PARAMETERIZATION OF SUBMESOSCALE AND LANGMUIR-SCALE PROCESSES AND INTERACTIONS Baylor Fox-Kemper (CU-Boulder & CIRES) with P.E. Hamlington (CU-Boulder), L. Van Roekel (Northland College), & P.P. Sullivan (NCAR)

OS2012 Abstract 10618, 016: Dynamics and Observations of Submesoscale Oceanic Processes Wed. 2/22/2012, 14:15-14:30, Room 251

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# Results

- Submesoscale mixed layer eddy restratification
- Search Langmuir turbulence mixing
  - Both are important to global surf. layer (10%)
  - Both are regionally more important (x2 to x4)
- Momentum=Langmuir, Stratification Change=Both
- Coupling between these scales exists! Parameterize?

# Parameterizations

 Anyone who doesn't take truth seriously in small matters cannot be trusted in large ones either.

--AlbertEinstein



### The Character of <sup>10</sup> km the Submesoscale

17.1

(NASA GSFC Gallery)

(Capet et al., 2008)



Longitude

Temperature on day:0



Fronts
Eddies
Ro=O(1)
Ri=O(1)
near-surface
1-10km, days

Eddy processes often baroclinic instability (Boccaletti et al '07, Haine & Marshall '98).





#### Physical Sensitivity of Ocean Climate to Submesoscale Eddy Restratification: MLE implemented in CCSM (NCAR), CM2M & CM2G (GFDL)

Bias

w/o

MLE









CM2M H<sub>mi</sub> Submeso-deBM (m) SEP



#### Deep ML Bias reduced

B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg,M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels.Parameterization of mixed layer eddies. III: Implementation and impact in

global ocean climate simulations. Ocean Modelling, 39:61-78, 2011.

Monday, March 5, 2012



#### Improves CFCs (water masses)

**Bias with MLE** 

#### Bias w/o MLE



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# The Character of

# the Langmuir Scale

- Near-surface
- Langmuir Cells & Langmuir Turb.
- Ro>>1
- Ri<1: Nonhydro</p>
- ⊘ 10–100m
- a 10s to mins
- w, u=O(10cm/s)
- Stokes drift
- Eqtns:Craik-Leibovich
- PARAMS IN DEVELOPMENT!

lmøge: NPR.org Deep Water Horizon Spill



Figure Ia Illustration of Langmuie circulations showing notation used in this review and surface and subsurface motions.

#### Langmuir Mixing Estimate from Climatology (Wind->Wave)



UNDERESTIMATES WAVE IMPACT



Crude estimate of the effect of Langmuir mixing in a forward ESM on MLD (m)

Data + LES, Southern Ocean mixing energy: Langmuir (Stokesdrift-driven) and Convective

But, how well do we know Stokes drift? (Turb. Lang. #=La = u<sup>\*</sup>/u<sub>s</sub>)



S.E. Belcher, A.A.L.M. Grant, K.E. Hanley, B. Fox-Kemper, L. Van Roekel, P.P. Sullivan, W.G. Large, A. Brown, A. Hines, D. Calvert, A. Rutgersson, H. Petterson, J. Bidlot, P.A.E.M. Janssen, and J.A. Polton. A global perspective on mixing in the ocean surface boundary layer. *Geophysical Research Letters*, 2011. In revision.

#### How well do we know Stokes Drift?

#### Reanalysis vs wave model

#### Altimetry vs wave model



Fig. 4. D<sub>2</sub> Comparison of ERA40 reanalysis and TOPEX satellite data with WW3 using eight year means (1994–2001).

#### Within a factor of 2. Assuming full-development (e.g., McWilliams & Restrepo, 1999) is worse

A. Webb and B. Fox-Kemper. Wave spectral moments and Stokes drift estimation. *Ocean Modelling*, 40(3-4): 273-288, 2011

#### Real World Forcing: Misaligned Wind & Waves



L. Van Roekel, B. Fox-Kemper, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. *Journal of Geophysical Research-Oceans*, 2012. Submitted.



