

Eulerian and Lagrangian-  
Novel Deep Connections  
between

Fronts, Waves and Turbulence  
a.k.a. "Esoteric" outcomes

Baylor Fox-Kemper (Brown)  
with CARTHE!!

GoMRI Synthesis, Tallahassee, FL

1/15-17/19

Sponsors: CARTHE/GoMRI

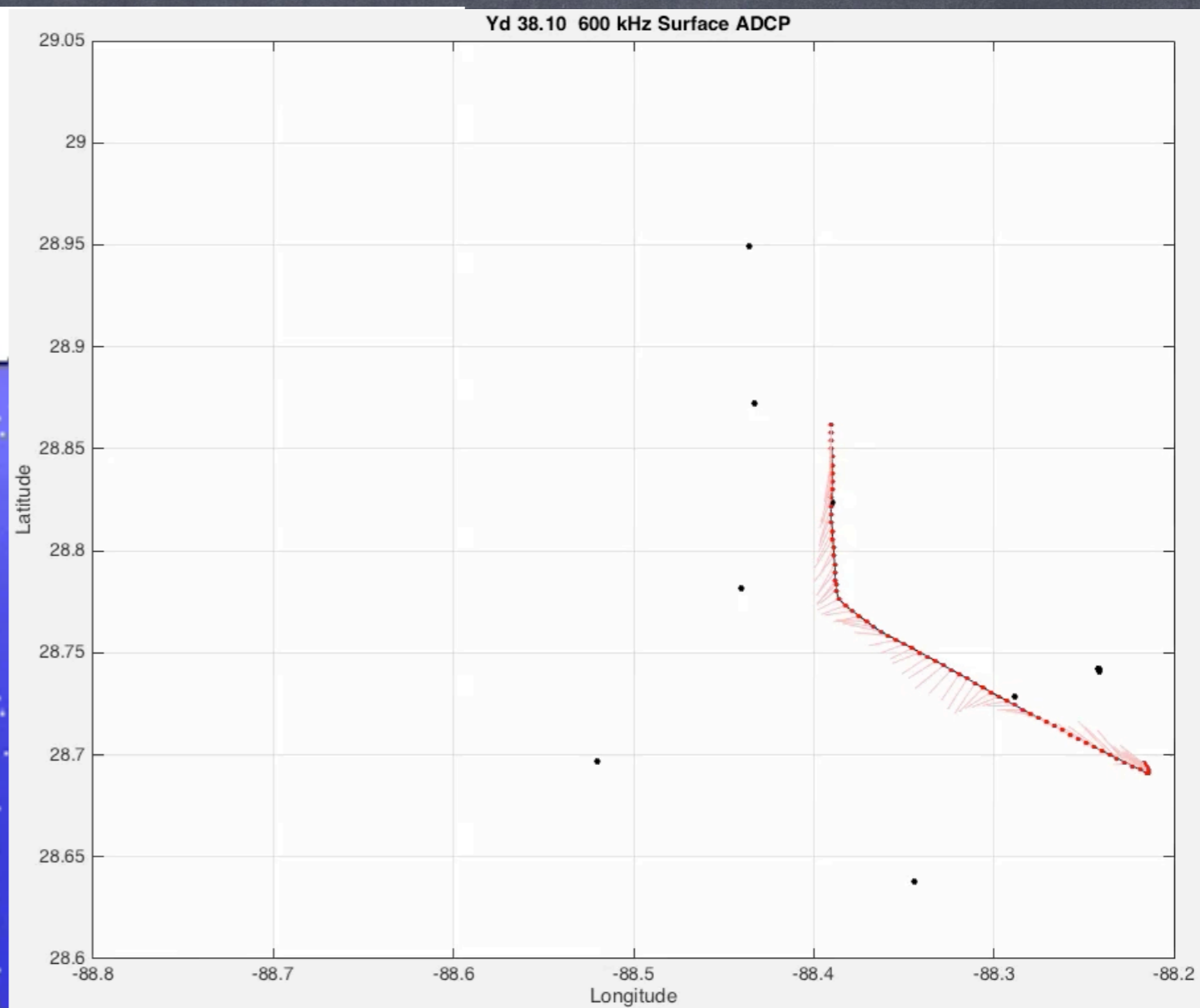
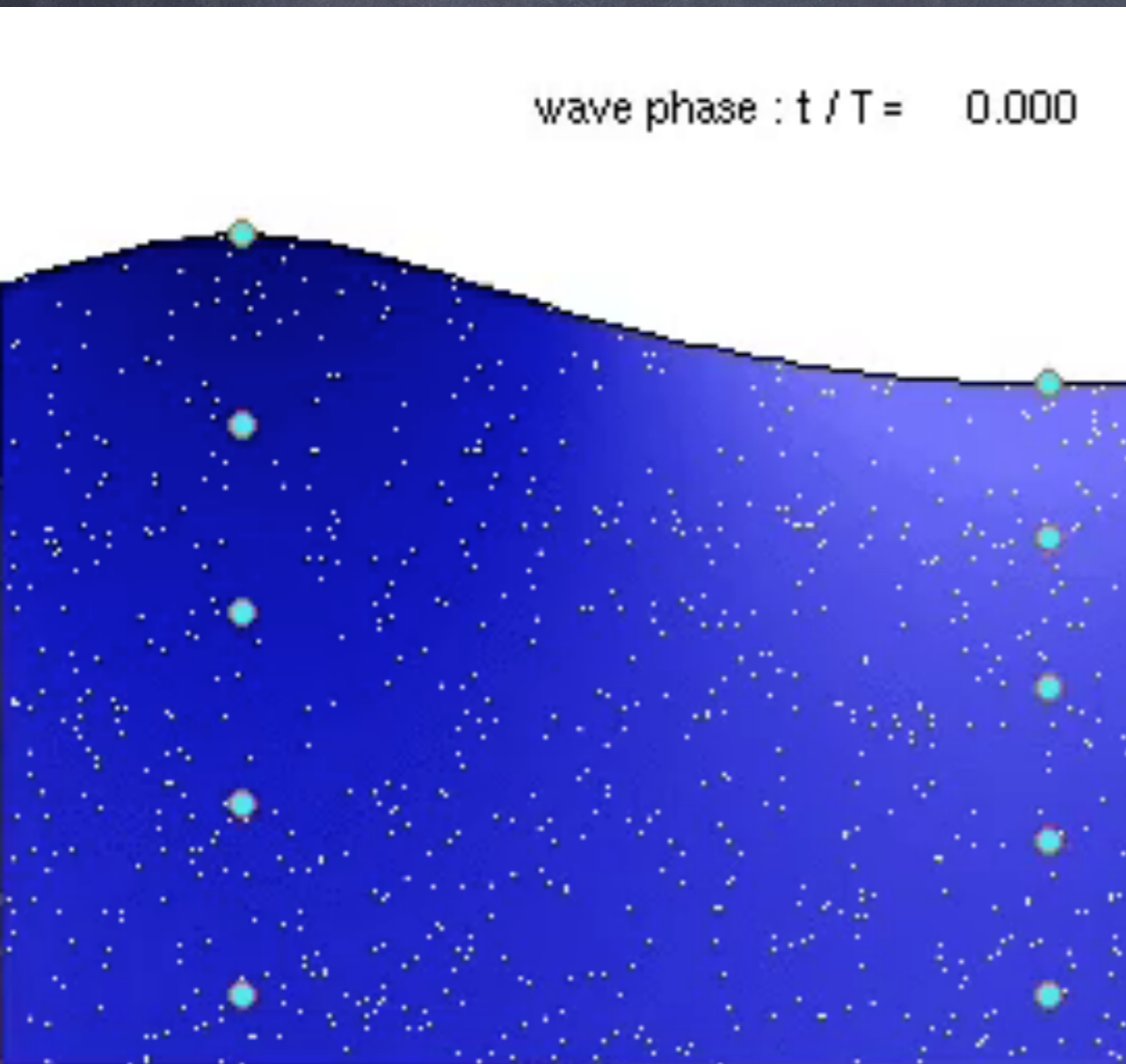
# Oil Modeling (Old)

- Coastal models (hydrostatic, 5km to 100m), e.g., versions of ROMS, NCOM, ADCIRC, UWIN-CM, etc, are circulation models with simple 1D mixing & simple diffusion coupled to wave models
- Large Eddy Simulations (nonhydrostatic, 1m res.) show that accumulation at Langmuir cells & buoyant effects are important.  
Lots of detail lost in coastal models.
- Thus, the old approach combined empirical oil spreading & coastal models to get at the upper ocean transport & dispersion.

# Eulerian & Lagrangian 2 FLAVORS in CARTHE

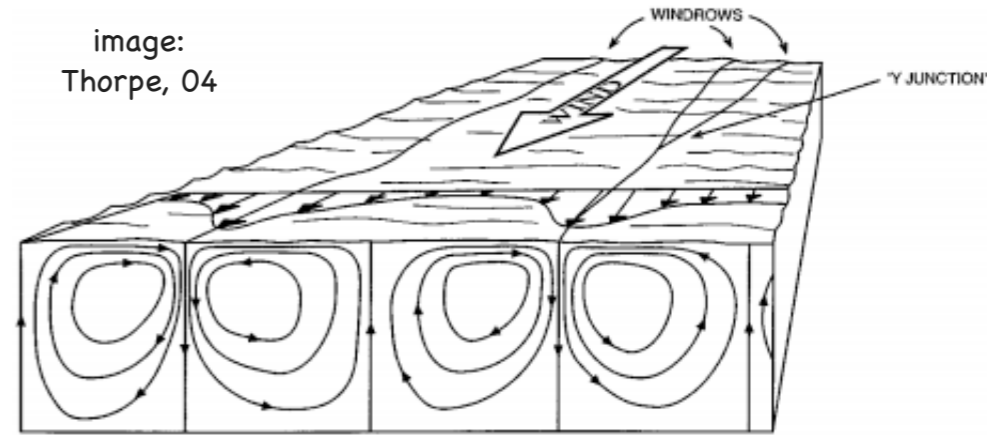
1: Waves

2: Drifters



# The Character of the Langmuir Scale

image:  
Thorpe, 04



**Figure 1** Sketch showing the pattern of mean flow in idealized Langmuir circulation. The windrows may be 2 m to 300 m apart, and the cell form is roughly square (as shown). In practice the flow is turbulent, especially near the water surface, and the windrows (Figure 2) amalgamate and meander in space and time. Bands of bubbles or buoyant algae may form within the downward-going (or downwelling) flow (see Figure 3).



