

Detection, Dynamics, and Consequences of Abyssal Ocean Variability

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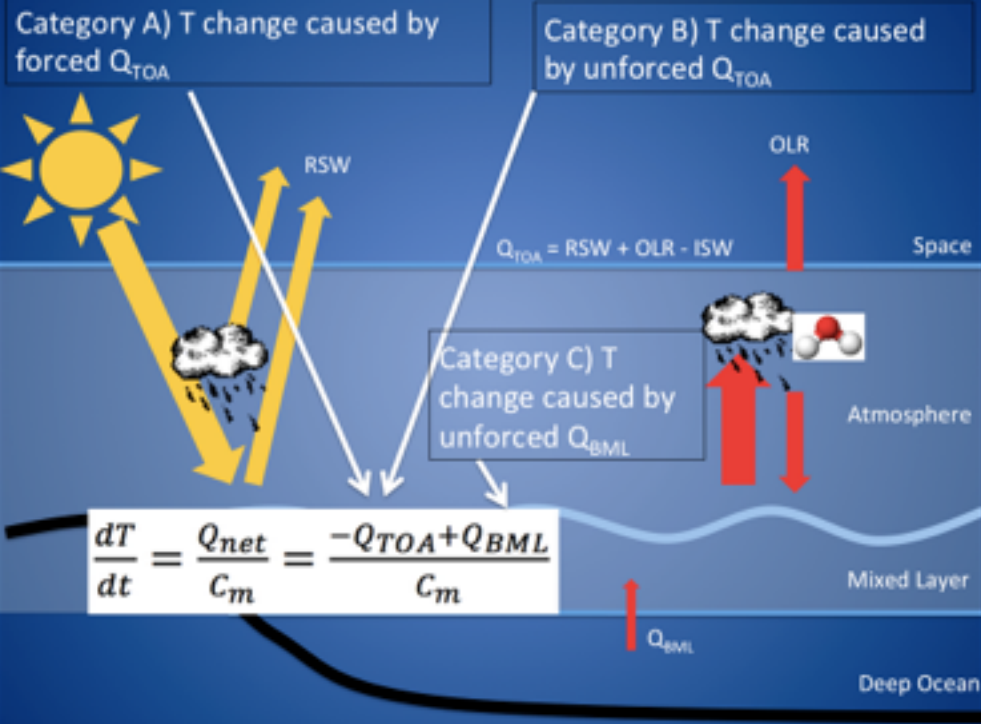
Atmosphere, Ocean, Planetary Physics Seminar
Oxford University 4/2/16

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and Society (IBES)

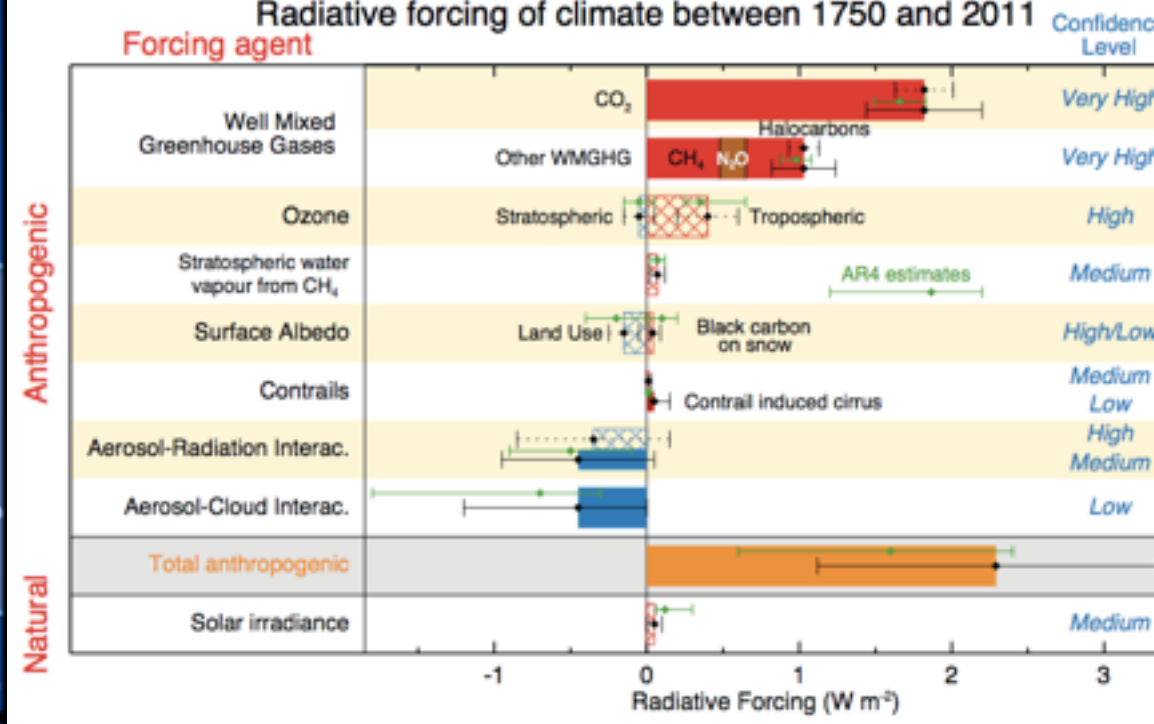


To understand ocean & climate variability,
it is important to distinguish:

- Presence of observable variability
- Understanding of past variability
- Modeling of variability
- Prediction of variability



Brown et al., 2014



IPCC AR5, 2013

Presence of observable variability

It is easier to observe the ocean consequences of air-sea exchange (ocean heat content (OHC), salinity) rather than exchanges (fluxes) themselves.

However, insufficient for prediction and attribution



Prediction & Attribution Goal: Effects of Anthropogenic Forcing



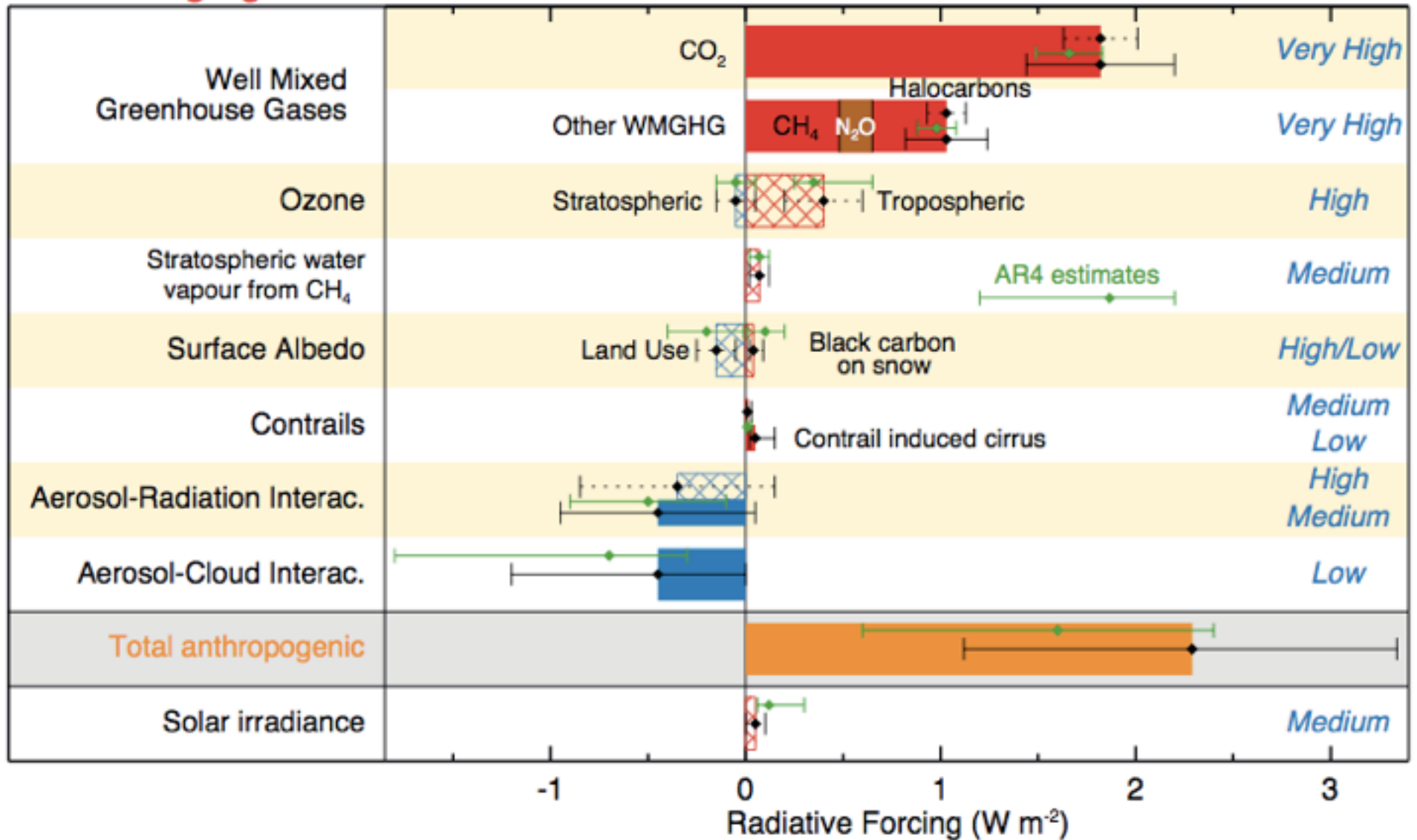
Radiative forcing of climate between 1750 and 2011

Forcing agent

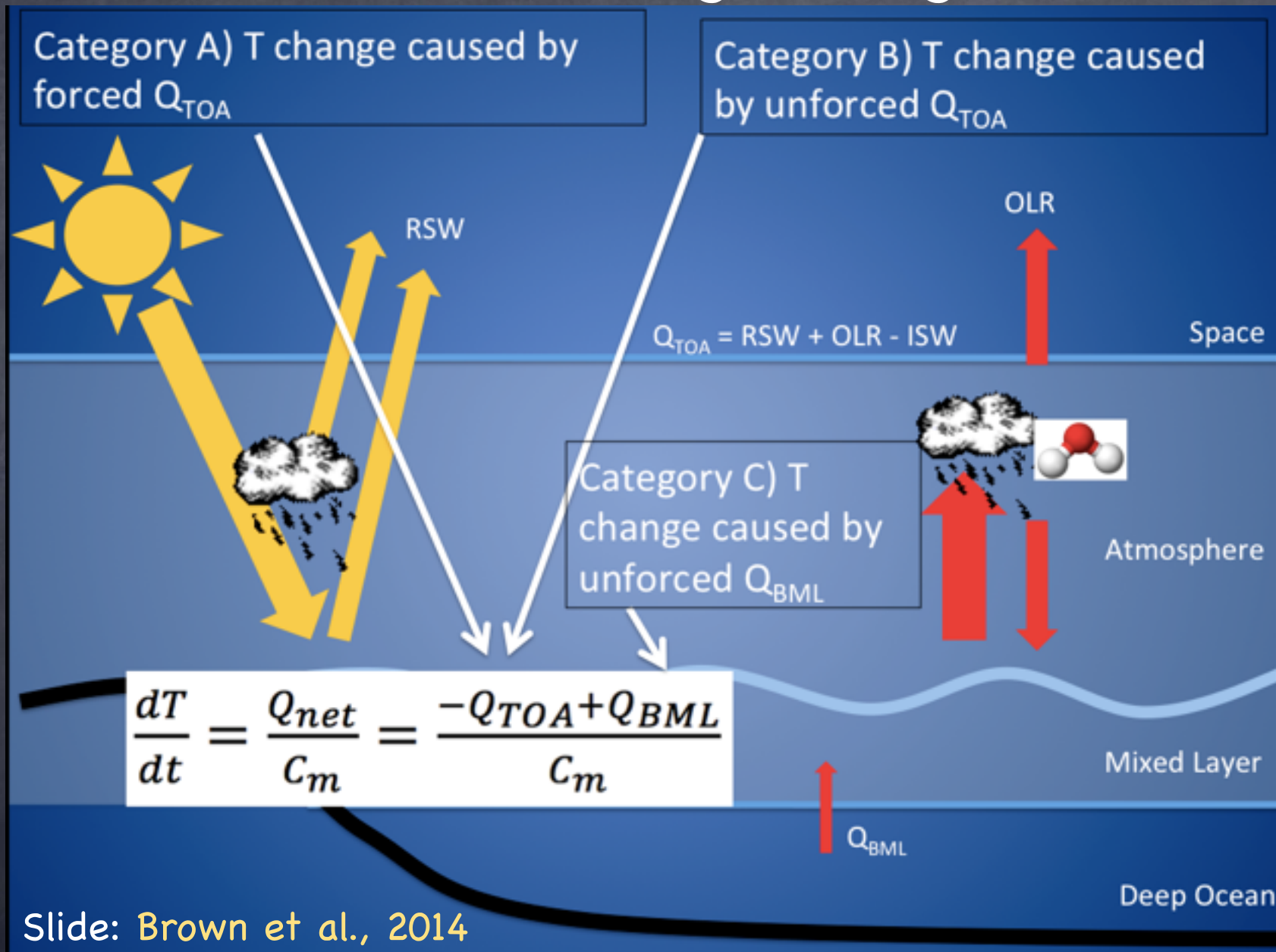
Confidence Level

Anthropogenic

Natural



Surface Energy Budget



Slide: Brown et al., 2014

- $O(2W/m^2)$ change to Q_{BML} as important as GHG
- Slight oversimplification—sensitivity + budget



What do hydrographic observations show?

Ocean Heat Content not fixed: Q_{BML} not zero (and varies)!

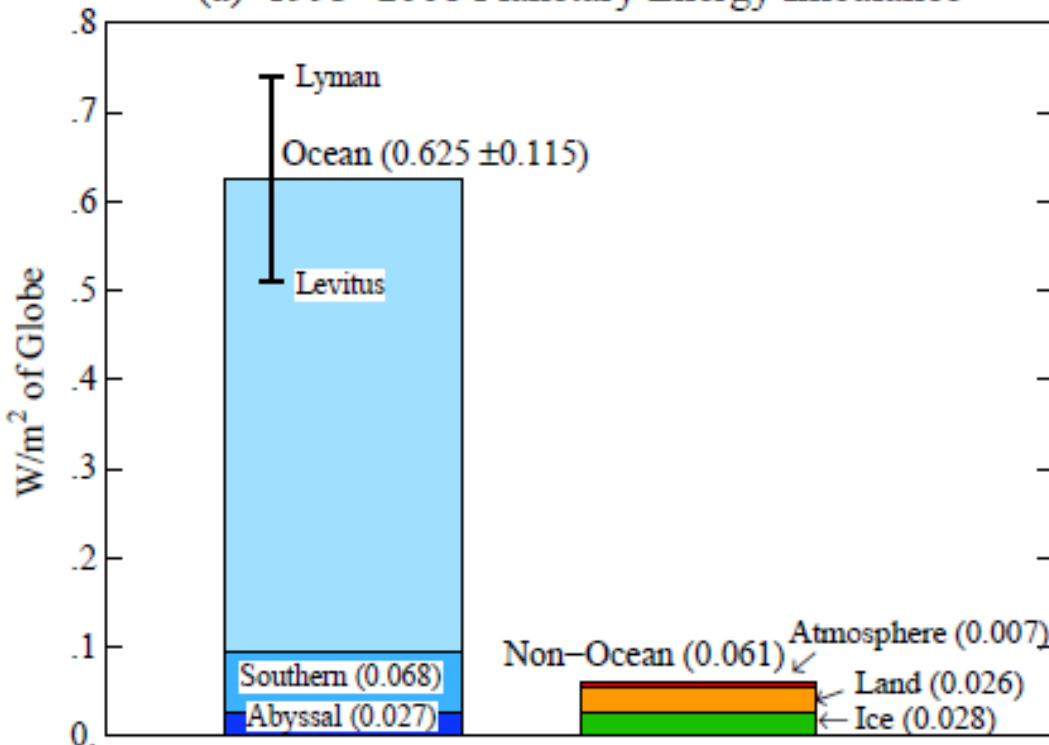
28% of anthropogenic forcing equals the warming in the oceans and about 70% goes back to space.

90% of anomalous warming is in the oceans.

0.7 W/m^2 to atmosphere only is about 1.5K/yr

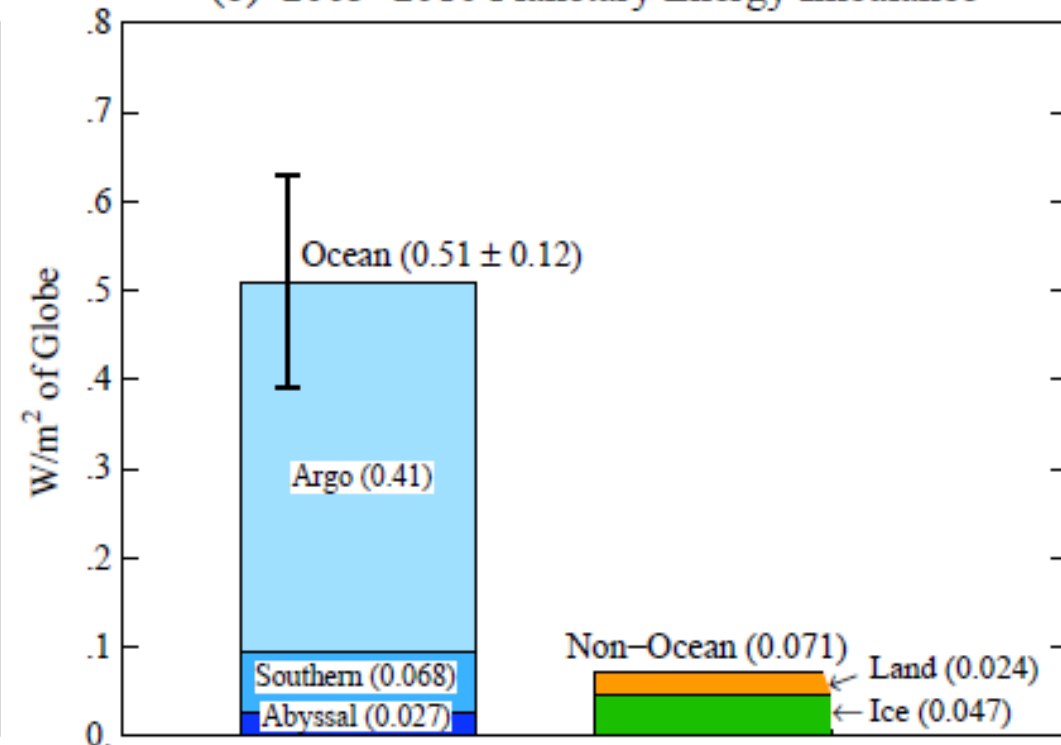
Trad. Hydrography

(a) 1993–2008 Planetary Energy Imbalance



From the Argo Era

(b) 2005–2010 Planetary Energy Imbalance

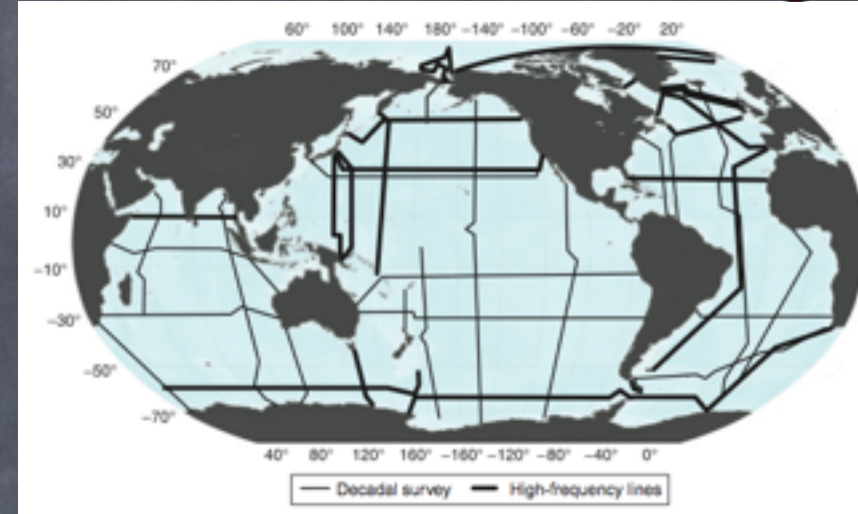


Hansen et al. (2011)

How do we know OHC?

