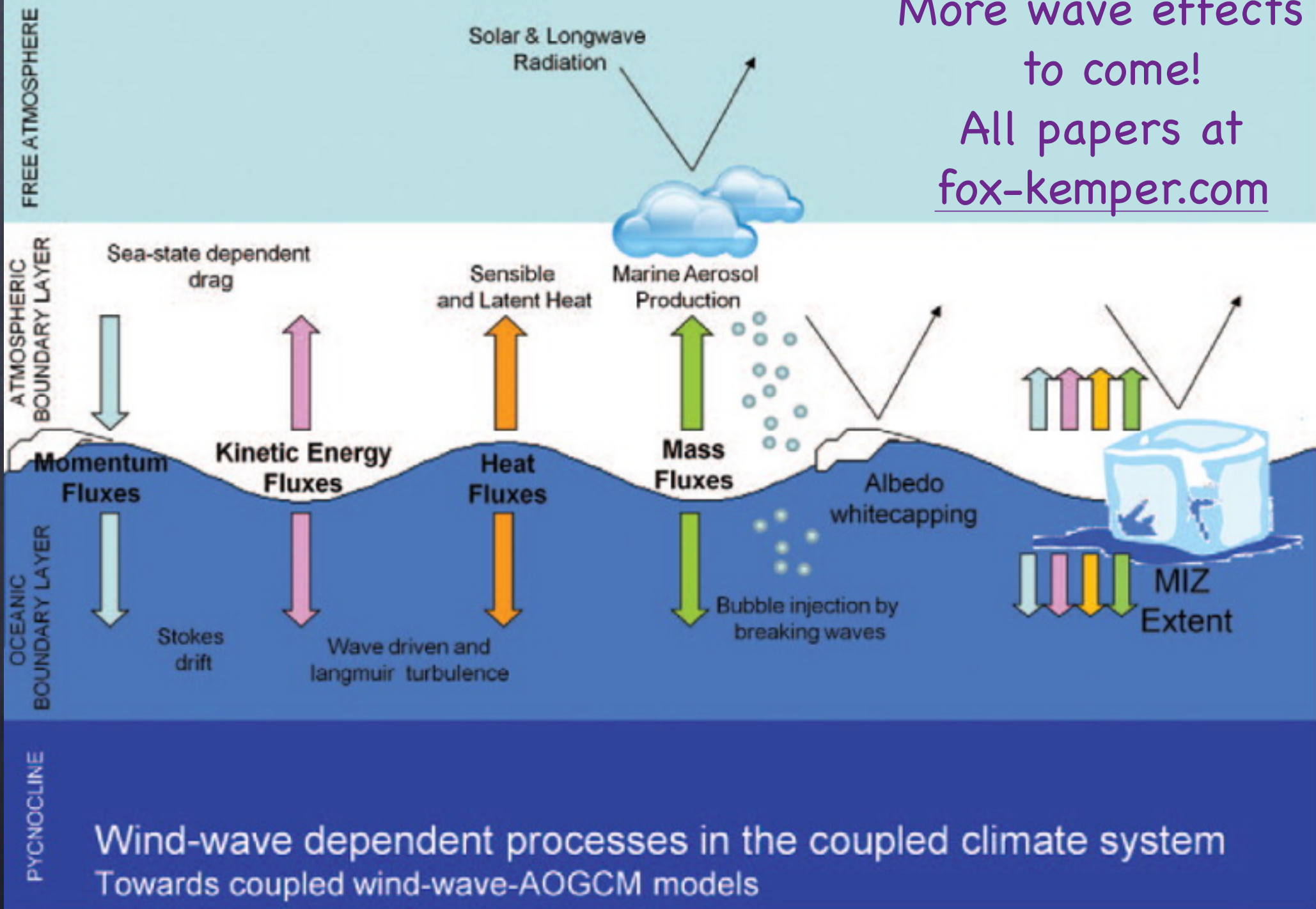
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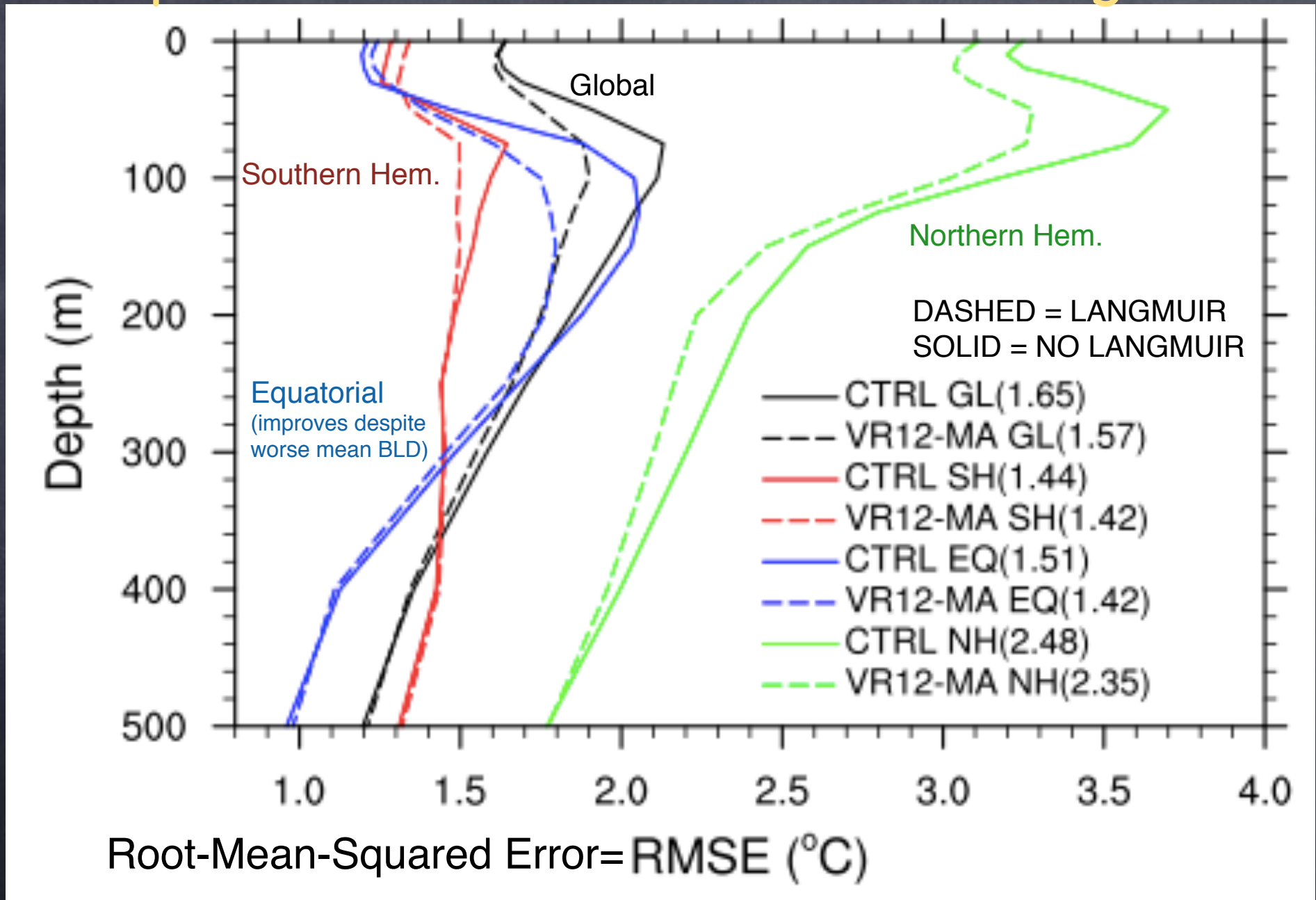
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More wave effects
to come!
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Subsurface Temperature: Improved vs. Observations with Langmuir