Carlson et al. 2018

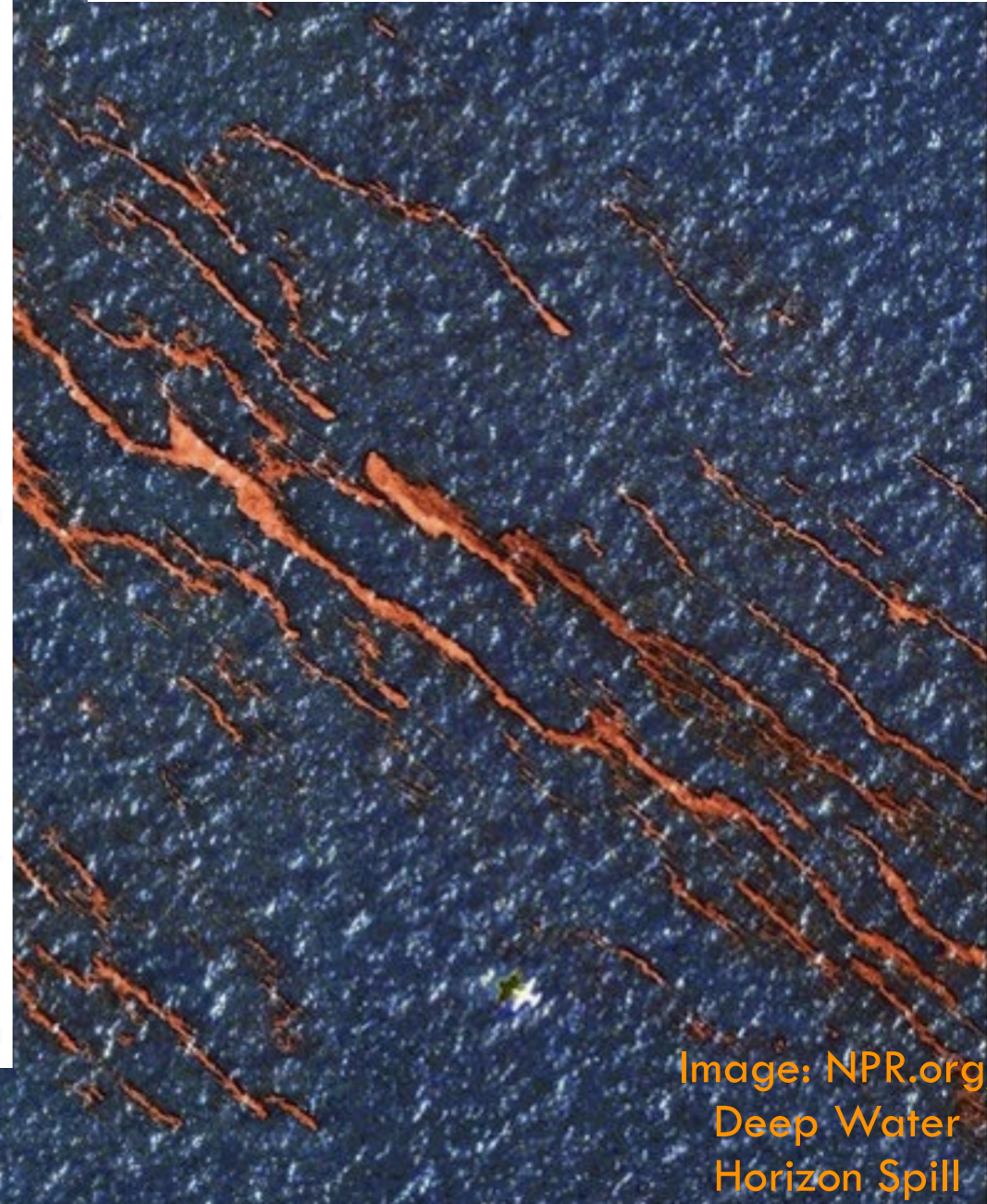


Image: NPR.org,  
Deep Water  
Horizon Spill

# 3 Stokes Forces appear in the “Wave-Averaged Equations”:

All rely only on Stokes drift of waves

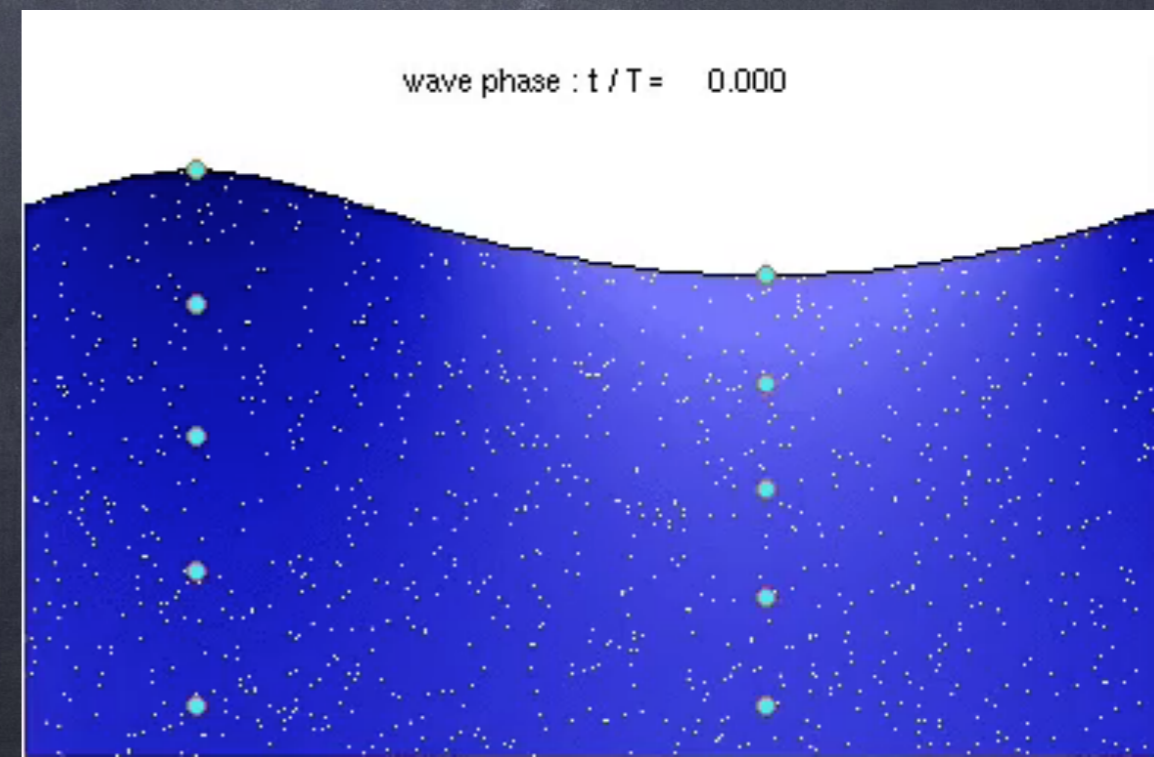
1: Stokes Advection: parcels, tracers, momentum move with Lagrangian, not Eulerian flow

2: Stokes Coriolis: water parcels experience Coriolis force during this motion

3: Stokes Shear Force

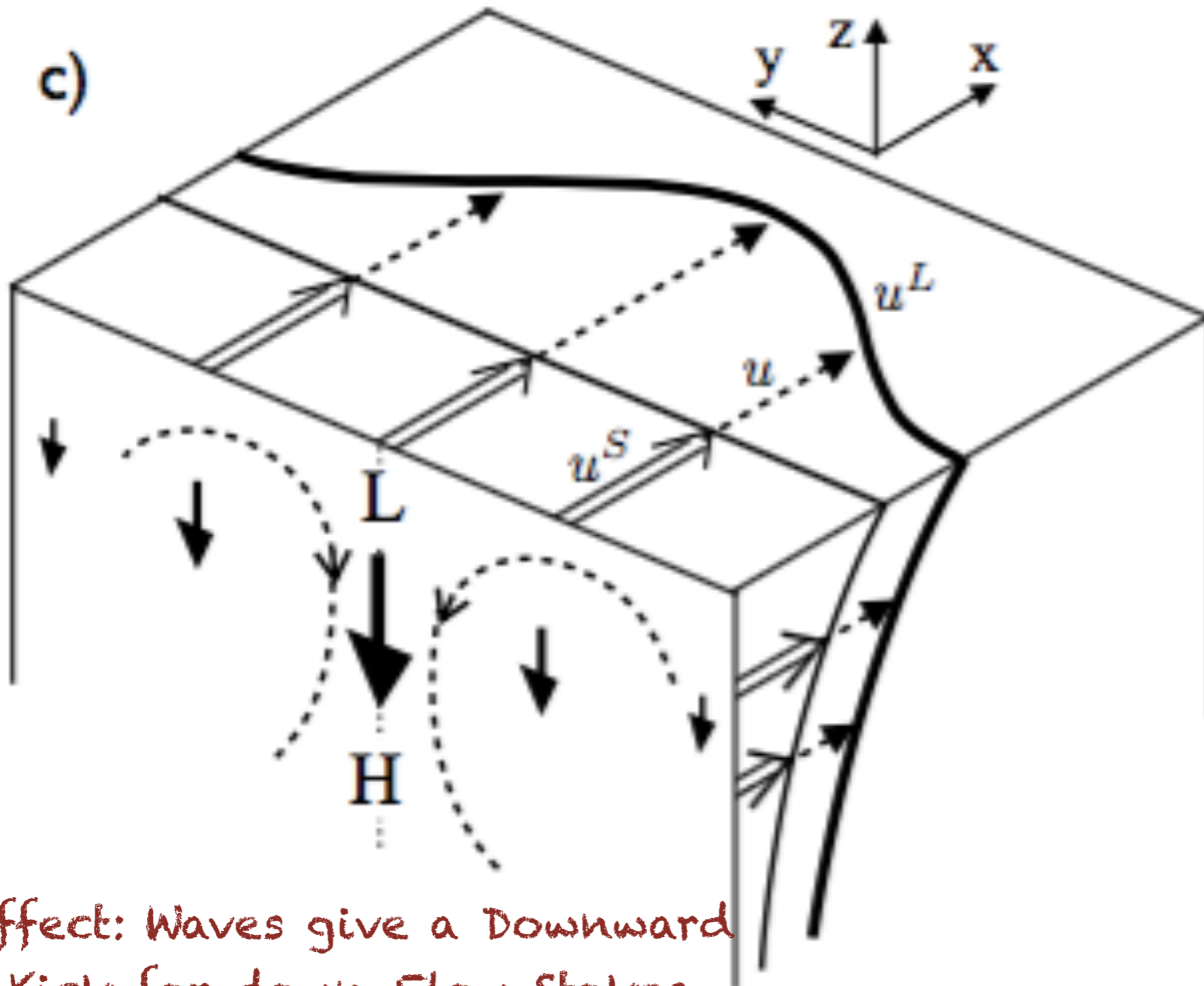
N. Suzuki and BFK. Understanding Stokes forces in the wave-averaged equations. *Journal of Geophysical Research-Oceans*, 121:1-18, 2016.

Notable: Wang & Ozgokmen (2017) examine these forces with wave-resolving models



# Even before we get to the good bits...

- First guess: Stokes drift moves stuff, including oil, on the surface, e.g., Curcic et al. (2016), Judt et al. (2016).
- Problem: Anti-Stokes Eulerian response is critical, which only results if the Stokes forces are included in momentum equations, too.
- Outcome: should include Stokes forces in ocean models—our new formulation allows in hydrostatic models.