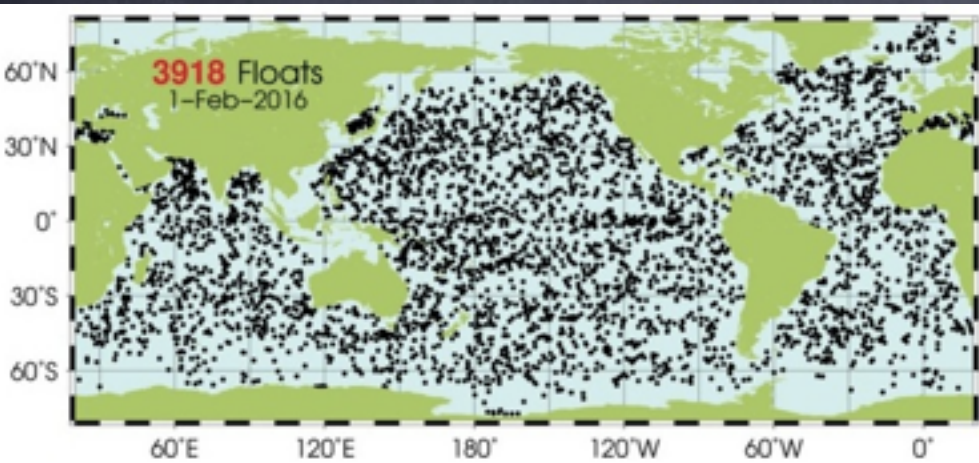


Traditional Hydrography (<http://www.ukosnap.org/>)



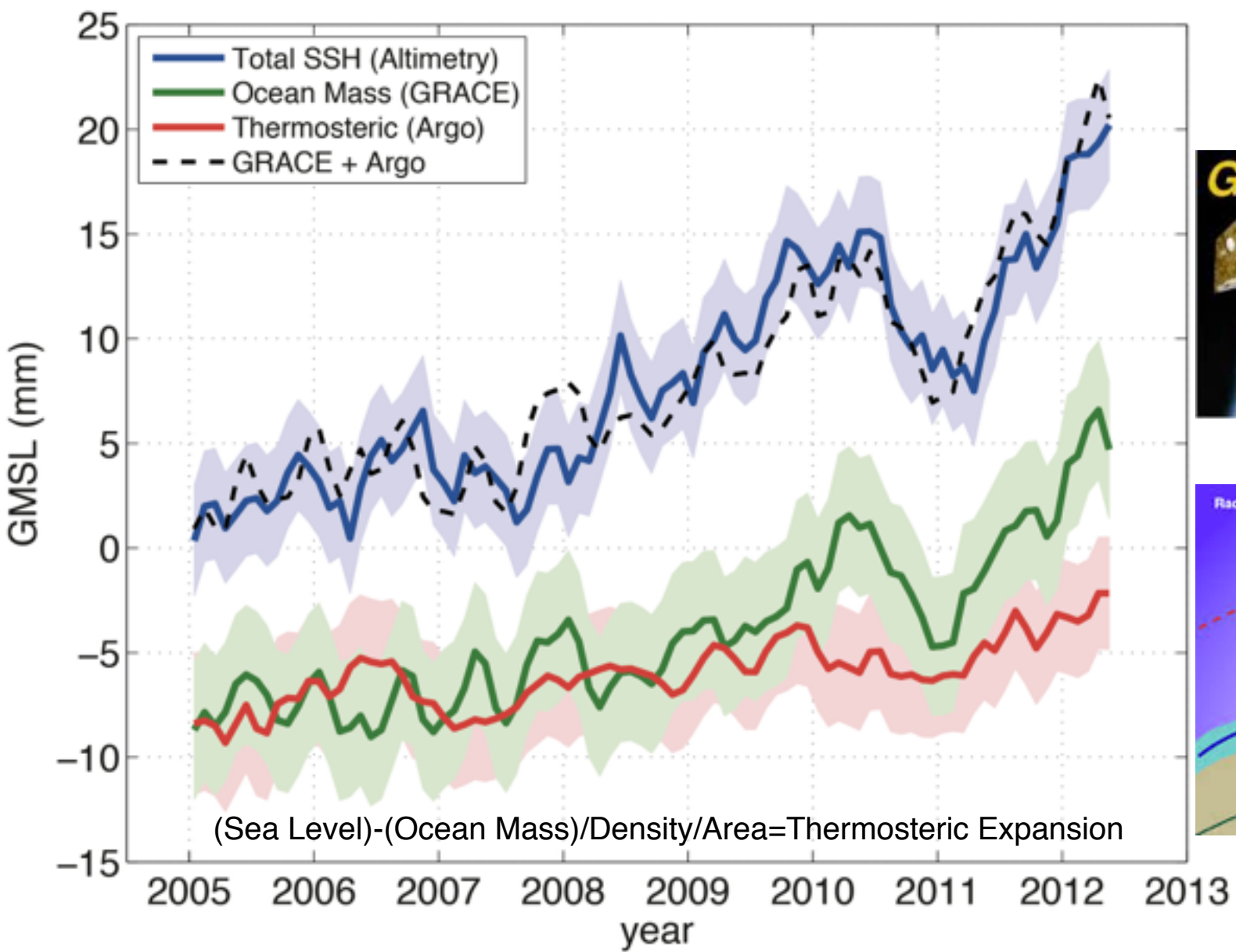
GO-SHIP repeat sections: Siedler et al. 2013

Autonomous: e.g., Argo and Satellites.
<http://www.argo.ucsd.edu/>

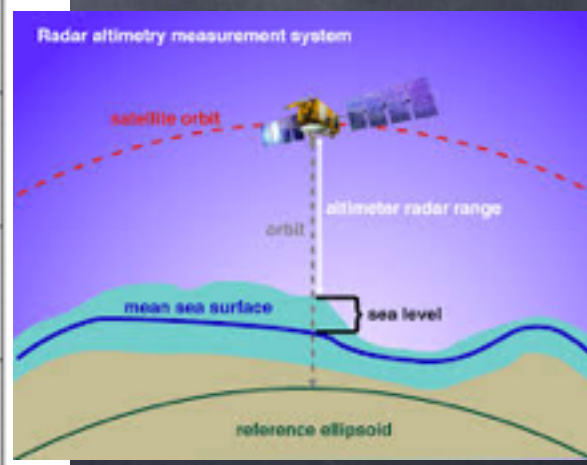


Argo floats presently active

Another reason to care about ocean warming—and to observe it (by subtraction): Sea Level Rise



podaac.jpl.nasa.gov



nesdis.noaa.gov



Surface, Mixed Layer, Seasons?

0.7 W/m²

=

Atmosphere:

1.5K/yr

=

3.4m Ocean:

1.5K/yr

=

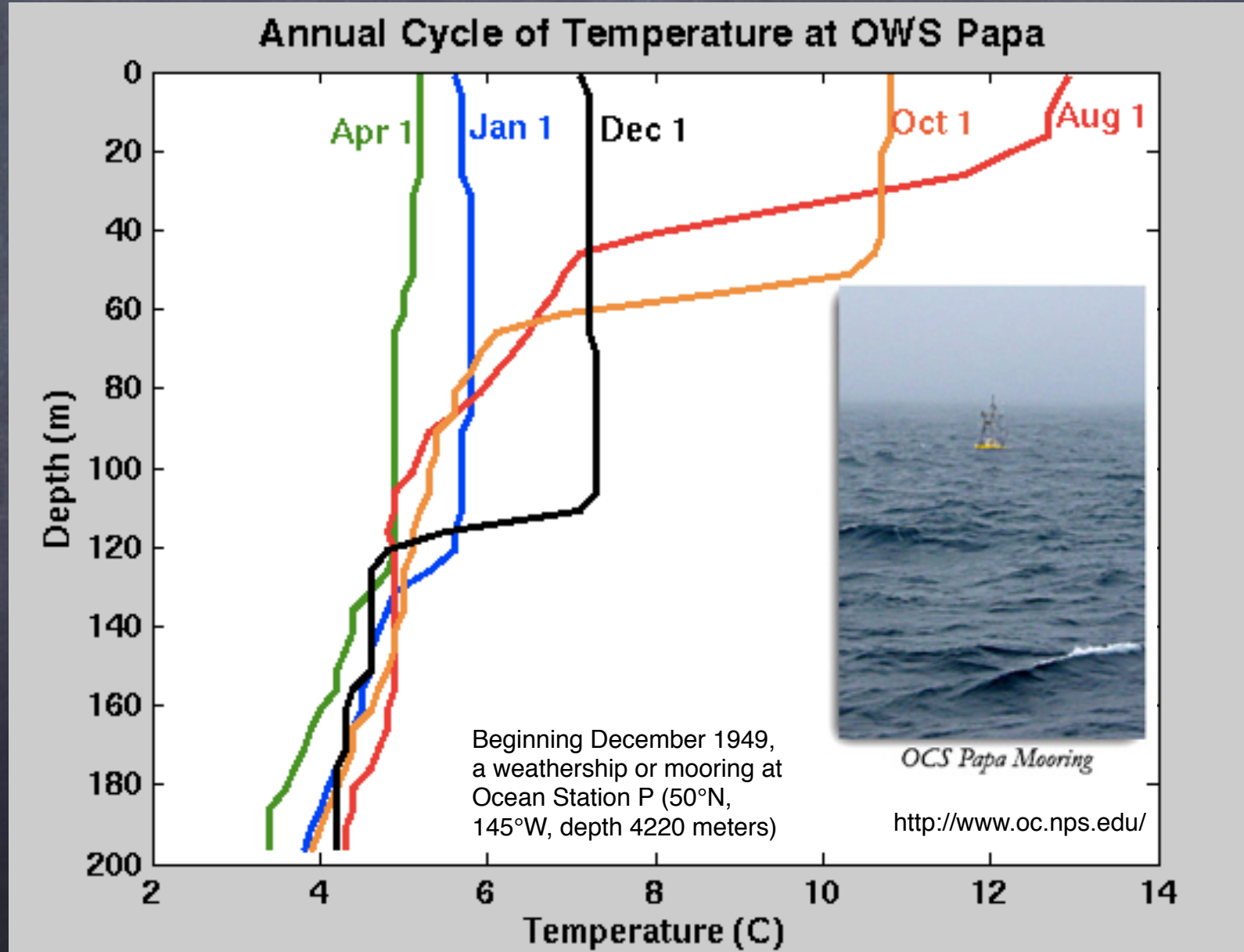
34m Ocean:

0.15K/yr

=1% of

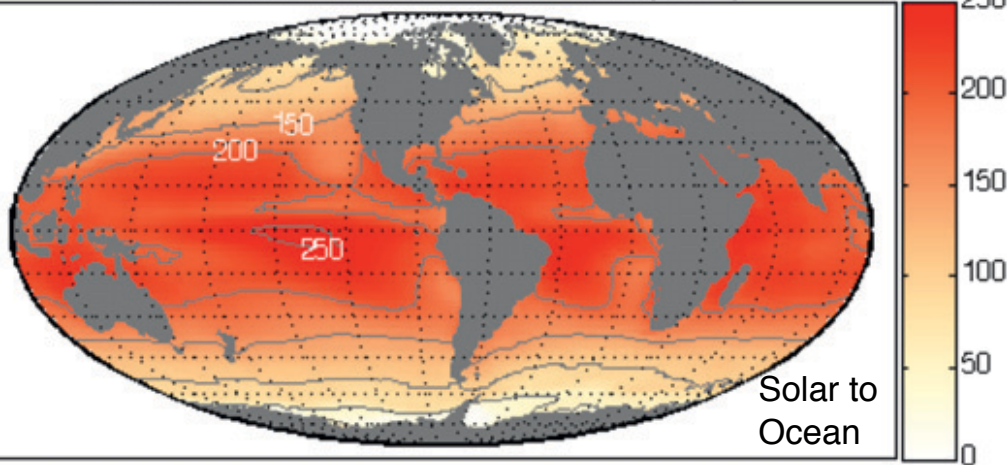
mixed layer

seasonality

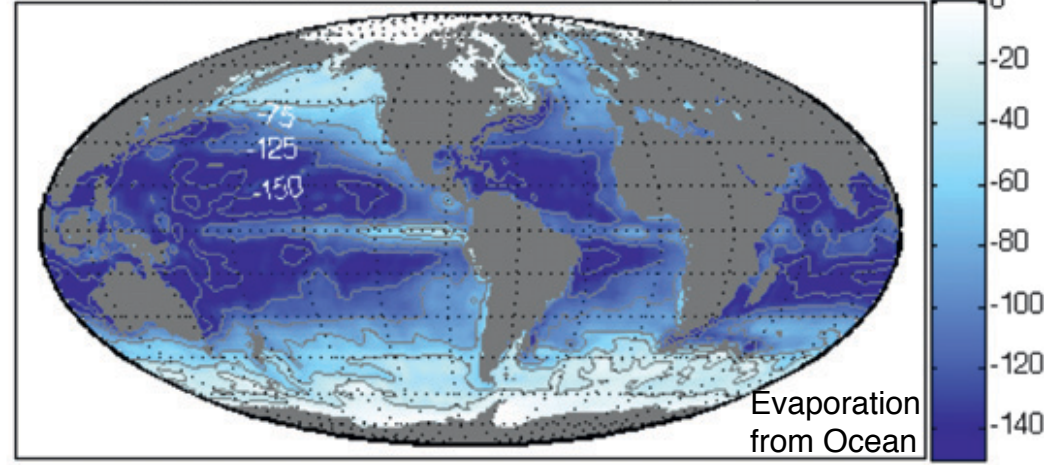


The net Q_{BML} is also about 1% of different flux components and about 1% of net spatial extremes

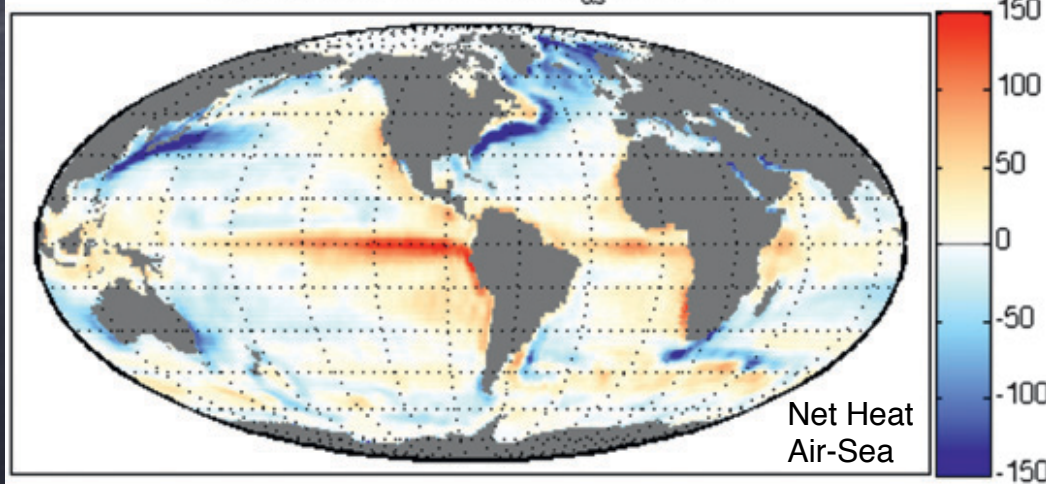
Mean of 1986-2005 CORE net sw heat flux (W/m^2)



Mean of 1986-2005 CORE latent heat flux (W/m^2)



Mean of 1986-2005 CORE Q_{as} (W/m^2)

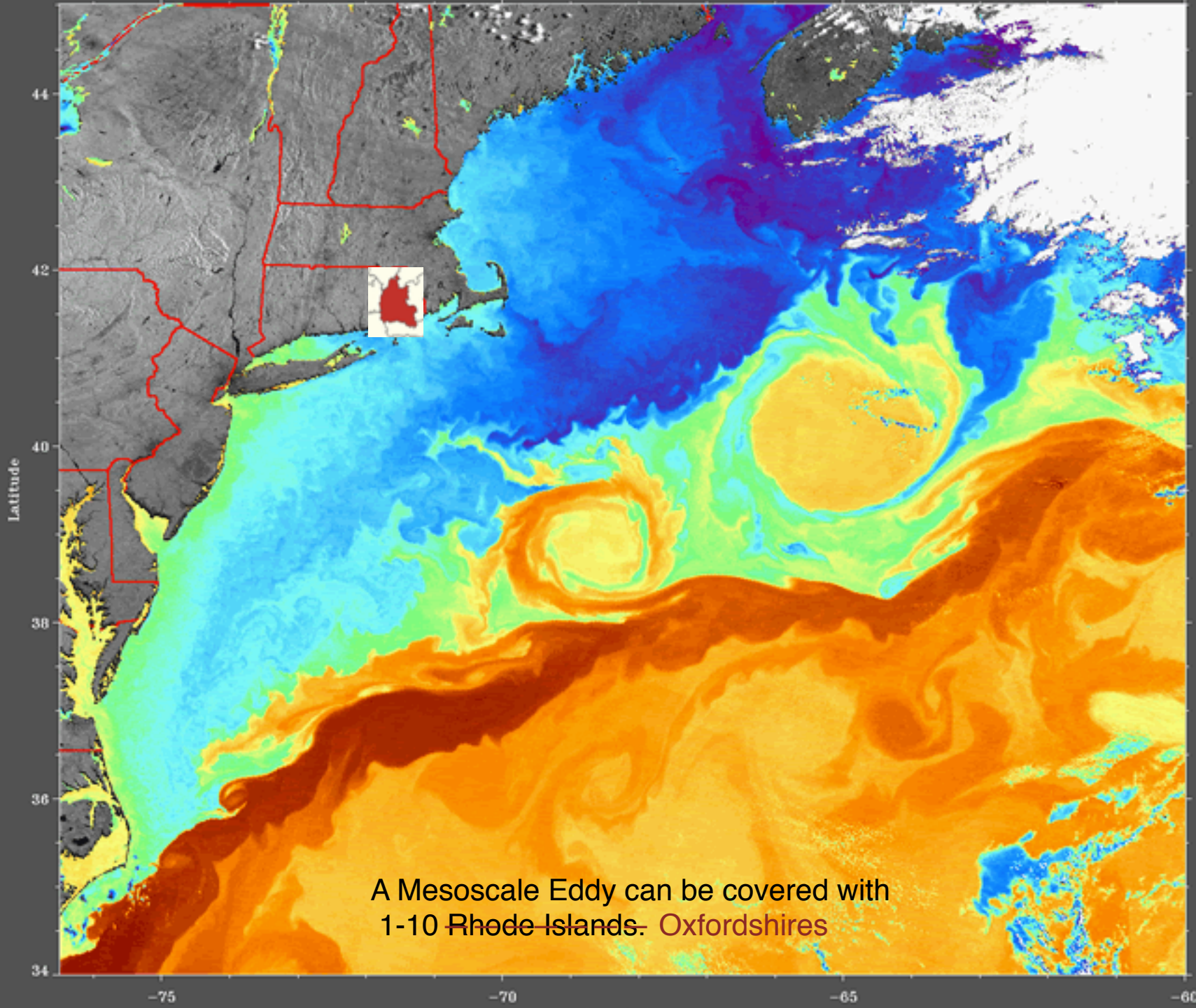
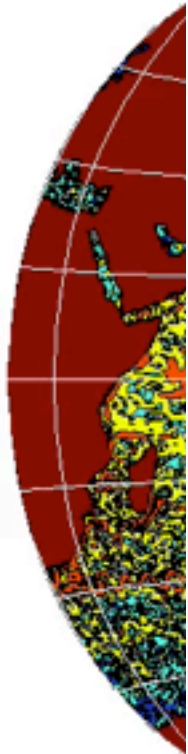