Enhancing ocean ventilation

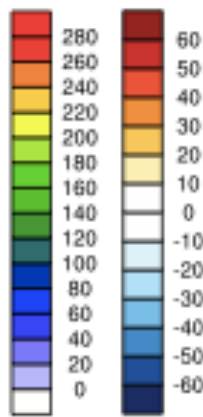
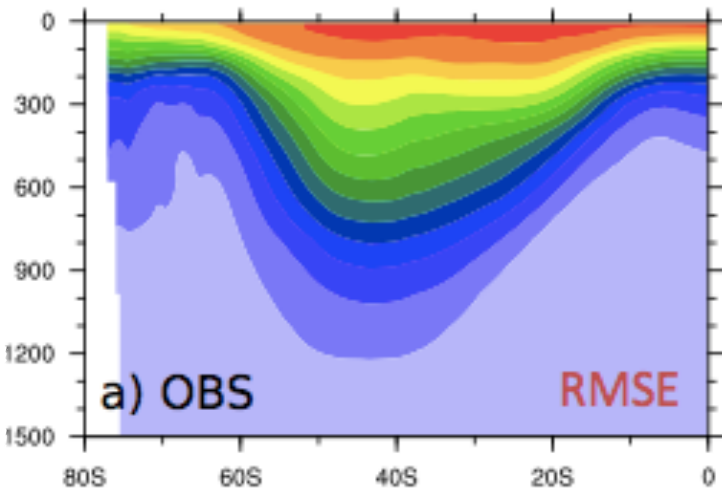
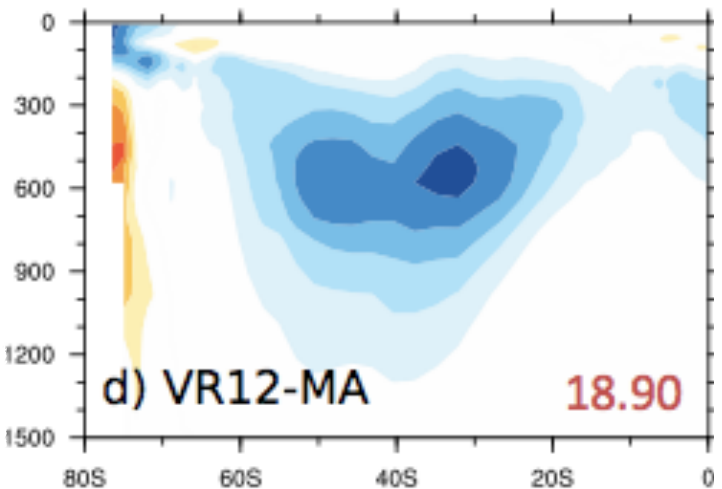
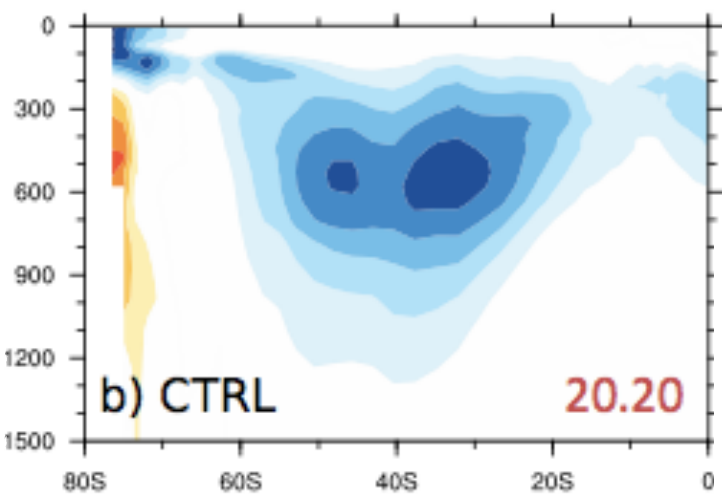


Fig. 3. Impact on the zonal mean pCFC-11 (patm) in the Southern Hemisphere.
 (a) Observation^[6] (GLODAP);
 (b) Biases in the control test without Langmuir mixing;
 (c) - (e) Biases in tests with Langmuir mixing.