**Figure 1**  
windrows  
practice th  
amalgama  
within the

Effect: Waves give a Downward  
Kick for down-Flow Stokes

$$M_{Ro} \equiv \max(1, Ro)$$

$$Re = \frac{UL}{\nu}$$

$$Ro = \frac{U}{fL}$$

$$Ri = \frac{N^2}{(U_z)^2}$$

$$\alpha = H/L$$

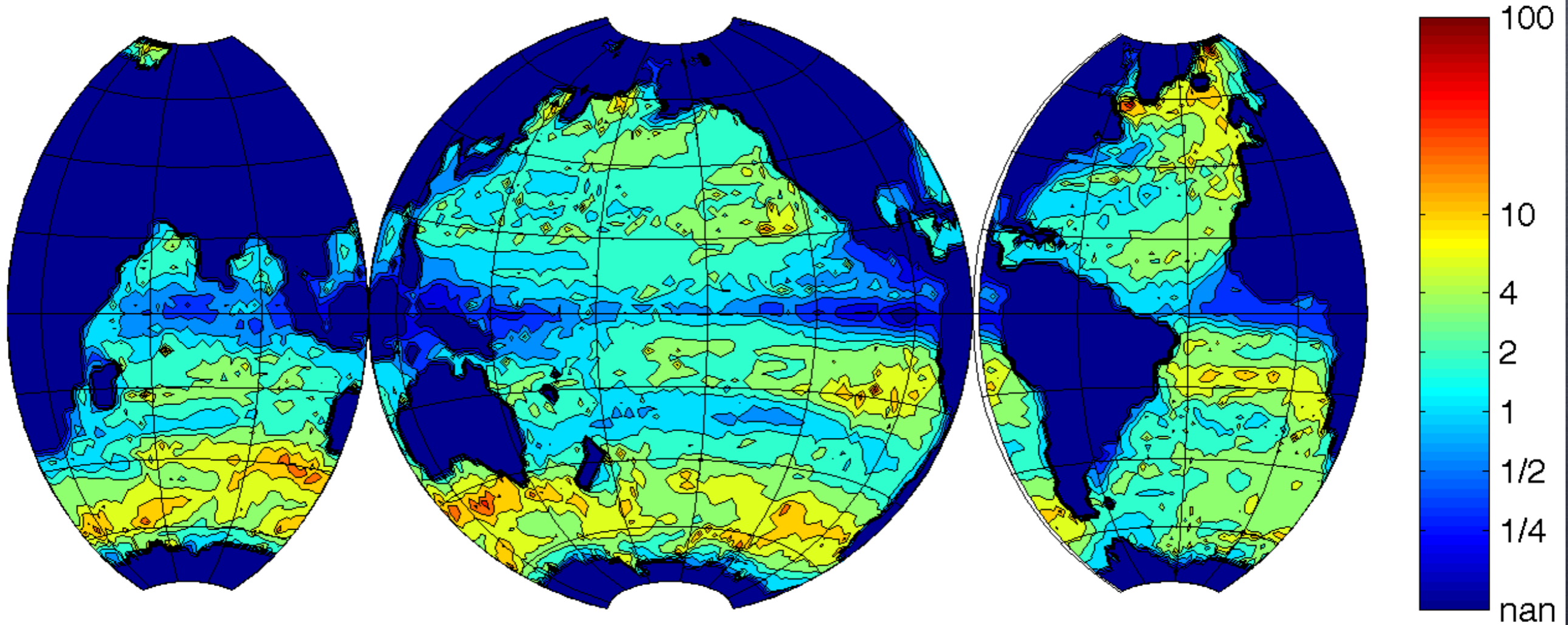
$$\epsilon = \frac{V^s H}{fLH_s}$$

"wavy hydrostatic" if  $\epsilon \gg 1$

$$\frac{\alpha^2}{Ri} \left[ w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b - \epsilon v_j^L v_{j,z}^s} + \frac{\alpha^2}{Re Ri} w_{,jj}$$

Do Stokes force directly affect larger scales?

$\epsilon/Ro$



“wavy hydrostatic”

$$\epsilon \gg 1$$

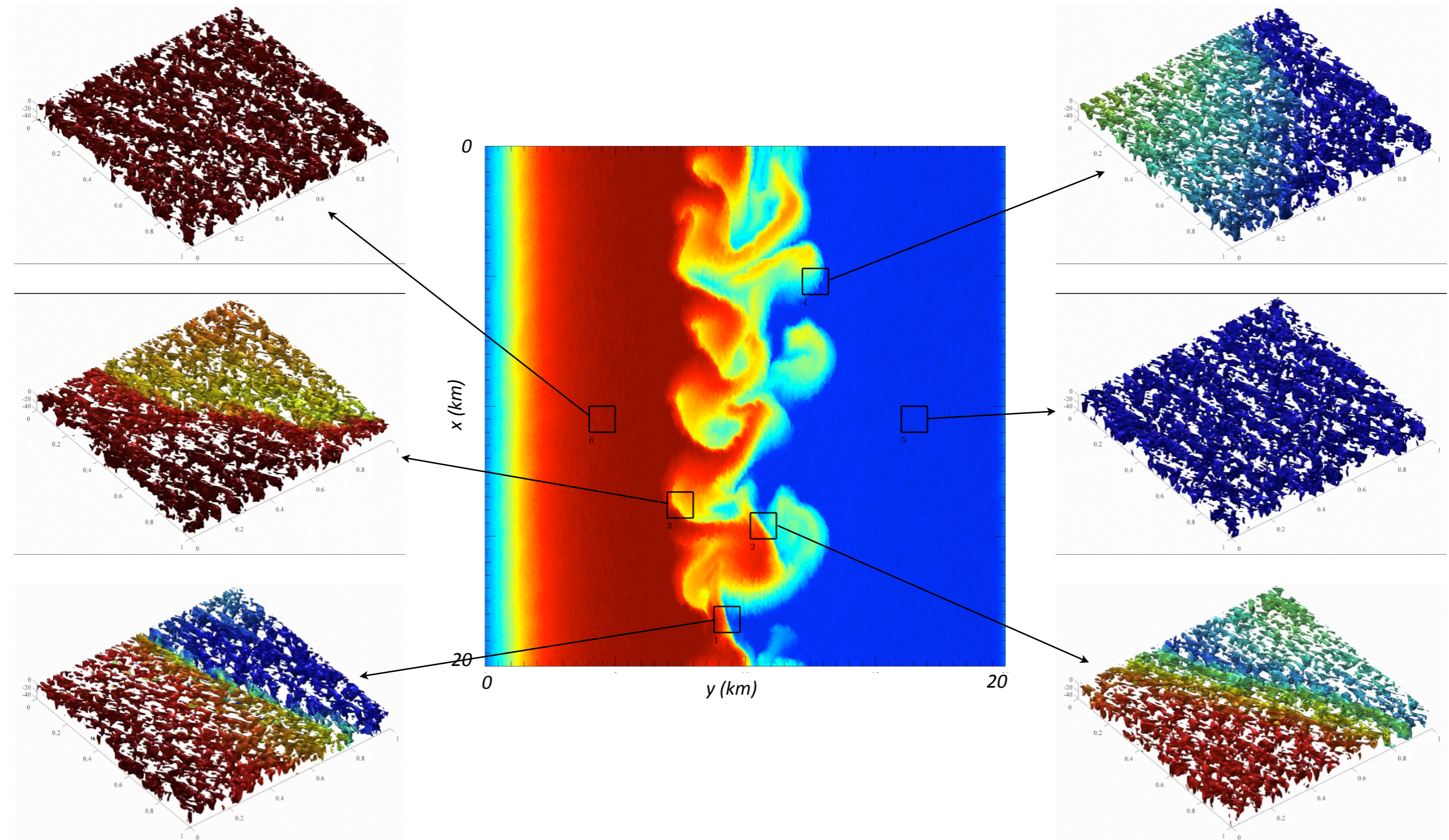
$$\epsilon = \frac{V^s H}{f L H_s}$$

$$Ro = \frac{U}{f L}$$

J. C. McWilliams and BFK. Oceanic wave-balanced surface fronts and filaments. *Journal of Fluid Mechanics*, 730:464-490, 2013.