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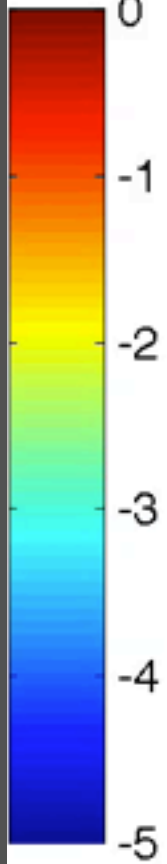
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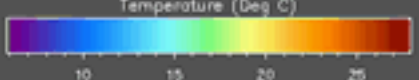
Satellite
view of
flows

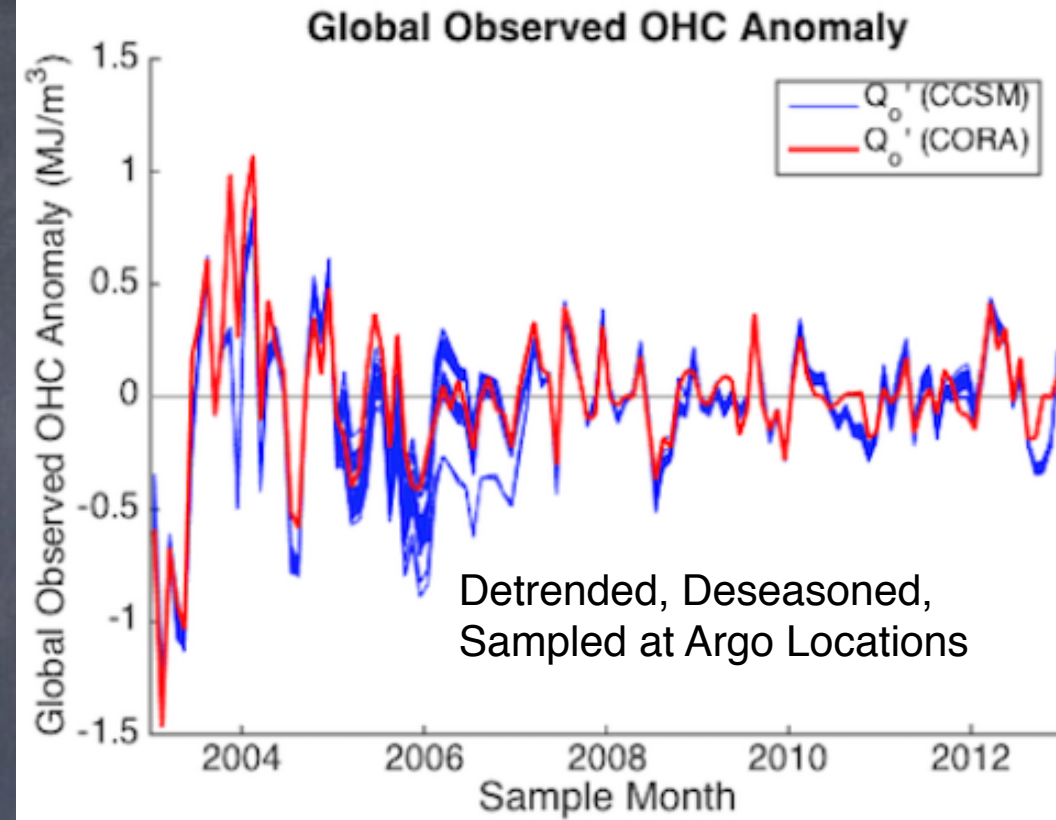
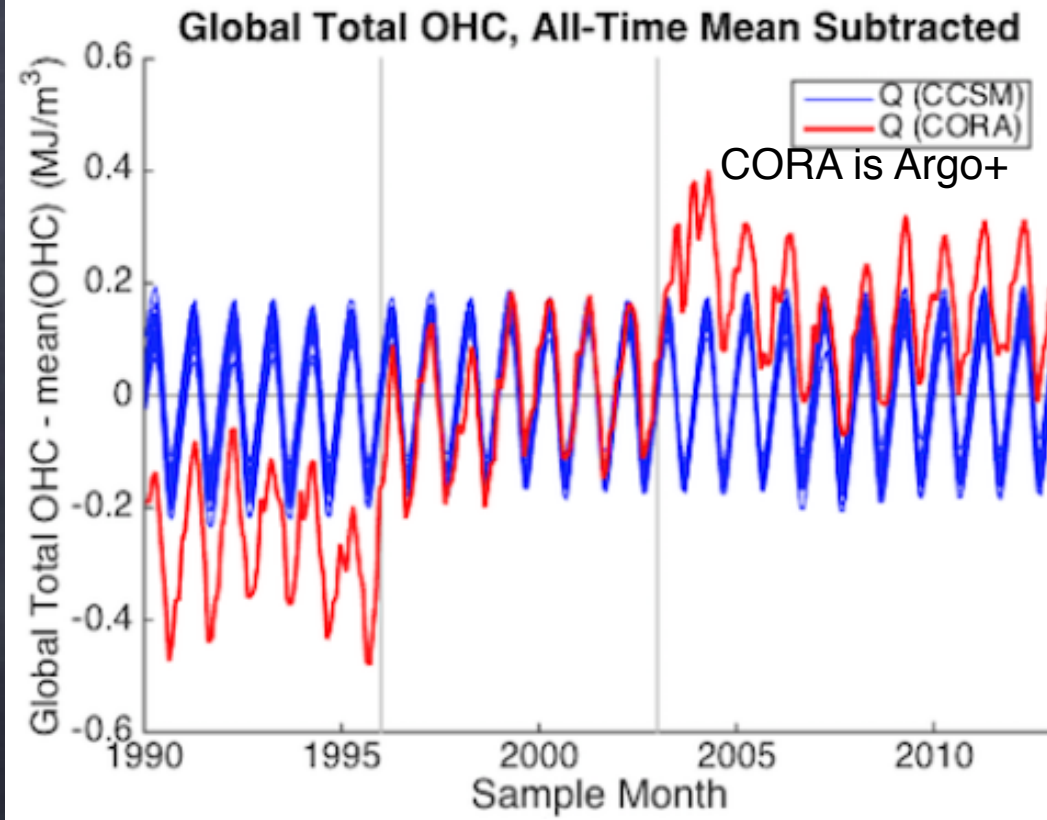


A Mesoscale Eddy can be covered with
1-10 Rhode Islands. Oxfordshires



WATER SURFACE TEMPERATURE
Land and Clouds from Channel 2
NOAA-12 AVHRR 1997 Jun 11 11:27 UT





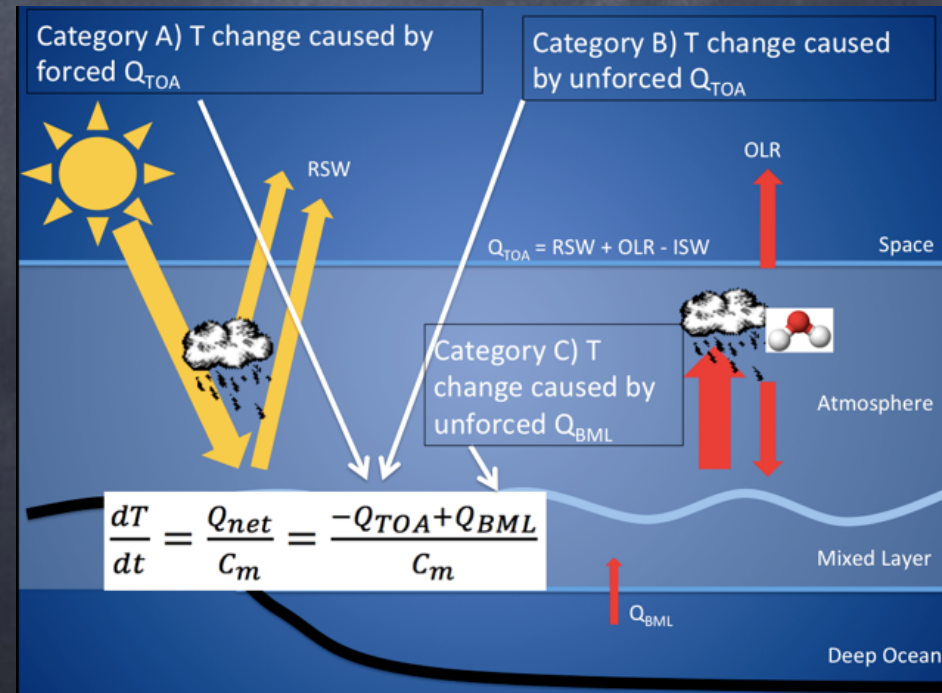
Sophisticated analysis to overcome Ship & Argo sampling problems—inherent uncertainty, $O(0.2W/m^2)$, on interannual to decadal timescales in global average. $O(10W/m^2)$ without analysis.



CU, soon Brown

Presence of observable variability

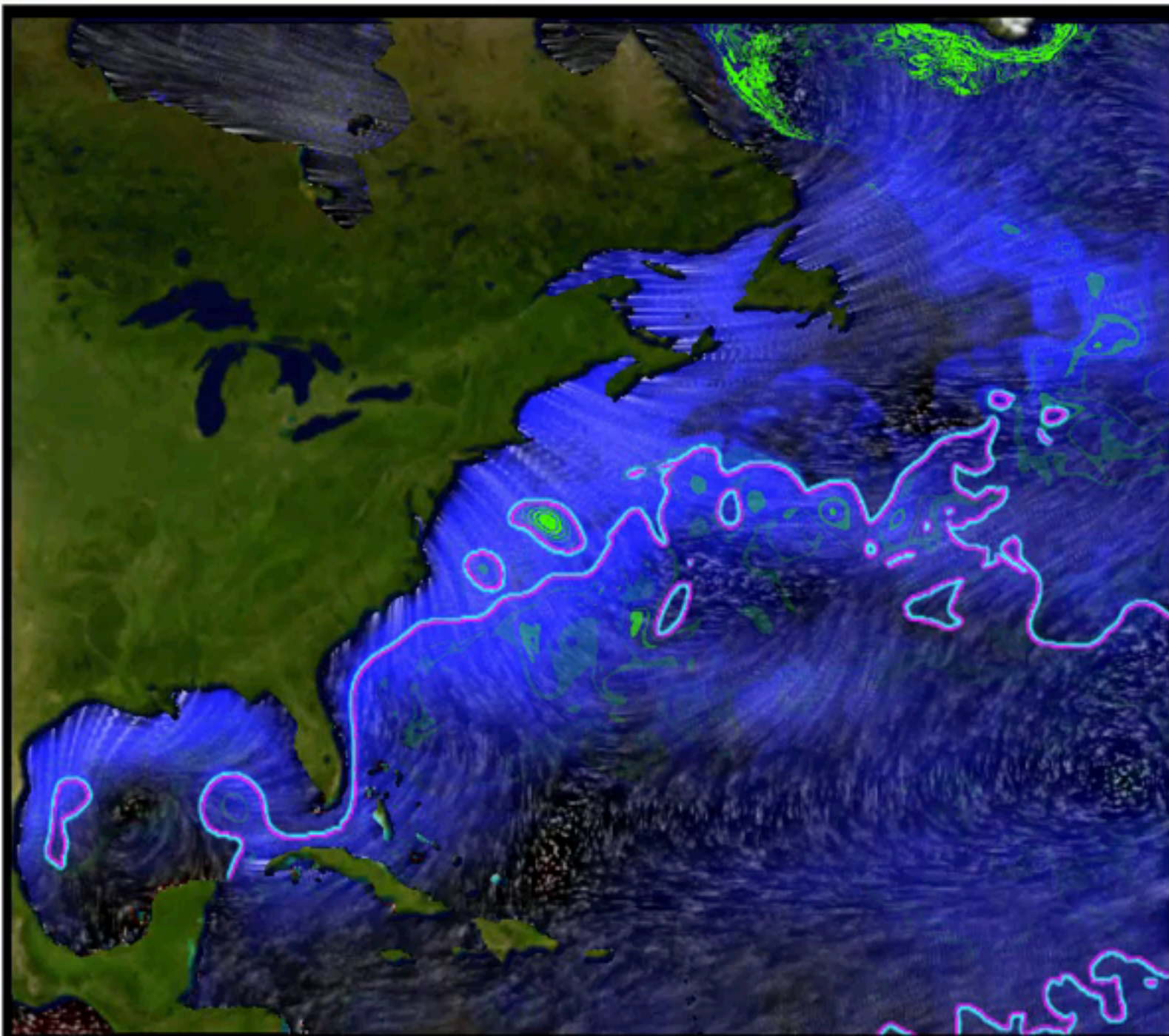
- There is observable (autonomous, satellite & ship) ocean heat content variability.
- The near surface seasonal cycle, regional variations, and individual flux components are $O(100 \text{ W/m}^2)$
- Imbalance of Q_{TOA} and net mixed layer entrainment Q_{BML} are $O(1 \text{ W/m}^2)$



- In Situ & SSH agree.



- Understanding of past variability
 - Monday Morning Quarterbacking abounds in variability analyses, e.g.:
 - You can't use 1998 as a start year for climate change—it was the biggest ENSO event of the past 100yr...
 - Phase of the PDO explains the recent warming hiatus, but we don't know what PDO is...
 - May explain and test our understanding, but it has little predictive power.