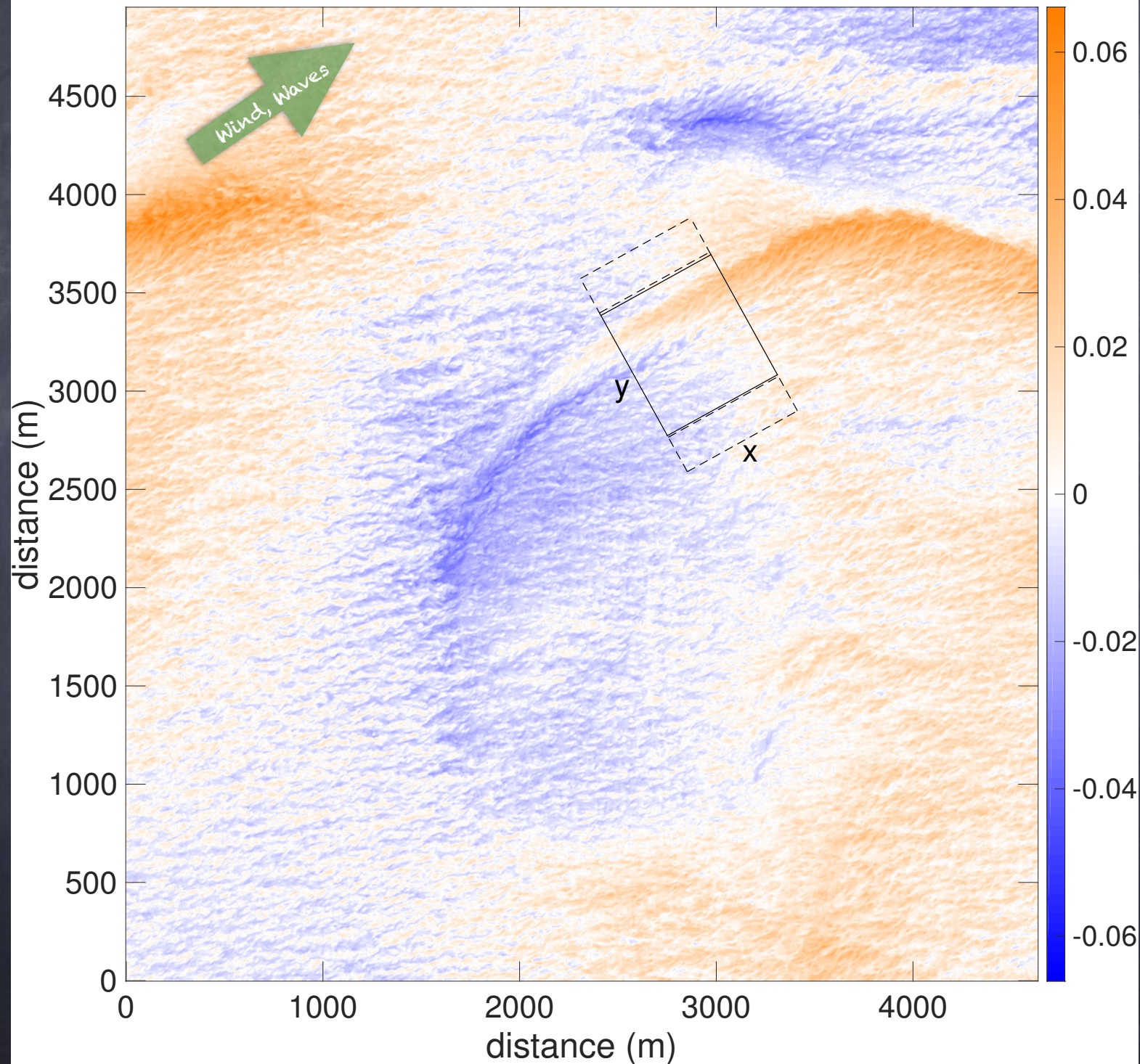


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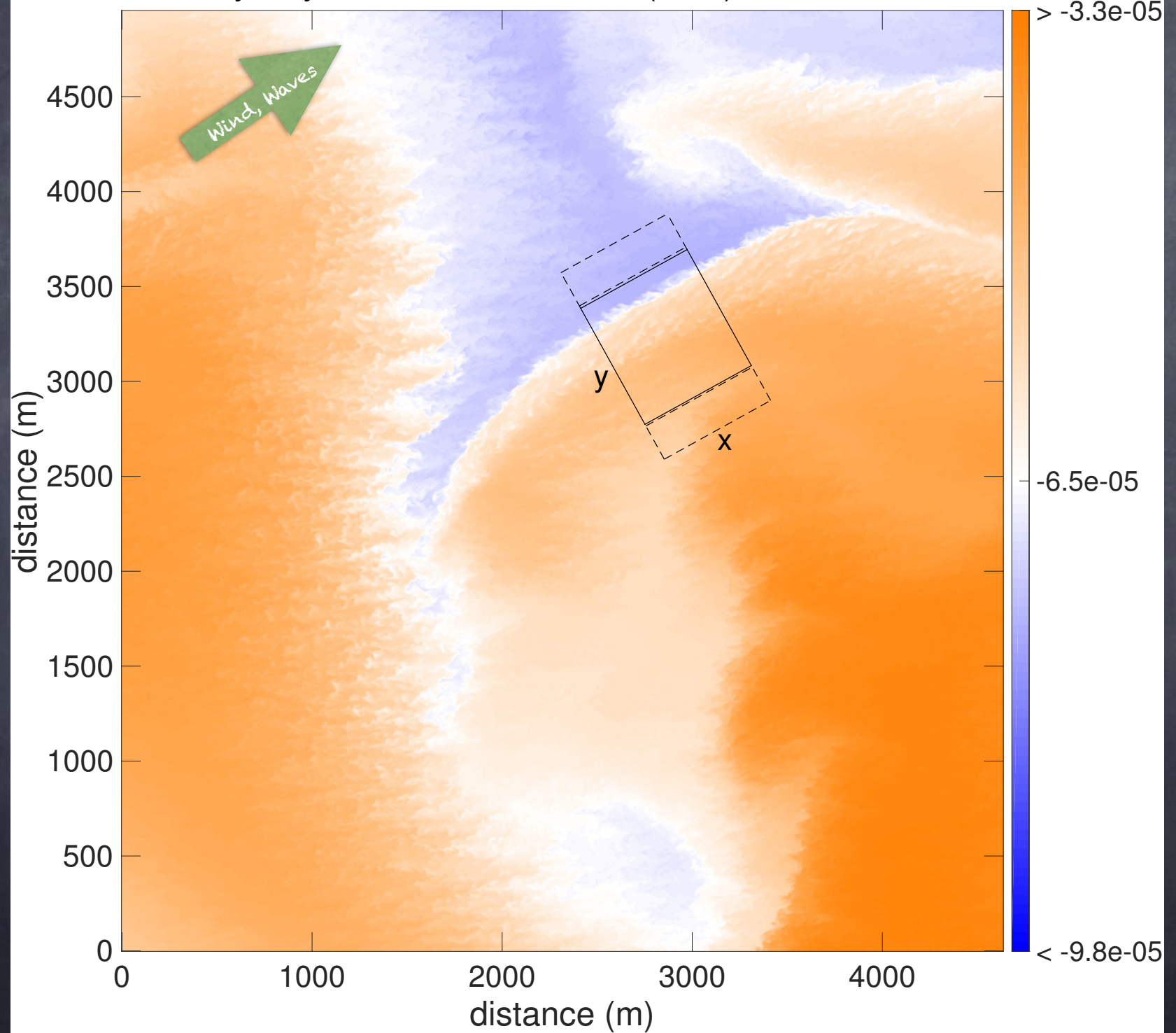
Ocean Uptake:
 Chlorofluorocarbons
 (manmade pollutant,
 detectable & known
 source)
 Improved vs.
 Observations with
 Langmuir Mixing

Case	Global	Southern Hemisphere
CTRL	23.90	20.20
MS2K	29.89	30.99
SS02	34.16	41.90
VR12-AL	22.14	18.53
VR12-MA	23.23	18.90
VR12-EN	20.67	16.44

velocity in the x-direction - the horizontal mean (ms^{-1}) at $z = -11.25\text{m}$



buoyancy - the horizontal mean (ms^{-2}) at $z = -11.25\text{m}$



vertical velocity (ms^{-1}) at $z = -11.25\text{m}$