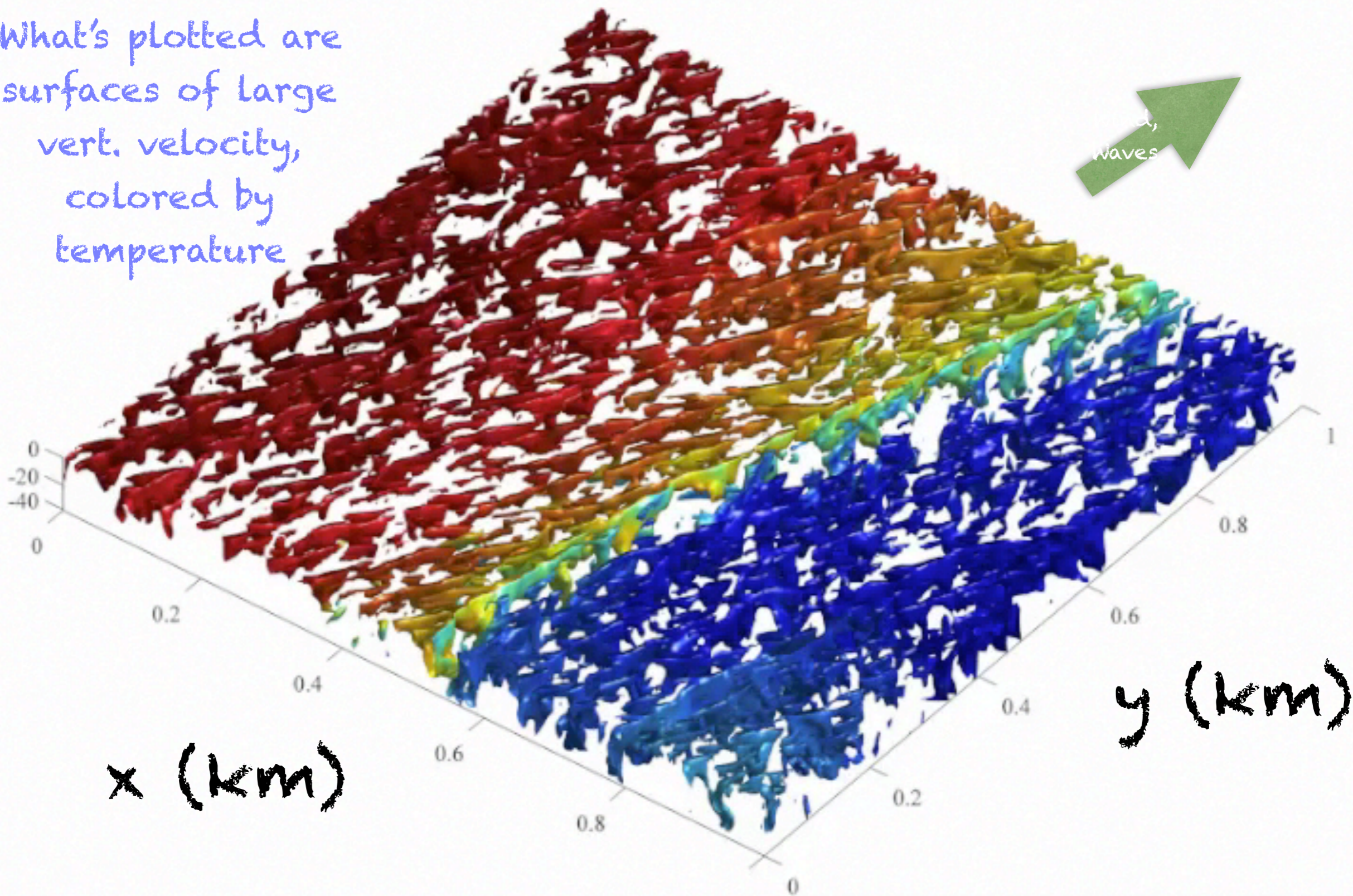


# A Large Eddy Simulation spanning the submesoscale (100m-10km) and boundary layer turbulence (5-50m)

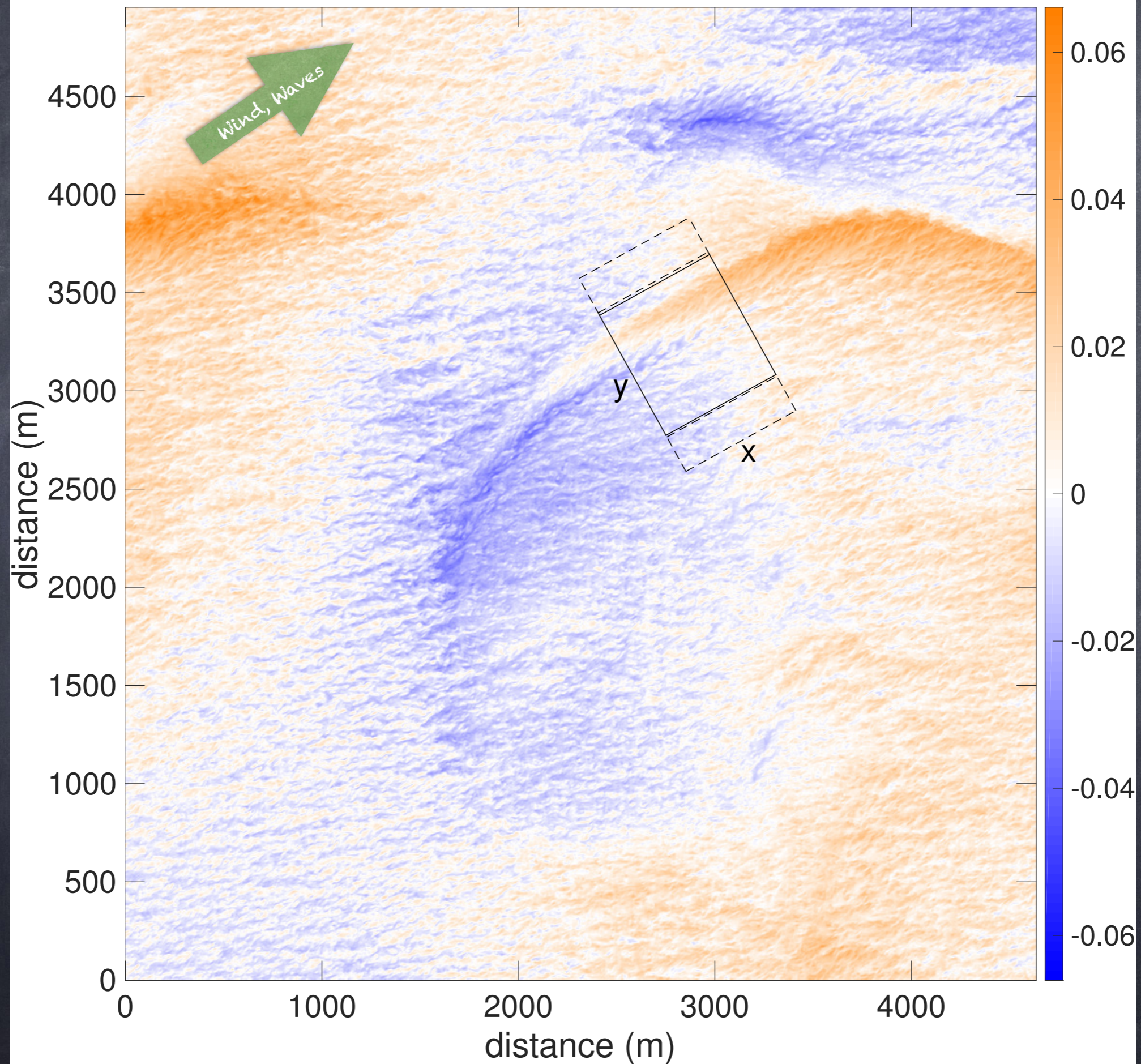


P. E. Hamlington, L. P. Van Roekel, BFK, K. Julien, and G. P. Chini. Langmuir-submesoscale interactions: Descriptive analysis of multiscale frontal spin-down simulations. *Journal of Physical Oceanography*, 44(9): 2249-2272, September 2014.

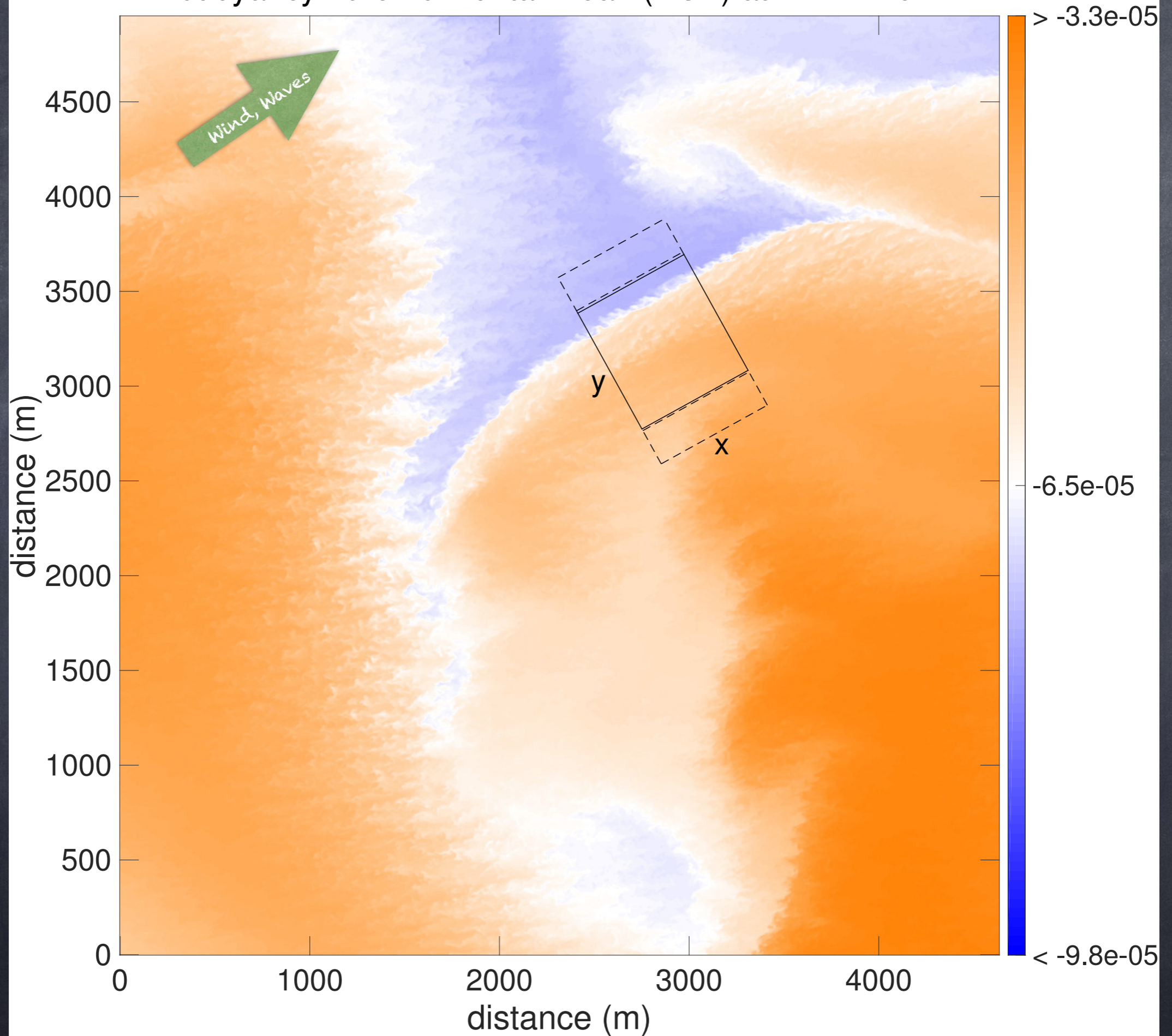
What's plotted are  
surfaces of large  
vert. velocity,  
colored by  
temperature



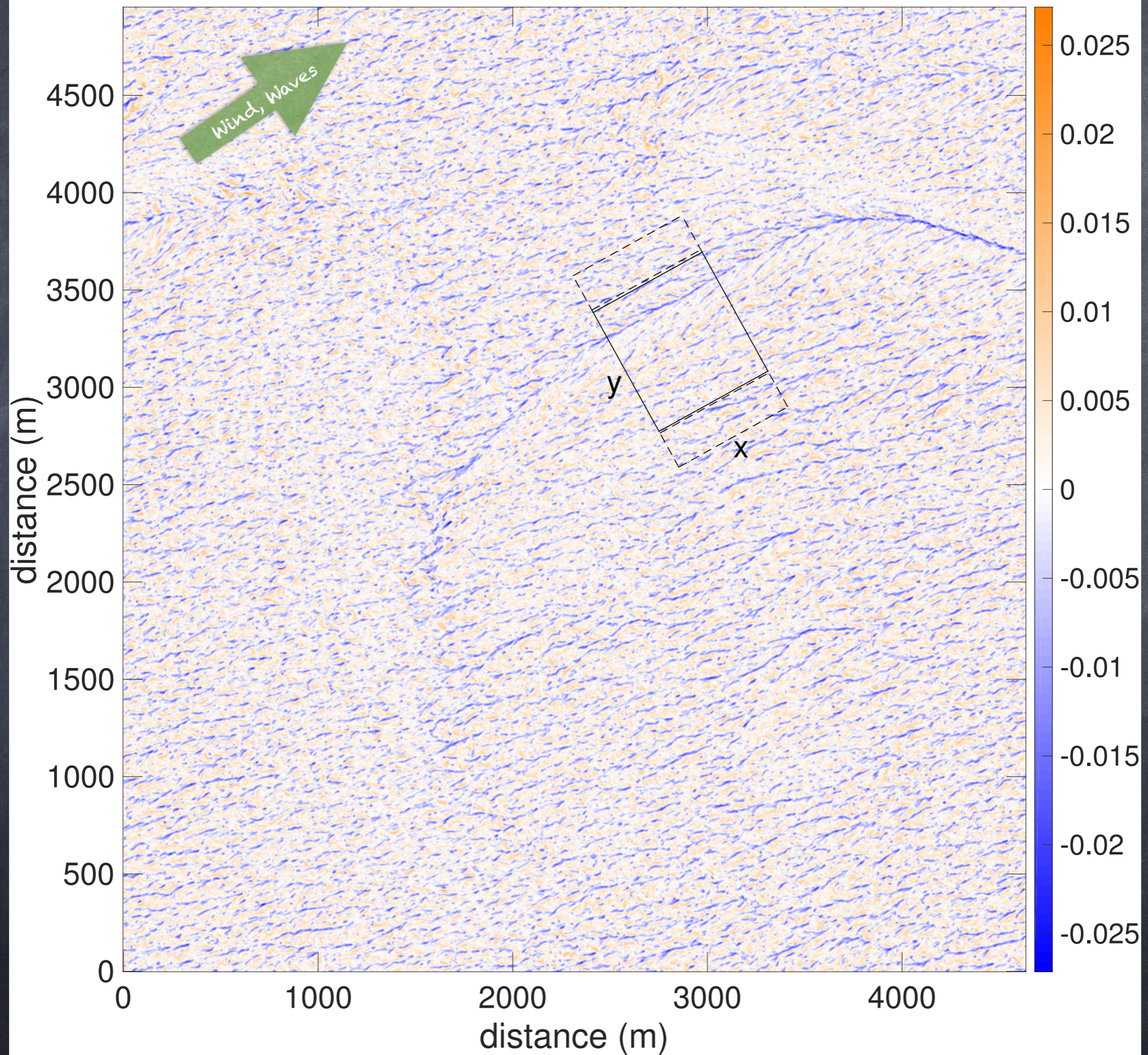
velocity in the x-direction - the horizontal mean ( $\text{ms}^{-1}$ ) at  $z = -11.25\text{m}$