Weather,
Atmosphere
Fast

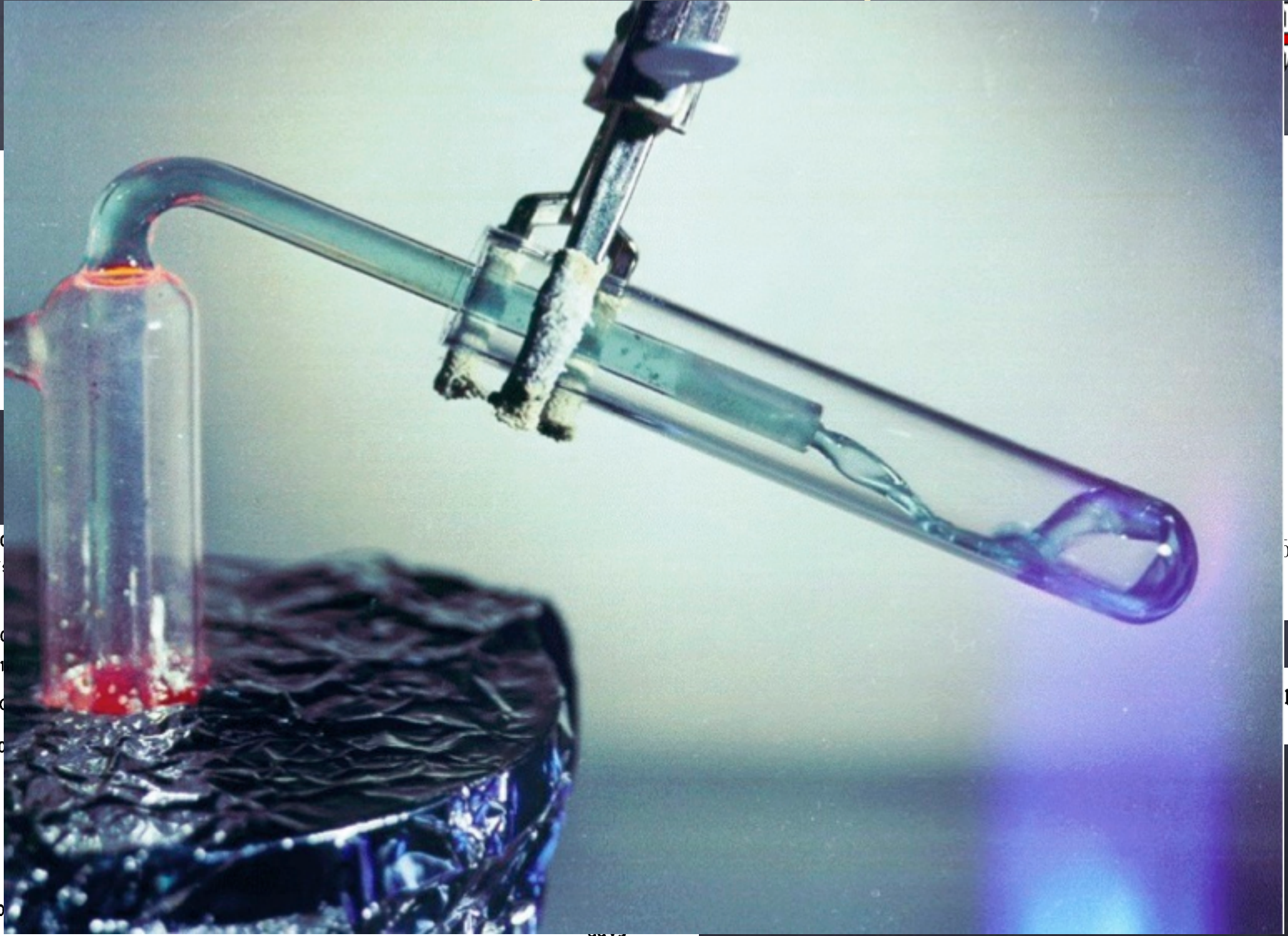
Ocean,
Climate
Slow

3.4m of ocean
water has
same heat
capacity as
the WHOLE
atmosphere

tau / qflux / theta200m / kppMLD

Jan 1 00:30 2001

Modeling of variability



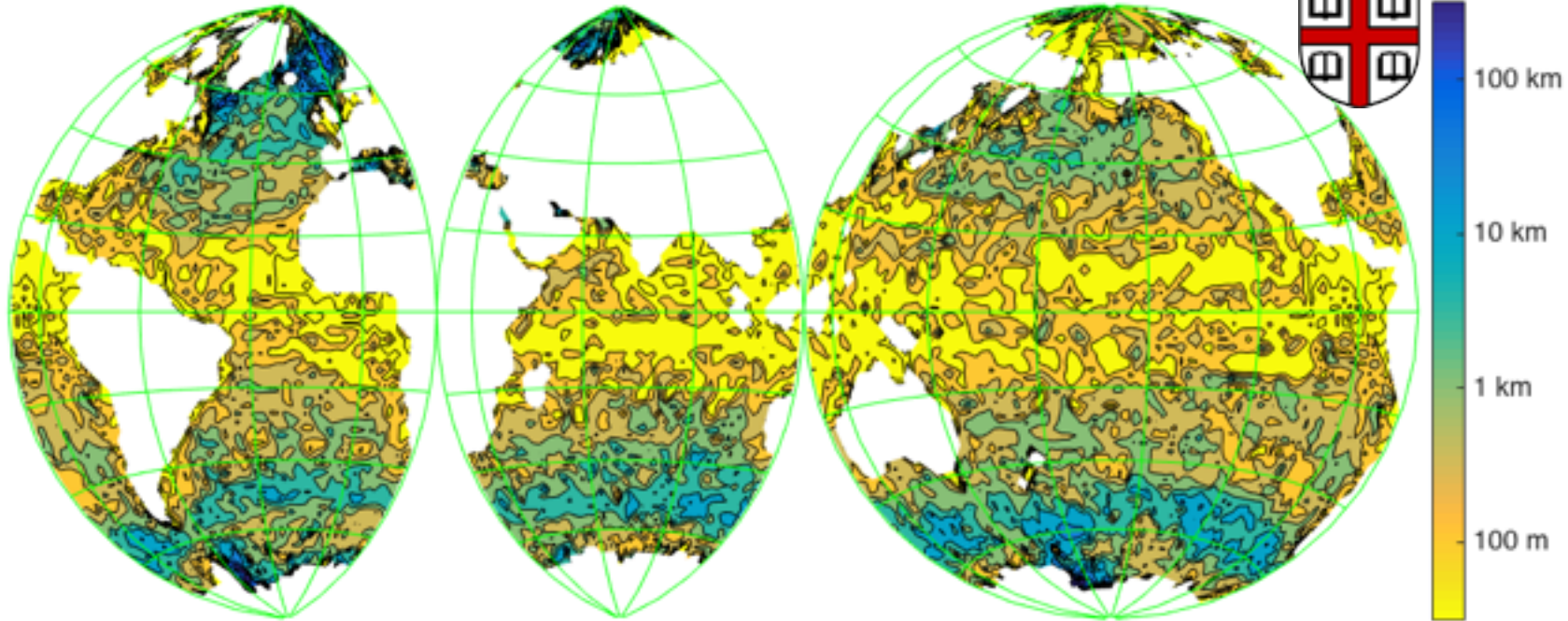
0.0
ly/
0.0
°C
0
-0

5
1

days



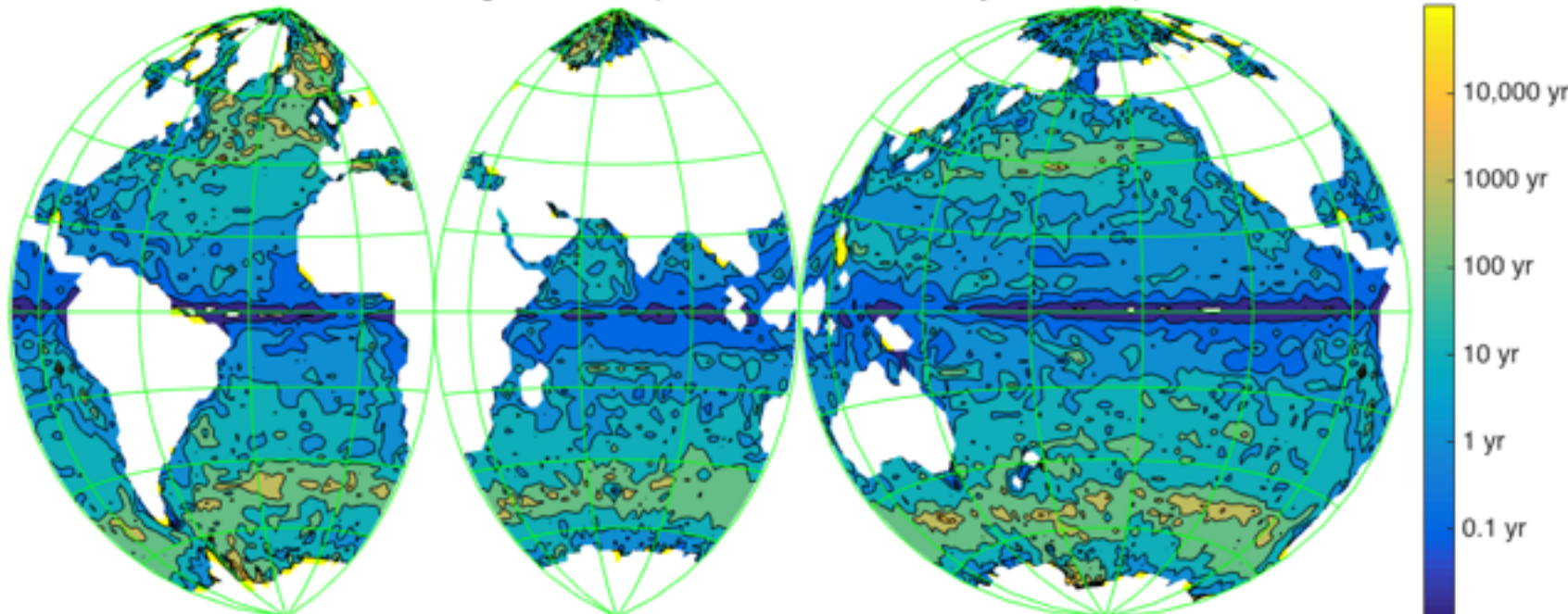
Equivalent Depths of Watermasses by Source (Gebbie & Huybers, 2011)



Consider lots of 1D Oceans: one per watermass

Wind (Ekman) flushing gives upper limit to λ^{-1} timescale

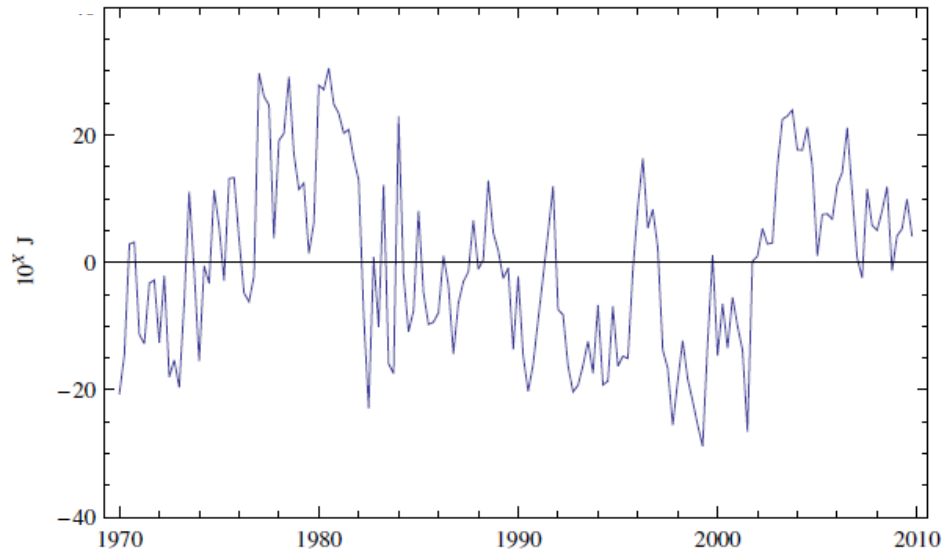
Ekman Flushing Timescale (ECCOv4 + Gebbie & Huybers, 2011)



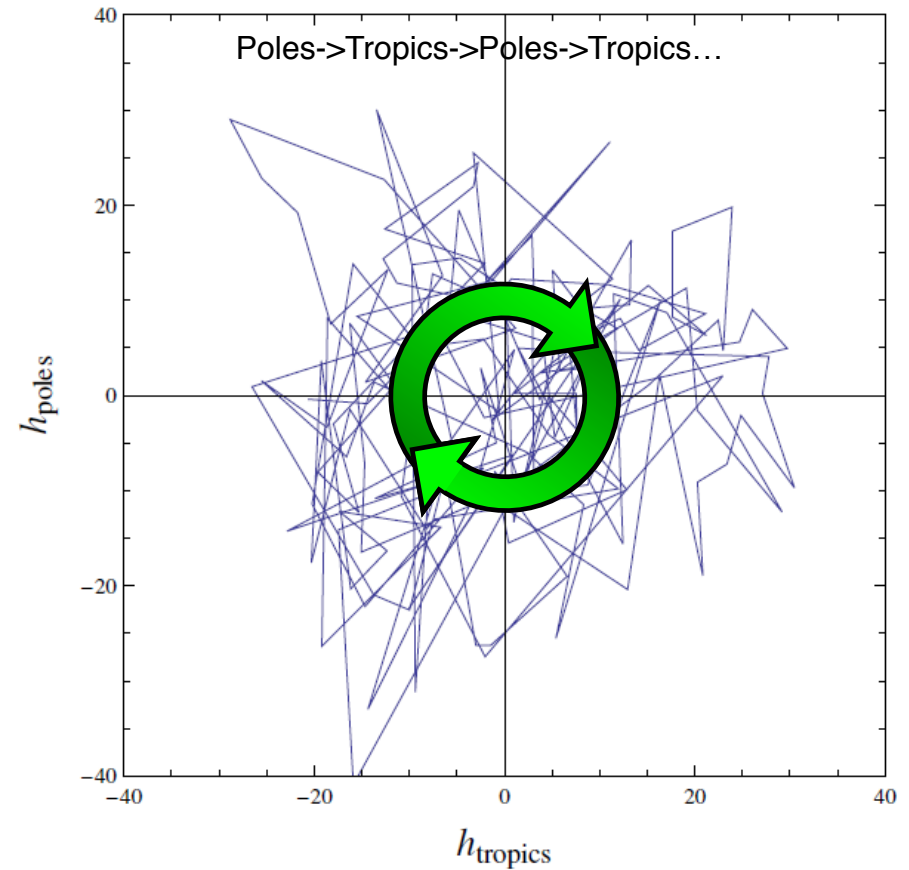
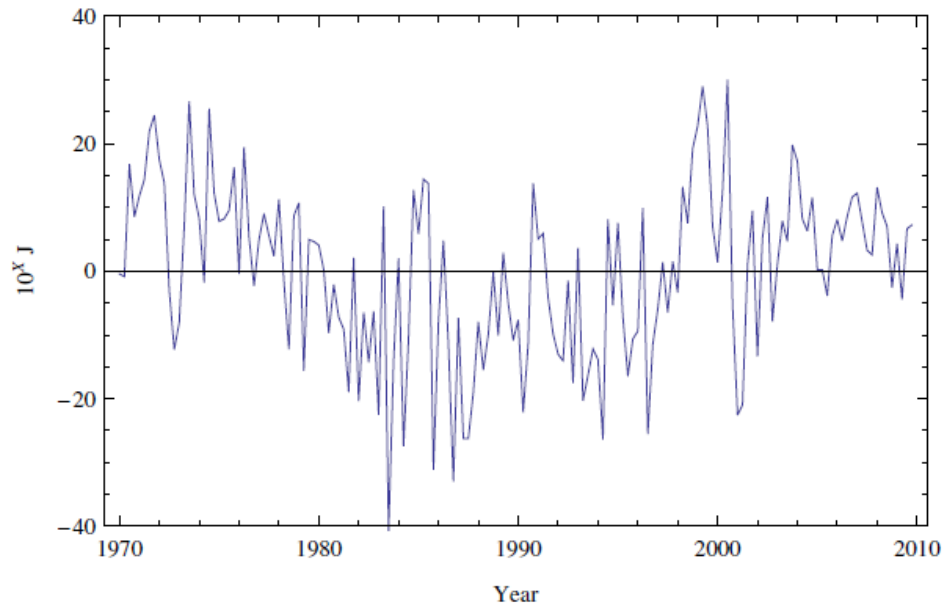
If Connections Occur Between Regions— Predictability Can Arise, Even in Stochastic Systems.



Tropical Ocean Heat Content h_{tropics}



Polar Ocean Heat Content h_{poles}



This is the root cause of
most stochastic model
predictability beyond persistence

Global climate models do pretty well at matching heat fluxes and watermasses.

Statistically significant differences in a few timescales & regions from obs.
(Ticks=10 W/m²)

Models get better every generation due to improved resolution and parameterizations

What does it take to make these improvements?

S. C. Bates, BFK, S. R. Jayne, W. G. Large, S. Stevenson, and S. G. Yeager. Mean biases, variability, and trends in air-sea fluxes and SST in the CCSM4. *Journal of Climate*, 25(22):7781-7801, November 2012.

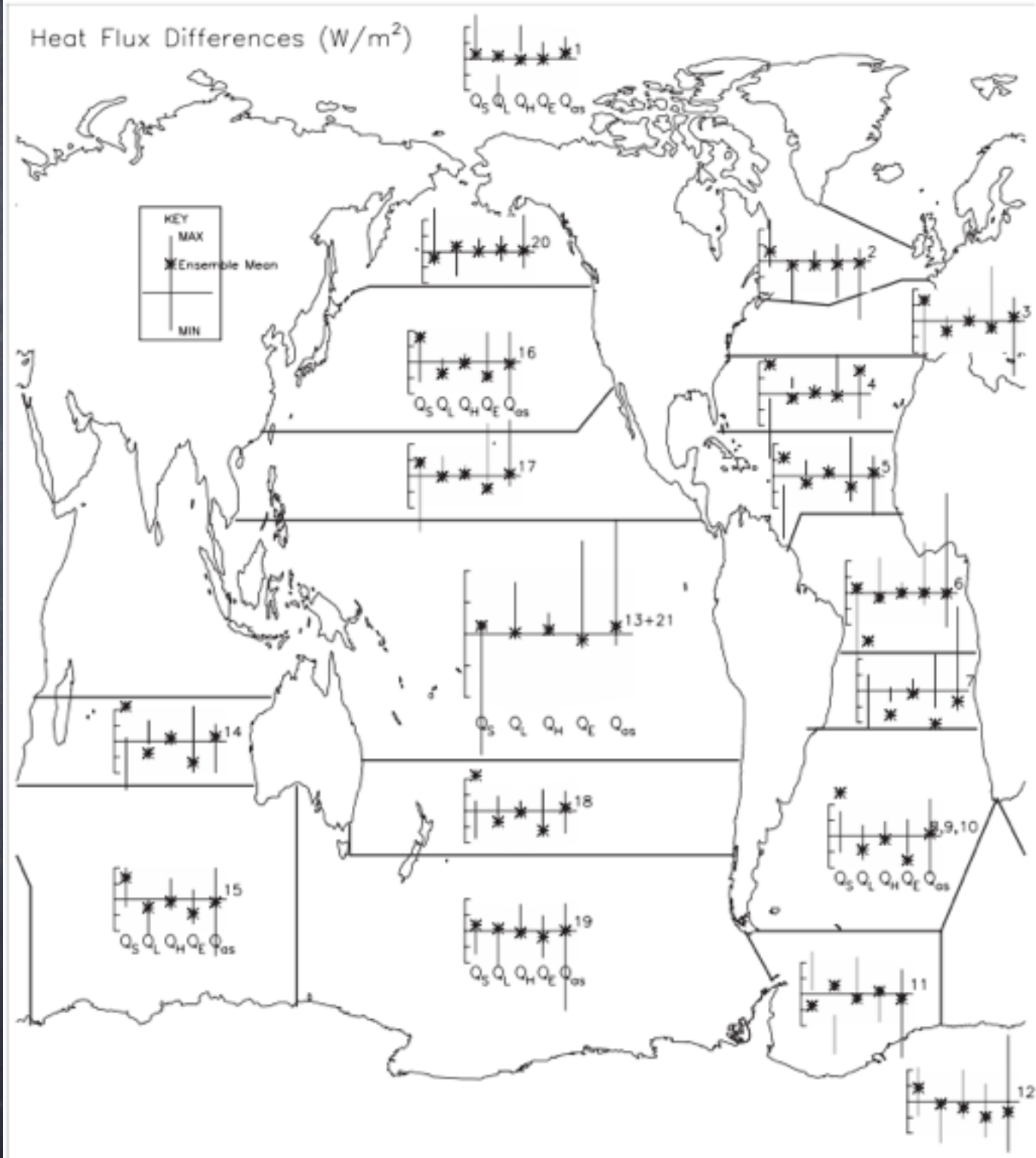
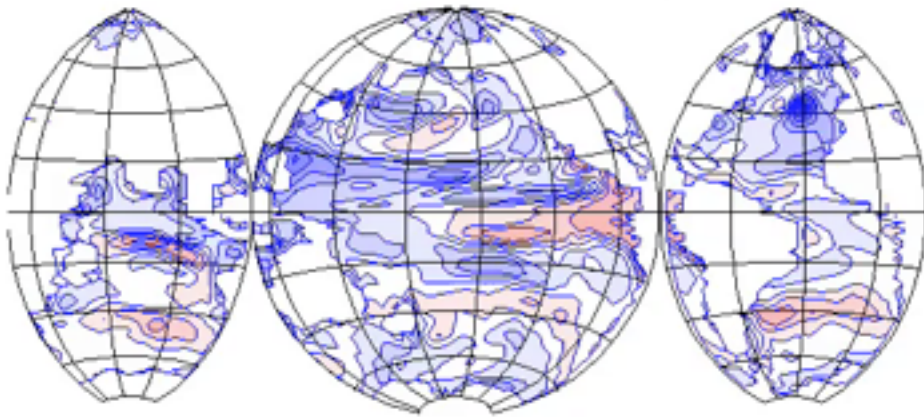


FIG. 4. Regional averages of the CCSM4 20C ensemble mean heat flux components differenced with the CORE

What does a climate model—WITHOUT WARMING—
look like in Ocean Heat Content Variability?