 $\overline{w'^2}/[u_*\cos(\alpha_{LOW})]^2$ 

Generalized Turbulent Langmuir No., Projection of u\*, u<sub>s</sub> into Langmuir Direction

$$\frac{\overline{w'^2}}{u_*^2} = 0.6 \cos^2 \left(\alpha_{LOW}\right) \left[1.0 + \left(3.1La_{proj}\right)^{-2} + \left(5.4La_{proj}\right)^{-4}\right],$$

$$La_{proj}^2 = \frac{|u_*|\cos(\alpha_{LOW})}{|u_s|\cos(\theta_{ww} - \alpha_{LOW})},$$

$$\alpha_{LOW} \approx \tan^{-1}\left(\frac{\sin\left(\theta_{ww}\right)}{\frac{u_*}{u_s(0)\kappa}\ln\left(\left|\frac{H_{ML}}{z_1}\right|\right) + \cos\left(\theta_{ww}\right)}\right)$$

# A theory for LC direction!

L. Van Roekel, B. Fox-Kemper, P. P. Sullivan, P. E. Hamlington, and S. R. Haney. The form and orientation of Langmuir cells for misaligned winds and waves. *Journal of Geophysical Research-Oceans*, 2012. Submitted.

Wind and wave forced, dying submeso filament  $Ro \approx 0.1$  $\overline{Ri} < 1$  $La_t \approx 0.3$ 

Computational parameters: Domain size: 20km x 20km x -160m Grid points: 4096 x 4096 x 128 Resolution: 5m x 5m x -1.25m

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#### Wind Only

Coupling between Langmuir and Submeso?

2 runs: Both spindown of submesoscale filament

Right --> Stokes & Wind

> <-- Left Wind Only

Stokes & Wind



1.08e-02

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Coupling between Langmuir and Submeso?

2 runs: Both spindown of submesoscale filament

Right --> Stokes & Wind

> <-- Left Wind Only





2 runs:

filament

Right -->

<-- Left

Wind Only

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#### Heat <wT>. Upper=Total, Lower=small-scales only



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#### Momentum: <uw>. Upper=Total, Lower=small-scales





Wind-wave dependent processes in the coupled climate system Towards coupled wind-wave-AOGCM models

L. Cavaleri, B. Fox-Kemper, and M. Hemer. Wind waves in the coupled climate system. *Bulletin of the American Meteorological Society*, 2012. Accepted.

# Results

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# All papers at: fox-kemper.com/research

B. Fox-Kemper, G. Danabasoglu, R. Ferrari, S. M. Griffies, R. W. Hallberg, M. M. Holland, M. E. Maltrud, S. Peacock, and B. L. Samuels. Parameterization of mixed layer eddies. III: Implementation and impact in global ocean climate simulations. *Ocean Modelling*, 39:61-78, 2011.

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L. Cavaleri, B. Fox-Kemper, and M. Hemer. Wind waves in the coupled climate system. *Bulletin of the American Meteorological Society*, 2012. Accepted. If you want to know about the symmetric tensor, see Scott Bachman's Poster!

 $\mathbf{u}'b' = \mathbf{R}\nabla\bar{b}$ 

 $\mathbf{R} = \mathbf{S} + \mathbf{A} = \begin{bmatrix} \kappa & \kappa S - \psi \\ \kappa S + \psi & \kappa S^2 \end{bmatrix}$ 

Mixed Layer Eddy Restratification Estimating eddy buoyancy/density fluxes:  $\overline{\mathbf{u}'b'} \equiv \Psi \times \nabla \overline{b}$ A submeso eddy-induced overturning:

 $\Psi = \frac{C_e H^2 \mu(z)}{|f|} \nabla \bar{b} \times \mathbf{\hat{z}}$ 





Mixed Layer Eddy Restratification Estimating eddy buoyancy/density fluxes:  $\overline{\mathbf{u}'b'} \equiv \mathbf{\Psi} \times \nabla \overline{b}$ 

A submeso eddy-induced overturning:  $\Psi = \frac{C_e H^2 \mu(z)}{|f|} \nabla \overline{b} \times \mathbf{\hat{z}}$ Ę

in ML only:  $\mu(z)=0 ext{ if } z < -H$ 



ł

channel

100

200

300

along-channel (km)

400

500

600

Temperature on day:0

Mixed Layer Eddy Restratification Estimating eddy buoyancy/density fluxes:  $\overline{{f u}'b'}\equiv {f \Psi} imes 
abla \overline{b}$ 

A submeso eddy-induced overturning:  $\Psi = \frac{C_e H^2 \mu(z)}{|f|} \nabla \overline{b} \times \mathbf{\hat{z}}$ 

in ML only:  $\mu(z) = 0 ext{ if } z < -H$ 

For a consistently restratifying,  $\overline{w'b'} \propto \frac{H^2}{|f|} \left| \nabla_H \overline{b} \right|^2$ 





20 40 40 Mixed Layer Eddy Restratification Estimating eddy buoyancy/density fluxes:  $\overline{{\bf u}'b'}\equiv {\bf \Psi}\times \nabla \overline{b}$ 