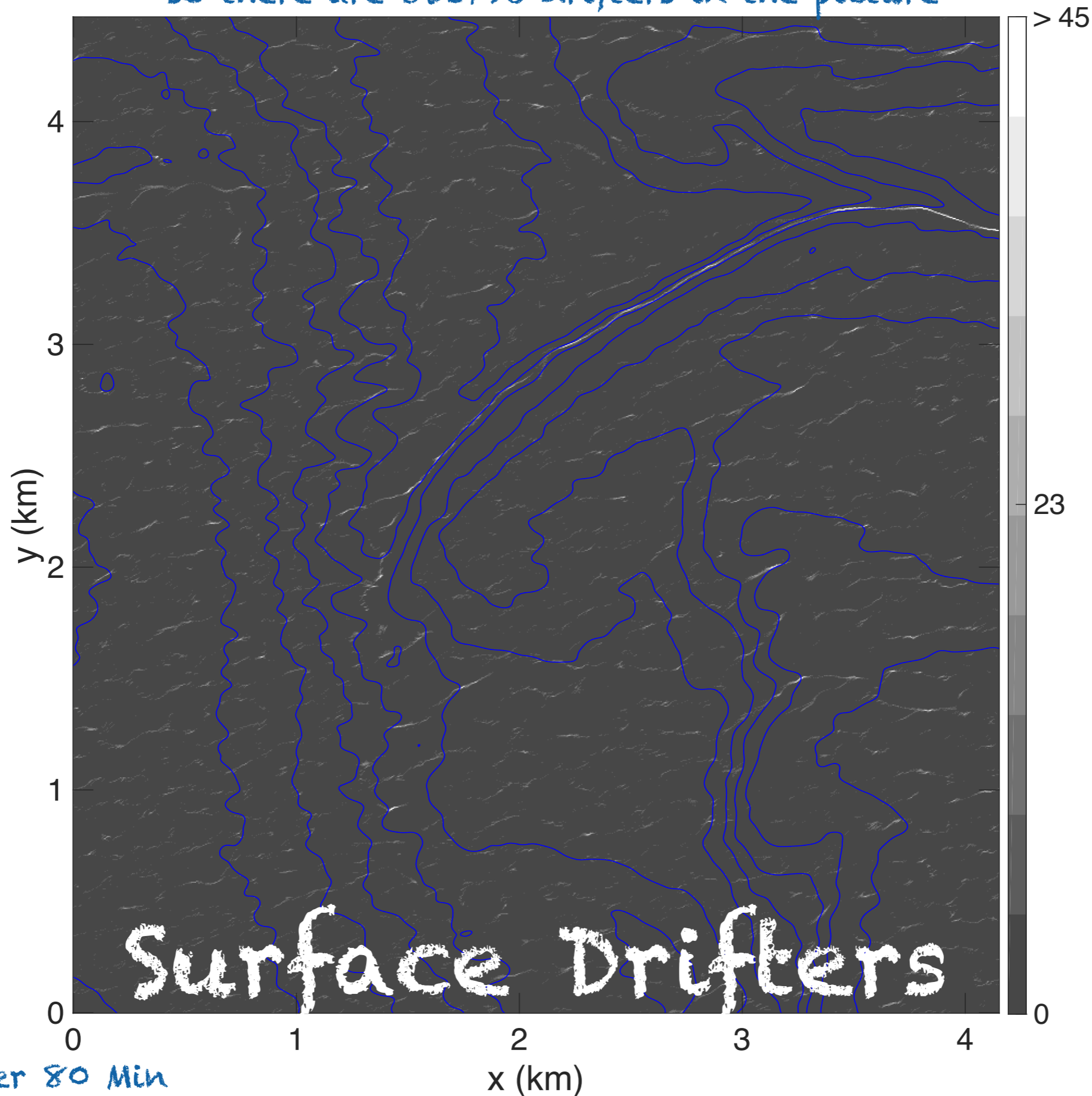
buoyancy - the horizontal mean ( $\text{ms}^{-2}$ ) at  $z = -11.25\text{m}$



vertical velocity ( $\text{ms}^{-1}$ ) at  $z = -11.25\text{m}$

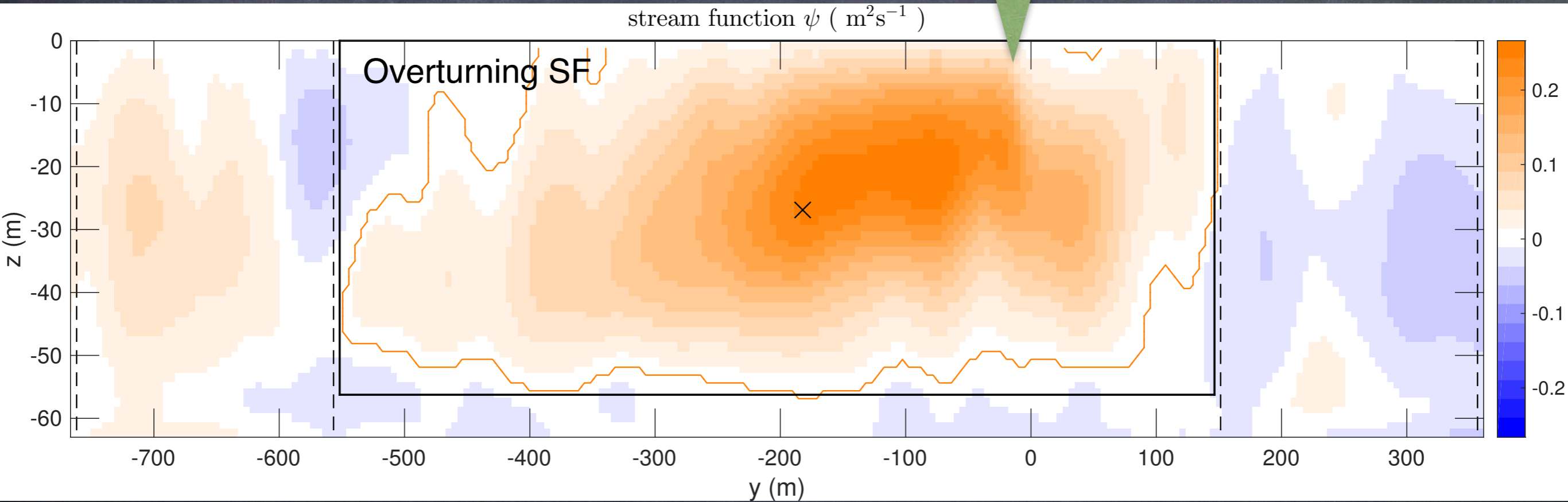
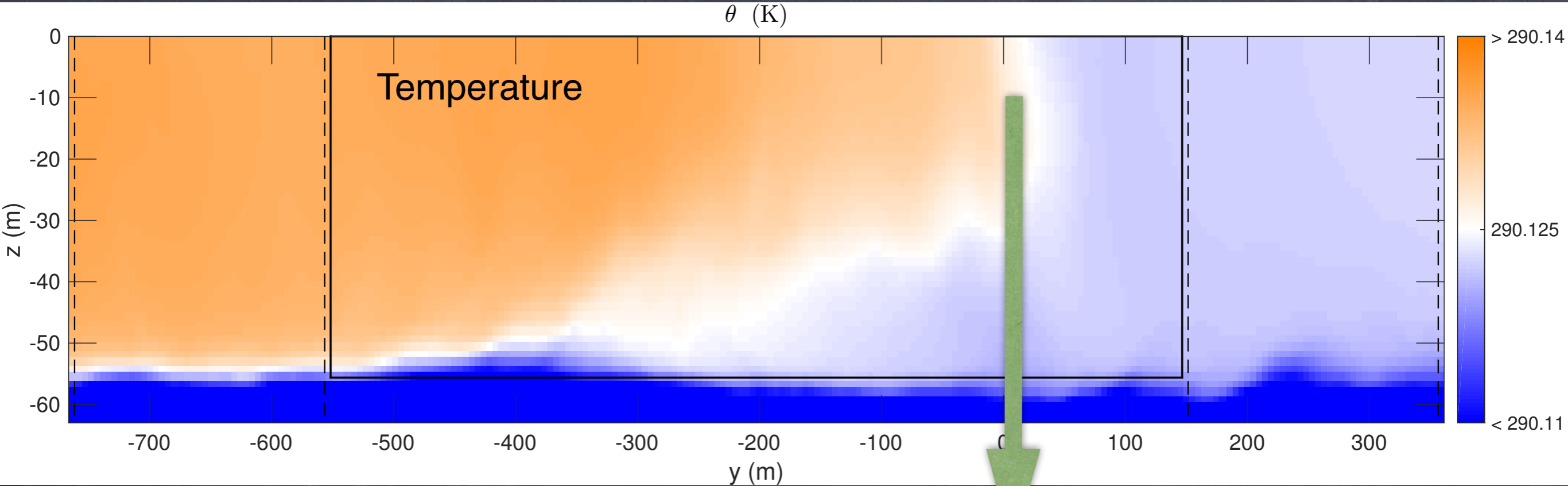


Initially every surface node has 1 drifter, so there are 851796 drifters in the picture



After 80 Min

See also Harcourt & Taylor (2018) buoyant LES.



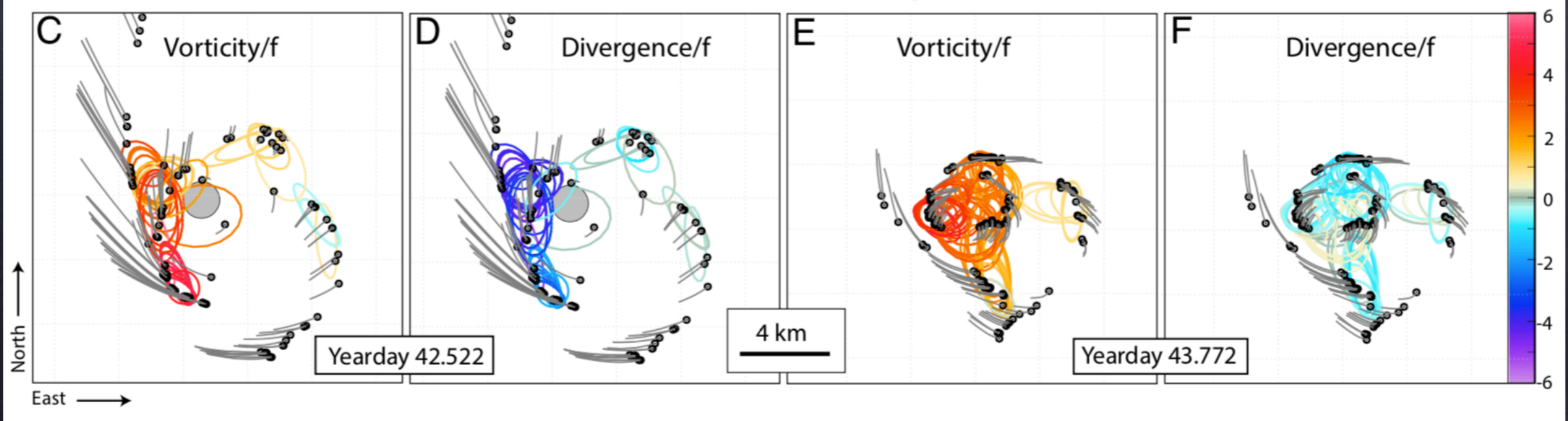
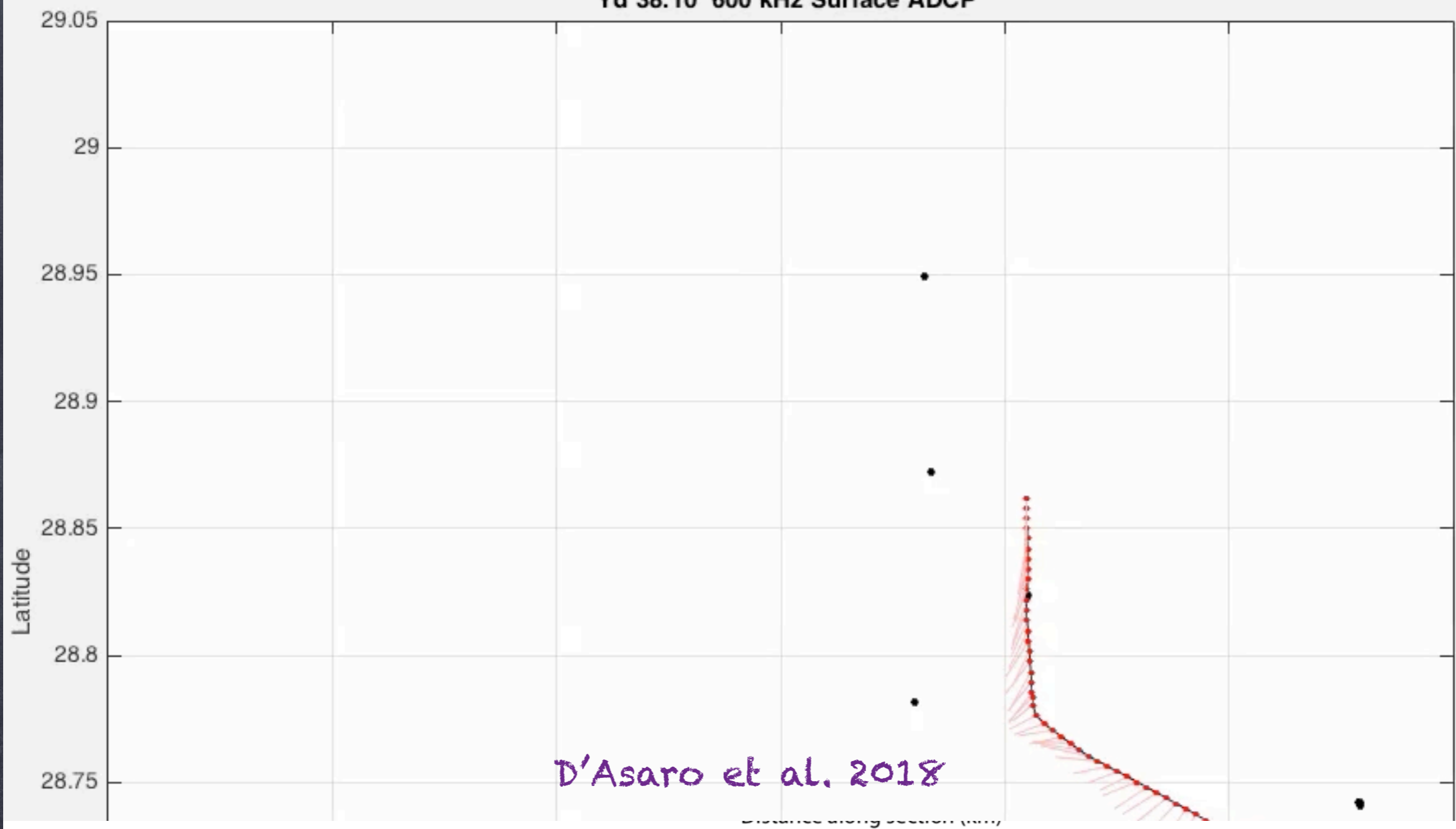
N. Suzuki, BFK, P. E. Hamlington, and L. P. Van Roekel. Surface waves affect frontogenesis. *Journal of Geophysical Research-Oceans*, 121:1-28, 2016.