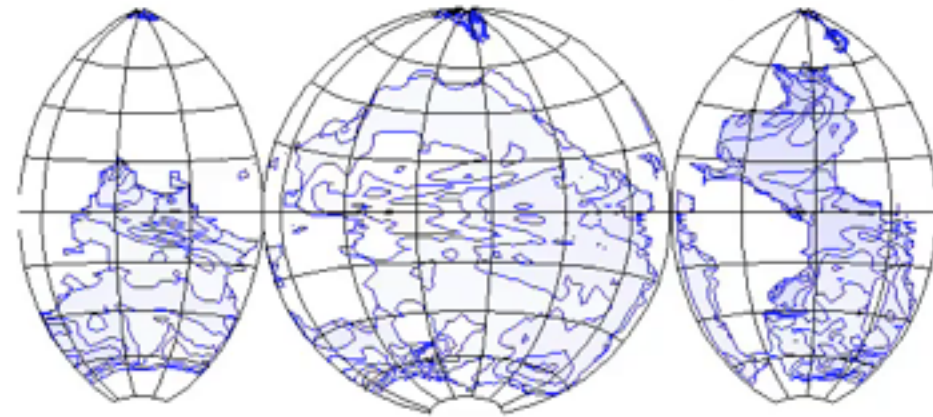
Doesn't even include mesoscale eddies

0-2km Depth Heat Content Anomaly (J) in year 200



Contours = 4 units

Below 2km Depth Heat Content Anomaly (J) in year 200



Contours = 1 unit

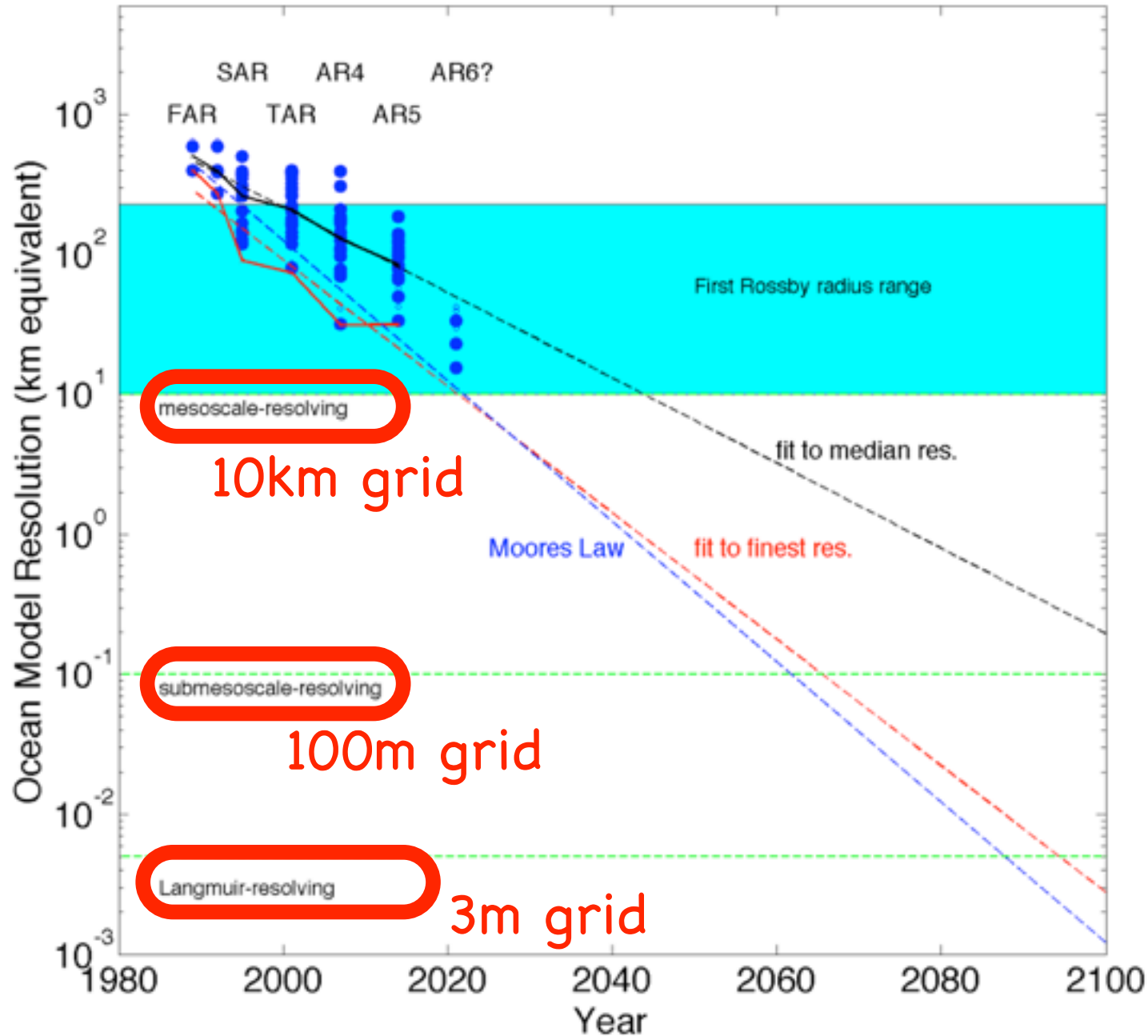
From the >1000yr steady forcing CCSM3.5
runs of Stevenson et al. 2012

S. Stevenson, BFK, and M. Jochum, 2012: Understanding the ENSO-CO2 link using stabilized climate simulations. *Journal of Climate*, 25(22):7917–7936.

Too Simple: What about directly modeling processes in climate models?
Don't we have big enough computers? or won't we soon?



Resolution of Ocean Component of Coupled IPCC models

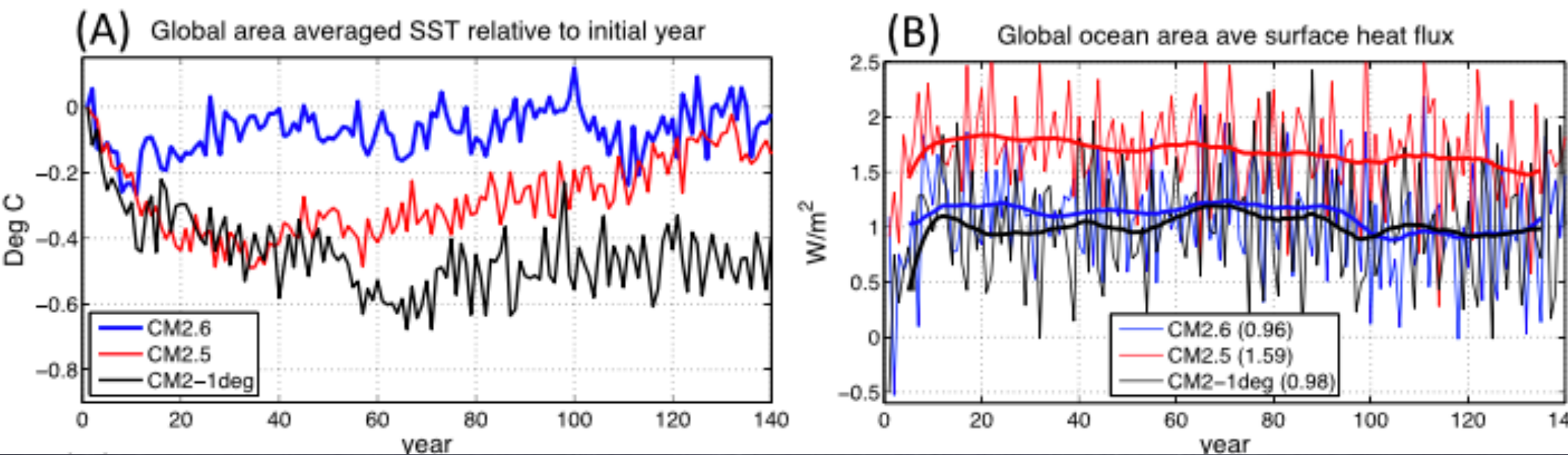


Here are the collection of IPCC models...

If we can't resolve a process, we need to develop a parameterization or subgrid model of its effect

By comparing resolved mesoscale eddies to parameterized ones (with same 50km atmosphere), we get another entry in the pile!

$O(0.7 \text{ W/m}^2)$ persistent and $O(0.4 \text{ K/century})$, i.e., significant warming to upper 1500m of ocean.



Stephen M. Griffies, Michael Winton, Whit G. Anderson, Rusty Benson, Thomas L. Delworth, Carolina O. Dufour, John P. Dunne, Paul Goddard, Adele K. Morrison, Anthony Rosati, Andrew T. Wittenberg, Jianjun Yin, and Rong Zhang, 2015: Impacts on Ocean Heat from Transient Mesoscale Eddies in a Hierarchy of Climate Models. *J. Climate*, 28, 952–977.

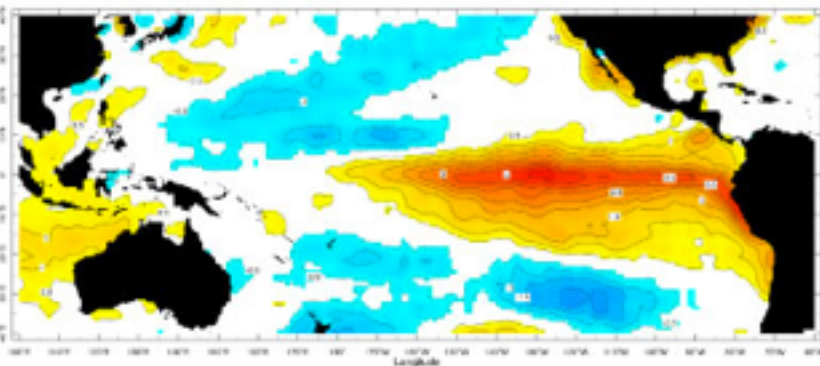
In global models, even with sophisticated parameterizations, numerics, and/or eddy-resolving, the overall heating of the abyss is *tunable* by choice among reasonable parameter values.

Prediction of variability

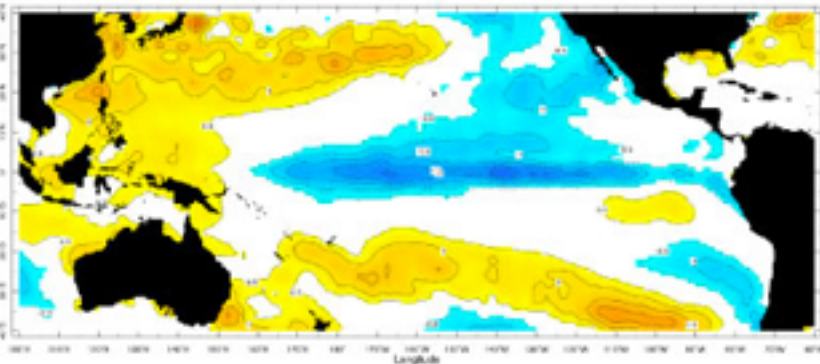
Predictability of ENSO events limited to < 1yr

ENSO statistics more predictable?

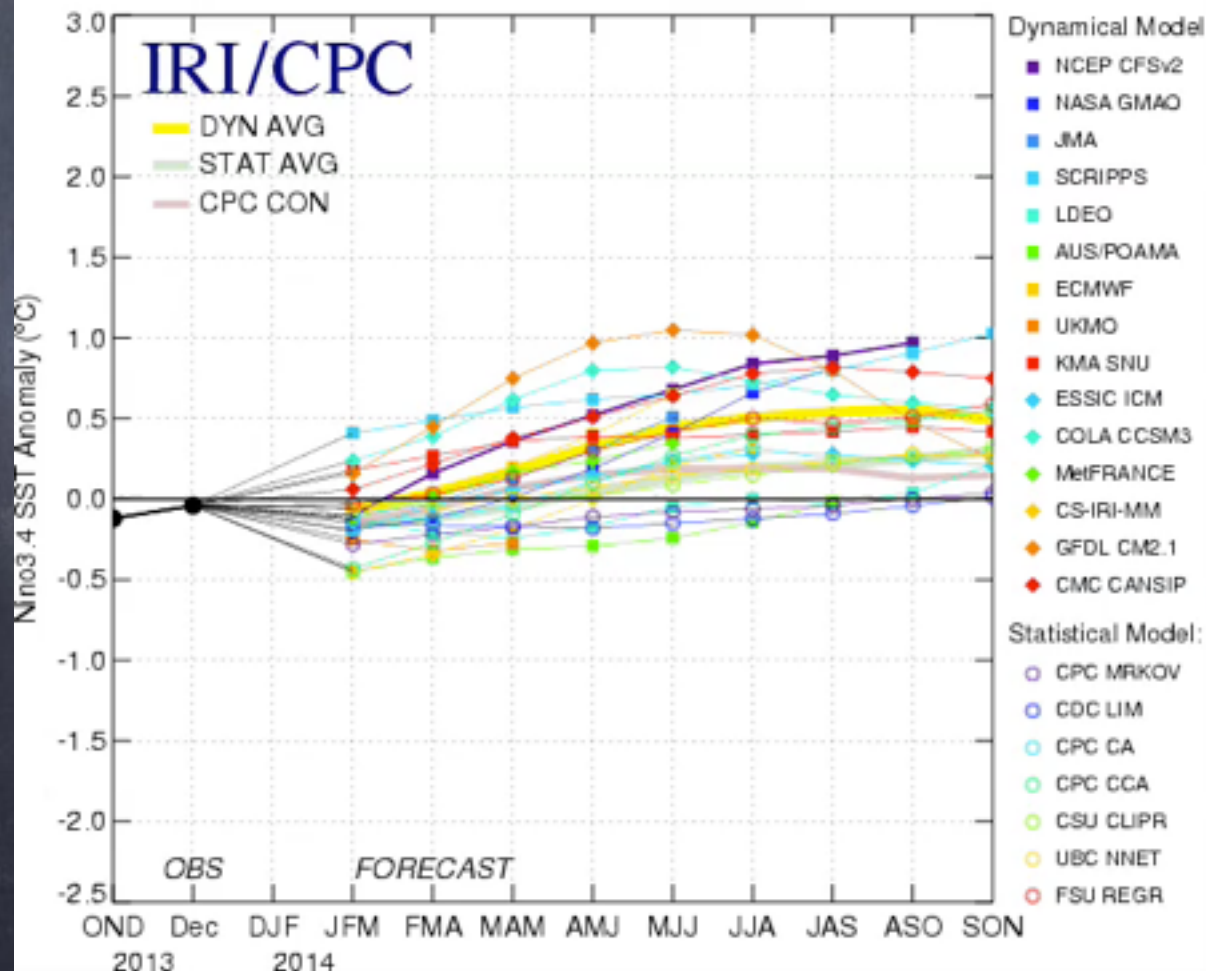
El Niño Episode Sea Surface Temperatures
Departure from average in degrees Celsius
Dec 1982 - Feb 1983



La Niña Episode Sea Surface Temperatures
Departure from average in degrees Celsius
Dec 1998 - Feb 1999



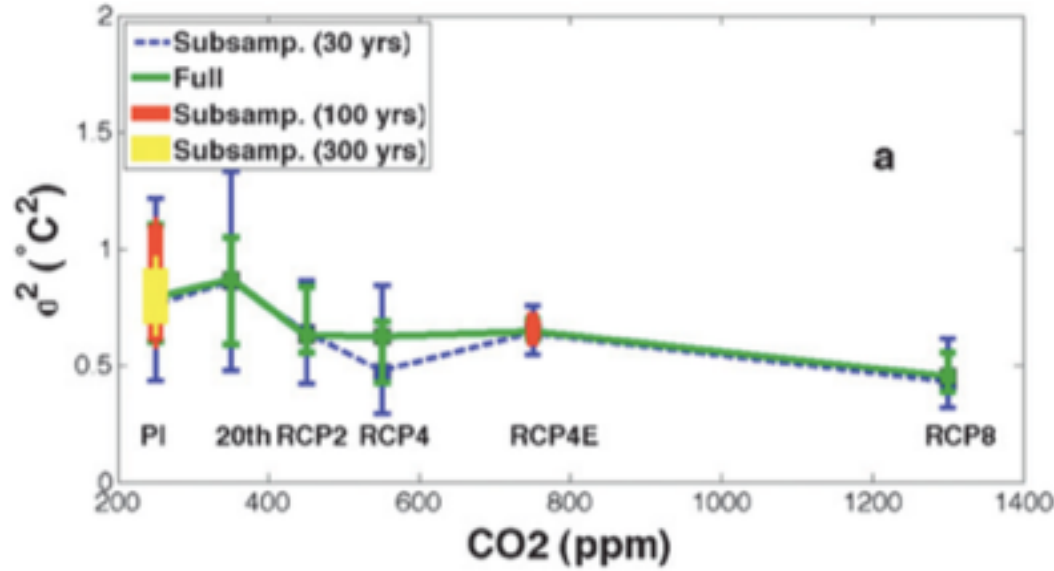
Mid-Jan 2014 Plume of Model ENSO Predictions



Almost no
change to Direct
ENSO variability
with GHG...
(>200 yr to detect)



Stevenson



Big GHG Change
to ENSO impacts!

INDIRECT Proxy
Reconstructions
won't work!!!