A submeso eddy-induced overturning:  $\Psi = \frac{C_e H^2 \mu(z)}{|f|} \nabla \overline{b} \times \hat{z}$ 

in ML only:  $\mu(z)=0 ext{ if } z < -H$ 

For a consistently restratifying,  $\overline{w'b'} \propto \frac{H^2}{|f|} \left| \nabla_H \overline{b} \right|^2$ and horizontally downgradient flux.  $\overline{\mathbf{u'}_H b'} \propto \frac{-H^2 \frac{\partial \overline{b}}{\partial z}}{|f|} \nabla_H \overline{b}$ 





Temperature on day:0 Mixed Layer Eddy Restratification Estimating eddy buoyancy/density fluxes: Ē Ň  $\overline{\mathbf{u}'b'} \equiv \mathbf{\Psi} \times \nabla b$ A submeso eddy-induced overturning:  $\Psi = \frac{C_e H^2 \mu(z)}{|f|} \nabla \overline{b} \times \hat{\mathbf{z}}$ y (km) x (km) Surface Temp. Day: 900 in ML only: Overturning Streamfunction  $\mu(z) = 0$  if z < -HMixed Layer For a consistently restratifyi z (m) Eddy  $\overline{w'b'} \propto rac{H^2}{|f|} \left| 
abla_H \overline{b} \right|^2$ Buoy. Flux ML Base and horizontally downgradient Pycno- $\overline{\mathbf{u'}_H b'} \propto rac{-H^2 rac{\partial b}{\partial z}}{\|f\|} 
abla_H$ cline y (km)



grid cell mean ice thickness MIN - -11.34 MAX - 14.44 MIN = -0.56 MAX = 3.76

ice area (aggregate)

MIN = -36.88 MAX = 41.64

Figure 10: Wintertime sea ice sensitivity to introduction of MLE parameterization (CCSM<sup>+</sup> minus CCSM<sup>-</sup>): January to March Northern Hemisphere a) ice area and b) thickness and July to September Southern Hemisphere c) ice area and d) thickness.

#### Affects sea ice

MN - -0.61 MAX - 0.66

0.12

-0.12 -0.16

NO RETUNING NEEDED!!!



May Stabilize AMOC Maximum AMOC at 45n in coupled MOM 28 26 ർ CM2.1 (mean=24.5, std=1.9) CM2Ma+(mean=23.9, std=1.6) CM2Ma-(mean=21.5, std=2.9) 50 100 150 200 250300

Year

# Submeso Param. in Data?

Submesoscale activity over the Argentinian shelf X Capet, EJ Campos, AM Paiva - Geophysical Research Letters, 2008

Impact of atmospheric coastal jet off central Chile on sea surface temperature from satellite observations (2000–2007), L Renault, B Dewitte, M Falvey, R Garreaud... - J. Geophys. Res, 2009

Interactions between a Submesoscale Anticyclonic Vortex and a Front\* C Chavanne, P Flament... - Journal of Physical ..., 2010

Mixing rates across the Gulf Stream, Part 2: Implications for nonlocal parameterization of vertical fluxes in the surface boundary layers R Inoue, RR Harcourt... - Journal of Marine ..., 2010

Spatial variability and temporal dynamics of surface water pCO2,[Delta] O2/Ar and dimethylsulfide in the Ross Sea, Antarctica PD Tortell, C Guéguen, MC Long, CD Payne... - Deep Sea Research ..., 2011

# Multiscale

- Langmuir &
   Submeso resolving
   LES
- 20kmx20kmx0.3km
- Grid4096x4096x128
- 5x5x<1m resolution</p>
- Compromises- wind, front, wave,
   size, etc

y

 Leif tells us how the winds wil







# The Scales, and the Sim

Day 6.5 of a Submeso Resolving run Vert. Velocity=w

Day 6.5 of a Submeso Resolving run Near Surf. Temp



# The Scales, and the Sim $f^2 < \left| f \frac{\partial \overline{v}}{\partial z} \right| = M^2 < (3f)^2$

 $Ro \approx 0.1$ 

Ri < 1

Wind & Fronts Only

#### No Stokes Drift



-0.015

#### Surf. Temp (K)

+0.015

Temp (K)

Wind & Fronts Only

#### No Stokes Drift



#### Surf. Temp (K)

 $w^2$  (m<sup>2</sup>/s<sup>2</sup>) <  $(400m/d)^2$ 

Wind & Fronts Only

#### No Stokes Drift



Surf.