# Stokes Drift = Lagrangian-Eulerian, Energizes Turbulence (old) & Fronts (new)

- ◉ **New:** Tracer & drifter advection by Stokes drift can be included in coastal (nonhydrostatic) models; care is needed to preserve anti-Stokes flow—i.e., consistent application of Stokes forces.  
**Future:** wider application of these approaches
- ◉ **New:** Stokes Shear force energizes fronts  
**Future:** modeled, but not yet observed directly
- ◉ **Improved:** Turbulent mixing can be parameterized into Langmuir schemes for turbulence effects.  
**Future:** include buoyant tracer effects



Yd 38.10 600 kHz Surface ADCP



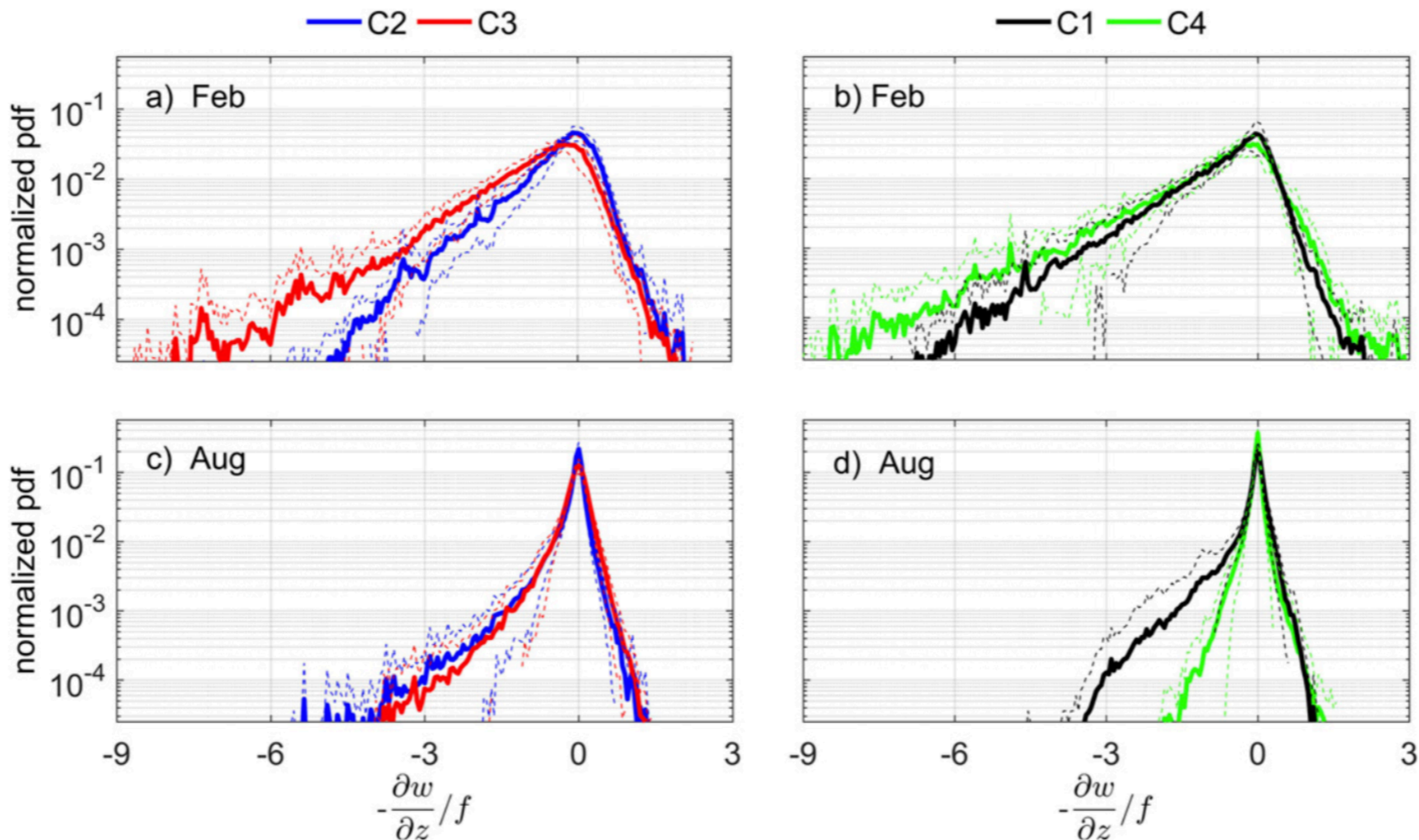


FIG. 12. Horizontal velocity divergence distributions for the four clusters in Fig. 1 in the HR With-River solution [(left) C2 (blue) and C3 (red); (right) C1 (black) and C4 (green)]. PDFs are averaged over eight deployments (thick lines) 3 days after release. The standard error, calculated as twice the standard deviation of the eight deployments, is also indicated (dashed lines). (a),(b) February and (c),(d) August.

# GLAD: Isotropic Forward Cascade from 30km?

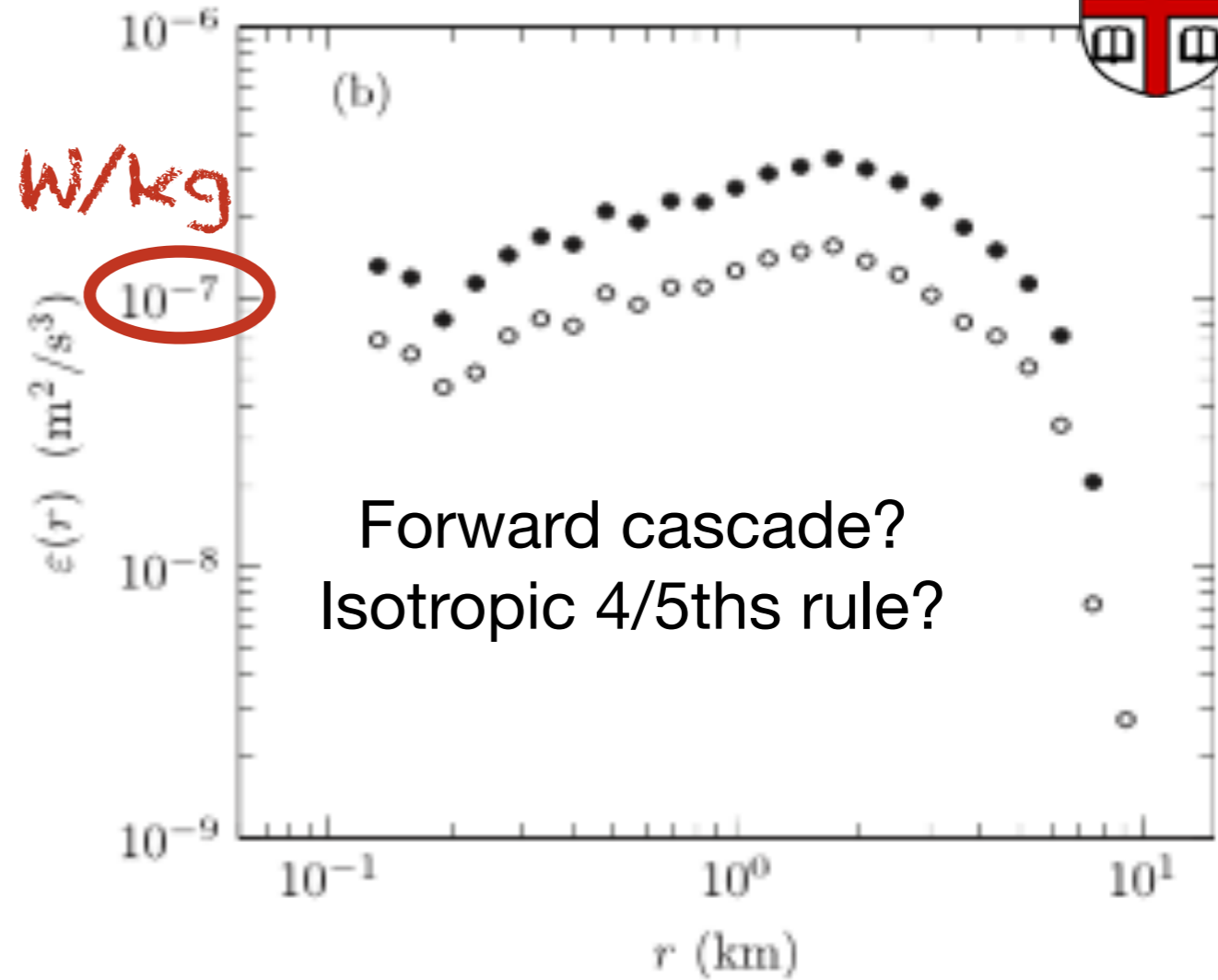
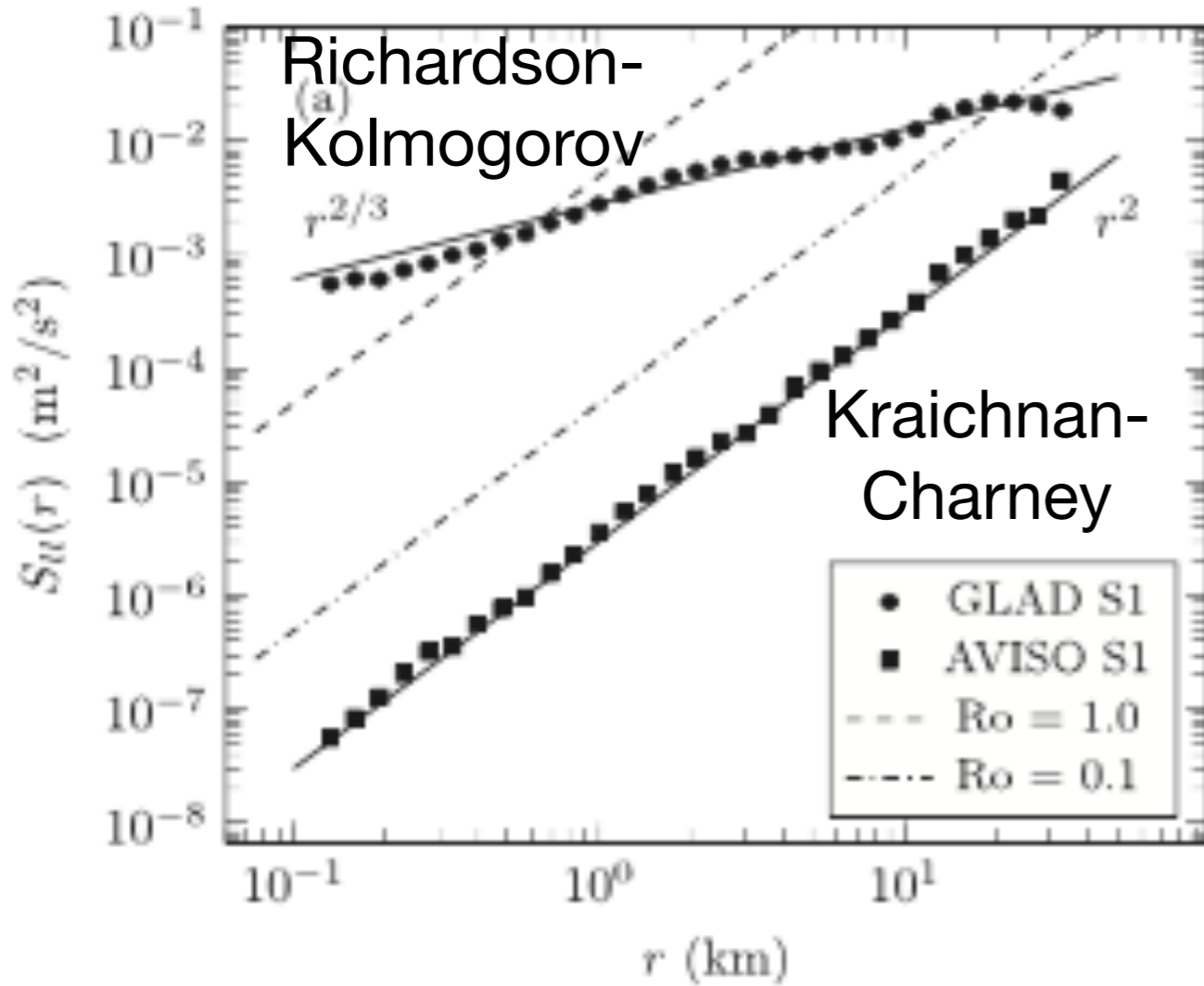