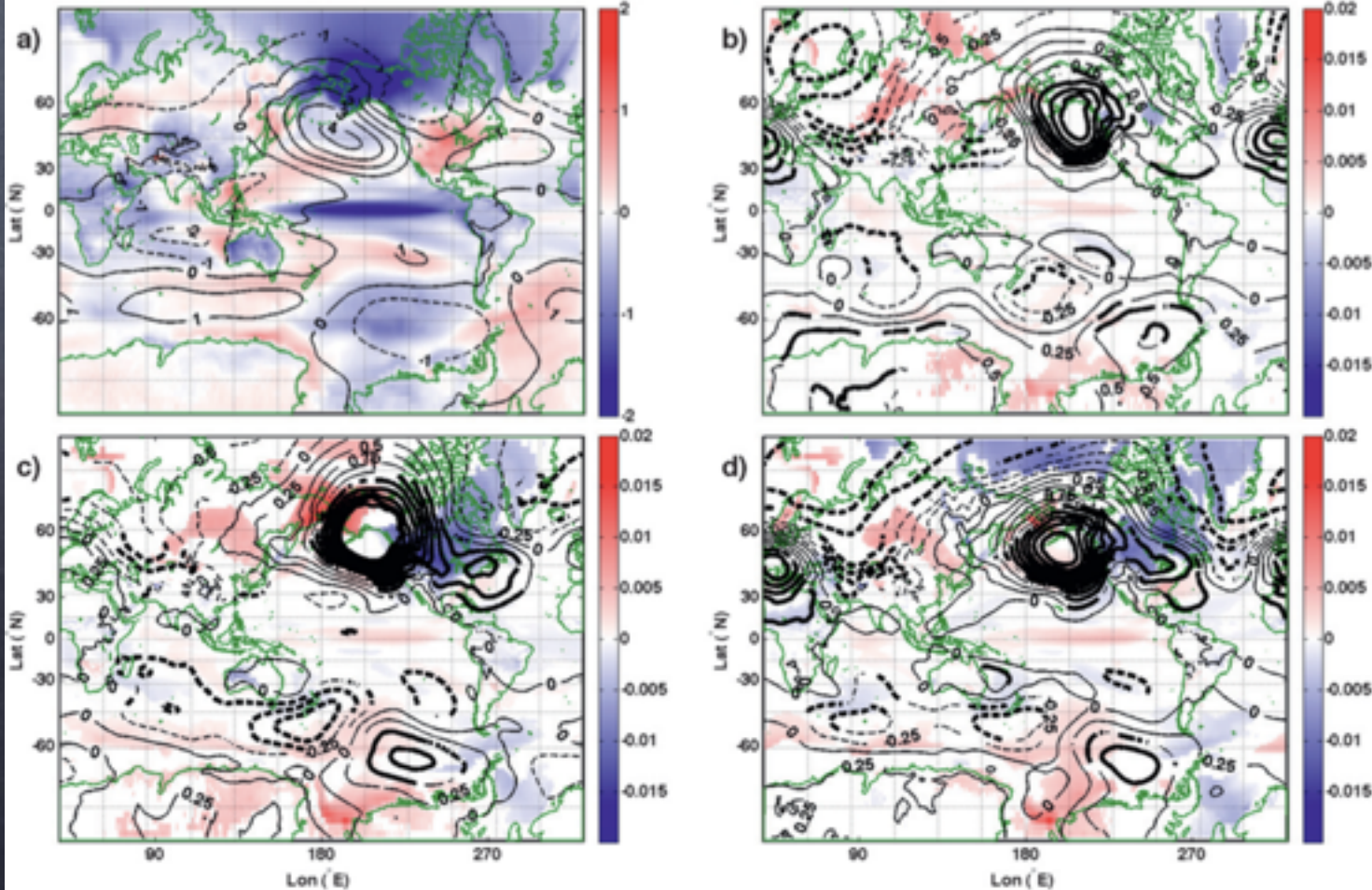
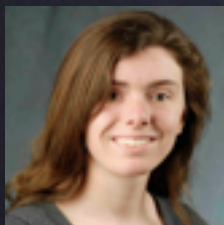
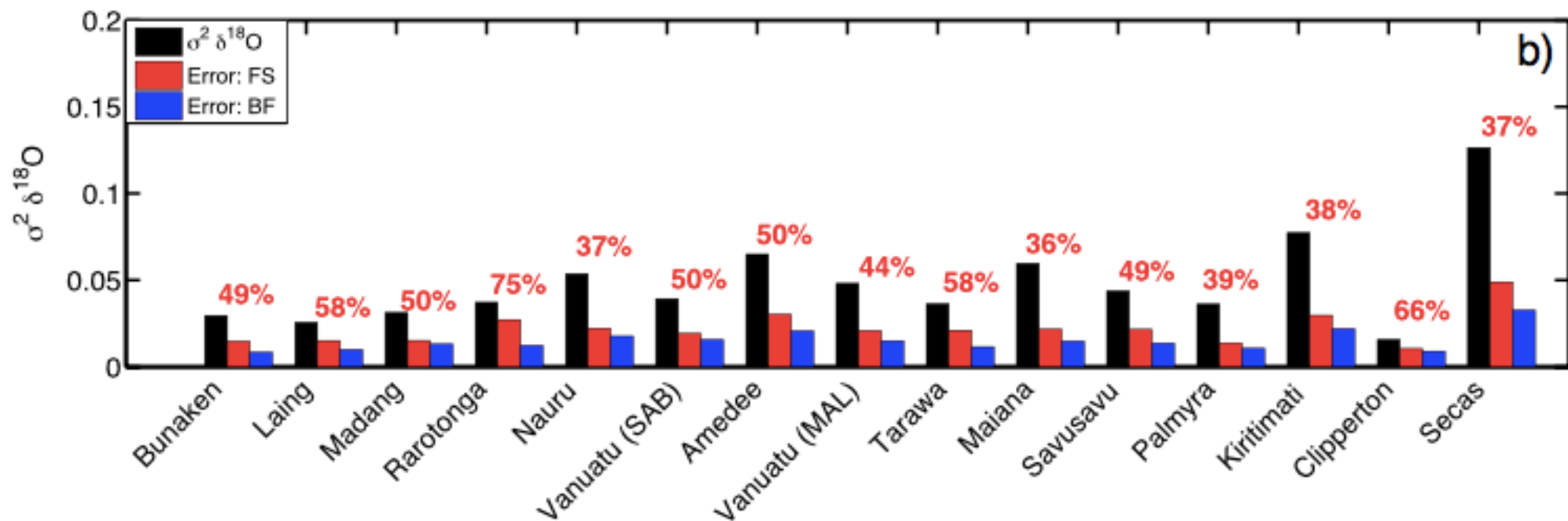
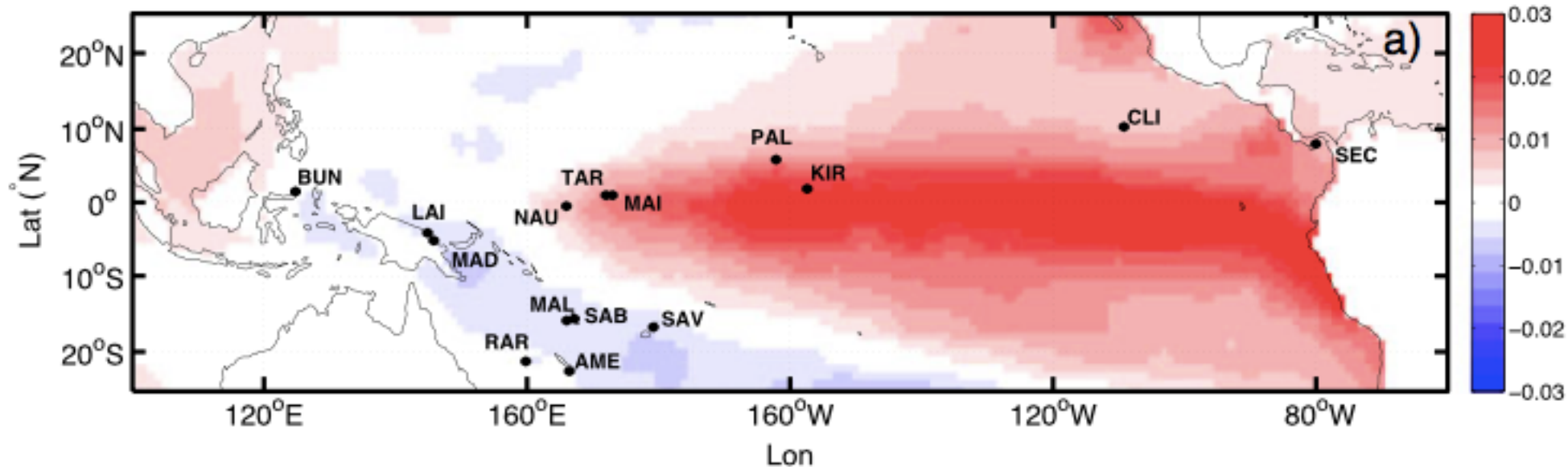


FIG. 6. As in Fig. 5, but for La Niña DIF.

S. Stevenson, BFK, M. Jochum, R. Neale, C. Deser, and G. Meehl.
Will there be a significant change to El Niño in the 21st century?
Journal of Climate, 25(6): 2129-2145, March 2012.



S. Stevenson, H. V. McGregor, S. J. Phipps, and B. Fox-Kemper. Quantifying errors in coral-based ENSO estimates: Towards improved forward modeling of $\delta^{18}\text{O}$. *Paleoceanography*, 28(4):633-649, December 2013.

Understanding of past variability



New: Abyssal Variability is the HARDEST!

- Stochastic damping very slow!
 - huge heat capacity (biggest watermasses on Earth)!
- Timescales may be very long!
 - Watermasses $O(1500\text{yr})$ old by radiocarbon
- Lengthscales may be very short!
 - (weak stratification implies a Rossby radius of $O(2\text{km})$ for modes trapped in AABW only)
- Water "formed" in very small areas!
 - Small-scale atmospheric & oceanic phenomena will be disproportionately important on air-sea effects
- Difficult to observe, IMPOSSIBLE TO MODEL = FUN!

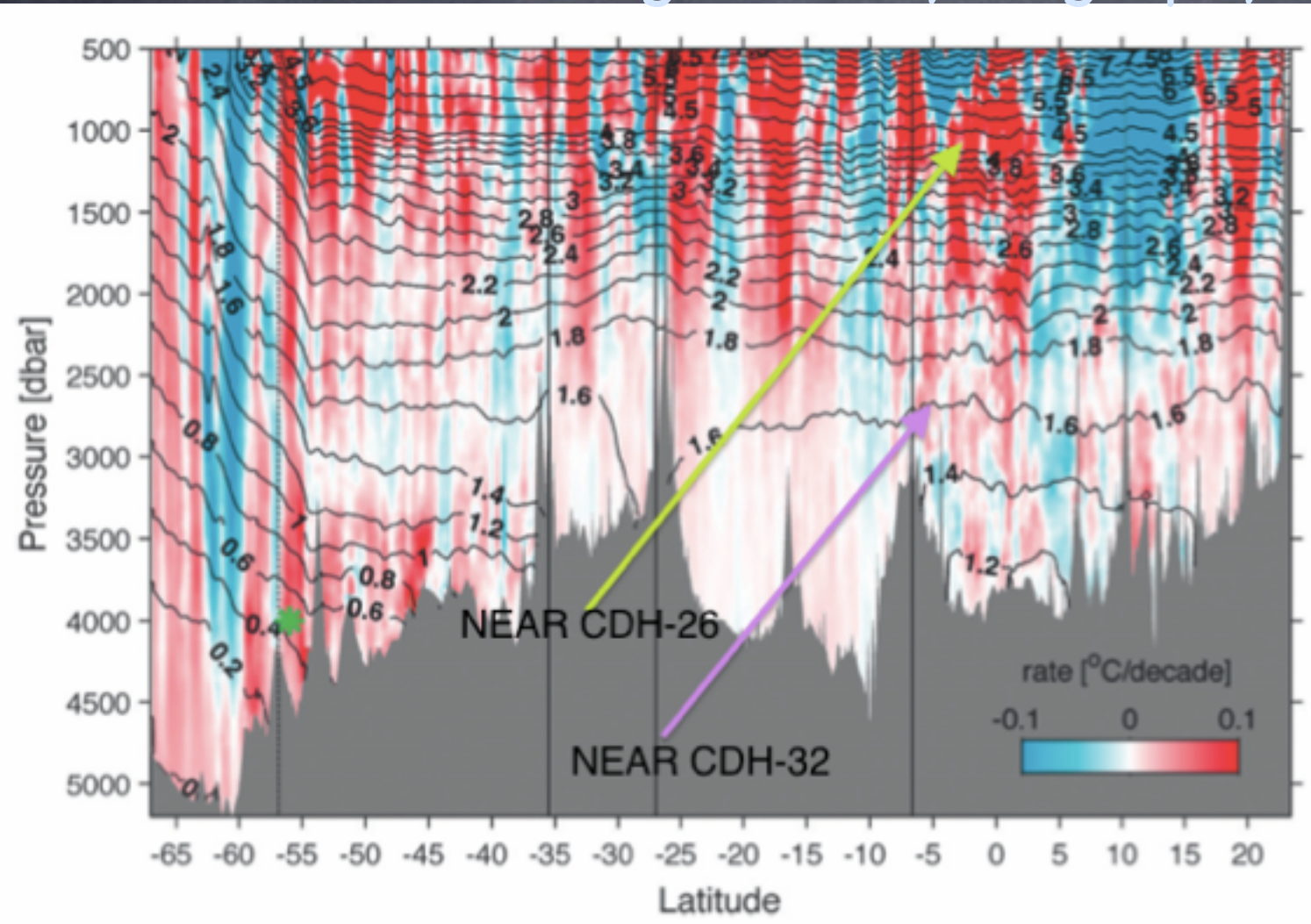
Understanding of past variability



Even with Argo, it will be a while until we have long timescale variability. What to do?

Pattern of Warming from Hydrography

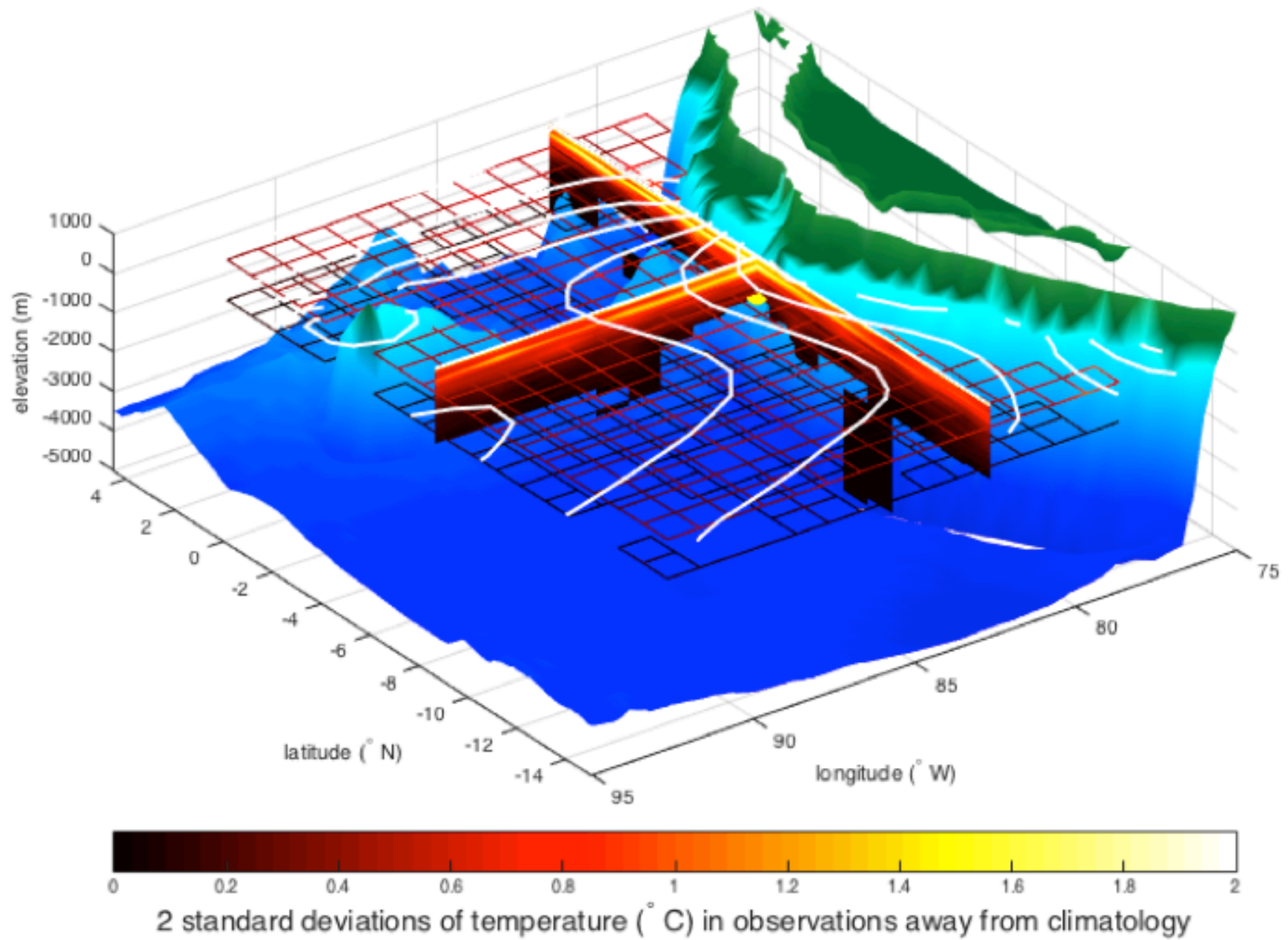
Examine CDH-26 sediment core from the Holocene indicated



Purkey & Johnson, 2010



Brown



S. Bova, T. D. Herbert, and B. Fox-Kemper. Deep ocean variability detected with individual benthic foraminifera. 2016. In preparation.

Understanding of past variability



Assessing variability using individual benthic foraminifera

$$\delta^{18}O = \left(\frac{\left(\frac{^{18}O}{^{16}O} \right)_{sample}}{\left(\frac{^{18}O}{^{16}O} \right)_{standard}} - 1 \right) * 1000 \text{ ‰}$$

- Benthic foraminiferal $\delta^{18}O$ values record temperature and salinity properties of ambient seawater

$$T (\text{°C}) = 21.6 - 5.50 \times (\delta^{18}O_c - \delta^{18}O_{sw})$$

Bemis et al. 2002

$$\delta^{18}O_{sw} = -14.38 + 0.42 * \text{salinity}$$

Conroy et al. 2014

- Individual foraminifera provide 2-3 week snapshots of seawater properties
- We analyze 30-40 individuals within 200 year windows to assess the mean and variance of foraminiferal $\delta^{18}O$ values On roughly decadal timescales

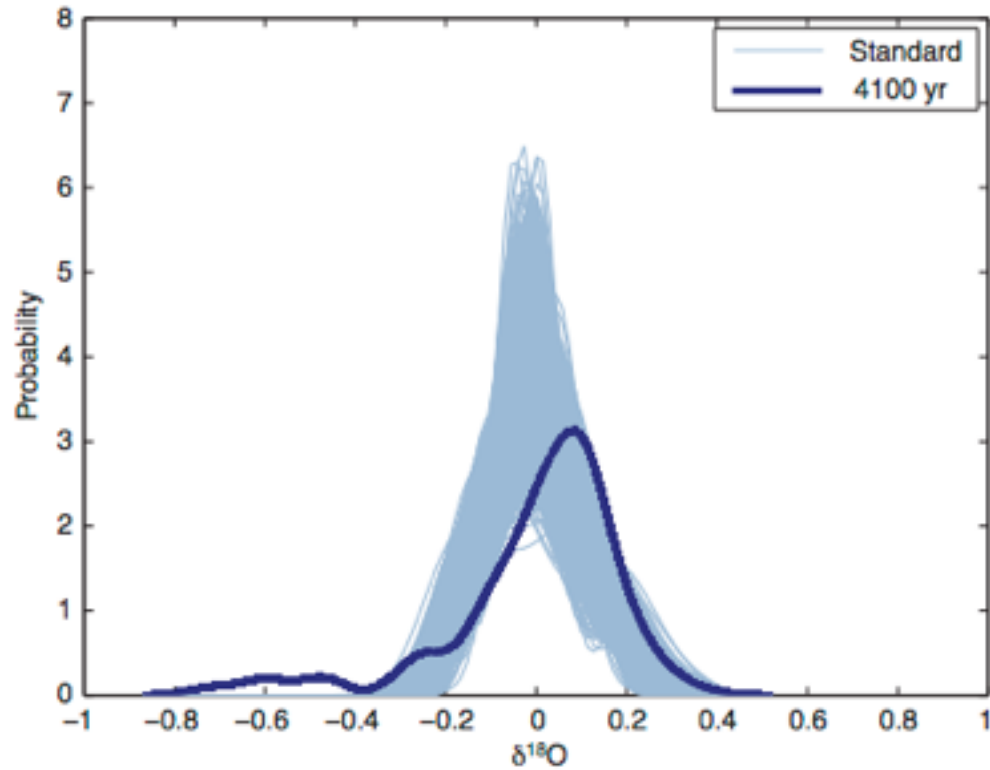


Uvigerina spp.



Brown

$p < 0.01$

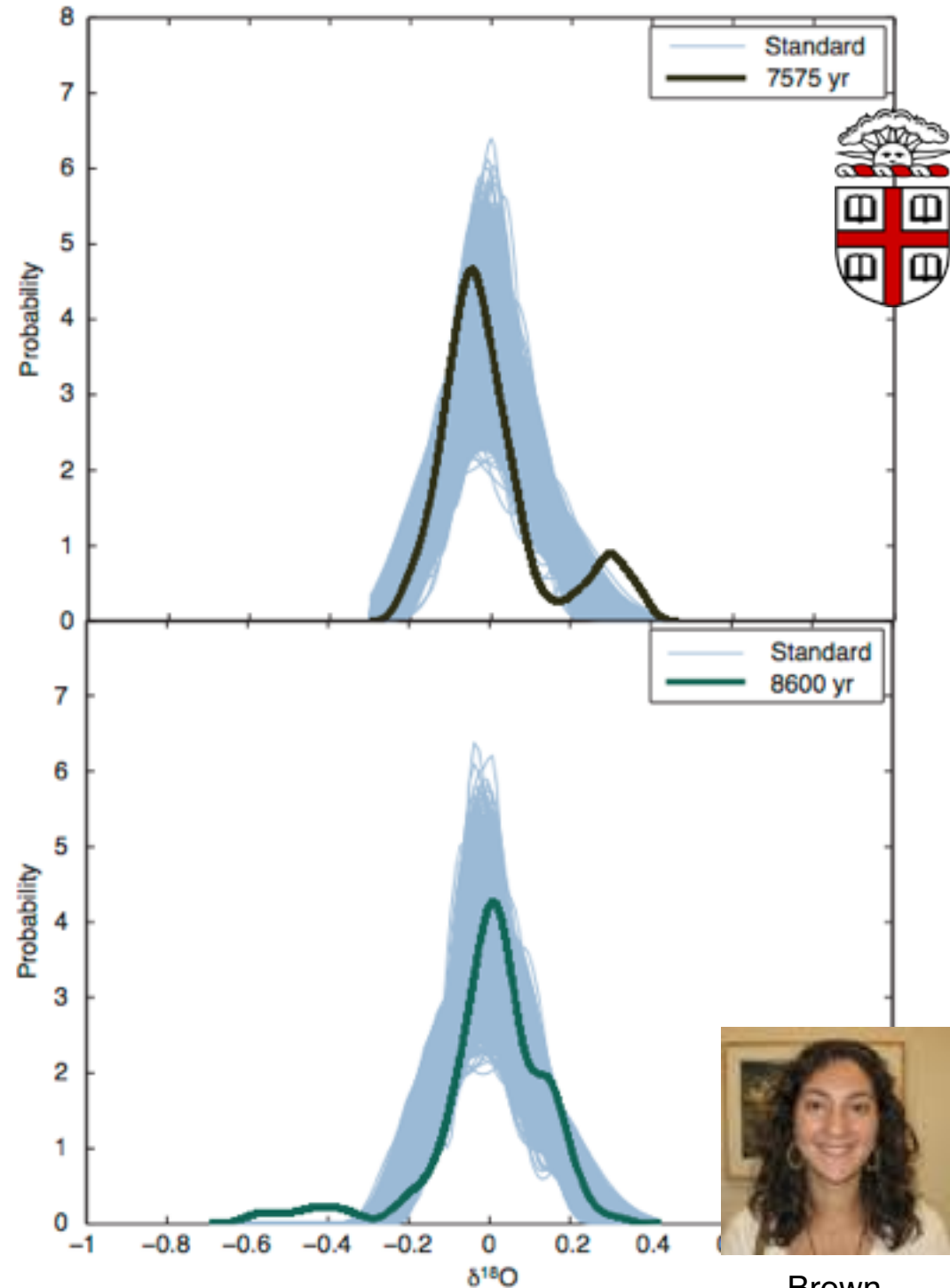


At these three time intervals, the spread of individual values exceeds a size-matched spread of instrumental standards.