FIG. 7. (a) Second order longitudinal structure function versus separation distance showing Richardson-Kolmogorov,  $r^{2/3}$ , energy cascade scaling for GLAD data and Kraichnan,  $r^2$ , enstrophy cascade scaling for AVISO-based synthetic trajectories. (b) Sign-reversed third order longitudinal structure function scaled by  $r$  for the GLAD observations.

$140TW \text{ (global)} / (1.4 \cdot 10^{21} \text{ kg}) = 10^{-7} \text{ W/kg}$   
 Winds:  $\sim 20TW$  global  
 Tides:  $3.5TW$  global

D'Asaro et al (2011): Enhanced @ Fronts:  $10^{-5}$  to  $10^{-6} \text{ W/kg}$

**Evidence of a forward energy cascade and Kolmogorov self-similarity in submesoscale ocean surface drifter observations**

Andrew C. Poje, Tamay M. Özgökmen, Darek J. Bogucki, and A. D. Kirwan, Jr.

# "Cascade" Scalings

3D: Richardson/Kolmogorov/Smagorinsky/Corrsin

$$E \propto \epsilon^{2/3} \ell^{5/3}, \quad S_2 \propto \epsilon^{2/3} r^{2/3}, \quad \epsilon \propto \nu \alpha^2, \quad \nu = \text{Pr} \kappa \propto \Delta x^2 |\alpha| \propto \epsilon^{1/3} \ell^{4/3}$$

Drifters?

2D/QG: Barnier/Kraichnan/Charney/Leith

$$E \propto \eta^{2/3} \ell^3, \quad S_2 \propto \eta^{2/3} r^2, \quad \eta \propto \nu (\nabla \omega)^2, \quad \nu \propto \Delta x^3 |\nabla \omega| \propto \eta^{1/3} \ell^2, \quad \kappa \propto ?$$

AVIS0?

Submesoscale: McWilliams/?/?F-K?

$$E \propto \ell^2, \quad S_2 \propto r^1, \quad d/dt(PE + KE) = ??, \quad \nu = ?, \quad \kappa = ?$$

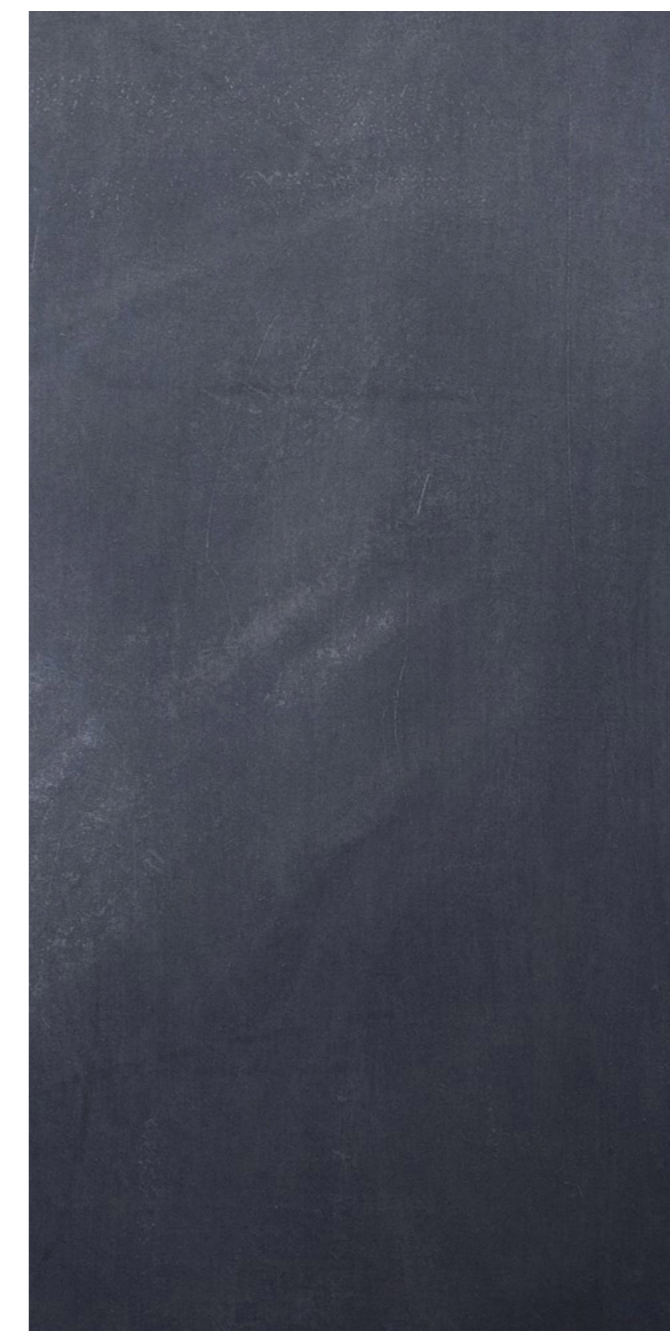
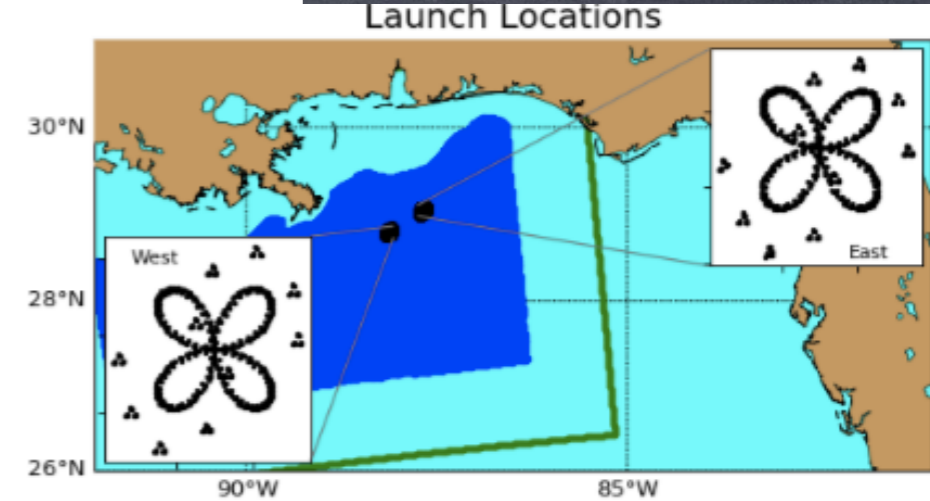
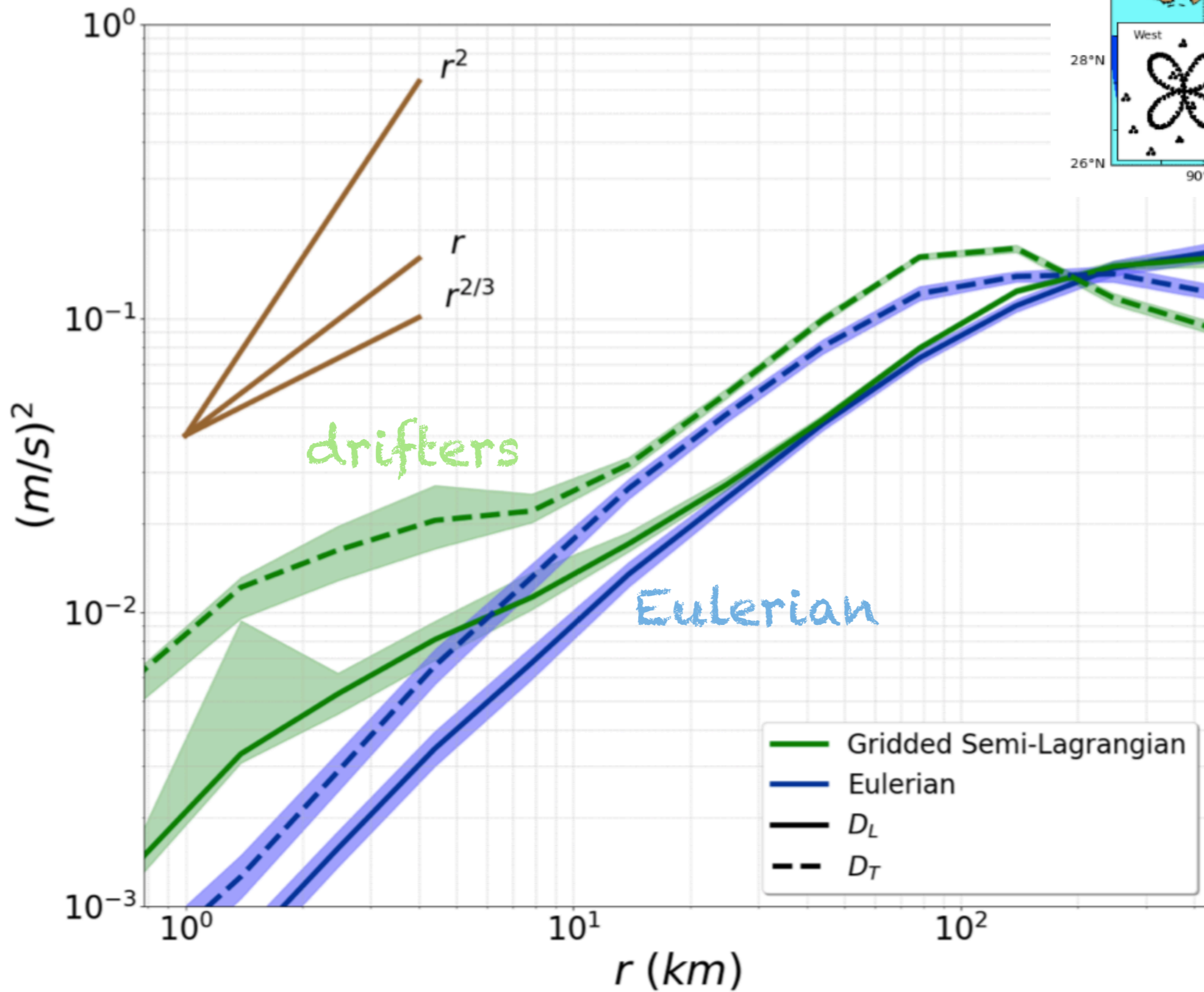
ROMS?