The statistical significance of this deviation is given by the p-values of a Kolmogorov-Smirnov test comparing the distributions.

If this is right—abyssal variability may have an **unexpectedly important role, intermittently** through the past!

$p < 0.10$



Some timescales from theory— What is this variability?

- Advective timescale—“water age”, estimated by Gebbie & Huybers from tracers.
- Baroclinic waves
 - <https://www.youtube.com/watch?v=oljinlD2yho>
 - Baroclinic Kelvin & Rossby

Advective Timescale

500–1500 yr
below 2000m

too slow for
global warming

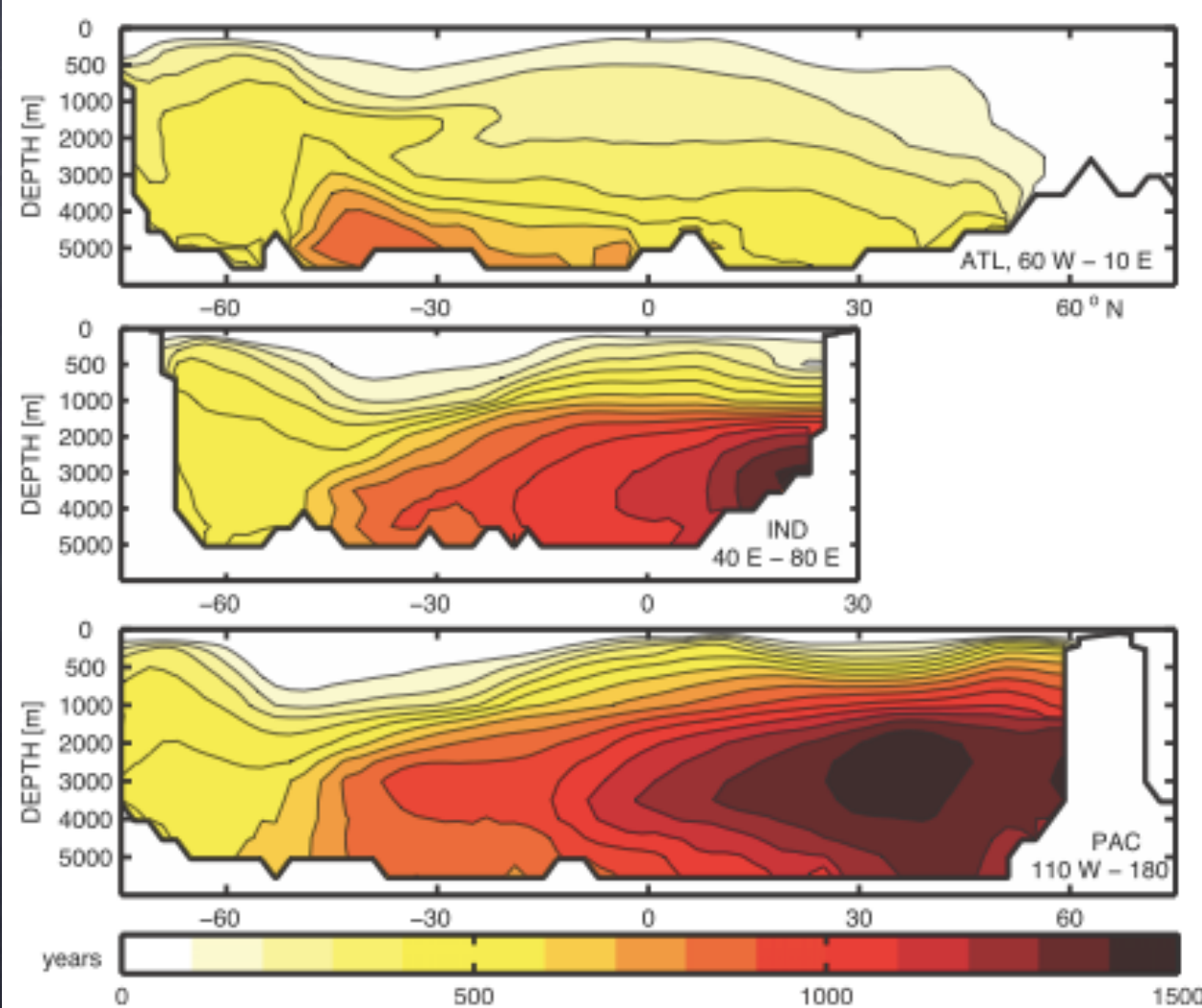


FIG. 8. Latitude–depth sections of mean age: (top) the Atlantic averaged between 60°W and 10°E, (middle) the Indian Ocean between 40° and 80°E, and (bottom) the Pacific between the date line and 110°W. The contour interval is 100 years in all panels.

**The Mean Age of Ocean Waters Inferred from Radiocarbon Observations:
Sensitivity to Surface Sources and Accounting for Mixing Histories**

GEOFFREY GEBBIE

Department of Physical Oceanography, Woods Hole Oceanographic Institution, Woods Hole, Massachusetts

PETER HUYBERS

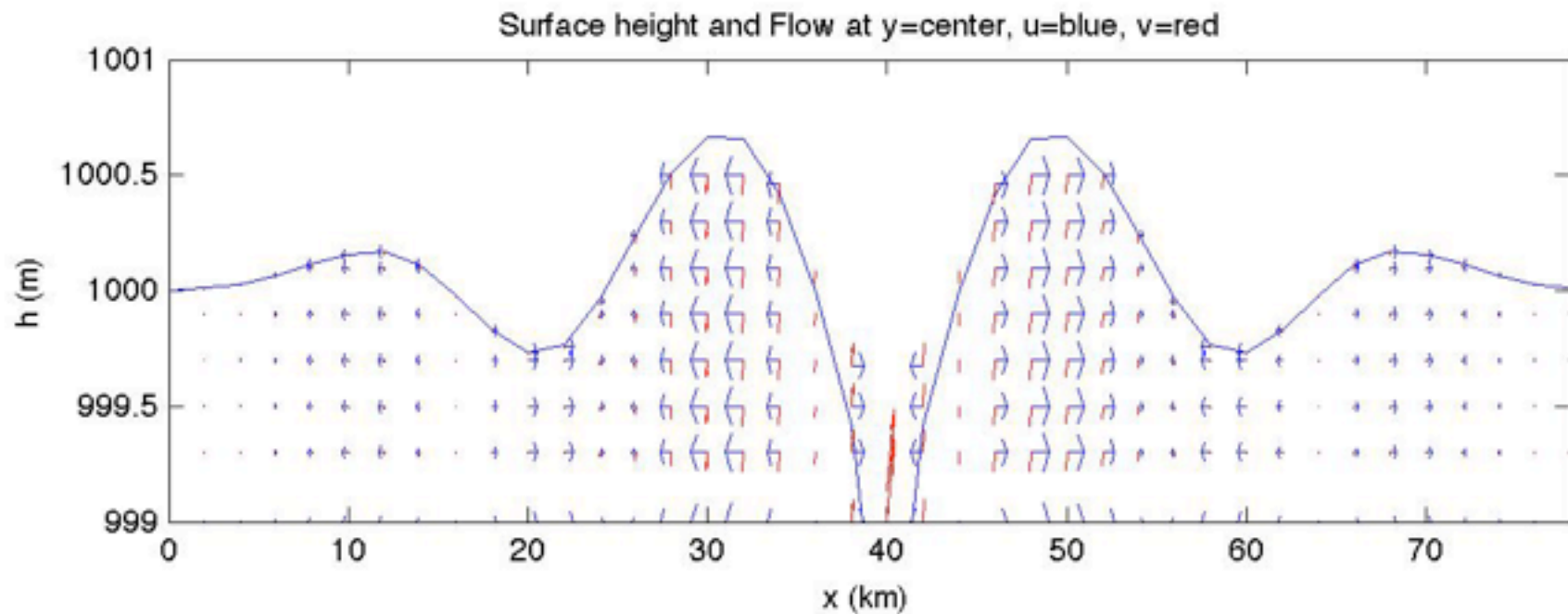
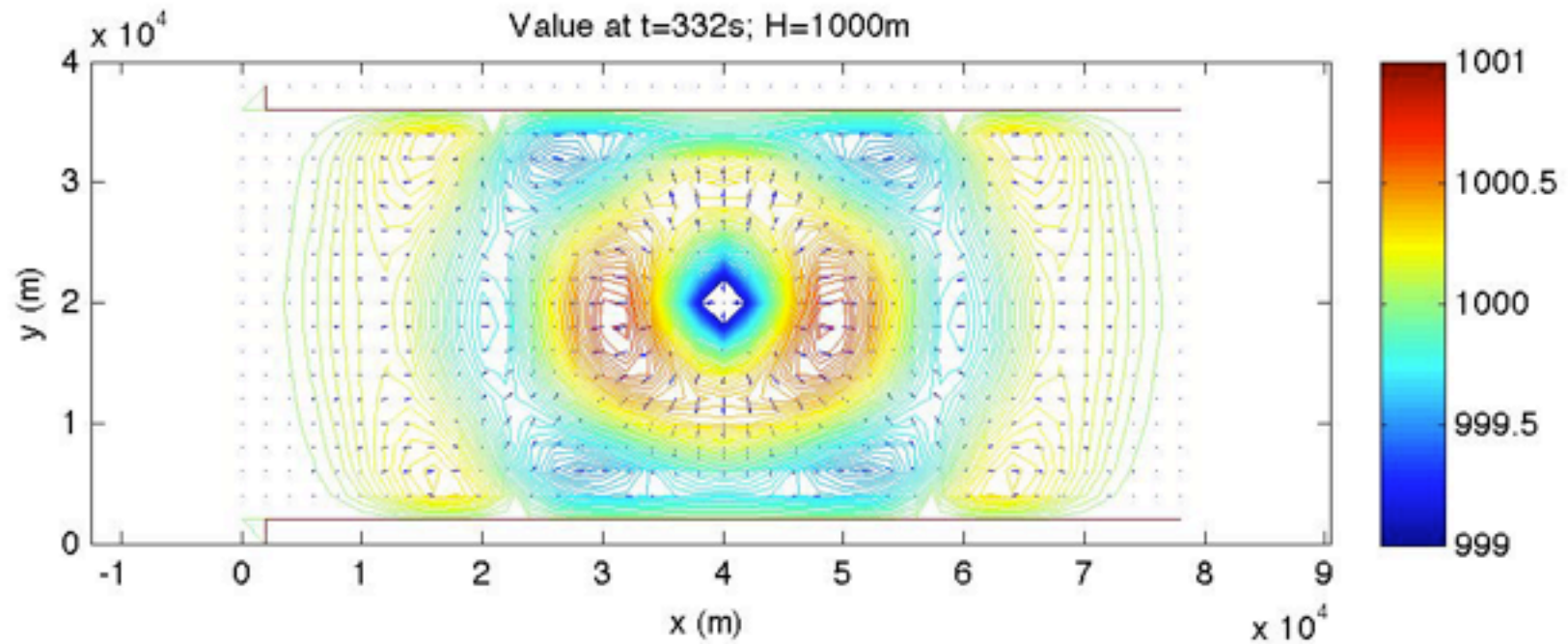
Harvard University, Cambridge, Massachusetts

Simpler: Reduced Gravity

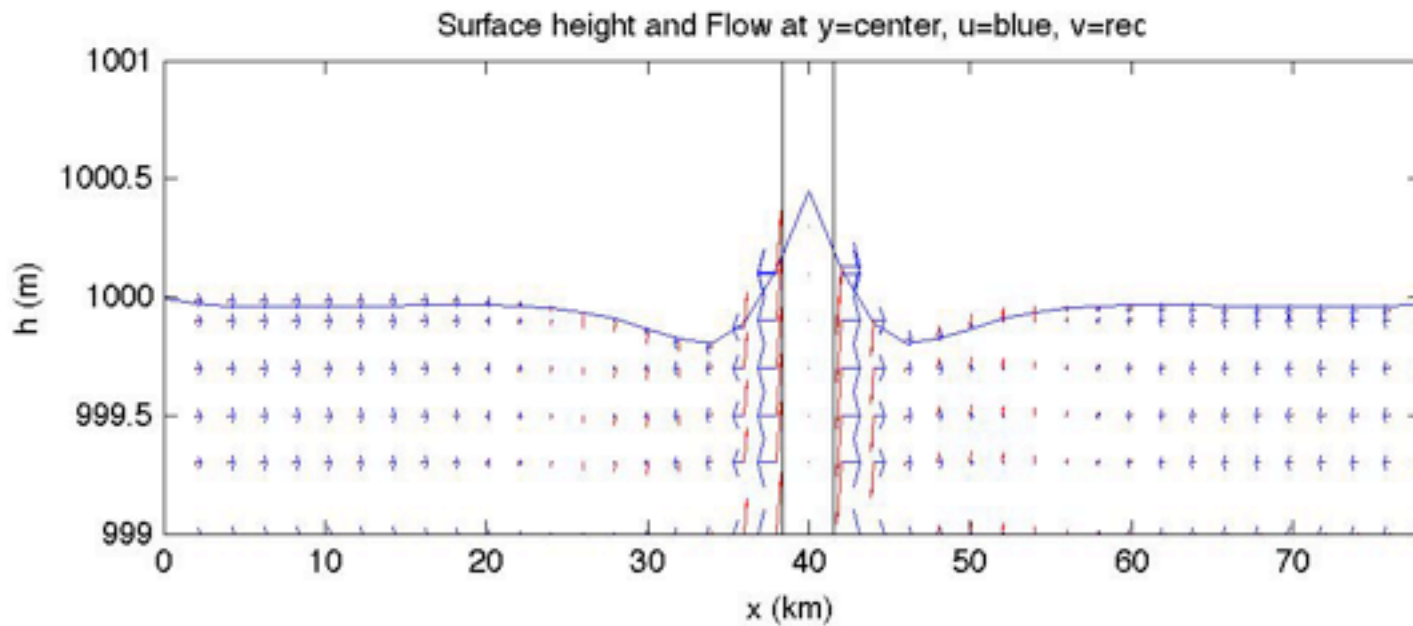
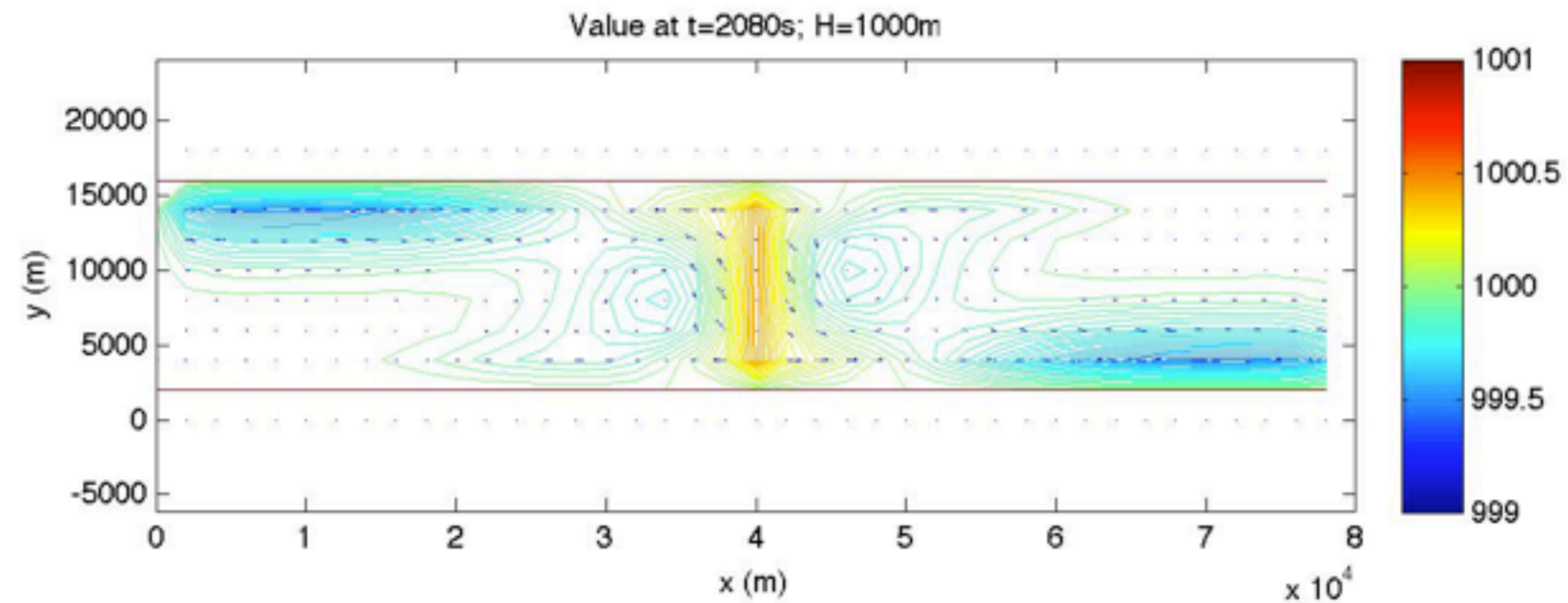
- If motions are coherent only below a given level $-h$, and zero above this level, then bottom layer dynamics are just the shallow water equations with g' as g .
- $g' = g$ (ratio of density difference to total density)
- For AABW vs overlying water—potential density ratio is about 0.02%. Layer is roughly 2km thick.
- Internal gravity wave speed is $c^2 = g'(H-h) = (2 \text{ m/s})^2$
- 4 months to traverse the Pacific N-S.