**Near Surface Vertical Velocity**

Courtesy of J. Molemaker & J. McWilliams

# Gridded Semi-Lagrangian and Eulerian Second Order Structure Functions

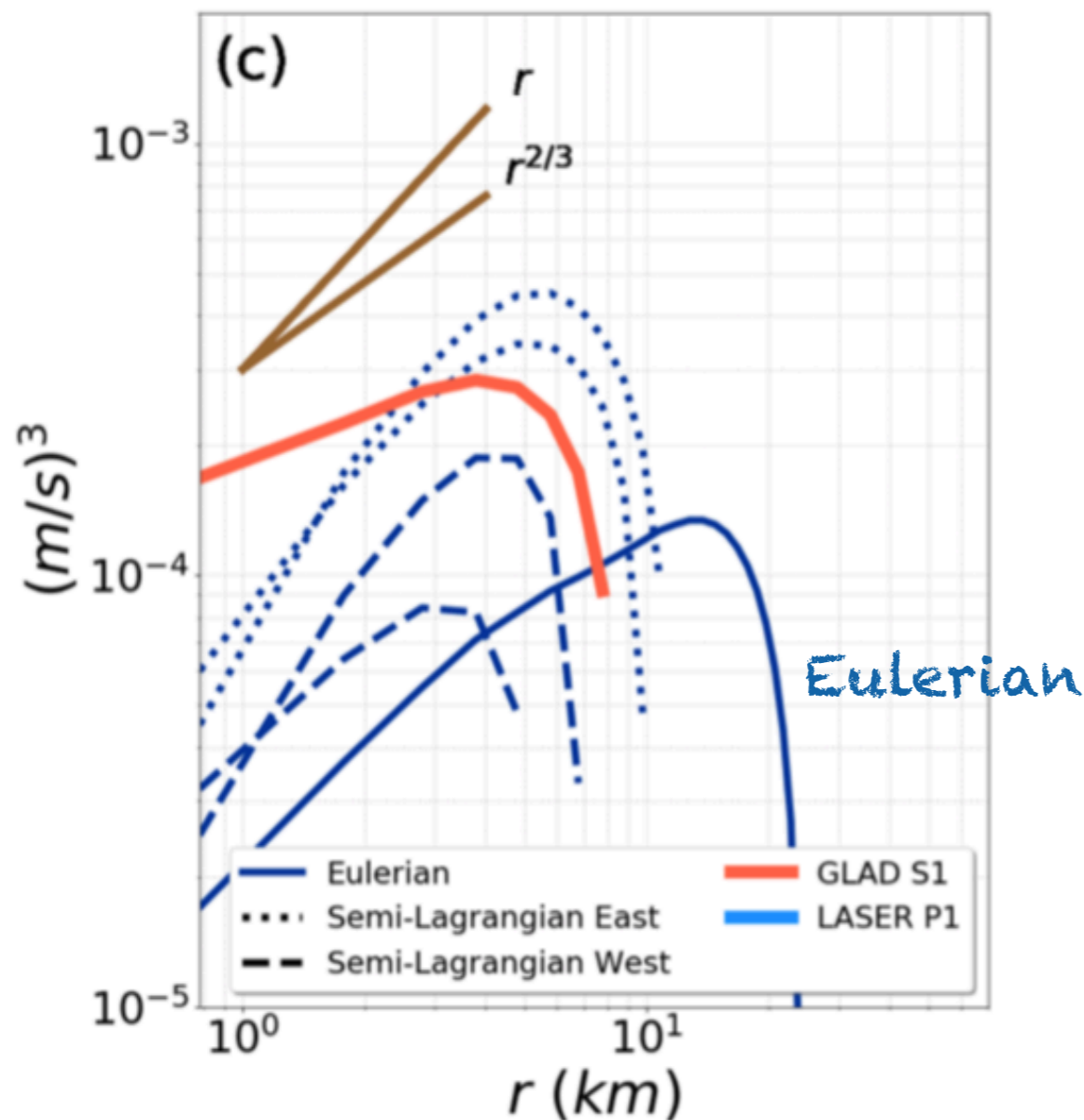


J. Pearson, B. Fox-Kemper, R. Barkan, J. Choi, A. Bracco, and J. C. McWilliams. Impacts of convergence on Lagrangian statistics in the Gulf of Mexico. *Journal of Physical Oceanography*, 2019. In press.

# Eulerian vs. Lagrangian: Lagrangian is biased toward sampling convergent, energetic fronts & etc.

$D_L^3 < 0$  "forward cascade"

$$\epsilon \propto \langle ([\mathbf{u}(\mathbf{x} + \mathbf{r}) - \mathbf{u}(\mathbf{x})] \cdot \mathbf{r} / |\mathbf{r}|)^3 \rangle / |\mathbf{r}|$$



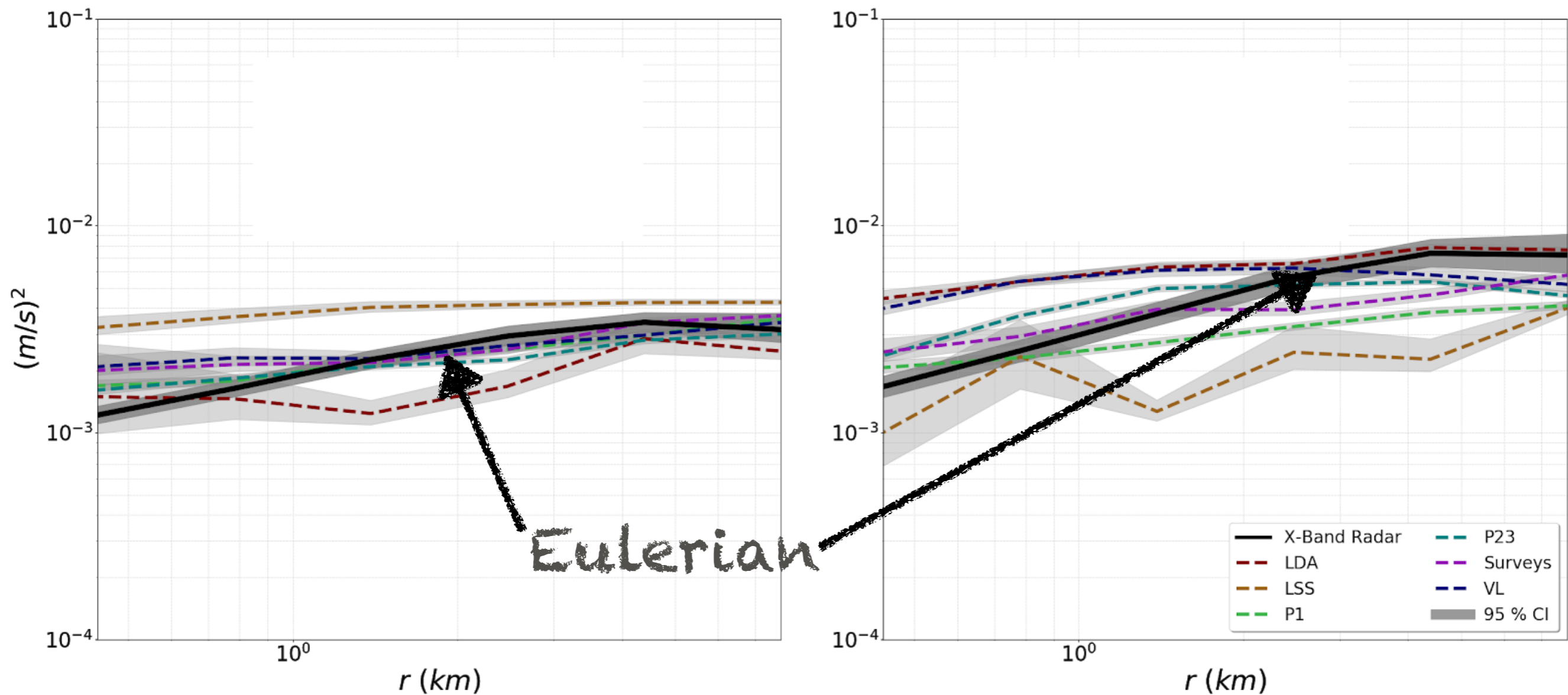
Lagrangian  $\epsilon$   
overestimates  
by factor of  
2 to 8 on  
1km to 10km

# Same Story in LASER Observations (w/ UDEL.)

(Note: compensated w/ 3D slope)

Longitudinal  $S2/r^{(2/3)}$

Transverse  $S2/r^{(2/3)}$





# Conclusions

- **Old & Improved:** Eulerian and Lagrangian flows, and the difference between them, have major implications for modeling, observation, and analysis.
- **Old:** Stokes drift advects, but (New:) Stokes forces affect Eulerian flows. Stokes forces energize Langmuir turbulence & down-Stokes fronts.  
**New and Future:** coastal model improvements
- **New:** Turbulence statistical scalings from drifters & Eulerian measurements differ on the submesoscales in models & obs, because drifters sample fronts more.  
**Future:** drifter theory & assimilation techniques including bias.
- **Conjecture:** the dynamics governing the drifters seems to resemble 3D turbulence, so perhaps energy dissipation rate at fronts is a key to open the scaling factor challenge.

# Wave-Averaged Equations

following Lane et al. (07), McWilliams & F-K (13)

and Suzuki & F-K (15)

(for horizontally uniform Stokes drift)

(Lagrangian) geostrophic

$$Ro [v_{i,t} + v_j^L v_{i,j}] + \frac{M_{Ro}}{Ri} w v_{i,z} + \boxed{\epsilon_{izj} v_j^L} = -M_{Ro} \pi_{,i} + \frac{Ro}{Re} v_{i,jj}$$

$$\frac{\alpha^2}{Ri} \left[ w_{,t} + v_j^L w_{,j} + \frac{M_{Ro}}{Ro Ri} w w_{,z} \right] = \boxed{-\pi_{,z} + b} - \epsilon v_j^L v_{j,z}^s + \frac{\alpha^2}{Re Ri} w_{,jj}$$

hydrostatic

$$b_t + v_j^L b_{,j} + \frac{M_{Ro}}{Ro Ri} w b_z = \frac{1}{Pe} b_{,jj}$$

Plus boundary condition

$$v_{j,j} + \frac{M_{Ro}}{Ro Ri} w_z = 0$$

$$M_{Ro} \equiv \max(1, Ro)$$

$$Re = \frac{UL}{\nu} \quad Ro = \frac{U}{fL} \quad Ri = \frac{N^2}{(U_{,z})^2}$$

$$\alpha = H/L \quad \epsilon = \frac{V^s H}{fLH_s}$$