But, PV not affected by gravity waves—
circulation will not adjust this fast

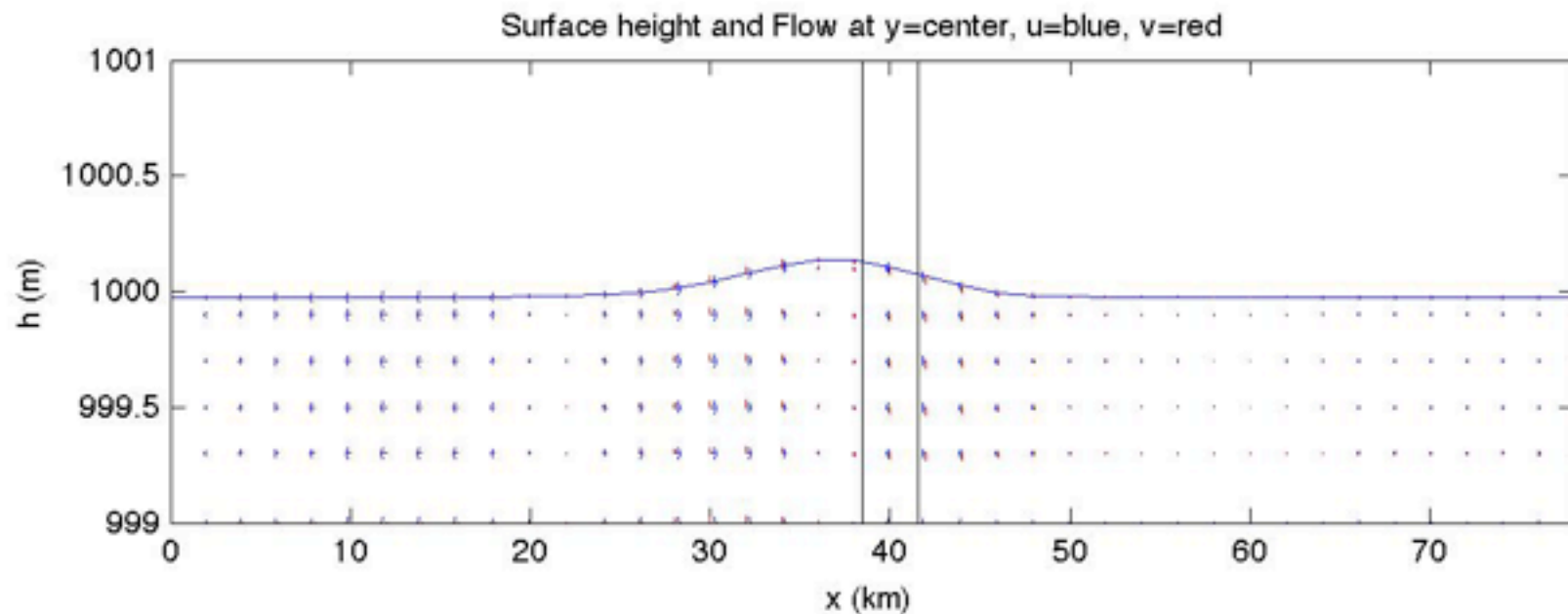
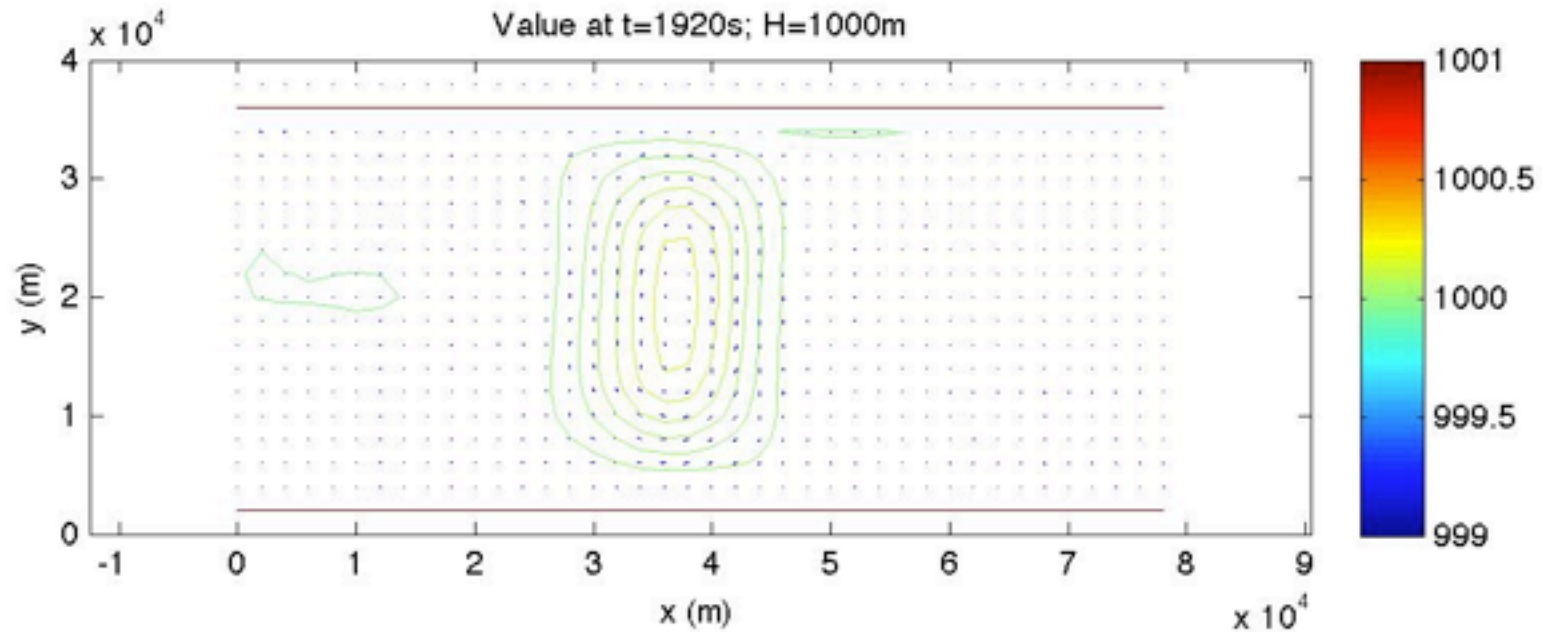
A plunger in a nonrotating channel



A plunger in a rotating channel drives a different wave



A blob of water in a beta-plane rotating channel leaves a rotating high--slow westward propagation



Kelvin & Rossby Waves

- Kelvin have same speed as gravity waves, but trapped against coastlines. Kelvin waves speed along the coastlines (<1 yr), and generate Rossby waves to adjust the basin interior. Based on the reduced gravity estimate, it will take at least 25 years.
- The low stratification of the abyssal water gives only a 20km Rossby radius, so speed is slow and waves are hard to resolve in numerical models.

$$T = \frac{L_B}{\beta L_D^2} \approx \frac{L_B |f|^2}{\beta g' (H - h)} \approx 25 \text{yr}$$

Converging over Bottom Topography & Downward Control

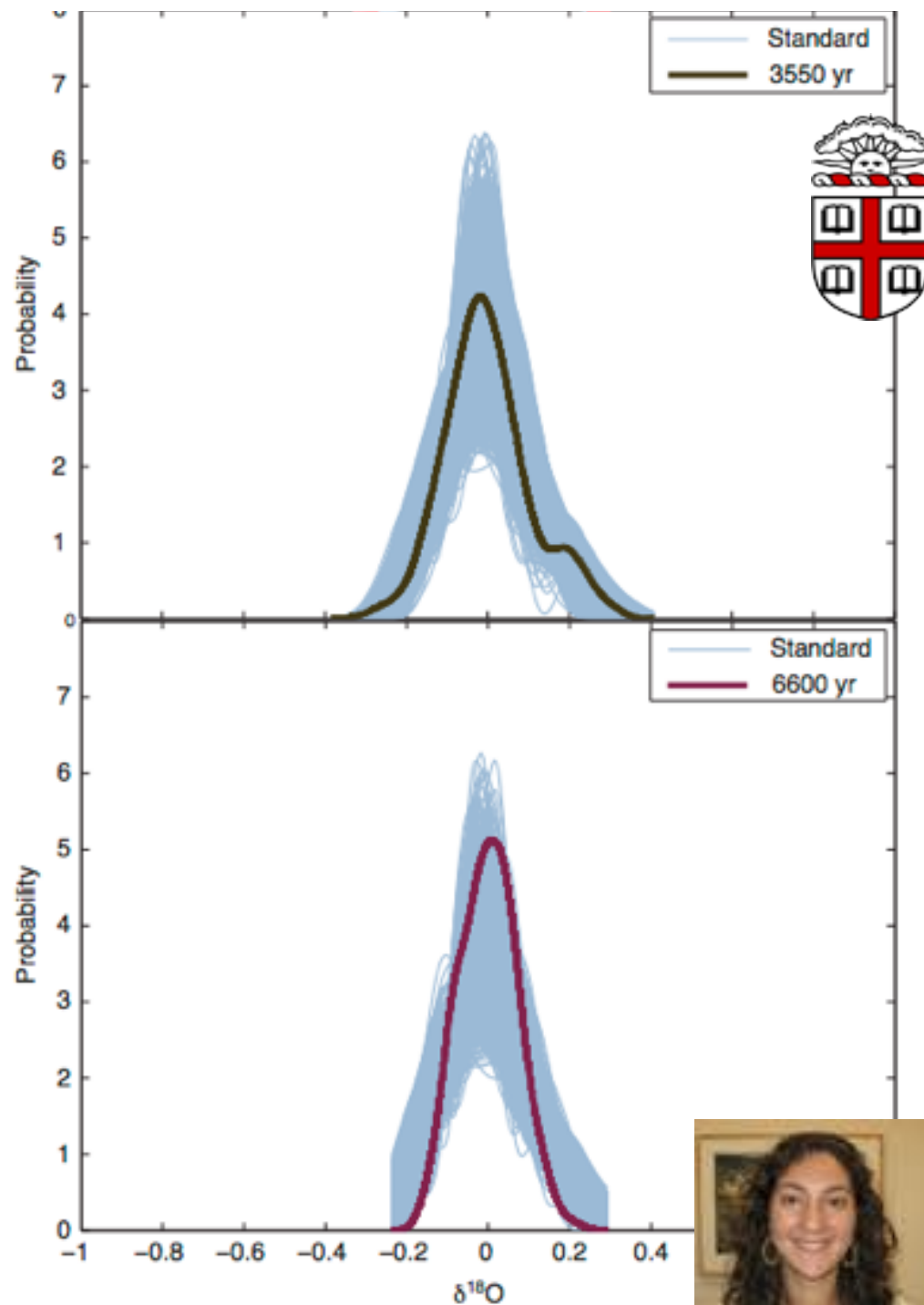
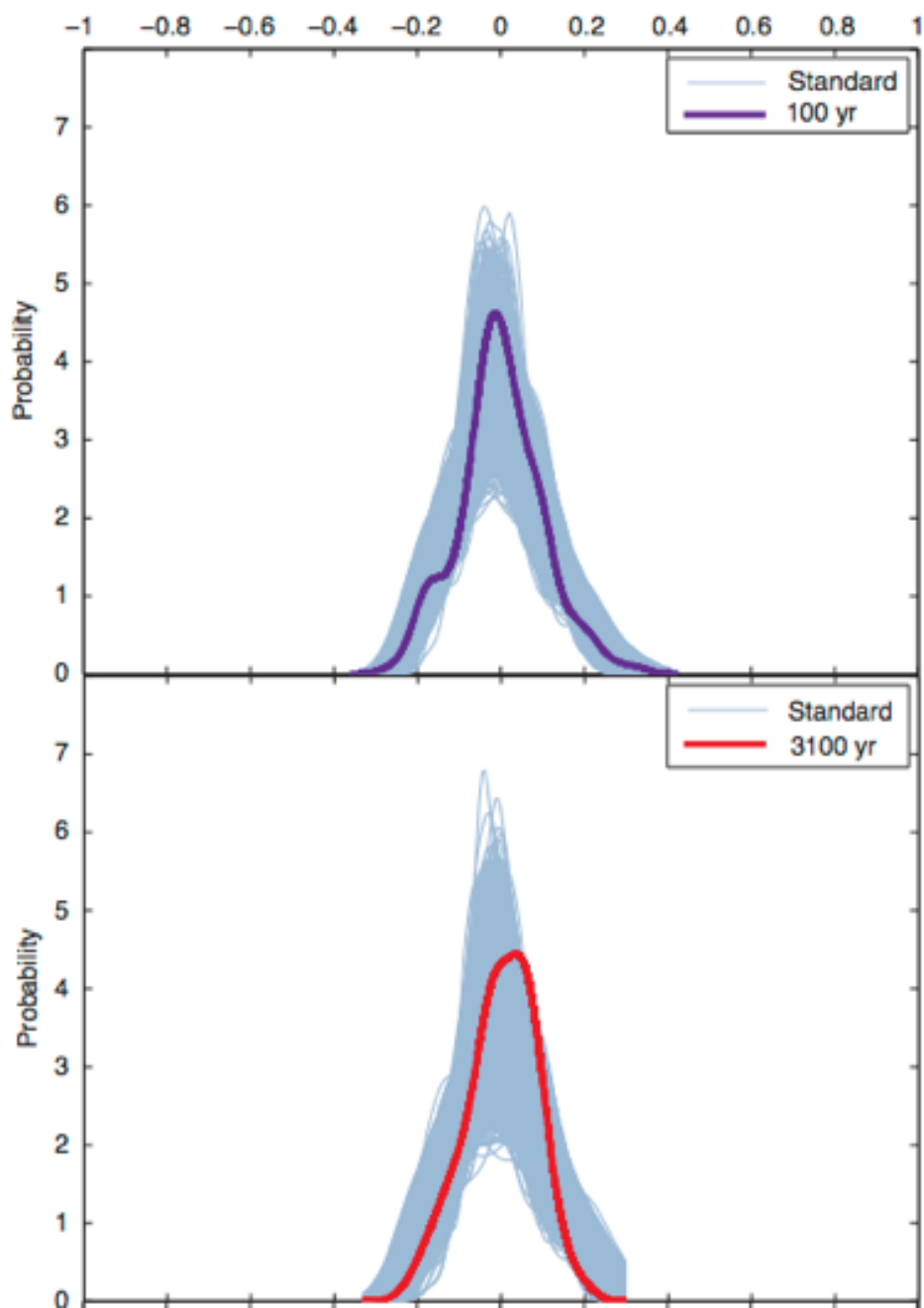
- Estimates were for flat bottom oceans, horizontally constant stratification, and non-equatorial rotation.
- As there is probably much more variability in the near surface, it is possible that including these effects will make the continuation of upper ocean variability create significant near bottom variability
- However, as the movies only depict heat content (and not PV), more to do to see exactly how this works and quantitatively estimate.

Conclusions



- **Presence of observable variability**
 - Regional $O(100 \text{ W/m}^2)$, Global Net $O(1 \text{ W/m}^2)$
 - Difficult due to sampling, obs. density & duration
 - Many problems require paleothermometry!
- **Understanding of past variability**
 - Not always a path to progress w/o models & predictions
 - But, discovery of new processes & unexpected variability is a way forward to better predictions!
- **Modeling of variability**
 - Stochastic models work—but not very predictive.
 - Deterministic models: discrepancies in tuning, params, resolution.
 - Fun to work on parameterizations & process understanding, though!
- **Prediction of variability**
 - Possible in some regions, chaos limits the forecast window.
 - Accurate global budgets need process-level understanding and modeling.

Null Results



At these four time intervals, the spread of individual values fits within a size-matched spread of instrumental standards.

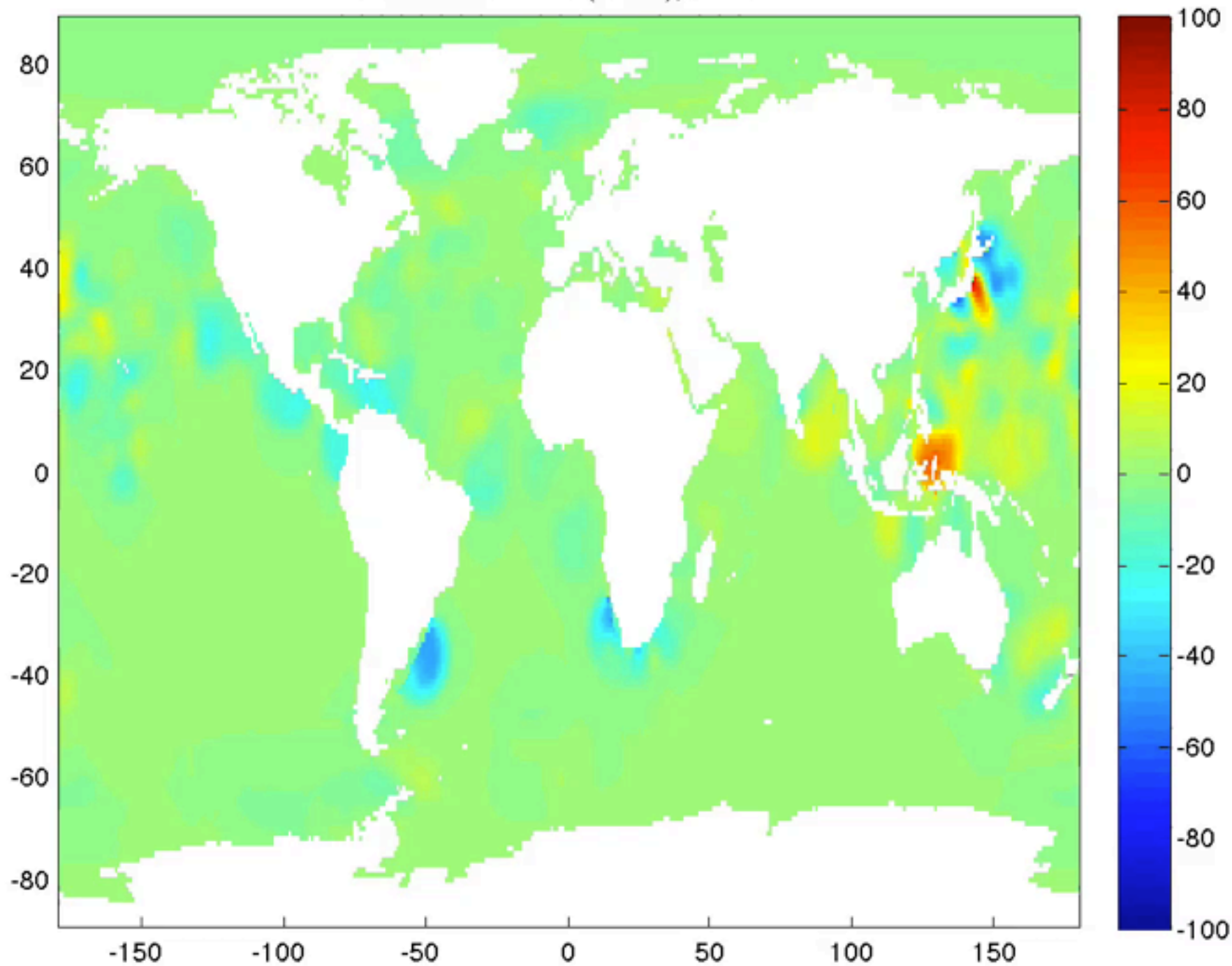
Figure Credit: Sam Bova

Brown



Compare to Observational Product

Heat Content 0-700m (10^{18} J), Feb55



Levitus, S., J. I. Antonov, T. P. Boyer, O. K. Baranova, H. E. Garcia, R. A. Locarnini, A.V. Mishonov, J. R. Reagan, D. Seidov, E. S. Yarosh, M. M. Zweng, 2012: World Ocean heat content and thermosteric sea level change (0-2000 m) 1955-2010. *Geophys. Res. Lett.*, 39, L10603, doi: 10.1029/2012GL051106"